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ANALYTICAL STUDY OF TIME-DEPENDENT
DEFORMATION IN PERMAFROST

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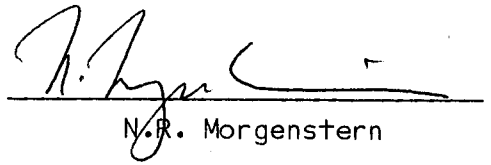
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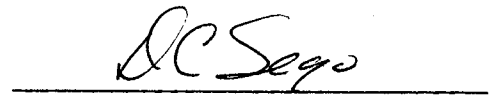
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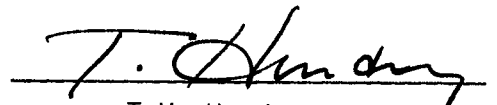
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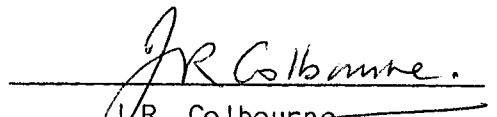
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The undersigned certify that they have read, and recommend to the Faculty of Graduate Studies and Research, for acceptance, a thesis entitled 'ANALYTICAL STUDY OF TIME-DEPENDENT DEFORMATION IN PERMAFROST' submitted by Paul C. Weerdenburg, in partial fulfillment of the degree of Master of Science in Civil Engineering.


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ABSTRACT

This thesis describes an analytical study of in-situ creep behavior in ice-rich permafrost. An incremental initial strain procedure was formulated to solve steady state creep problems using the finite element method. Two case histories were analyzed where naturally occurring creep had been monitored for several years.

The numerical analyses of the left bank of the Great Bear River at the proposed Arctic Gas crossing has shown that the steady state creep in the ice-rich glaciolacustrine clay at the site can be modelled by a simple power law with an exponent of 3.0 and a coefficient of $3.33 \times 10^{-9} \text{ kPa}^{-3} \text{ yr}^{-1}$. This strain rate is six times slower than the value for polycrystalline ice at an equivalent temperature. The exact form of the constitutive relationship for the glaciodeltaic sand overlying the clay remains unclear.

A review of the available in-situ deformation studies carried out at the Fox Tunnel near Fairbanks, Alaska, showed that attenuating creep behavior extended well over one year and contributed significantly to the total room closure. In this case, the flow law for polycrystalline ice did not yield an upper bound solution to the observed room closure measurements. However, it is felt that the difference lies in time-dependent failure caused by stress release which was not accounted for in the numerical model.

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CHAPTER I

INTRODUCTION

1.1 General

Depletion of conventional oil and gas reserves has led to an increase in resource exploration activity in the northern frontier lands of Canada and the United States during the past decade and into the foreseeable future. Rising world energy prices has enhanced the economic feasibility of exploiting these reserves and transporting them to market. The largest such project completed to date is the Alyeska Pipeline which transports Alaskan North Slope oil to a shipping terminal at Valdez. The first major northern oil pipeline to be completed in Canada will be the 900 km long pipeline forecast to be in place by 1985 to transport oil south from the expanded oil field in Norman Wells, N.W.T. to Zama, Alberta where it will connect with an existing pipeline.

Resource exploration and development will require major civil works to be constructed on or within permafrost. These structures will present unique problems to engineers as the strength and deformation behavior in frozen soil will respond in both the time and thermal domain. The presence of ice as either discrete segregated structures or in the pore spaces, imparts a major influence on the creep behavior that characterizes frozen soil.

The in-situ creep behavior of frozen soil will have a major influence on establishing geotechnical design guidelines that ensure serviceable foundation performance for structures founded in permafrost. Ladanyi (1972) has developed an engineering theory for creep in frozen soils. The finite

element technique is now routinely used for solving a large class of engineering problems. The aim of the research reported herein was to use the finite element technique to predict the in-situ creep deformation patterns in ice-rich permafrost foundation soils.

To date, these are only two documented case histories comparing long term in-situ and laboratory creep behavior of frozen soil. An analytical study was carried out by Thompson and Sayles (1972) in conjunction with a field program of the U.S. Army Cold Regions Research and Engineering Laboratory (USACREL) in the Fox Tunnel north of Fairbanks, Alaska. In-situ measurement of creep deformations were carried out for three and one half years. Savigny (1980) reported the results of an extensive field and laboratory program on the proposed Arctic Gas crossing of the Great Bear River near Fort Norman, N.W.T. Naturally occurring creep movements in the left bank of the Great Bear River were recorded for a two year period.

1.2 Scope of Thesis

Chapter II presents a review of the rheological behavior of frozen soil. Since the deformation behavior of frozen soil is governed by the amount of ice present, a detailed review of the deformation behavior of polycrystalline ice is presented. This is followed by a review of the deformation behavior for reconstituted frozen soil and natural permafrost. The existing constitutive relationships for creep in natural permafrost samples are examined. An upper bound flow law for the creep in ice-rich fine-grained permafrost is presented at the conclusion of this chapter.

Chapter III presents the theoretical basis for the incremental initial strain procedure for solving for the stress state in a non-linear viscous medium. The treatment is restricted to steady state creep problems since the empirically developed constitutive relationship for frozen soil is given by a simple power law relating strain rate to stress. The development of the plane strain finite element programme CREEP is also presented.

Finite element analyses of the in-situ creep behavior of the left bank of the Great Bear River at the proposed Arctic Gas crossing is presented in Chapter IV. These analyses were carried out in an attempt to verify the simple power law as the constitutive equation governing the behavior of frozen ice-rich, fine-grained permafrost soils.

The in-situ deformation behavior at the Fox Tunnel at Fairbanks, Alaska is examined in Chapter V. A synthesis of all the in-situ deformation studies reported in the literature for the Fox Tunnel was carried out in an attempt to gain some insight into the deformation processes surrounding underground openings in frozen ground.

The final chapter presents a summary of the results of this research as well as the limitations of the current analytical approach. Recommended suggestions for future analytical and laboratory creep studies in frozen soils are also presented.

The formulation of the finite element equations for solving creep problems with the incremental initial strain procedure is presented in Appendices A and B. A user's manual and programme listing for the finite element programme CREEP are given in Appendices C and D, respectively. The accuracy of the finite element programme is compared with a thick wall cylinder closed form solution in Appendix E.

CHAPTER II

RHEOLOGICAL BEHAVIOR OF FROZEN SOILS

2.1 General

An understanding of the deformation behavior of frozen ground over a wide range of stress and temperature conditions is required before a numerical analysis can be carried out. Deformation behavior of frozen ground is generally governed by the amount of ice present and its temperature.

Laboratory studies over the past two decades have focussed attention on deriving empirical constitutive relationships describing the stress-strain-time behavior of frozen soils. Many of the early studies were restricted to short duration creep tests on ice and remoulded frozen soils at stresses and temperatures well beyond the range of interest of practical geotechnical problems. As the major factors influencing the creep behavior of frozen soil have been delineated, specifically the soil temperature and ice structure, the need for high quality, long term, low stress creep tests on natural permafrost samples has arisen.

The following sections will present a brief review of the time-dependent load-deformation behavior of ice, remoulded frozen soil and natural permafrost soils as it relates to low stress creep behavior. Unless otherwise mentioned, stress and temperature conditions commonly encountered in geotechnical engineering are assumed.

2.2 Composition of Frozen Soil

Frozen soil can be considered as a complex, four phase, natural formation, consisting of solid mineral particles, ice, unfrozen water and gases. The solid mineral particles present in a frozen soil have an important influence on its geotechnical properties. The primary factors are the grain size and shape as well as the physicochemical nature of the mineral surfaces, which are determined by the mineralogical composition and the cations present. The gaseous phase consists of water vapour and has an insignificant influence on the behavior of frozen soils.

Unfrozen water is present in two states as strongly bound and loosely bound water. The strongly bound water is adjacent to the soil particle. The very high electro-molecular forces present in this layer of water suppresses the formation of ice crystals, even at very low temperatures. The intermolecular forces are also present in the loosely bound water surrounding the strongly bound water. However, this layer is capable of releasing the heat of crystallization at temperatures below 0°C.

The amount of unfrozen water present in a frozen soil depends on its temperature, specific surface and type of soil mineral present and chemistry of the pore water. Each frozen soil is characterized by a specific curve relating its unfrozen water content and temperature.

Of all the phases present in frozen soil, ice is the most important component in the rate, time and temperature dependent properties which characterize frozen soil. Ice can be present both as discrete segregated structures and in the pore spaces of the soil mass. Ice present in frozen soils is generally polycrystalline.

The reader is referred to Anderson and Morgenstern (1973) for a complete summary of conditions affecting the state of frozen ground.

2.3 Creep of Ice

The presence of ice dominates the time-dependent characteristics of frozen soil. Ice 1H is the predominant ice type found in frozen soil. The structure of ice 1H is hexagonal. Each crystal consists of layers of hexagonal rings, referred to as the basal planes. The perpendicular direction to the basal plane is the optic or 'C' axis.

Single crystals of ice deform readily at low stresses along discrete bands parallel to the basal plane of the crystal structure. This type of deformation is referred to as basal glide (easy glide). Single ice crystals will also deform in the perpendicular or non-basal direction (hard glide). However, for a given strain rate, the applied stress can be as much as 20 times greater for hard glide than for easy glide.

Polycrystalline ice is composed of randomly oriented single ice crystals. Physical processes observed during the deformation of polycrystalline ice are dislocation climb, grain boundary slip, cavity formation at the grain boundaries, recrystallization and microcracks developing into ice grains (Michel, 1978).

The deformation response of polycrystalline ice subjected to a load consists of four distinct regions as illustrated by curve II in Figure 2.1. Each of these regions consist of:

- 1) instantaneous elastic strain (OA)
- 2) creep strain in a primary mode at a decreasing strain (AB)

- 3) secondary creep at a constant minimum strain rate (BC)
- 4) tertiary creep at an accelerating strain rate leading to failure (CD)

The shape of the creep curve is a function of the stress level and ice temperature. For a given temperature, secondary creep is suppressed at high stress levels (undamped creep) while primary creep (damped creep) dominates at low stress levels.

Creep tests on polycrystalline ice in compression, tension, shear and other specialized loading configurations have been carried out to evaluate the form of the constitutive equation. Glen (1975) has recognized two possible ways of identifying a flow law. The first flow law would relate stress to strain rate at a very long time under the application of load allowing recrystallization to occur. The second form of the flow law relates the secondary strain rate to stress before any recrystallization of the ice. At higher stresses, the first flow law would produce deformation rates greater than the minimum (or secondary) rate because recrystallization may produce a grain size and orientation more favorable for plastic flow.

The constitutive equation of polycrystalline ice is most commonly represented empirically by the simple power law. Experimental evidence has shown that the power creep law represents the steady state creep data in the low to intermediate stress range. From a practical viewpoint, this mathematical representation of creep behavior is simpler to use than a physical theory (i.e. rate process theory) because the material parameters are kept to a minimum.

The simple power law relating strain rate to stress is given by:

$$\dot{\epsilon} = A\sigma^n \quad (2.1)$$

where $\dot{\epsilon}$ = axial strain rate (time)⁻¹
 σ = axial stress
 A = coefficient (function of temperature and ice type)
 n = creep exponent

In order to extend uniaxial creep data to multi-axial stress states, it is common practice to adopt the concept of effective stress, σ_e , and effective creep strain rate, $\dot{\epsilon}_e$. Several definitions of effective stress and strain rate have been put forward in the metal and ice creep literature. Some of these are listed in Table 2.1.

The definition of effective stress is defined by the square root of the second invariant of the deviatoric stress tensor multiplied by a constant. Likewise, the effective strain rate is defined in a similar manner except that the second invariant of the strain rate tensor is used. The difference between the alternate definitions lies in the value of the constant. Since the stress is raised to an exponent in the simple power law, the magnitude of the constant will influence the effective strain rate in a non-linear manner as shown in Figure 2.3. As shown in the figure, the effective strain rate can change by a complete order of magnitude depending on which definition of effective stress is chosen. The steady state creep law for ice at -2°C (Morgenstern et al., 1980) is plotted in Figure 2.3.

TABLE 2.1 DEFINITIONS OF EFFECTIVE STRESS AND
EFFECTIVE STRAIN RATE

Source	Effective Stress	Effective Strain Rate
Odqvist (1966) Dorn et al (1945) Ladanyi (1972) Finnie & Heller (1959)	$\sqrt{3} \sqrt{J_2^\sigma}$	$\frac{2}{\sqrt{3}} \sqrt{J_2^{\dot{\epsilon}}}$
Nye (1957)	$\sqrt{J_2^\sigma}$	$\sqrt{J_2^{\dot{\epsilon}}}$
Meier (1959)	$\sqrt{\frac{2}{3}} \sqrt{J_2^\sigma}$	$\sqrt{\frac{2}{3}} \sqrt{J_2^{\dot{\epsilon}}}$
Vialov (1962)	$\sqrt{J_2^\sigma}$	$\sqrt{2} \sqrt{J_2^{\dot{\epsilon}}}$

where J_2^σ = second invariant of deviatoric stress tensor

$J_2^{\dot{\epsilon}}$ = second invariant of strain rate tensor

In this work, the effective stress and strain rate as defined by Odqvist (1966) has been used, i.e.:

$$\sigma_e = \sqrt{3} \sqrt{J_2^\sigma} = \sqrt{\frac{3}{2} s_{ij} s_{ij}} \quad (2.2)$$

$$\dot{\epsilon}_e = \sqrt{\frac{2}{3}} \sqrt{J_2^{\dot{\epsilon}}} = \sqrt{\frac{2}{3} \dot{\epsilon}_{ij} \dot{\epsilon}_{ij}} \quad (2.3)$$

These definitions are adopted because they recover Norton's law, i.e. equation 2.1, for the special case of uniaxial stress. The flow law now becomes:

$$\dot{\epsilon}_e = A \sigma_e^n \quad (2.4)$$

Since the effective stress is a function of the deviatoric stress only, the strain rate is independent of the hydrostatic state of stress.

An extensive review of the secondary creep data of polycrystalline ice was carried out by Morgenstern et al. (1980). Based on this review, the authors presented the values shown in Table 2.2 for the parameters A and n in the power law.

A more detailed review of the rheological characteristics of polycrystalline ice has been prepared by Glen (1975), Roggensack (1977) and Sego (1980).

TABLE 2.2 CREEP PARAMETERS FOR POLYCRYSTALLINE ICE

Temperature (°C)	$A(\text{kPa}^{-3}\text{yr}^{-1})$	n
-1	4.5×10^{-8}	3.0
-2	2.0×10^{-8}	3.0
-5	1.0×10^{-8}	3.0
-10	5.6×10^{-9}	3.0

2.4 Creep of Frozen Soil

Vialov (1963) described the physical processes involved in the time dependent plastic deformation of frozen soils. The application of a load causes local stress concentrations at soil-ice contacts leading to pressure melting of the ice. A pressure gradient resulting from differences in surface tensions causes the water to migrate to regions of lower stress where it then refreezes. The pressure melting process is accompanied by a breakdown of the ice and structural bonds of the soil grains and plastic deformation of the ice, causing a rearrangement of the mineral particles. The end result is the process known as creep.

Many laboratory studies have been carried out in order to determine the empirical creep parameters for reconstituted frozen soils and natural permafrost soils. The creep tests carried out on remoulded frozen soils are helpful in determining the various factors that influence the creep rate of frozen soils. However, many of these tests were carried out beyond the temperature and stress range of practical interest in geotechnical engineering.

A second difficulty encountered when applying the creep parameters of reconstituted frozen soils to natural permafrost soils is reproducibility of the ice facies in the remoulded samples. In natural permafrost samples, the ice is present in the pore spaces as well as in the form of reticulate and segregated ice. Laboratory evidence of Savigny (1980) shows deformations localized along segregated ice in natural permafrost soils. Since the reticulate ice structure can not be reproduced in reconstituted frozen samples, natural permafrost samples must be tested.

Vialov (1959) describes the stress-strain response of a frozen soil as damped or undamped creep. Damped deformations occur when the applied stress is less than the long term strength of the frozen soil. Damped creep behavior is characterized by an instantaneous elastic displacement followed by primary creep where the deformation rate continuously decreases with time. Curve III in Figure 2.1 is typical of damped creep behavior. Undamped deformations occur when the applied stress exceeds the long term strength. In undamped creep, the transient stage gives way to a steady state period where the strain rate reaches a minimum value. The deformations continue to grow until they reach a limiting value or the onset of tertiary creep where the deformation rate will accelerate to failure. A typical undamped deformation response is given by curves I and II in Figure 2.1.

Damped creep deformations are prevalent in dense, ice-poor frozen soils. Since ice has a zero long term strength, its deformation behavior is characterized by undamped creep. The deformation response of ice-rich frozen soils will lie somewhere between these two limiting cases.

Weaver (1979) proposed a classification system for frozen soils based on frozen bulk density that relates to low stress creep behavior. This classification system is shown in Table 2.3.

A generalized theory of creep of frozen soil put forward by Ladanyi (1972) advocates a mathematical rather than a physical treatment of undamped creep behavior. The proposed uniaxial power creep law describing steady state creep is defined as:

$$\dot{\epsilon}_e = \dot{\epsilon}_c (\sigma_e / \sigma_c)^n \quad (2.5)$$

where: σ_c = temperature dependent creep coefficient
 $\dot{\epsilon}_c$ = secondary strain rate corresponding
to the application of the stress, σ_c
 n = creep exponent

This form of the power law is similar to that given by equation 2.4 except that it is given in a normalized form. Making use of the definitions of effective stress (equation 2.2), the strain rate for an axially symmetric state of stress (i.e. $\sigma_2 = \sigma_3$) is given by:

$$\dot{\epsilon}_e = \dot{\epsilon}_c [(\sigma_1 - \sigma_3) / \sigma_c]^n \quad (2.6)$$

where: $\sigma_1 - \sigma_3$ = principal stress difference

TABLE 2.3 FROZEN SOIL CLASSIFICATION SYSTEM
(Proposed by Weaver, 1979)

SOIL TYPE	DESCRIPTION
Dirty Ice	<ul style="list-style-type: none"> - applies to ice that has a low solids concentration - $\gamma = 0.9 - 1.0 \text{ Mg/m}^3$ - the soil particles present reduce the average grain size of the ice crystals resulting in higher creep rates than pure ice
Very Dirty Ice	<ul style="list-style-type: none"> - applies to ice that has medium to high solids concentration - $\gamma = 0.9 - 1.0$ and $1.6 - 1.8 \text{ Mg/m}^3$ - very little grain to grain contact between soil particles - lower secondary creep rates than polycrystalline ice because soil impedes dislocation movement
Ice-Poor Frozen Soil	<ul style="list-style-type: none"> - applies to saturated frozen soil whose deformation patterns are characterized by primary creep - $\gamma = 1.7 - 1.8$ and $1.9 - 2.0 \text{ Mg/m}^3$
Ice-Rich Frozen Soil	<ul style="list-style-type: none"> - applies to soils that have a continuous network of segregated ice - the overall creep response is complex and is very sensitive to the reticulate structure of the segregated ice, bulk density grain size and ground temperature

The strain rate derived from equation 2.6 is independent of the hydrostatic pressure. The hydrostatic stress can be expected to influence the stress-strain rate behavior up to failure of unconsolidated frictional earth materials. Ladanyi (1972) extended the stress-strain rate constitutive law to account for the mean normal stress using a two or three parameter failure theory.

Vialov (1962) proposed that the strength of a frozen soil can be represented by a series of Mohr failure envelopes, each one corresponding to a specific time to failure. Assuming validity of the Mohr-Coulomb failure criterion in the pre-failure state, the dependence of strain rate on mean normal pressure for an axially symmetric stress state is expressed as:

$$\dot{\epsilon}_e = \dot{\epsilon}_c \left[\frac{(f+2)(\sigma_1 - \sigma_3) - 3(f-1)\sigma_m}{3\sigma_c} \right]^n \quad (2.7)$$

where: $f = (1 + \sin\phi)/(1 - \sin\phi)$
 $\sigma_m = (\sigma_1 + \sigma_2 + \sigma_3)/3$
 ϕ = angle of internal friction

For a frictionless soil, $f = 1$, equation 2.7 simplifies to equation 2.6. Equation 2.7 assumes full mobilization of internal friction over the pre-failure state. This assumption leads to a non-zero strain rate at a zero stress difference. Thus, application of equation 2.7 should be restricted to strains close to failure. This limitation may also be overcome by assuming a time dependent angle of internal friction.

Ladanyi (1972) used the extended von Mises failure criterion to account for effect of mean normal stress on strain rate for frozen soils with low internal friction. For an axially symmetric state of stress, the strain

rate can be expressed in terms of mean normal stress and principal stress difference as:

$$\dot{\epsilon}_e = \dot{\epsilon}_c \left[\frac{(r+1)(\sigma_1 - \sigma_3) - 3(r-1)\sigma_m}{2\sigma_c} \right]^n \quad (2.8)$$

where r = ratio of uniaxial compressive
strength to uniaxial tensile strength

Equation 2.8 is subject to the same limitations as equation 2.7 in that a zero stress difference gives a non-zero strain rate assuming a constant strength ratio, r . The limitation of equations 2.7 and 2.8 are overcome by simply assuming the strain rate in a pre-failure state is independent of hydrostatic stress and only the creep strength is made a function of the mean normal stress. This is consistent with the experimental findings of Sayles (1973) at higher strain rates and confining pressures.

From a practical viewpoint, the engineering theory is more favourable than a theory describing creep in terms of more fundamental laws of physics because the required number of experimental material parameters are kept to a minimum.

Tests on ice-rich remoulded frozen soils have been carried by Vialov (1962), Perkins and Ruedrich (1973) and Sayles and Haines (1974) and others. Although the existence of a true steady state in the shorter duration tests is questionable. The longer duration tests generally show creep exponent values approaching a value of 3.0.

A comprehensive study of creep of frozen sands carried out by Sayles (1968) showed that strain rates continued to attenuate with time for tests

carried out at stress levels below the long term strength. Damped creep behavior was also observed by Sayles and Haines (1974) in unconfined compression creep on fine-grained remoulded frozen samples when the strain did not exceed 20% over the test duration.

The influence of confining pressure on the creep rate of remoulded frozen Ottawa sand was studied by Sayles (1973). The findings of this study reported that the amount of creep strain can be reduced by increasing the confining pressure. Primary creep was still evident at strains exceeding 20%. A separate study carried out by Alkire and Andersland (1973) showed an exponential decrease in creep strain rate with an increase in confining pressure. Application of confining pressure enhances the frictional behavior of the sand.

Many recent laboratory studies have focussed attention on measuring the empirical parameters in the constitutive equation for natural permafrost soils. Thompson and Sayles (1972) studied the in-situ horizontal and vertical closure of the Fox Tunnel in Alaska. The roof of the tunnel was overlain with 17 m of frozen ice-rich Fairbanks silt. Samples of the silt were tested in unconfined compression at stress levels between 250 and 2000 kPa. The test temperature was -1.7°C . They concluded the constitutive equation for the frozen silt has the same form as equation 2.4 with an exponent of 4. Numerical analysis of the room predicted a closure rate more than 3 times faster in the field than that measured in the laboratory.

Roggensack (1977) tested ice-rich fine-grained glaciolacustrine clay from the Mackenzie River Valley. Triaxial constant stress creep tests were carried out at deviatoric stress levels between 20 and 400 kPa and temperatures close to -1°C . This data is shown in Figure 2.3. Although the data showed considerable scatter, especially at stresses below 100 kPa, a

bilinear flow law was fit to the data. A power law fit to the data above 100 kPa gives an exponent of 2.75 which is slightly faster, but consistent with the value for polycrystalline ice. In the low stress range, transient processes play a dominant role in the overall deformation response of the sample. Increasing the confining pressure did not have as a significant effect on decreasing the creep rate as was reported for coarse-grained soils (e.g. Sayles, 1973; Alkine and Andersland, 1973).

McRoberts et al. (1978) tested undisturbed ice-rich silt from Norman Wells, N.W.T. Constant stress creep tests were carried out in a triaxial cell at deviatoric stress levels between 10 and 690 kPa. Test temperatures ranged from -1 to -3°C . Their results are shown in Figure 2.4, adjusted to a reference temperature of -1°C . An upper bound to their data is given in terms of a bilinear flow law:

$$\dot{\epsilon}_a = \frac{1.6 \times 10^{-7}}{(1-T)^{1.8}} \sigma_d^3 + \frac{1.5 \times 10^{-14}}{(1-T)^{1.8}} \sigma_d^6 \quad (2.9)$$

where $\dot{\epsilon}_a$ = axial strain rate (yr^{-1})

σ_d = deviatoric stress (kPa)

T = temperature ($^{\circ}\text{C}$)

All but one of the samples tested at a deviatoric stress exceeding 400 kPa terminated in failure.

In the low stress range the value of the creep exponent is 3.0, consistent with polycrystalline ice. Tertiary creep may be influencing the strain rates in the high stress range.

Savigny (1980) tested fine-grained, ice-rich glaciolacustrine clay samples from the proposed Arctic Gas crossing of the Great Bear River, N.W.T. Constant stress triaxial creep tests were carried out at stress and temperature conditions simulating field behavior. The laboratory results reported showed no correlation with the flow law for polycrystalline ice (Morgenstern et al., 1980). Savigny (1980) attributed this scatter to the ground ice structure and the generation of a non-uniform stress fields within the specimens in response to application of confining pressure.

Nixon (1978) reviewed the data of Roggensack (1977) and McRoberts et al. (1978). He proposed a tentative upper bound to the creep rates of McRoberts et al. (1978) at a temperature of -2.5°C and stress less than 100 kPa:

$$\dot{\epsilon}_z = 2.135 \times 10^{-8} \sigma^3 \quad (2.10)$$

where the units are in $(\text{years})^{-1}$ and kPa.

2.5 Conclusions

The combined results of Roggensack (1977), McRoberts et al. (1978) and Savigny (1980) are shown in Figure 2.5. The two flow laws shown in this figure are:

- 1) flow laws for polycrystalline ice (Morgenstern et al., 1980)
- 2) tentative upper bound to soil data of fine-grained ice-rich soil proposed by Nixon (1978) extended above 100 kPa.

With a few exceptions, the laboratory data fall within a narrow band at stress levels exceeding 75 kPa. Below this stress, transient processes may still dominate.

Based on the data presented in Figure 2.5, it can be concluded that the creep law of polycrystalline ice represents a reasonable upper bound to the constitutive behavior of fine-grained, ice-rich permafrost. Exceptions to this occur in the high stress range where the creep tests terminated in failure.

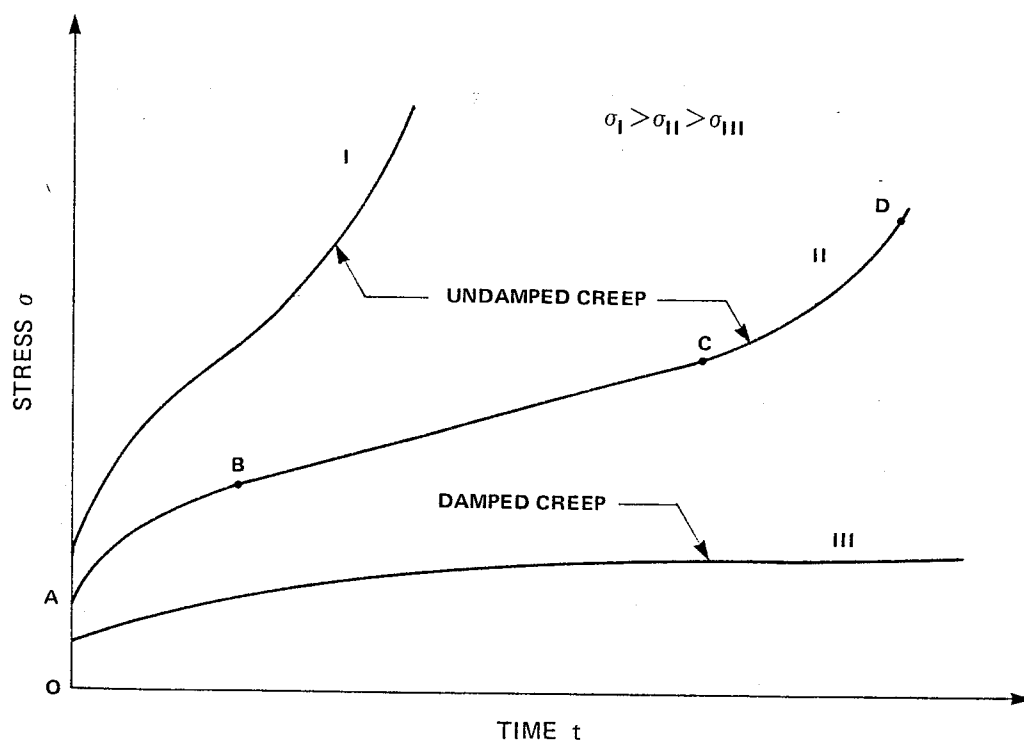


FIGURE 2.1

TYPICAL CONSTANT STRESS
CREEP CURVES

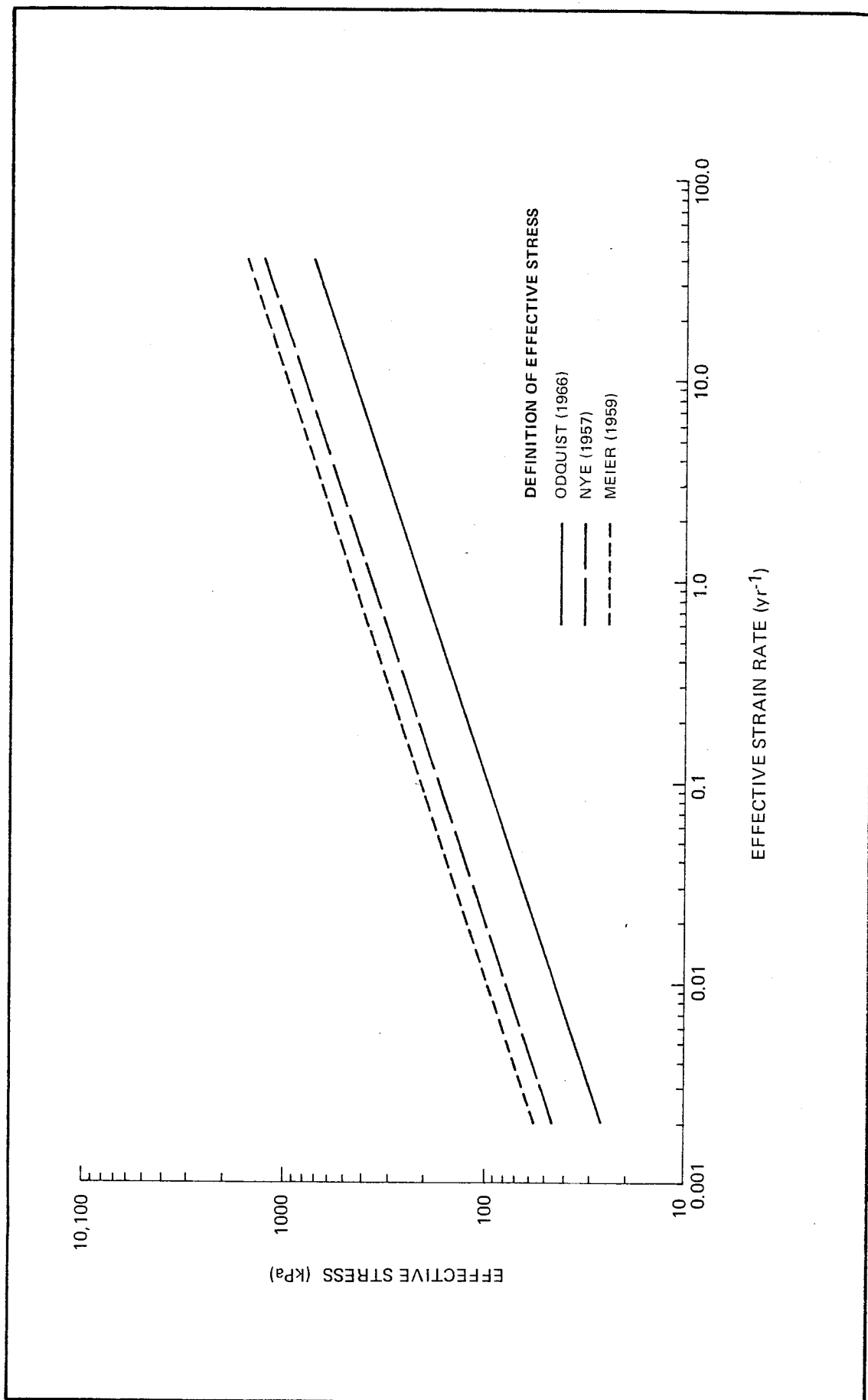


FIGURE 2.2 COMPARISON OF DEFINITION OF EFFECTIVE STRESS

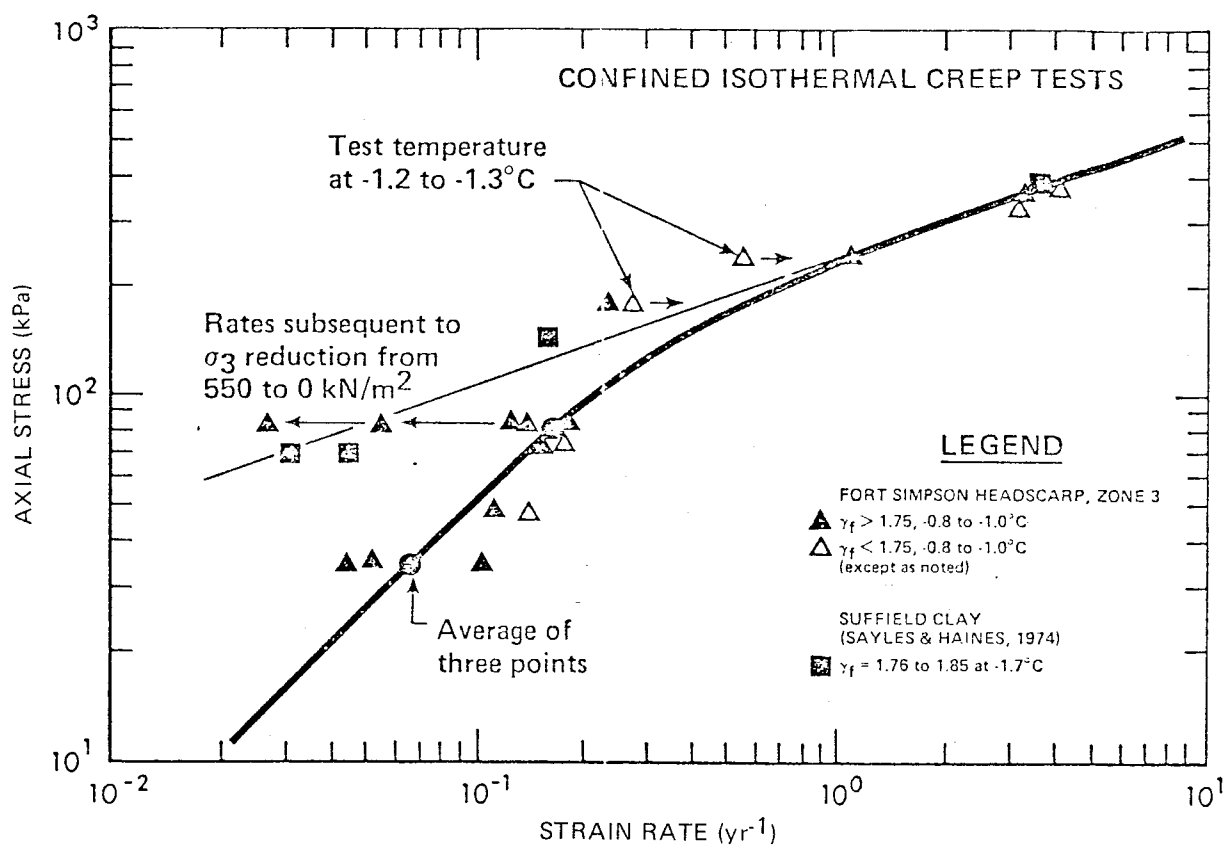


FIGURE 2.3 AXIAL STRESS VERSUS STRAIN RATE
(Roggensack, 1977)

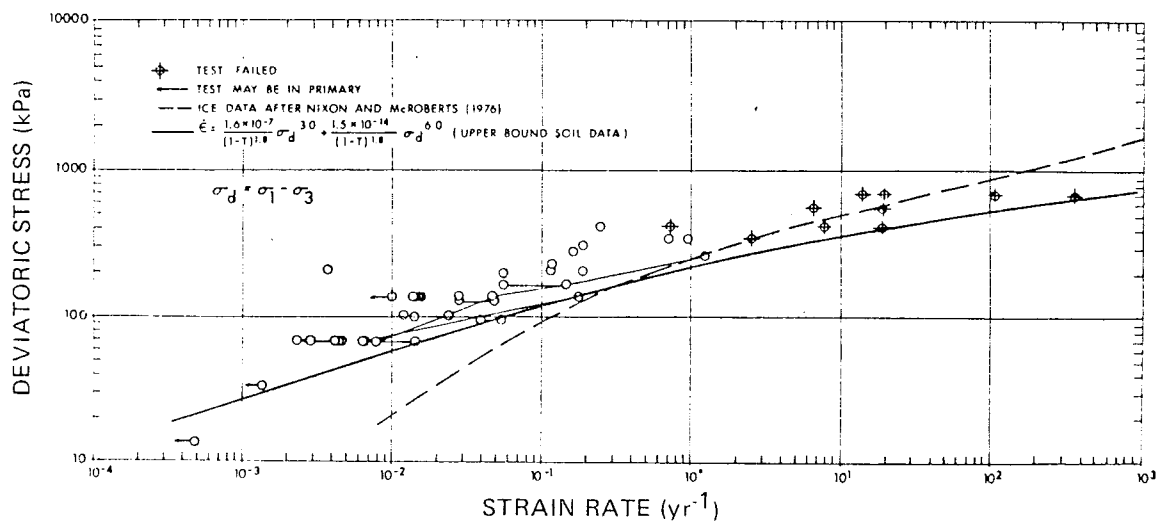


FIGURE 2.4 DEVIATORIC STRESS VERSUS STRAIN RATE
(McRoberts et al., 1978)

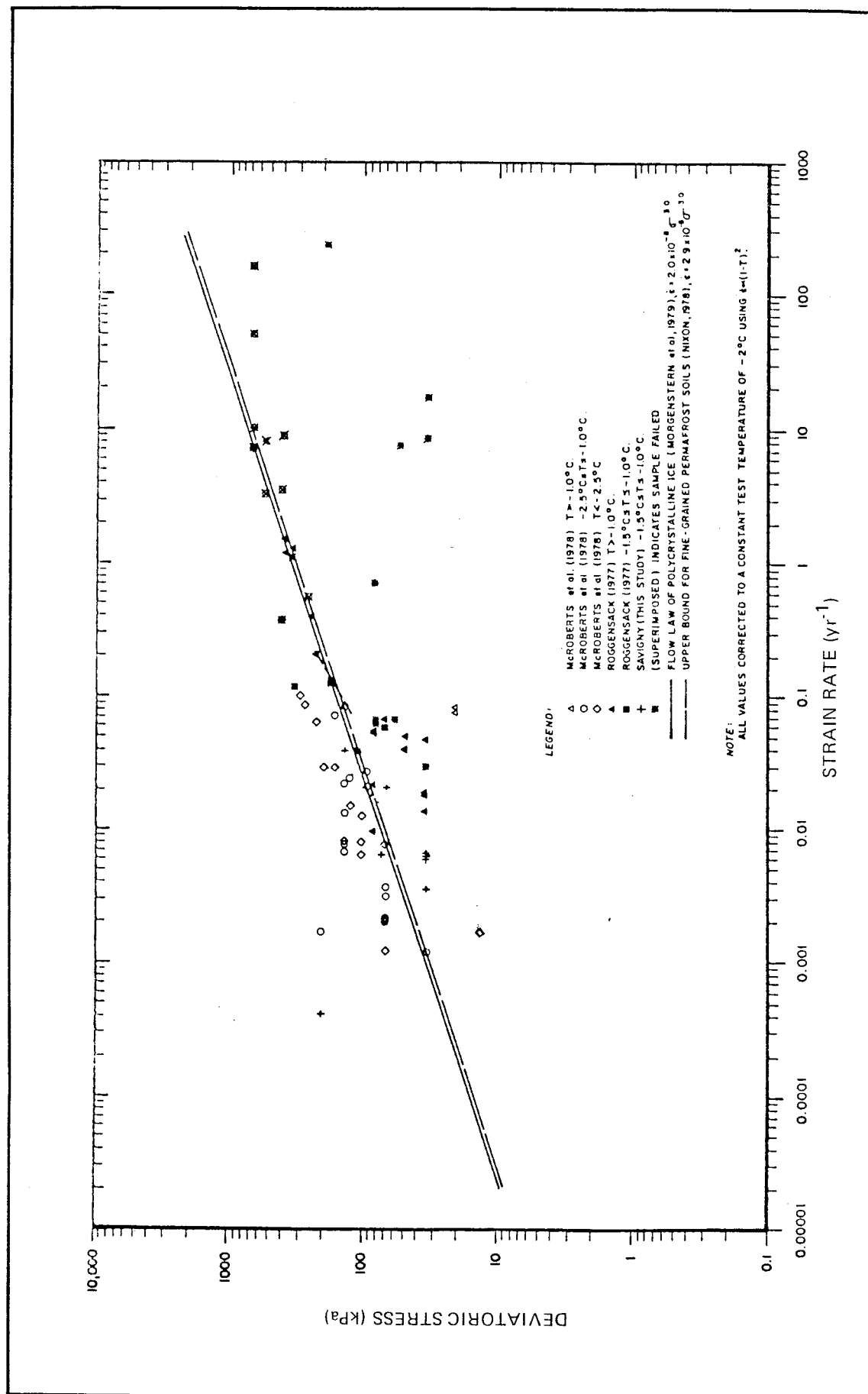


FIGURE 2.5 STEADY-STATE CREEP DATA FOR ICE-RICH GLACIOLACUSTRINE CLAY FROM VARIOUS LOCATIONS IN THE MACKENZIE RIVER VALLEY, N.W.T. (Savigny, 1980)

CHAPTER III

FINITE ELEMENT CREEP ANALYSIS

3.1 General

The primary problem faced by geotechnical engineers concerned with the design and construction of large structures founded on permafrost is the determination of the stress state in the frozen soil. Experimental evidence has shown that time-dependent deformations govern the behavior of frozen soils. The viscous nature of the ice present in discrete layers and in the pore spaces will cause ice-rich soils to creep at very low deviatoric stress levels.

Hoff (1954) simplified the determination of the stress state in a non-linear viscous medium by proposing the elastic analogue. The analogue simply states that:

"The stress distribution in a body whose deformations are governed by a generalized version of a non-linear creep law is the same as that in a non-linear perfectly elastic body provided that the stress-strain law and the boundary conditions are suitably chosen".

There are two methods which can be used to solve creep problems:

- 1) direct iteration procedure - the final solution is obtained by an iterative solution of the non-linear equations
-

- 2) incremental procedure - the final solution is obtained by solving a number of linearized problems between the initial and final time.

In the finite element method, the inclusion of creep behavior is most easily handled by using an incremental procedure and treating the creep strains as initial strains.

The formulation of the incremental procedure was presented by Mendelson et al (1959). In their approach, they extended the solution of plastic flow problems using the method of successive approximations to include creep behavior with arbitrary creep laws. In this method, the elastic solution is used to calculate the first approximation to the creep strains. These strains are then used to calculate a new stress and total strain state within the body. An improved creep strain approximation is then computed and the procedure is repeated until convergence is obtained for that time interval. Essentially, this solution technique is iterative within each time increment. Creep solutions were presented for uniaxial creep in a flat plate and a biaxial case of rotating disks.

Lin (1962) also presented an incremental procedure for analyzing plates subject to bending with arbitrary creep characteristics. However, in this approach the creep strain increments were treated as equivalent loads and edge moments. This procedure is much less complicated than the method presented by Mendelson et al. (1959) since the method of successive approximations may require several iterations within the same time interval.

The application of the finite element method for solving non-linear creep problems using an incremental procedure where the creep strains are treated

as initial strains has been studied by many researchers over the past 15 years. The incremental initial strain procedure is widely used because:

- 1) it is independent of the type of creep law used;
- 2) it provides a description of the intermediate stages of creep
- 3) it is quite simple to extend basic finite element codes to include nonlinear effects arising from creep.

Gerstenkorn and Kobayashi (1966) used the method to solve for creep deformations in a thick-walled cylinder subjected to an internal pressure. Greenbaum and Rubinstein (1968) developed a finite element procedure to study the time-dependent behavior of axisymmetric bodies such as thick-walled cylinders fitted with different types of end closures. Sutherland (1970) extended the work of Greenbaum and Rubinstein (1968) to include problems subject to plane stress and plane strain conditions.

The incremental initial strain procedure was used by Emery (1971) to study the creep behavior of slopes and excavations in rock, cohesive soils and ice. Nair and Boresi (1970) used the incremental solution method to study the time-dependent behavior of circular cavities in rock subjected to a hydrostatic pressure. This procedure was also used by Van Winkel et al. (1972), to predict time-dependent deformation of circular openings in salt media.

A variable stiffness method for creep analysis was presented by Kim and Kuhlemeyer (1977). This method allows the use of relatively large time intervals for which the initial strain procedure would become unstable. However, the disadvantage of this method is the additional computer time associated with regeneration of the structure stiffness matrix at the end of each time interval. This method has proven useful in soil-structure interaction problems where the structure exhibits material non-linearity.

Thompson and Sato (1976) and Dawson and Tillerson (1977) reported a finite element formulation for simulating the creeping flow of an incompressible material. The stress-strain rate response of the material can either be linear (i.e. Newtonian fluid) or non-linear (i.e. non-Newtonian fluid). This method allows the ability to follow the flow through very large changes in material geometry. Velocity components are chosen as the primary variables.

The velocity field is constrained by the material incompressibility condition. This constraint is incorporated into the potential energy functional through the use of a Lagrange multiplier. The multiplier is physically interpreted as the mean normal stress or the pressure within the element. Thus, approximating functions for the pressure as well as the velocity for each element must be chosen. This adds greatly to the solution time required to solve the field equations.

Numerical difficulties arise when trying to satisfy the material incompressibility constraint. To ensure complete incompressibility everywhere within the element, the approximating functions chosen for the pressure within the element must be the same degree as that used for the element strain. Thus, if a linear approximation to the pressure field is chosen, a full quadratic approximating function must be chosen for the components of the velocity field. The complete incompressibility restriction on the velocity field has been shown to give slow convergence for coarse element grids.

Material incompressibility can be satisfied in an average sense over the entire element. This is accomplished by assuming a constant pressure and linear variation for the strain rates and dilatation within each element. In this approach, the total volume of the element remains

unchanged although local incompressibilities within the element are still free to exist. Relaxation of the incompressibility constraint will give more accurate results for coarse element layouts. However, the constant pressure assumption requires a fine element layout in situations where an accurate evaluation of element stresses are desired.

3.2 Incremental Initial Strain Procedure

As a material undergoes creep, the stresses, σ_{ij} , in the body will change with time as shown schematically in Figure 3.1. In the incremental procedure, the smooth stress-time curve is replaced by a series of incremental steps. Each step consists of a constant stress period, Δt , followed by an instantaneous increment of stress.

The fundamental assumption in the incremental creep solution technique is that the time can be subdivided into sufficiently small time intervals such that the stress can be assumed constant within each time interval. If valid, the non-linear creep problem can be solved as a series of linear problems for each time interval using the stresses of the previous time interval to calculate new increments of creep strain and treating them as initial strains for the current time period.

The following assumptions are made in order to establish a stress-strain relationship for a creeping material:

- 1) there is no volume change associated with the creep strains
- 2) for an isotropic material the principal directions of the strain rate and stress tensors coincide;
- 3) the creep rate is independent of any superimposed hydrostatic state of stress;

- 4) the generalization of the uniaxial creep laws to the multiaxial state of stress should recover the uniaxial relationship for the case of uniaxial stress.

Assumptions 1, 2 and 3 can be stated mathematically as:

$$\Delta \dot{\epsilon}_{ij} = \alpha s_{ij} \quad (3.1)$$

where: α = constant of proportionality

This equation states that the increments of creep strain rate are proportional to the instantaneous values of the deviatoric stress tensor and are independent of stress history.

Substituting equation 3.1 into Odqvist's definition of effective strain rate (i.e. equation 2.3) and making use of the definition of effective stress yields:

$$\alpha = \frac{3}{2} (\Delta \dot{\epsilon}_e^C / \sigma_e) \quad (3.2)$$

Substituting equation 3.2 into 3.1 gives the stress-strain rate relationship for a creeping material:

$$\Delta \dot{\epsilon}_{ij}^C = \frac{3}{2} (\Delta \dot{\epsilon}_e^C / \sigma_e) s_{ij} \quad (3.3)$$

For the plane strain case, the following equations are valid:

$$\sigma_e^2 = \frac{1}{2} [(\sigma_x - \sigma_y)^2 + (\sigma_y - \sigma_z)^2 + (\sigma_z - \sigma_x)^2 + 6\tau_{xy}^2] \quad (3.4)$$

$$\Delta \dot{\epsilon}_e^2 = \frac{4}{3} [(\Delta \dot{\epsilon}_x^C)^2 + (\Delta \dot{\epsilon}_y^C)^2 + \Delta \dot{\epsilon}_x^C \Delta \dot{\epsilon}_y^C + (\Delta \dot{\gamma}_{xy}/2)^2] \quad (3.5)$$

$$\Delta \dot{\epsilon}_x^C = (\Delta \dot{\epsilon}_e^C / \sigma_e) (2\sigma_x - \sigma_y - \sigma_z) / 2 \quad (3.6)$$

$$\Delta \dot{\epsilon}_y^C = (\Delta \dot{\epsilon}_e^C / \sigma_e) (2\sigma_y - \sigma_z - \sigma_x) / 2 \quad (3.7)$$

$$\Delta \dot{\epsilon}_z^C = (\Delta \dot{\epsilon}_e^C / \sigma_e) (2\sigma_z - \sigma_x - \sigma_y) / 2 \quad (3.8)$$

$$\Delta \dot{\gamma}_{xy}^C = 3(\Delta \dot{\epsilon}_e^C / \sigma_e) \tau_{xy} \quad (3.9)$$

The essential feature of the incremental procedure is to proceed in small intervals of time and relate an increment of strain to an increment of stress. The final stress and strain state in the body is obtained by summing each increment.

At the beginning of each time interval, the stresses, elastic strains and creep strains are known from the calculations of the previous time increments. At time $t=0$, the elastic solution is used as the starting point. The incremental procedure for each increment of time can be summarized by the following five steps:

- 1) obtain the value of the effective stress (σ_e) from the stresses (σ_{ij}) determined from the previous time step;
- 2) calculate the effective creep strain increment ($\Delta\epsilon_e$) by substituting the effective stress (σ_e) (determined from step 1) into the appropriate creep relationship;
- 3) calculate creep strain increments ($\Delta\epsilon_{ij}^C$) in each direction with the stresses (σ_{ij}) of the preceding time step, the effective stress (σ_e) (from step 1) and the effective creep strain increment ($\Delta\epsilon_e$) (determined in step 2);
- 4) treating the creep strain increments ($\Delta\epsilon_{ij}^C$) (calculated in step 3) as initial strains and using the constitutive equations, boundary conditions and equilibrium equations for the particular problem, calculate the increment of stress ($\Delta\sigma_{ij}^C$) at the end of the time interval; (the initial strains are converted to fictitious creep forces applied at the nodal points);
- 5) the incremental stresses ($\Delta\sigma_{ij}^C$) (obtained in step 4) are added to the stress of the previous time increment to obtain a new stress distribution for the current time interval.

This procedure is repeated for each time interval until either the final time is reached or until the stress distribution does not change (i.e. a steady state condition has been achieved). As long as the incremental stresses are small compared to the previous stresses, the basic assumption of the incremental procedure is not violated. Generally, this condition can be met by selecting time increments small enough to yield the desired accuracy without employing iterative techniques.

3.3

Cumulative Creep Rule

In general, the effective incremental creep strain rate is a function of the effective stress, total effective creep strain, temperature and the strain history of the material. Thus, a general functional form of $\Delta \epsilon_e^C$ is expressed as follows:

$$\Delta \epsilon_e^C = F(\sigma_e, \epsilon_e^C, T, t) \quad (3.10)$$

where: T = temperature
 t = time

Materials exhibiting damped creep response have a strain rate that is a function of time as well as applied stress. An expression must be obtained from constant stress creep tests that is able to predict the strain rate under varying states of stress. Assume the damped constant stress creep data can be presented by an empirical relationship of the form:

$$\epsilon_e = A \sigma_e^n t^m \quad (3.11)$$

where: m = time exponent

The creep strain rate can be obtained by differentiating equation 3.11 with respect to time:

$$\dot{\epsilon}_e = f(\sigma_e, t) = m A \sigma_e^n t^{m-1} \quad (3.12)$$

where: $f(\sigma_e, t)$ = general function form expressing the creep strain rate as a function of applied stress and time.

The time variable can be eliminated from the creep strain rate expression by substituting equation 3.11 into equation 3.12:

$$\dot{\epsilon}_e = g(\sigma_e, \epsilon_e) = mA^{1/m} \sigma_e^{n/m} \epsilon_e^{(1-1/m)} \quad (3.13)$$

where: $g(\sigma_e, \epsilon_e)$ = general function form expressing the creep strain rate as a function of applied stress and total creep strain.

It is clear that integration of equation 3.12 and 3.13 will not give the same result for the general case. Two basic rules that are currently used to describe the time dependency of strain for creep in metals, plastics and concrete are the strain hardening law and the time hardening law.

The strain hardening law for a set of isothermal constant stress curves is illustrated in Figure 3.2. This hypothesis implies the creep strain rate is a function only of the instantaneous value of stress and the total accumulated creep strain. A change in stress state is represented by a horizontal line from one constant stress curve to the new stress curve at the same total creep strain. The total time is given by the sum of all constant stress time intervals. This concept assumes a mechanical-equation-of-state exists for a creeping material relating the variables stress, strain and strain rate. The behavior is independent of the stress history the material was subjected to in the early stages of creep. This is analogous to the theorem proposed by Odquist (1966) which stated:

"If a test piece is subjected to a series of positive stress values, $\sigma_1, \sigma_2, \dots, \sigma_n$, each acting during a time period $\Delta t_1, \Delta t_2, \dots, \Delta t_n$, then the resulting total creep is independent of the order in which the stress values were applied, each during its respective time period".

The time hardening concept for a set of isothermal constant stress creep curves is illustrated in Figure 3.3. This cumulative rule assumes the creep strain rate is governed by the instantaneous value of the stress and total time from the beginning of the test and is independent of the stress history of the material prior to any point in time. A change in stress state is represented by a vertical line from one constant stress curve to the new stress curve at the same total elapsed time. The total accumulated creep strain is given by the sum of the creep strain increments during each time interval.

The validity of these two hypothesis have not been specifically studied in creep tests on permafrost samples. Experience gained in metal and plastic creep as well as undrained creep behavior of cohesive soils suggests that the strain hardening cumulative creep law adequately describes the creep strain rate under varying stress conditions. Ladanyi and Johnston (1973) used the strain hardening hypothesis to determine the in-situ creep parameters for a frozen varved silty clay from multistage pressuremeter data. The strain hardening cumulative creep law can also be used to predict the creep rate of a material under varying temperature conditions (Dorn, 1961).

Derivation of a cumulative creep law which included all the factors governing the creep rate of a material would be a formidable task. Cumulative creep laws have been put forward in the metal creep literature which take stress and temperature history into account. However, these can lead to untractable mathematical expressions.

3.4 Finite Element Formulation of Incremental Initial Strain Method

The basis of the finite element method is to represent the continuum as an assemblage of finite elements. These elements are interconnected at a finite number of points referred to as nodal points. Approximating functions are chosen to represent the variation of the field variable over the element. A variational principle of mechanics is used to derive the equilibrium equations for each element. The equilibrium equations for the entire body are obtained by summing all the individual element equations with proper regard to displacement continuity at the nodal points. The boundary conditions are applied and the entire set of equations are solved to obtain the unknown displacements at the nodal points.

In the present study, the Theorem of Minimum Potential Energy is used to formulate the equilibrium equations. The elements used are constant strain triangles. Within each element, displacement functions are chosen such that inter-element compatibility is maintained. The displacements are assumed to be linear functions of the coordinates. The potential energy of an element is calculated by subtracting the work done by the external forces from the stored strain energy. Minimization of the potential energy yields the equilibrium equations for the element. The equilibrium equations for the entire assemblage are obtained by summing the individual element contributions. The resulting set of linear algebraic equations are solved for the unknown displacements. A detailed derivation of the equilibrium equations for the constant strain triangle element can be found in numerous standard finite element texts [see e.g. Desai and Abel (1972) and Zienkiewicz (1977)].

The creep problem is solved by assuming the change in total strain during one time interval is composed of a change in elastic and creep strains.

$$\Delta \epsilon_X^T = \Delta \epsilon_X^E + \Delta \epsilon_X^C = \frac{1}{E} [\Delta \sigma_X - \nu (\Delta \sigma_Y + \Delta \sigma_Z)] + \Delta \epsilon_X^C$$

$$\Delta \epsilon_Y^T = \Delta \epsilon_Y^E + \Delta \epsilon_Y^C = \frac{1}{E} [\Delta \sigma_Y - \nu (\Delta \sigma_Z + \Delta \sigma_X)] + \Delta \epsilon_Y^C \quad (3.14)$$

$$\Delta \epsilon_Z^T = \Delta \epsilon_Z^E + \Delta \epsilon_Z^C = \frac{1}{E} [\Delta \sigma_Z - \nu (\Delta \sigma_X + \Delta \sigma_Y)] + \Delta \epsilon_Z^C$$

where the superscript T, E and C denote total, elastic and creep respectively.

Plane strain problems have a kinematic constraint placed on the strain in the Z direction (i.e. $\Delta \epsilon_Z^T = 0$). Thus, the third expression in equation 3.14 gives the value for the stress in the Z direction:

$$\Delta \sigma_Z = \nu (\Delta \sigma_X + \Delta \sigma_Y) - E \Delta \epsilon_Z^C \quad (3.15)$$

Substituting equation 3.15 into the first two expressions in equation 3.14 yields:

$$\Delta \epsilon_X^T = \frac{1}{E} [(1-\nu^2) \Delta \sigma_X - \nu(1+\nu) \Delta \sigma_Y] + \Delta \epsilon_X^C + \nu \Delta \epsilon_Z^C \quad (3.16)$$

$$\Delta \epsilon_Y^T = \frac{1}{E} [(1-\nu^2) \Delta \sigma_Y - \nu(1+\nu) \Delta \sigma_X] + \Delta \epsilon_Y^C + \nu \Delta \epsilon_Z^C$$

Equation 3.16 shows that to account for plane strain, a strain equal to $\nu \Delta \epsilon_Z^C$ must be added to the strains in the X and Y directions. The increments of creep strain $\Delta \epsilon_{ij}^C$ are calculated from equation 3.3.

The strain-displacement relations and application of the Theorem of Minimum Potential Energy yields the equilibrium equations for one time interval as:

$$[K]\{\Delta q\} = \{P\} + \{F_C\} \quad (3.17)$$

where: $[K]$ = structure stiffness matrix
 $\{\Delta q\}$ = incremental nodal displacements
 $\{P\}$ = nodal load vector
 $\{F_C\}$ = vector of 'fictitious' nodal loads due to an increment of creep strain.

The introduction of creep strain increments into the equilibrium equations using the initial strain method is given in Appendix A.

At the end of each time interval, the vectors $\{P\}$ and $\{F_C\}$ and boundary conditions are known. Equation 3.17 is solved for the unknown displacement increments by the Choleski square root method for symmetrically banded matrices. For linear elastic analysis, the structure stiffness matrix does not change with time and need only be generated once for the initial time increment and stored to be used for each successive time increment. Should the magnitude of the displacements become excessive, the geometry may need to be updated and the stiffness matrix regenerated to reflect the new grid geometry. This was not necessary for the small displacement problems considered in this research.

3.5 Time Increment Selection

The selection of a time increment to ensure the solution process remain stable is very important in the incremental procedure. During the early or transient stages of creep, the stresses are changing very rapidly so it is necessary to select small time increments. As the solution approaches a steady state condition, the magnitude of the time interval can be increased.

Experience of Greenbaum and Rubinstein (1968), Sutherland (1970) and Emery (1971) shows that the solution process becomes unstable if the maximum change in creep strain is larger than the effective creep strain. The magnitude of time intervals chosen is a function of the stress state in the body and the specific form of the creep law.

The initial time increment is chosen such that the maximum ratio of the effective creep strain increment to the effective elastic strain is equal to or less than η_0 , i.e.:

$$[\Delta \epsilon_e^C / \Delta \epsilon_e^E]_{\max} < \eta_0 \quad (3.18)$$

where: $.04 < \eta_0 < .10$

Successive time increments are computed on the basis of the fractional change in maximum effective stress allowed per time interval, i.e.:

$$\Delta t_{i+1} = [\omega / (\Delta \sigma_e / \sigma_e)_{\max}] \Delta t_i \quad (3.19)$$

where: i) $.035 < \omega < .10$
or ii) $1.2 < \omega / (\Delta \sigma_e / \sigma_e)_{\max} < 2.0$

The maximum ratio of effective creep strain increment to the effective elastic strain for successive time increments must also satisfy:

$$[\Delta \epsilon_e^C / \Delta \epsilon_e^E]_{\max} < \eta_1 \quad (3.20)$$

where: $0.1 < \eta_1 < 1.0$

The optimum values for η_0 , η_1 and ω are a function of the specific creep law and the stress state that exists in the body. Thus, it is advantageous to compute the value for the time interval internally.

The manner in which the program calculates the time interval is as follows:

- 1) Assume for generality the creep behavior of the material is governed by the creep relationship given by equation 3.11, i.e.:
(the steady state creep law is derived from this expression by setting $m = 1.0$), i.e.:

$$\Delta \epsilon_e^C = A \sigma_e^n t^m$$

- 2) The initial time interval is computed by combining equation 3.18 with the above primary creep relationship:

$$\Delta \epsilon_1 = [\eta_o (\sigma_e)_{\max}^{(1-n)} / AE]^{1/m} \quad (3.21)$$

where: E = Young's modulus

The effective stress is calculated based on the elastic stress distribution at $t = 0$.

- 3) Succeeding time intervals are obtained by computing the maximum ratio $(\Delta \sigma_e / \sigma_e)_{\max}$ and satisfying the conditions given by equation 3.19. A further restriction on succeeding time intervals to prohibit divergence is satisfaction of equation 3.20.

A fictitious time, t_f , is introduced for materials governed by a primary creep law and obeying a strain hardening cumulative creep law. The fictitious time is required to calculate the effective creep strain increments for all increments i (where $i = 2, 3, \dots, n$) total number of increments):

$$\Delta \epsilon_e^C = A \sigma_e^n [(t_f + \Delta t)^m - t_f^m] \quad (3.22)$$

where the fictitious time is given by:

$$t_f = [\epsilon_e^C / (A \sigma_e^n)]^{1/m} \quad (3.23)$$

The fictitious time concept is illustrated in Figures 3.4 for an increasing and decreasing stress case.

Once a steady state condition is reached the stresses do not change with time and the creep strains can be extrapolated to the final time. Two tests can be made to check for the existence of steady state creep:

- 1) Check the maximum fractional change in effective stress, i.e.:

$$[\Delta\sigma_e/\sigma_e]_{\max} < \rho \quad (3.24)$$

- 2) Check the maximum change in effective stress per unit time, i.e.:

$$[\Delta\sigma_e/\Delta t_i]_{\max} < \chi \quad (3.25)$$

Values of ρ and χ depend on the specific problem to be solved.

The steady state condition checks can also be used to monitor the stability of the solution procedure. Erratic fluctuations of the maximum fractional change in effective stress or the maximum change in effective stress per unit time would indicate solution instability. To alleviate this problem, smaller values of η_0 , η_1 , and ω are required.

3.6 Finite Element Programme

The finite element programme for time dependent creep deformation analysis was originally developed by Emery (1971) at the University of British Columbia. Minor changes were made to the programme after it was obtained by the University of Alberta.

The incremental initial strain method described in a previous section, is used to solve the creep response. The programme is capable of solving isothermal creep problems with linear elastic material behavior in plane strain situations only. The element type used is the constant strain triangle. External loads consist of nodal point loads and/or gravity loading.

The general form of the flow law implemented into the programme is given by:

$$\dot{\epsilon}_e = [A_1 \sigma_e^{n_1} + A_2 \sigma_e^{n_2}] t^m \quad (3.26)$$

where:

- $\dot{\epsilon}_e$ = effective strain rate
- σ_e = effective stress
- t = time
- A_1, A_2 = creep coefficients
- n_1, n_2 = creep stress exponents
- m = creep time exponent

The parameters A_1 , A_2 , n_1 , n_2 and m are material constants. This general form of flow law was selected because it degenerates to a simple power law, a bilinear flow law and a primary (i.e. transient) power law.

A stop and restart capability has been implemented into the finite element program. This feature allows the creep solution to be interrupted after a specified time or number of increments. Solution results up to that point may be examined, or material properties and/or externally applied loads altered to reflect changes in the creep simulation.

A users' manual for the program CREEP is presented in Appendix C. The programme listing is given in Appendix D. A comparison of the closed-form thick wall cylinder solution with the finite element results is presented in Appendix E.

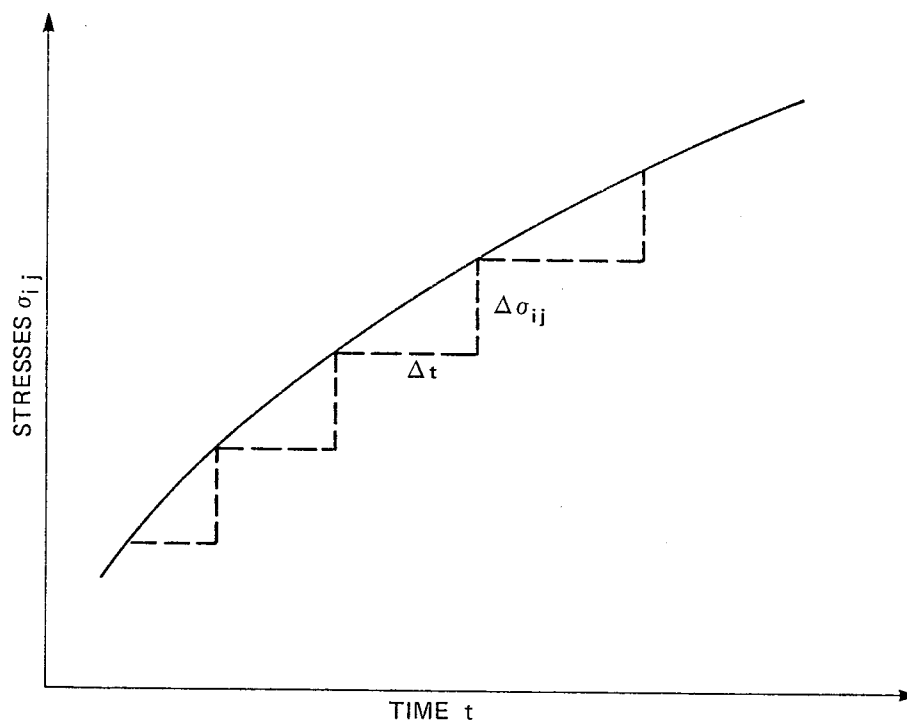


FIGURE 3.1

APPROXIMATION OF A SMOOTH STRESS-TIME
CURVE FOR INCREMENTAL PROCEDURE

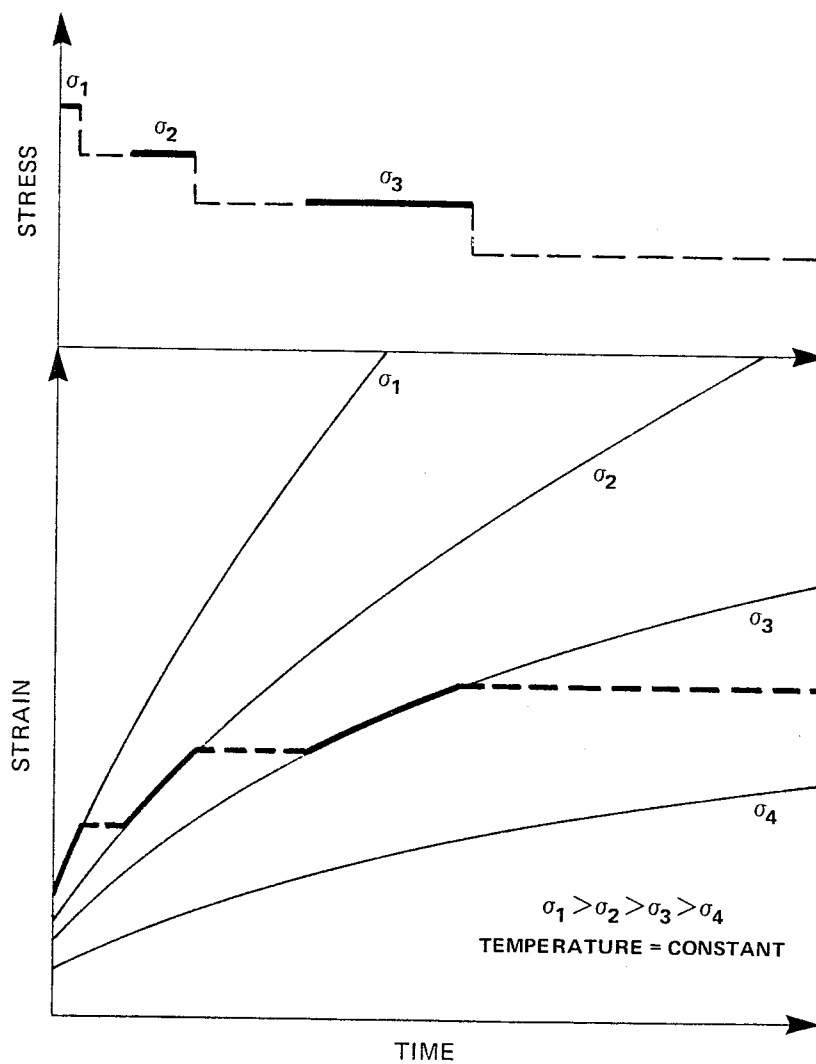


FIGURE 3.2

STRAIN HARDENING CUMULATIVE CREEP RULE

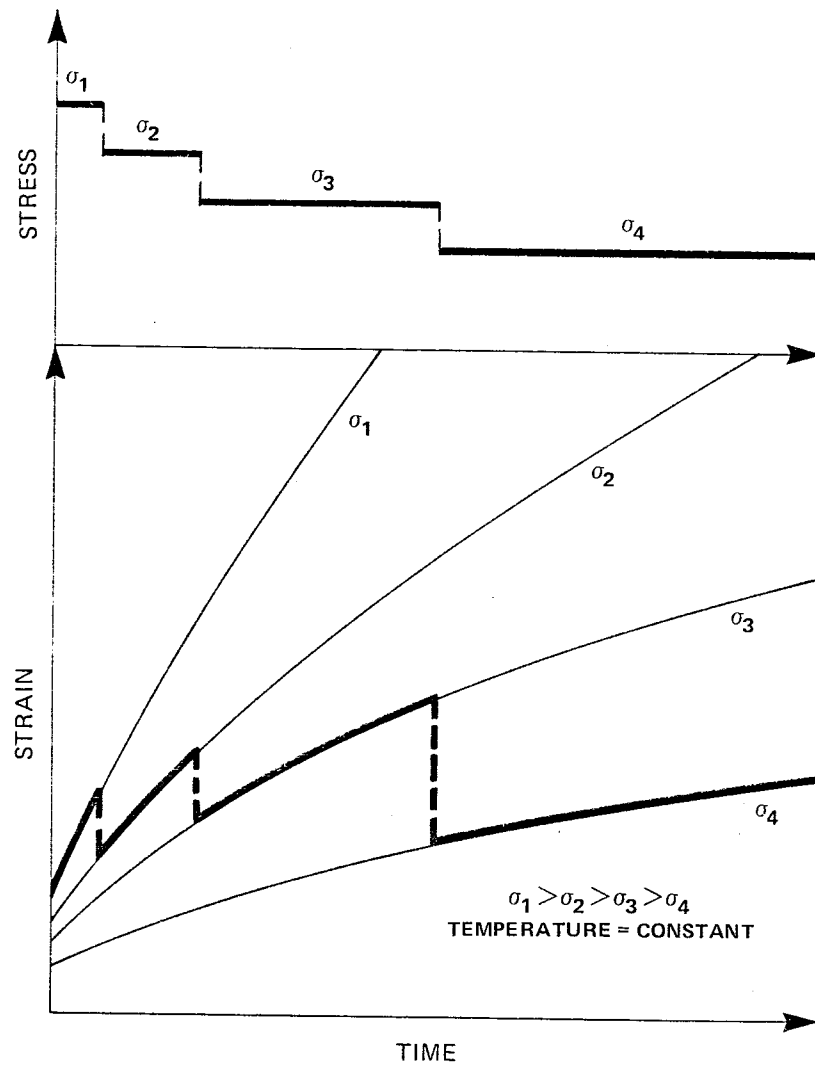
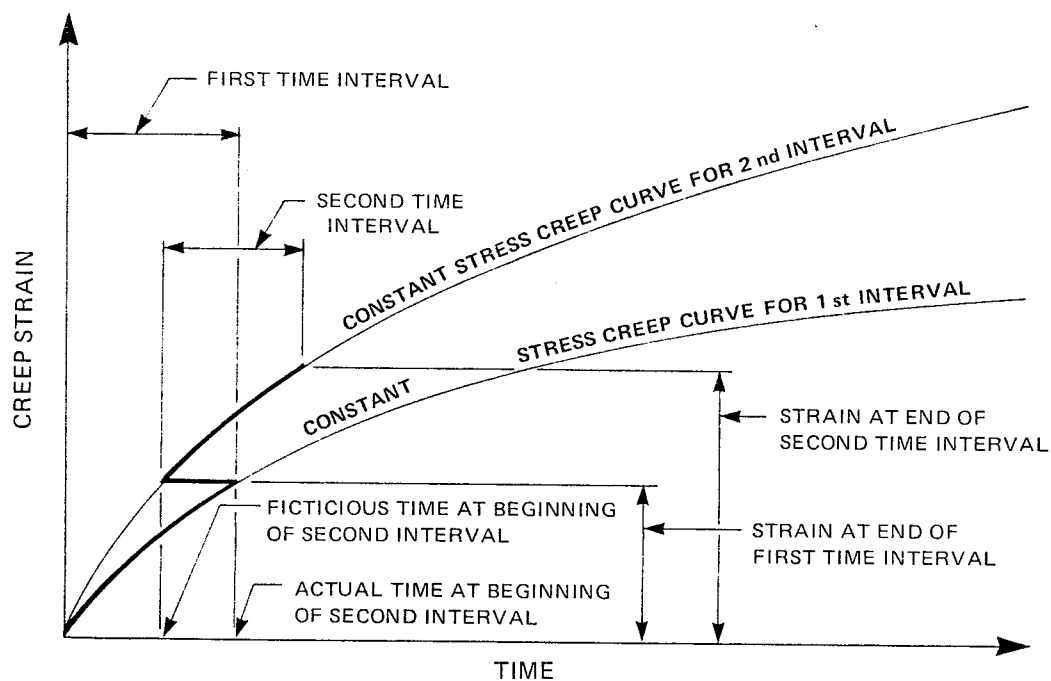
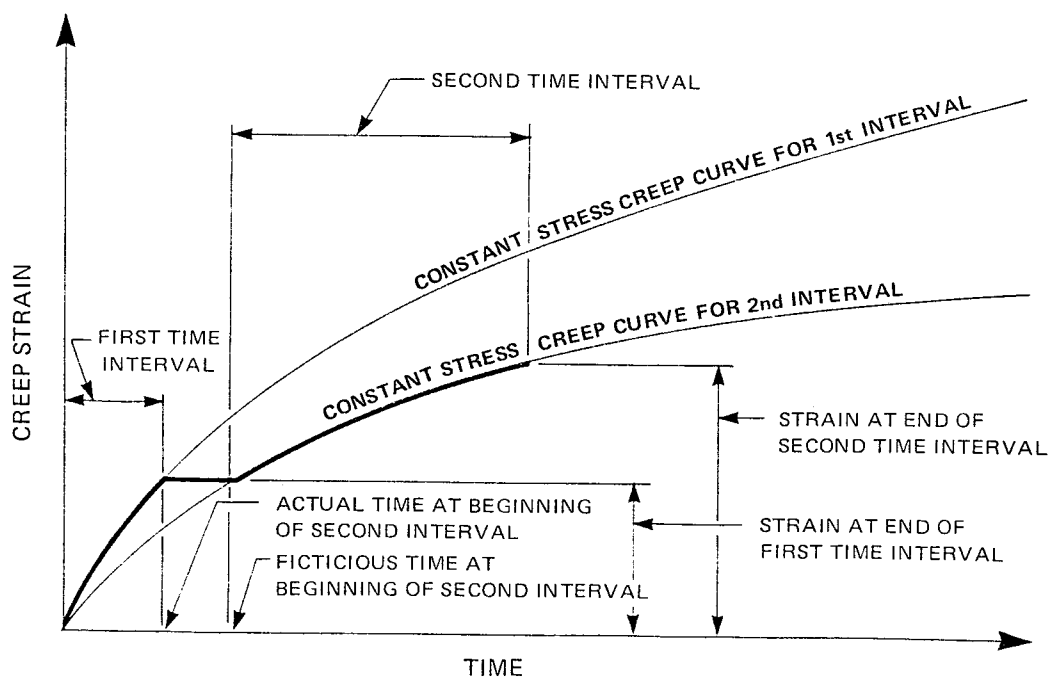


FIGURE 3.3

TIME HARDENING CUMULATIVE CREEP RULE



a) INCREASING STRESS



b) DECREASING STRESS

FIGURE 3.4

CALCULATION OF INCREMENT OF CREEP STRAIN
FOR CHANGING STRESS CONDITIONS

CHAPTER IV

IN SITU CREEP ANALYSIS OF A SLOPE IN ICE-RICH PERMAFROST

4.1 General

The proposed Arctic Gas pipeline route crosses the Great Bear River 7.2 km above the confluence of the Mackenzie and Great Bear Rivers (see Figure 4.1). The left bank of the Great Bear River is consistently steeper than the right bank for several kilometres upstream or downstream from the crossing. Naturally occurring creep was a concern at this particular site because the left bank rises at an average slope 22 degrees over 46 m from river level to the local plain level (see Figure 4.2). Mean river elevation is 56 m a.m.s.l. The local plain elevation is 102 m a.m.s.l. The top of the plain has very little topographic relief south of the slope crest.

The left bank of the Great Bear River at the proposed Arctic Gas crossing was studied by Savigny (1980). An extensive site investigation programme including drilling, sampling and instrumentation was carried out. Undisturbed permafrost core samples were obtained for high quality laboratory testing. The instrumentation was carefully monitored for a two year period. This chapter presents the results of a finite element prediction of the in-situ creep behavior of the proposed Arctic Gas crossing of the Great Bear River.

4.2 Site Geology

The sequence of geologic sediments encountered at the boreholes drilled at this site are shown in Figure 4.3. The slope is composed mainly of sand and clay. Till outcrops near river level. The river has formed in interbedded clay and sand which overlies shale and siltstone bedrock.

The shale and siltstone bedrock are highly weathered and soft. These rocks can be likened to an overconsolidated soil. The bedrock is overlain by an alluvial deposit of interbedded silty clay and sand approximately 6.7 m thick. The two major components of this unit are highly plastic silty clay and clayey silt. Ground ice is present in either reticulate or segregated forms.

A dense glacial till of Wisconsin age lies unconformably on the alluvial deposits. The soil matrix consists of a low to medium plastic sand-silt-clay mixture. The till is highly fissured where it outcrops along the valley wall. Pore ice and reticulate ice are the two most common types of ground ice present.

A fine-grained glaciolacustrine clay rests on the till and is overlain by glaciodeltaic sand. The glaciolacustrine clay was deposited in ice dammed lakes in late Wisconsin time. The soil matrix consists of a medium to highly plastic silty clay. This deposit is approximately 18 m thick at the proposed crossing site. The glaciolacustrine clay is ice-rich. Reticulate ice is most common with primary vertical veins and secondary horizontal veins. Ice veins up to 20 cm thick have been observed. Stratified ice is also common at the till contact. The ice content of this unit increases where the clay outcrops along the valley wall.

The uppermost unit at this site is a 20 m thick stratum of glaciodeltaic sand. This material was deposited following the disappearance of the large ice dams upstream on the Mackenzie River near Fort Good Hope, N.W.T. The sand is medium to fine grained with thin layers of low plastic silt common throughout. Excess pore ice is the dominant type of ground ice.

Some steeply dipping ice veins are also present but not to any great extent. The thickness of the active layer in the sand varies anywhere from .3 to 5 m.

4.3 Field Instrumentation

The field instrumentation installed on the left bank of the Great Bear River at the proposed pipeline crossing consisted of:

- 1) borehole inclinometers to record in-situ downslope creep velocities in ice-rich permafrost soils (boreholes GB1A, GB2 and GB3);
- 2) thermistors to measure the in-situ geothermal gradient (boreholes GB1A, GB2 and GB3);
- 3) piezometers to measure pore pressures below the base of the permafrost (boreholes GB1B and GB3A).

The location of all the boreholes is shown on Figure 4.2.

The pore pressures measured at borehole GB3A show a close correspondence between river level elevation and hydrostatic pore pressure. This is to be expected because the coarse alluvium that lies on top of the bedrock at this site provides a free drainage path.

The ground temperature contours below the depth of zero mean temperature fluctuation for the left bank of the Great Bear River at the proposed pipeline crossing are shown in Figure 4.4. These contours were inferred from the thermistors installed in boreholes GB1A, GB2 and GB3. The temperature data indicate that the permafrost at this site is warm (i.e. warmer than -3°C).

At the top of the slope (GB1A) the active layer is approximately 3 m thick and the depth of annual zero mean temperature fluctuation is 9 to 10 m. The depth of permafrost inferred from the geothermal gradient is at least 61 m. The temperature readings over the two year period show that removal of the surface vegetal cover has initiated a warming trend.

The thermal regime of the slope differs from that on top of the slope. This is primarily due to the aspect of the slope and the difference in vegetation. Along the slope the active layer is approximately 1.2 m thick and the depth of annual zero mean temperature fluctuation is 6 to 7 m. At approximately mid-height of the slope, the mean ground surface temperature is -3.3°C and near the toe of the slope, the warming effect of the Great Bear River increases the mean ground surface temperature to -2.5°C . The steeper thermal gradient measured near the toe of the slope (borehole GB3) is also a result of the thermal disturbance offered by the Great Bear River. The base of the permafrost rises to the river bed elevation as permafrost is absent below the Great Bear River.

Three inclinometers were installed to monitor downslope creep movements:

- i) inclinometer GB1A was installed at the top of the slope through the glaciofluvial sand and glaciolacustrine clay and terminates in the silty clay till;
- ii) inclinometer GB2 was installed at mid-height of the slope where the glaciolacustrine clay outcrops along the valley wall and terminates at the base of the till;
- iii) inclinometer GB3 was installed near the toe of the slope through the glaciolacustrine clay and till and terminates at the base of the interbedded clay and sand.

The location of each inclinometer is shown in Figure 4.5. The inclinometers were read on twelve occasions over a two year period extending from May 6, 1975 to June 14, 1977.

The inclinometer data show a strong correlation between deformations and ice lenses in the glaciolacustrine clay. In zones containing large ice lenses separated 1 to 2 m apart, the movements were large and abrupt and are concentrated at the ice lense location. In zones where the ice lenses are spaced closer than 1 m, the movements are generally smaller and more gradual through the zone. No net transverse slope movements of any practical value were recorded.

Movements recorded in the overlying sand are much more inconsistent than those measured in the clay. Savigny (1980) attributes this to drilling difficulties encountered in the granular material causing disturbed zones around the hole which lead to non-uniform stress distributions in the frozen sand.

The velocity versus depth profiles for all three inclinometer locations are shown in Figure 4.6. A net downslope velocity was recorded through the medium to high plastic glaciolacustrine clay at hole GB1A. The maximum velocity was recorded near the sand-clay contact. Erratic movements were recorded between the 29 m and 34 m depth.

Examination of the GB1A borehole stratigraphy reveals that this depth interval corresponds to a zone containing large pervasive ice lenses spaced greater than 1 m apart. The velocity maxima within this zone coincide with the location of large ice lenses. Above the 29 m depth and below the 34 m depth, there is a pseudo-linear increase in velocity with increasing distance away from the base of the clay layer. The steeper velocity gradient above 29 m is a result of large ice lenses spaced closely together. Below 34 m, the ice lenses are much smaller.

A uniform downslope velocity was recorded in the upper 12 m of the overlying sand at hole GB1A. The magnitude of this velocity is approximately one half of the velocity recorded in the clay just below the sand/clay contact. The 12 to 20 m depth interval of the sand predominantly shows an upslope velocity. The abrupt change at 12 m from upslope to downslope velocity could possibly be due to a large ice vein. At borehole GB1, a large ice vein was observed 13 m below the surface.

No net creep movements were recorded in the transverse-to-slope direction at the GB1A inclinometer.

The velocities recorded through the medium to high plastic clay are very erratic at the GB2 inclinometer location. Savigny (1980) describes two causes for the erratic in-situ velocity:

- 1) difficulties encountered in installation and monitoring of instrumentation
- 2) existence of in-situ sheared material as evidenced by the numerous slickensides in the core samples retrieved from this hole.

A small net downslope creep movement is evident in the silty clay till at the GB2 location. There is a pseudo-linear increase in creep velocity from the base of the till up to the till/glaciolacustrine clay contact.

A net downstream creep velocity was recorded in the glaciolacustrine clay at the GB3 inclinometer location. The velocity gradient appears to be fairly uniform. Erratic fluctuations in velocity exist between 2 and 4 m. The GB3 borehole description reveals that this depth interval coincides with a zone of small ice lenses that are widely spaced.

4.4 Finite Element Analyses

The stratigraphy of the left bank of the Great Bear River at the proposed crossing site was discretized for the finite element analysis as shown in Figure 4.7. The finite element grid consisted of 22.5 m of sand overlying 15.5 m of medium to high plastic glaciolacustrine clay. The stiff glacial till that underlies the clay was not discretized. The finite element grid was fixed at the clay/till interface to account for the very high in-situ modulus of the till.

The finite element programme CREEP computes movements in the plane of the slope only. Transverse-to-slope movements are assumed to be zero. The inclinometer data reveals that no net transverse-to-slope movements were recorded at hole locations GB1A and GB3. At hole GB2, the transverse-to-

slope velocity recorded in the glaciolacustrine clay is greater than the parallel-to-slope velocity. However, for reasons noted in the previous section, the data for hole GB2 is not considered reliable.

Elastic material properties were assigned on the basis of data presented in the literature. Mean values of bulk unit weight for the sand and clay were assigned on the basis of measured values obtained from the frozen core samples. The elastic material properties for the sand and clay are summarized in Table 4.1.

TABLE 4.1

ELASTIC MATERIAL PROPERTIES FOR FROZEN GLACIODELTAIC
SAND AND GLACIOLACUSTRINE CLAY

PROPERTY	SAND	CLAY
Young's Modulus	4900 MPa	785 MPa
Poisson's Ratio	.495	.495
Unit Weight	1.84 Mg/m ³	2.04 Mg/m ³

Creep properties of the sand and clay were altered to obtain close agreement with the in-situ behavior. Two finite element creep simulations were carried out.

The first creep analysis assumed that the sand and clay would creep at the same rate. A review of the rheological behavior of frozen soils presented in Chapter II concluded that the stress versus strain-rate relationship for polycrystalline ice represents an upper bound to the constitutive behavior of fine-grained, ice-rich permafrost soils. Using this argument, the flow law used in the first analysis was:

$$\dot{\epsilon} = 2.0 \times 10^{-8} \sigma^3 \quad (4.1)$$

It was expected that this flow law would overestimate the creep velocities in the sand and clay. The degree of overestimation would be greatest in the sand. Laboratory samples of the sand could not be obtained for laboratory creep testing because of sampling difficulties.

The predicted creep velocity profiles for all three inclinometer locations are shown in Figure 4.8. The in-situ creep velocities are also shown in this figure. At inclinometer locations GB1A and GB3, the predicted velocity profile generally has a similar form to the in-situ profiles. However, the magnitude is much greater. As expected, there is no correlation between measured and observed creep velocities at inclinometer location GB2.

A close examination of the magnitude of predicted velocities versus the observed values revealed that the difference between the two is a constant factor. The creep strain rate varies linearly with the coefficient in the simple power law. Hence, a proportional change in the coefficient will reduce the predicted velocities to the observed values.

The predicted velocities were approximately six times greater than the in-situ values. To reduce the predicted velocities proportionately, the coefficient in the simple power law was set equal to one sixth of its original value. Thus, the flow law for the sand and clay now becomes:

$$\dot{\epsilon} = .33 \times 10^{-8} \sigma^3 \quad (4.2)$$

The exponent was left unchanged from its original value of 3.0.

The predicted velocities reduced by a factor of six are shown in Figure 4.9. There is good agreement between the predicted and in-situ velocity profiles for the GB1A and GB3 inclinometer locations through the clay stratum. There is some discrepancy in the upper 2 to 3 m of the GB3 hole due to the influence of the steel casing (Savigny, 1980). Also, the datum for the finite element zero velocity at the GB1A and GB3 holes were translated horizontally to account for the in-situ movement which occurred in the underlying till.

The predicted in-situ creep law is shown in Figure 4.10 along with the flow law for ice (Morgenstern et al., 1980) and the upper bound for soil data as proposed by Nixon (1978). The predicted flow law shows good agreement with the laboratory data. As shown in this figure, the laboratory data overestimate the in-situ creep rate.

The accumulated creep displacement pattern after a steady state condition was achieved is shown in Figure 4.11. The displacement profiles versus depth are indicative of non-Newtonian fluid behavior.

The preliminary finite element creep analysis has shown that the in-situ creep behavior of the ice-rich glaciolacustrine clay can be modelled with the simple power law. However, the agreement between predicted and observed velocities through the sand at inclinometer GB1A is less convincing. At the outset, there was some doubt regarding the form of the constitutive behavior of the sand and whether the sand would creep at all. The ice-rich glaciolacustrine clay was expected to creep much faster than the overlying sand. This would set up tensile stresses in the sand near the sand/clay interface. Evidence of a complex stress state at this interface is revealed by the observed erratic movements at the base of the sand.

Selection of a suitable constitutive relationship for the sand is difficult due to the lack of long term low stress creep data for natural coarse-grained permafrost soils. Creep tests on reconstituted frozen sands have shown that the deformations continue to attenuate with time at stress levels below the long term strength (Sayles, 1968). Creep deformations in sand have also been shown to be a function of confining pressure (Sayles, 1973; Alkire and Andersland, 1973). There is also the question of the form of the constitutive relationship for the sand under tensile stresses. Since unfrozen sand has zero tensile strength, all the tensile stresses must be transmitted through the ice phase. Hawkes and Mellor (1972) have shown that the uniaxial strength of polycrystalline ice is similar in tension and compression at slow strain rates. However, the creep behavior of coarse-grained ice-poor frozen soils may differ considerably in tension and compression. In a compression creep test on a coarse-grained frozen soil, there are two components which resist deformation. First, there are the time-dependent deformations occurring in the ice phase and secondly, inter-particle friction must also be overcome. In a tension creep test only the ice phase will resist the application of load since the interparticle friction has been reduced to zero. To investigate these points, a second

creep analysis was carried out in which the sand was restricted to deform elastically. The clay was assigned the flow law given by equation 4.2.

The predicted velocity profiles for this analysis at all three inclinometer locations is shown in Figure 4.12. Restricting the sand to deform elastically impedes the creep deformations in the clay at the GB1A and GB2 inclinometer locations. At these two inclinometer locations, the predicted velocities do not agree either in form or magnitude with the observed data. The predicted velocities at the GB3 inclinometer are only slightly reduced from the analysis assuming uniform creep properties because this region is well beyond the influence of the sand stratum.

The displacement pattern for the second creep analysis is shown in Figure 4.13. In this case, the clay is being extruded out between the sand and underlying till. As the distance downslope from the point the sand pinches out increases, the displacement pattern begins to resemble non-Newtonian fluid behavior as it did in the first creep analysis.

The extrusion of the clay leads to very high horizontal tensile stresses in the sand as shown in Figure 4.14. The magnitude of the tensile stresses indicate the formation of tensile failure zones propagating through the sand as it restricts the downslope movement of the underlying clay.

4.5 Assessment of Analytical Results

On the basis of the finite element creep analysis of the left bank of the Great Bear River at the proposed Arctic Gas crossing, it can be concluded that the in-situ creep behavior of the glaciolacustrine clay can be accurately represented by the simple power law. The flow law given by equation 4.2 provided the best fit between observed and predicted

velocities. The exponent in the power law is the same as that for polycrystalline ice, however, the coefficient is one sixth of the ice value (Morgenstern et al. 1980).

The in-situ creep behavior of the sand may also be represented by the simple power law. However, this remains inconclusive until further insight is gained into the in-situ stress state in the sand. Confining the sand to behave elastically poses much too serious a restriction on creep movements in the clay. The in-situ creep behavior indicates the clay creeps at a faster rate than the sand. The dissimilar in-situ creep behavior in the clay and sand would set up a complicated stress field at the sand/clay interface. It is believed that the faster creep movements in the clay is causing tensile failure in the sand. Thus, to more accurately model the in-situ behavior of the sand, inclusion of a tensile failure criterion would be necessary. Approximate numerical techniques are available for treating soil and rock as either a 'no tension' or 'limited tension' material. Application of such procedures is severely limited in frozen soil mechanics since very little is known about their constitutive behavior under tensile stress conditions.

The largest discrepancies between observed and predicted velocity profiles appear to coincide with regions that have experienced failure. Two such zones are the base of the sand at inclinometer GB1A and the entire thickness of clay at GB2 inclinometer location. It is believed that the abrupt change in the in-situ creep properties between the clay and sand has lead to the development of a zone of tensile failure in the lower 8 m of sand. The upper 12 m of the sand, which is well beyond the zone of influence of the sand/clay interface, generally shows a uniform downslope velocity.

Another such zone occurs at the GB2 inclinometer location. The sand deposit pinches out just upslope of this borehole. The stiffer sand appears to impede the creep velocity in the underlying clay. Downslope of the GB2 inclinometer, the in-situ creep behavior of the clay is not influenced by the overlying sand and creeps at a faster velocity than the clay underlying the sand. The upslope restraint offered by the sand and the downslope pull of the unrestrained creep in the clay would lead to horizontal stress relaxation at GB2. It is believed that numerous slickensided surfaces observed in the core of borehole GB2 occurred as a result of continued horizontal stress relief while the vertical stress remained constant at the overburden pressure.

The in-situ creep behavior of the clay at inclinometer GB3 is not influenced by the sand. The steady state stress field in the clay in the vicinity of GB3 lies well beyond the disturbing influence of the sand.

4.6 Modelling Ice-Rich Permafrost as a Two-Phase Continuum

At the proposed Arctic Gas crossing of the Great Bear River, the data gathered from inclinometer locations GB1A and GB3 clearly show that velocities are erratic and more movement is associated with the large ice lenses where they are widely separated (Savigny, 1980). The constant stress creep tests carried out on samples obtained from the field program all showed that the deformation behavior was affected by the ground ice structure. Savigny (1980) reported that samples tested with a reticulate ice structure all failed along the soil/ice interface with virtually no shearing through the soil. One sample tested with stratified ice oriented perpendicular to the direction of load application failed by radial tensile failure in the soil. It is clear from the foregoing points, the ice structures present in an ice-rich permafrost play an important role in

establishing the overall deformation behavior of the soil mass. Thus, there is good reason to believe that the segregated ice features can be treated as discontinuities in a frozen soil matrix.

In the field of rock mechanics, the presence of discontinuities has long been recognized by engineers designing structures in the rock. The behavior of a rock mass is not only a function of the intact rock but also the nature and extent of the discontinuities present.

Generally, all rock masses contain planes of potential weakness which come in all lengths and spacings and have varying degrees of influence on the overall rock mass properties. A geological field investigation can accurately map the orientation and absolute position of the more important discontinuities so that their influence on the rock mass behavior can be adequately evaluated.

Conventional finite element analysis utilizes the continuum approach which enforces material compatibility throughout the body being analyzed. The continuum approach may be entirely adequate when dealing with small joint spacings relative to the size of the structure being analyzed. Situations where rocks deform along preexisting planes of weakness present additional degrees of freedom to the rock mass. Thus, the analytical model must allow relative movement to obtain a realistic solution.

When dealing with a single very important discontinuity whose orientation and position are known with confidence, it is possible to represent plane of weakness explicitly and calculate the resulting stresses and deformations in the rock mass. Several interface elements have been developed over the past ten years for solving problems that involve relative movement.

Goodman et al. (1968) introduced a four node displacement discontinuity element. Zienkiewicz et al. (1970) presented a continuous isoparatric interface element with linear strain along the element and uniform strain across its thickness. Ghaboussi et al. (1973) advocates the use of relative displacement as an independent variable to prevent adjacent blocks of continuous elements from penetrating into each other. Simmons (1981) used a thin linear strain isoparametric continuum element to model shearband yielding and strain weakening in dense, structured soils.

The essential feature in using these interface elements is to be able to predict the location of a shearzone or map a through-going discontinuity. Little is known about the aerial extent of individual lenses of segregated ice or reticulate ice veins. The ice lenses may vary from hairline cracks to massive ice bodies with vast horizontal and vertical extent. The factors governing the amount and distribution of ground ice are the type of host material, availability of moisture, rate of freezing and geologic history of events. The most common form of segregated ice in the glaciolacustrine clay at the proposed Arctic Gas Crossing site was reticulate ice with smaller amounts of segregated and stratified ice. The thicker primary veins in the reticulate ice structure are typically vertical making lateral correlation of these ice features very difficult.

The variability of ground ice is of major significance when studying the stability of buried, warm pipelines operating through permafrost. Determination of the amount and distribution of ground ice is essential for evaluating the configuration of the thawed ditch bottom profile. The lateral variation of ground ice can be determined by visual examination of ditch walls or detailed logging of continuous undisturbed samples obtained from closely spaced boreholes.

Mapping of ground ice structures found on ditch walls will provide the most reliable estimate of the continuity of individual ice features. However, an undertaking of this sort would be very costly. Ditch wall disturbance as a result of overbreak would preclude the use of a blast and backhoe operation to excavate the ditch. Large diameter bucket wheel ditchers produce a smooth and relatively undisturbed ditch wall ideal for observing the ground ice structure. Very few ditching trials have been carried out and reported in the public literature. Also, the few trials that have taken place were primarily concerned with the ability of the ditcher to excavate frozen permafrost rather than mapping the ground ice structures.

Detailed logging of the ice structures observed in undisturbed permafrost core samples obtained from closely spaced boreholes would be much more expedient. The ground ice may appear to be much more variable in the undisturbed core samples than that actually present in the soil mass. However, an estimate of the variability of ground ice in the soil mass can be obtained by a statistical evaluation of the observations at individual boreholes.

Use of statistical analysis to solve variability problems in geotechnical engineering is relatively recent. Quality control on construction projects and to a lesser extent the analysis of laboratory test data are two areas where much of the application of statistical techniques has taken place. Ideally, a known statistical model would be used to describe the population under consideration. Holtz and Krizek (1971) studied several foundation case histories and found that the laboratory test data essentially conform to the Gaussian or normal distribution.

Speer et al. (1973) reported the results of a ground ice variability study undertaken by Mackenzie Valley Pipeline Research Limited (MVPRL). The purpose of the ice variability study was to assess the amount of ground ice and its distribution in the thawed zone below a warm oil pipeline buried in ice-rich permafrost. The thaw settlement potential of representative geologic units through which the proposed pipeline might be buried is required to compute pipe stresses under operating conditions.

When fine-grained ice-rich permafrost soil thaws, water is released and settlements develop as the water is expelled from the pore space due to the soils self weight and applied loads. Laboratory studies have shown that for light loadings in relatively ice-rich soils, a major portion of the thaw settlement occurs during the thaw stage. This component of settlement is independent of external loading and is caused by excess water draining away. A rough estimate of the thaw settlement for ice-rich soils can be obtained from the thickness of visible ice lenses. Thus, it would seem reasonable to assume that the variation in thaw settlement provides an estimate of the variability of ground ice.

Field and laboratory studies were carried out by MVPRL to study ground ice variability in silt and clay soils. The lateral variation of ground ice content was studied in rectangular arrays of closely spaced borings. An estimate of the total potential settlement was calculated from an empirical correlation between frozen bulk density and thaw strain derived from laboratory tests on undisturbed core samples obtained during the field investigation.

Two arrays were drilled in lacustrine silts and clays deposited in a proglacial lake. These deposits are considered texturally similar to those encountered at the proposed Arctic Gas crossing of the Great Bear River.

The Norman Wells study site is located approximately 3 km northwest of the airport runway at Norman Wells. The Landing Lake array is located approximately 53 km northwest of Norman Wells on the east bank of the Mackenzie River. At both sites, reticulate and segregated ground ice structures were present in the lacustrine silt and clay. Massive ice structures were not observed.

The estimated total thaw settlement for the Norman Wells and Landing Lake study sites are shown in Figure 4.15. The depth interval that thaw settlements were calculated was 2 to 13 m. There is good reason to believe that for a uniform stratigraphic unit, the ground ice will be a normally distributed variable. The combined data for the Norman Wells and Landing Lake arrays plotted in terms of a thaw settlement probability distribution on normal probability paper is shown in Figure 4.16. The observed data shows little scatter about the theoretical normal distribution. Also shown in the figure are the results of the Kolmogorov-Smirnov goodness of fit test evaluated for normality at the 20% significance level. This test indicates that the thaw settlement in lacustrine silts and clays of the Mackenzie River Valley can be described by a normal distribution.

Although the foregoing has shown that the variability of ground ice in a uniform stratigraphic unit can be considered to be normally distributed, it has done little to elucidate the lateral extent of individual ice lenses. One method of estimating the applicable horizontal scale for a given depth interval is to observe the surface profile of ground which was once frozen but has subsequently thawed such as seismic cutlines, highways and airstrips.

A mathematical description of the overall ground ice correlation between pairs of points at various arbitrary spacings is through the concept

of an auto-correlation function. This function is defined as the product of the deviations from the mean thaw settlement measured at pairs of points spaced an arbitrary distance, ℓ , apart averaged over the entire population, ie:

$$R(k) = \frac{1}{n-k} \left[\sum_{i=1}^{n-k} (x_i - \bar{x})(x_{i+k} - \bar{x}) \right] \quad (4.3)$$

where: $R(k)$ = auto-correlation at spacing k
 n = total number of observations
 x_i = observation at point i
 \bar{x} = mean

For the thaw settlement data, the auto-correlation function expresses the dependence of the correlation between settlements measured at boreholes i and $i + k$ on the distance, $k\ell$, between boreholes. The distance at which the auto-correlation falls to zero indicates the horizontal scale over which the thaw settlements can be correlated. The product of the deviations from the mean thaw settlement spaced further than this distance can be expected to be as often positive as it is negative. Thus, the expected value of the sum of these products should be zero beyond the correlation distance for a very large population.

The auto-correlation function determined from the Norman Wells and Landing Lake thaw settlement data is shown in Figure 4.17. The correlation is almost zero for 15 m and fluctuates about zero at greater spacings between boreholes. This suggests that thaw settlements are only weakly correlated at boreholes spaced 15 m apart. No data is available at spacings less than 15 m since the closest borehole spacing at these two ice variability test sites was 15 m. Field observations of disturbed areas in lacustrine silts and clays of the Mackenzie River Valley indicate that 15 m

represents an upperbound for the distance at which ground ice content is correlated.

The thaw settlement data represents a one-dimensional integration of the ground ice conditions at a particular borehole over a given depth interval. The data derived from the MVPRL ice variability test sites at Norman Wells and Landing Lake suggest that for a uniform stratigraphic unit, the overall ground ice content of the soil mass can be described as a normally distributed variable. The ground ice content at a particular point is completely independent of neighboring points separated a distance of 15 m. Thus, the time-dependent deformation behavior of ice-rich frozen soil is most appropriately treated analytically by assuming homogeneity for the ice-rich permafrost until a better understanding of the lateral extent of individual ice lenses is obtained.

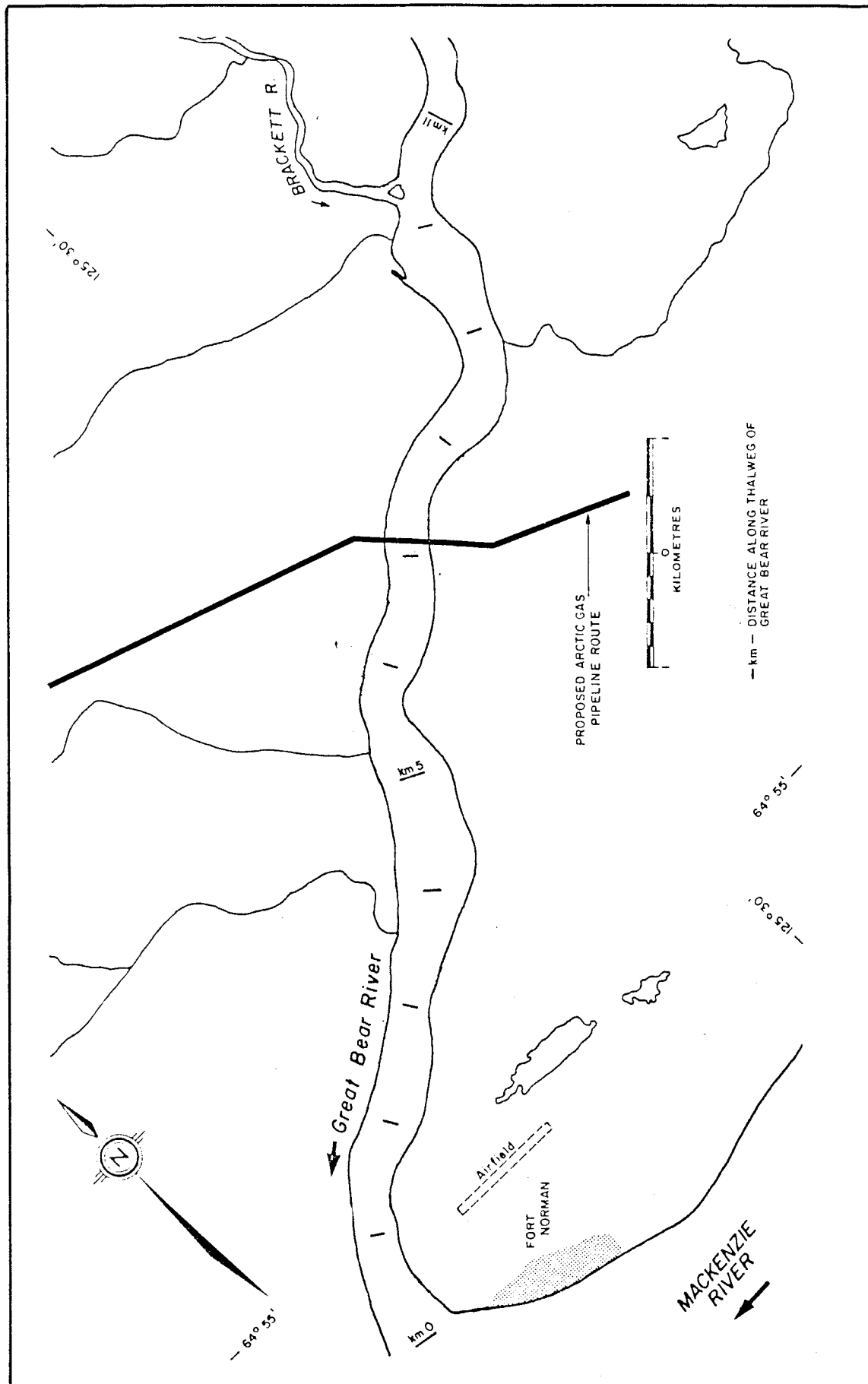


FIGURE 4.1 SITE PLAN OF PROPOSED ARCTIC GAS CROSSING OF GREY BEAR RIVER, N.W.T. (Savigny, 1980)

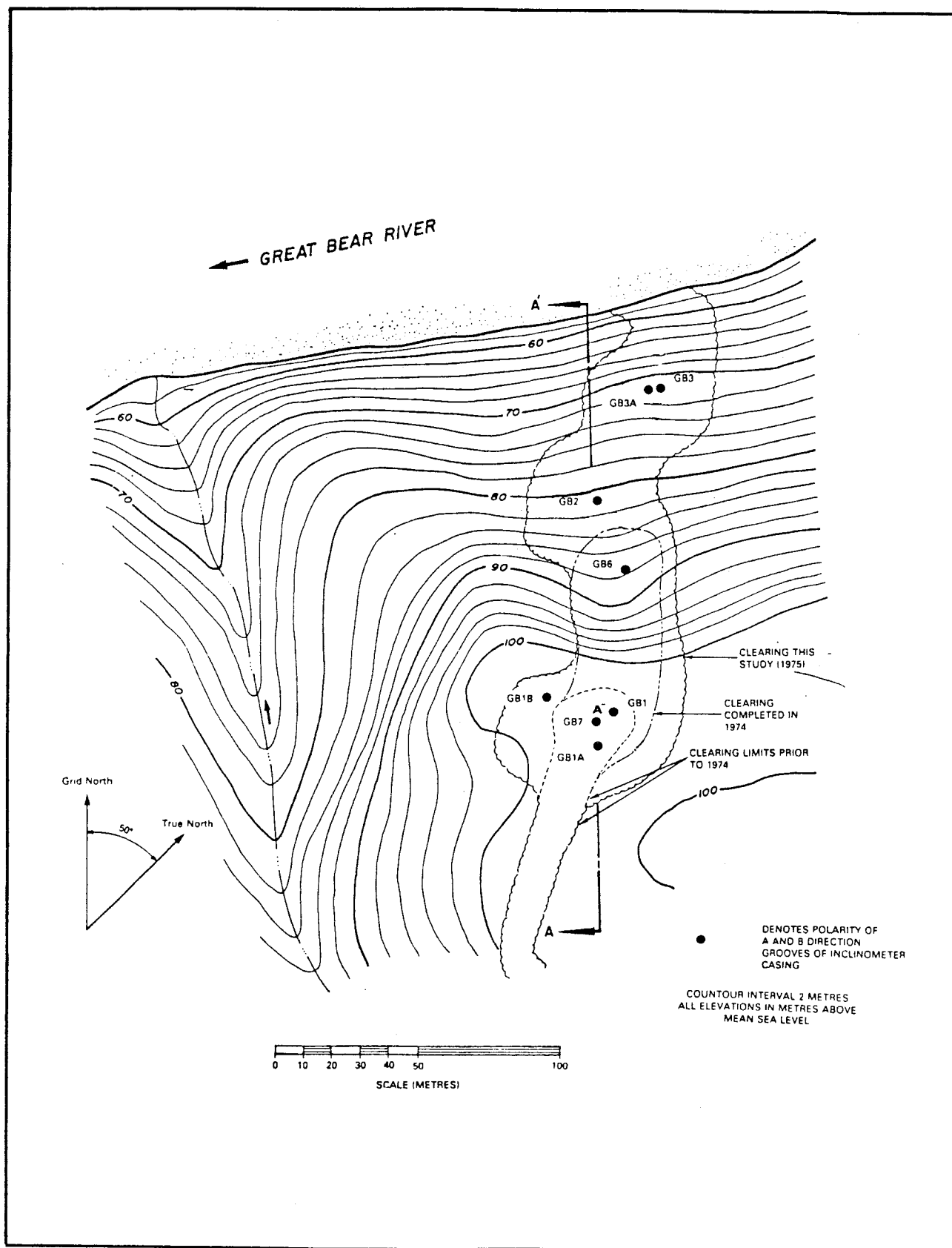


FIGURE 4.2

SITE PLAN OF LEFT BANK OF GREAT BEAR RIVER AT
PROPOSED ARCTIC GAS CROSSING
(Savigny, 1980)

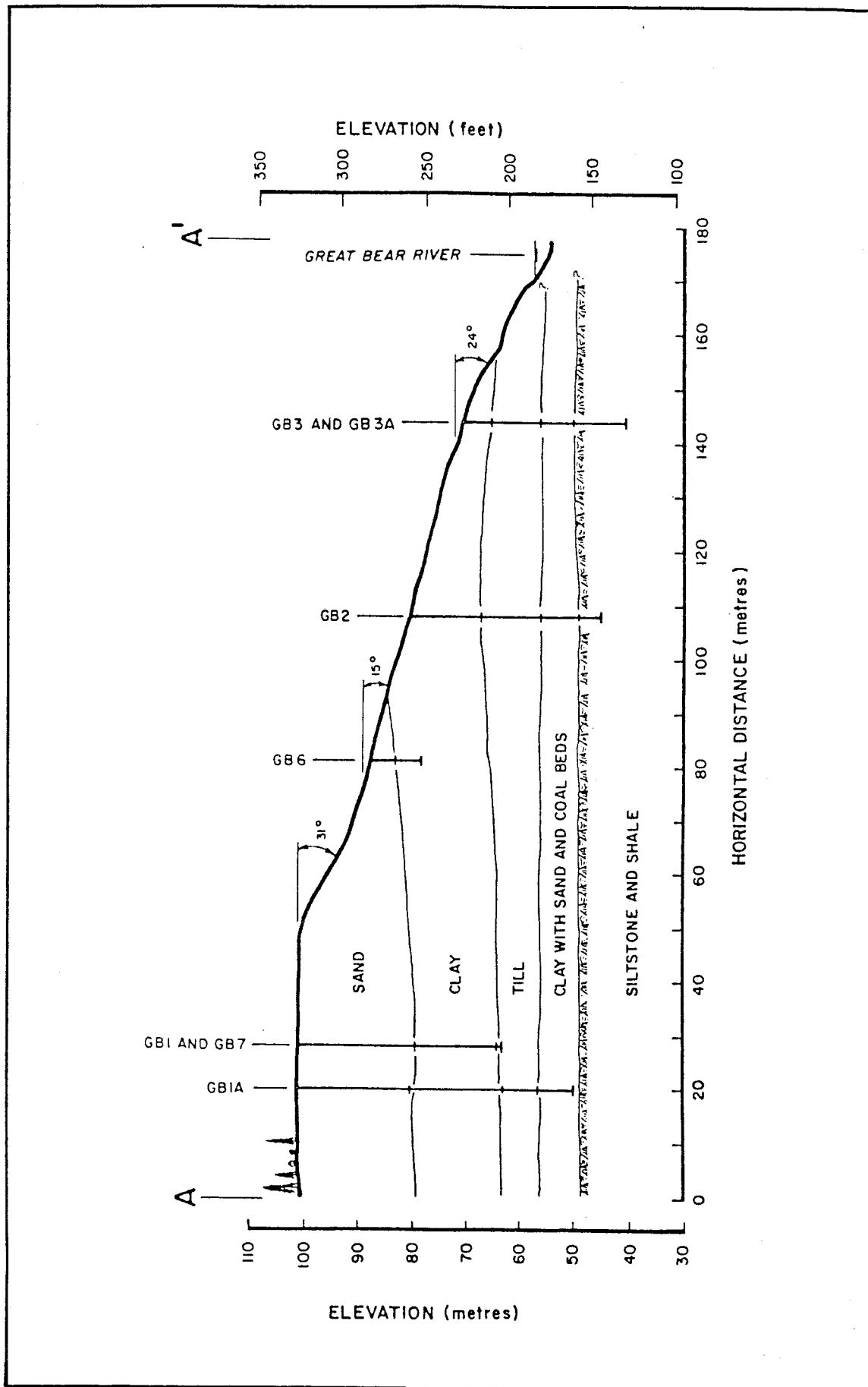


FIGURE 4.3 STRATIGRAPHIC CROSS SECTION; LEFT BANK OF GREAT BEAR RIVER AT PROPOSED ARCTIC GAS CROSSING (Savigny, 1980)

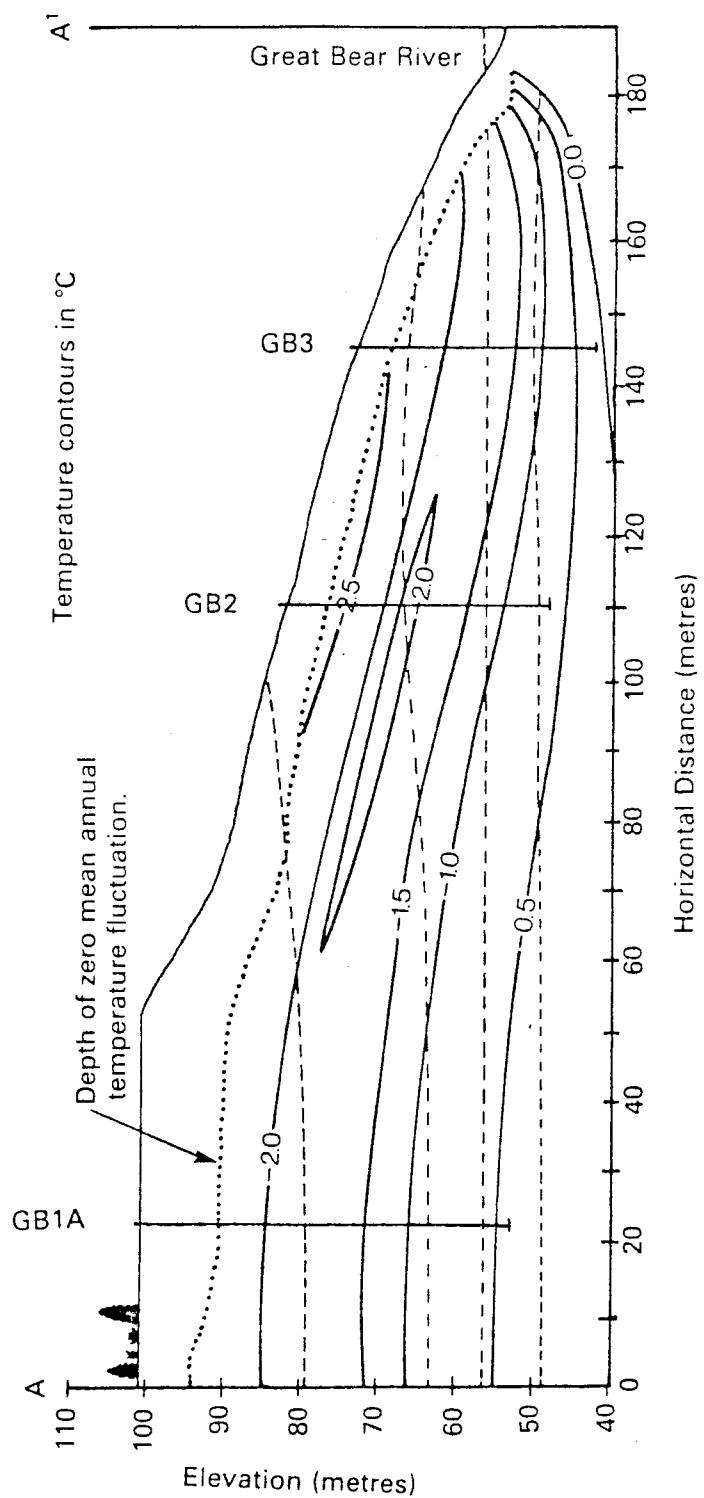


FIGURE 4.4 THERMAL CROSS SECTION; LEFT BANK OF GREAT BEAR RIVER AT PROPOSED ARCTIC GAS CROSSING (Savigny, 1980)

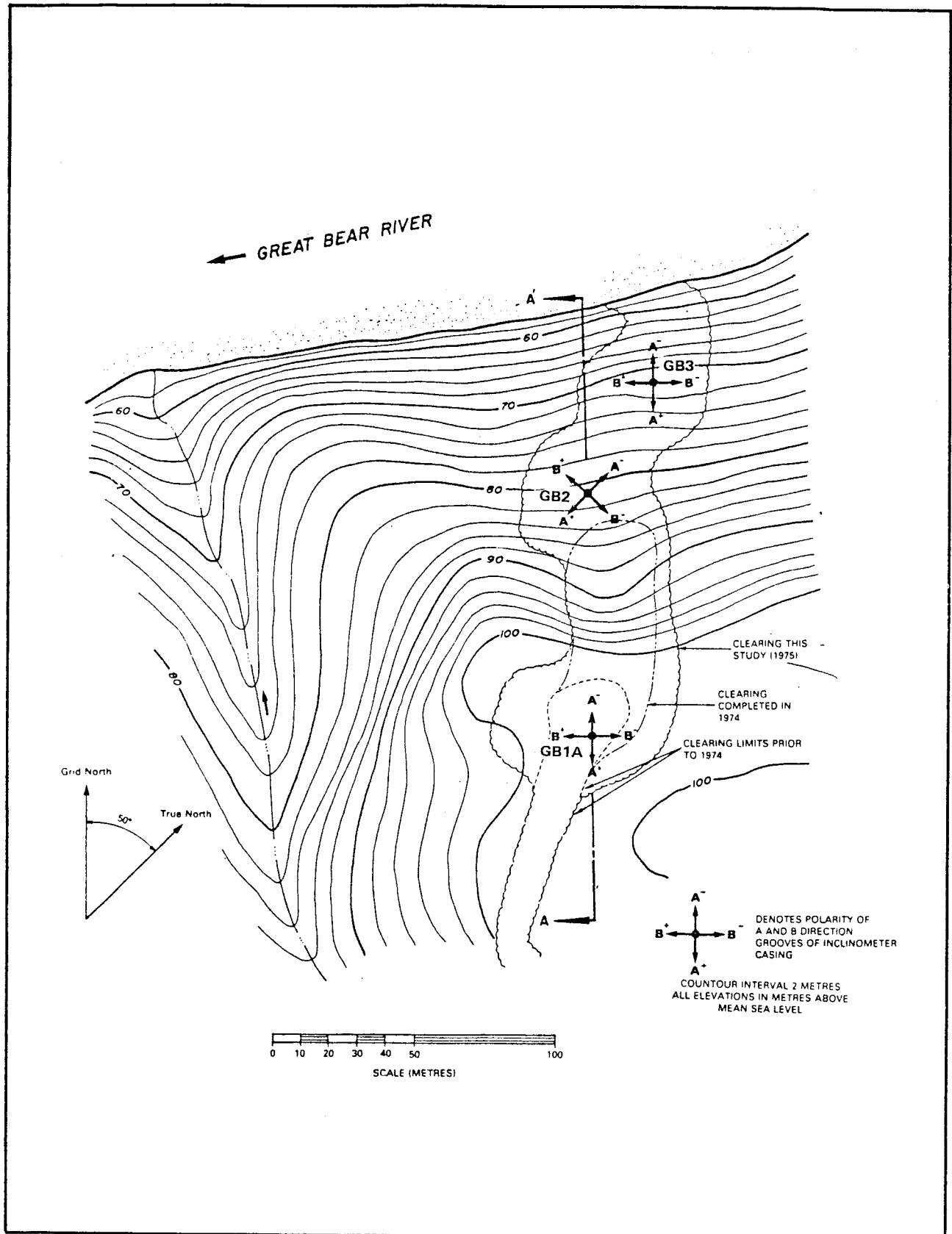
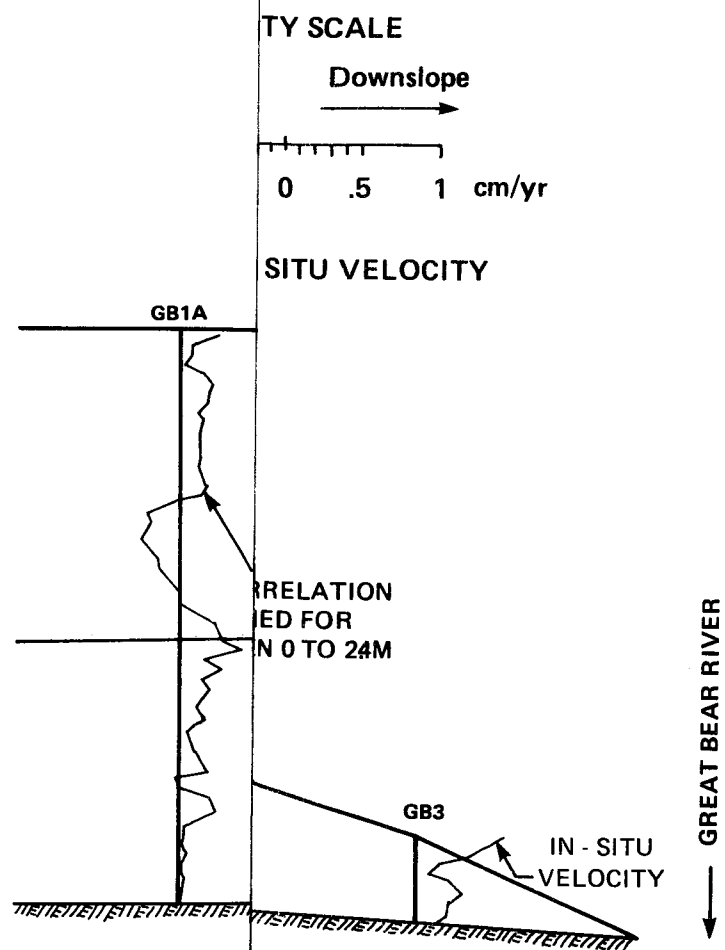
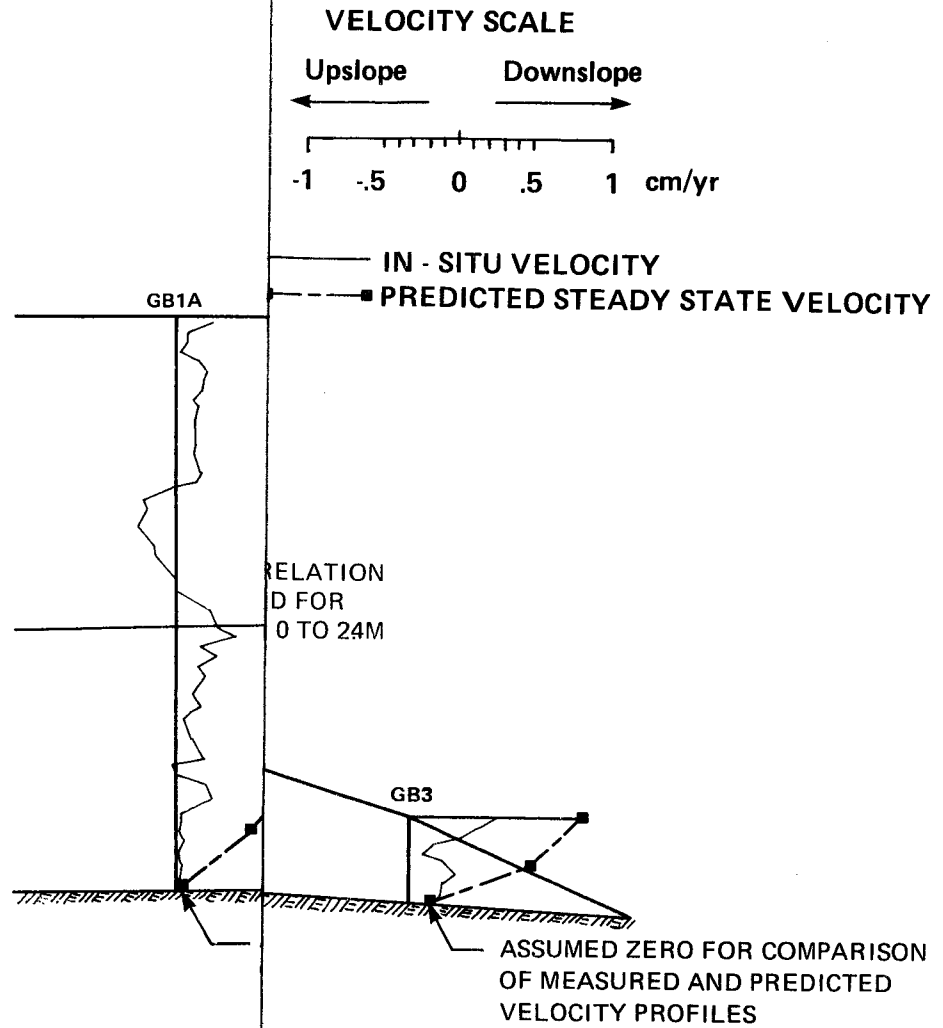


FIGURE 4.5

SITE PLAN OF INCLINOMETER LOCATIONS ON LEFT BANK OF GREAT BEAR RIVER AT PROPOSED ARCTIC GAS CROSSING (Savigny, 1980)





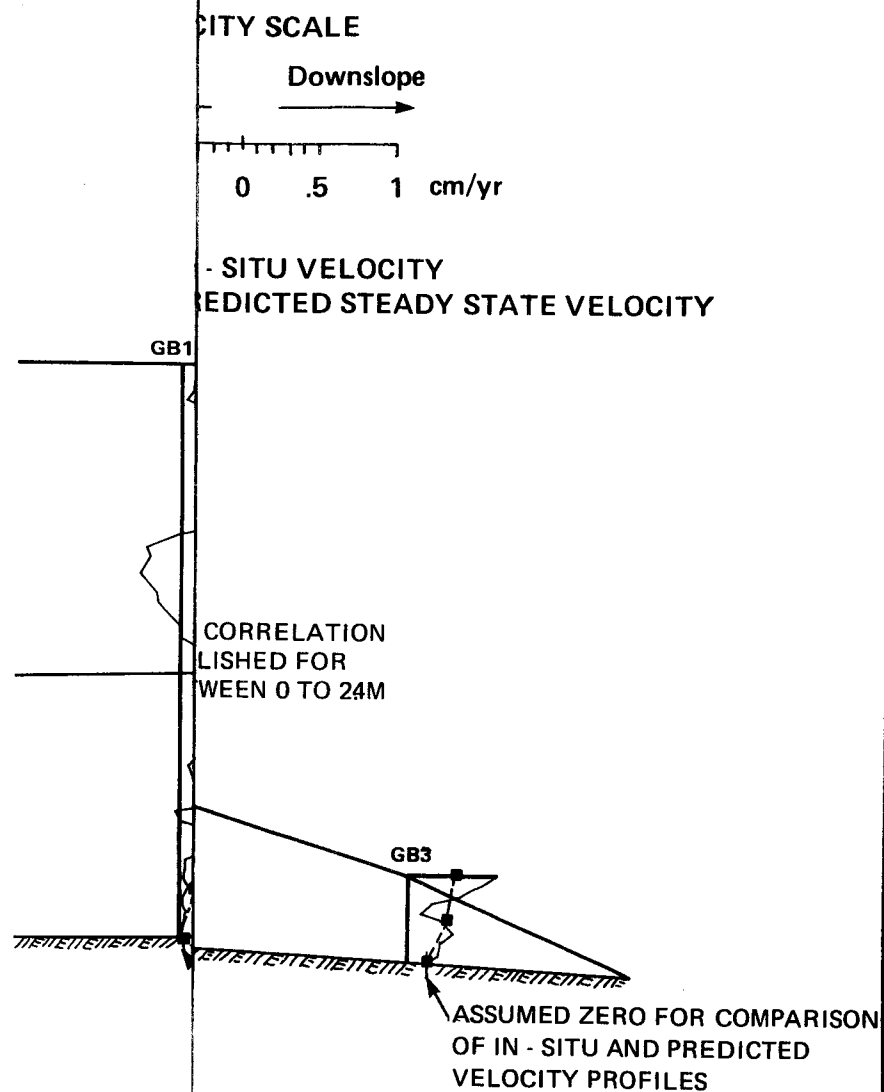


FIGURE 4.10
COMPARISON OF IN-SITU CREEP LAW AND STEADY-STATE CREEP DATA FOR
ICE-RICH GLACIOLACUSTRINE CLAY FROM VARIOUS LOCATIONS IN THE
MACKENZIE RIVER VALLEY, N.W.T.

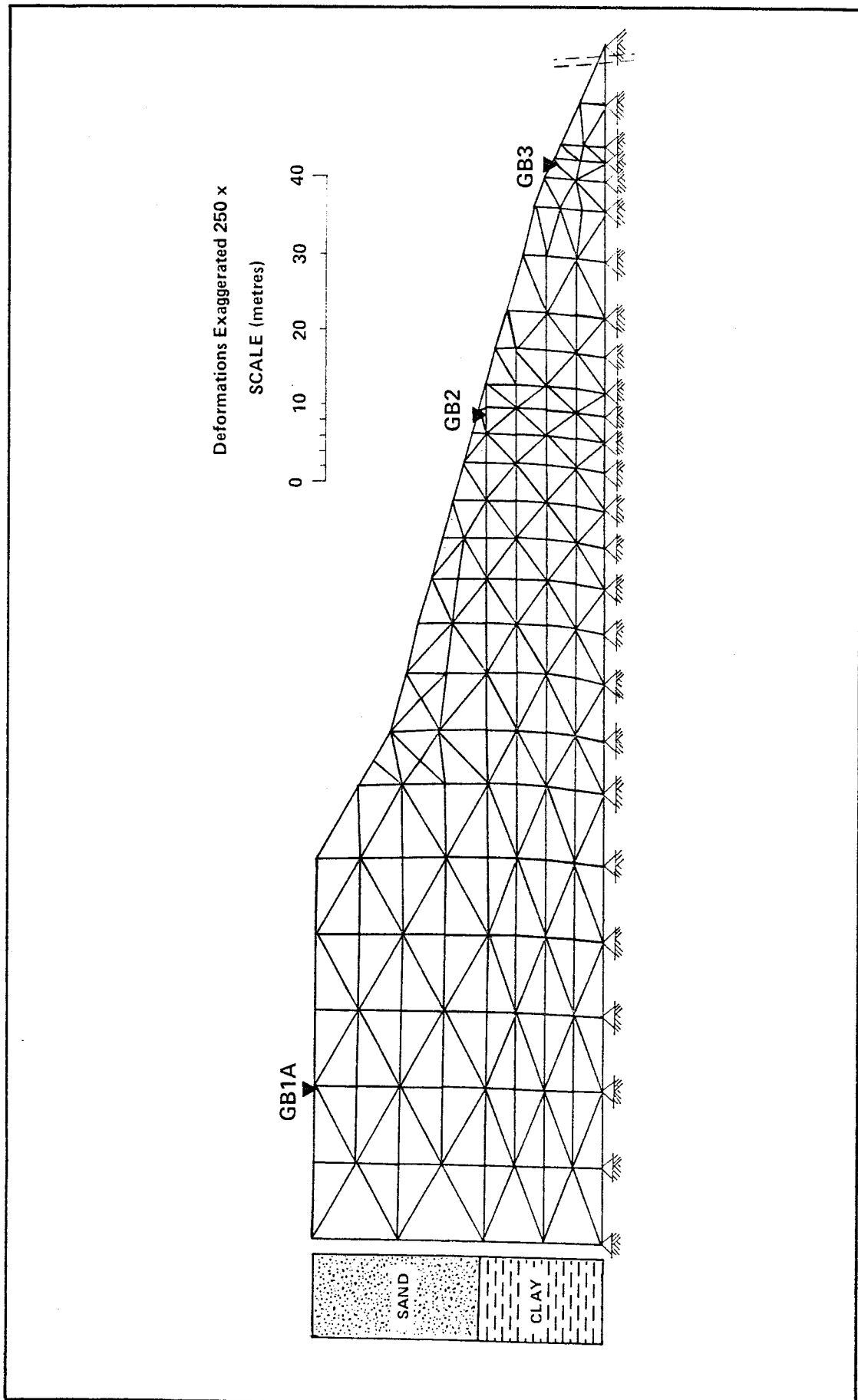
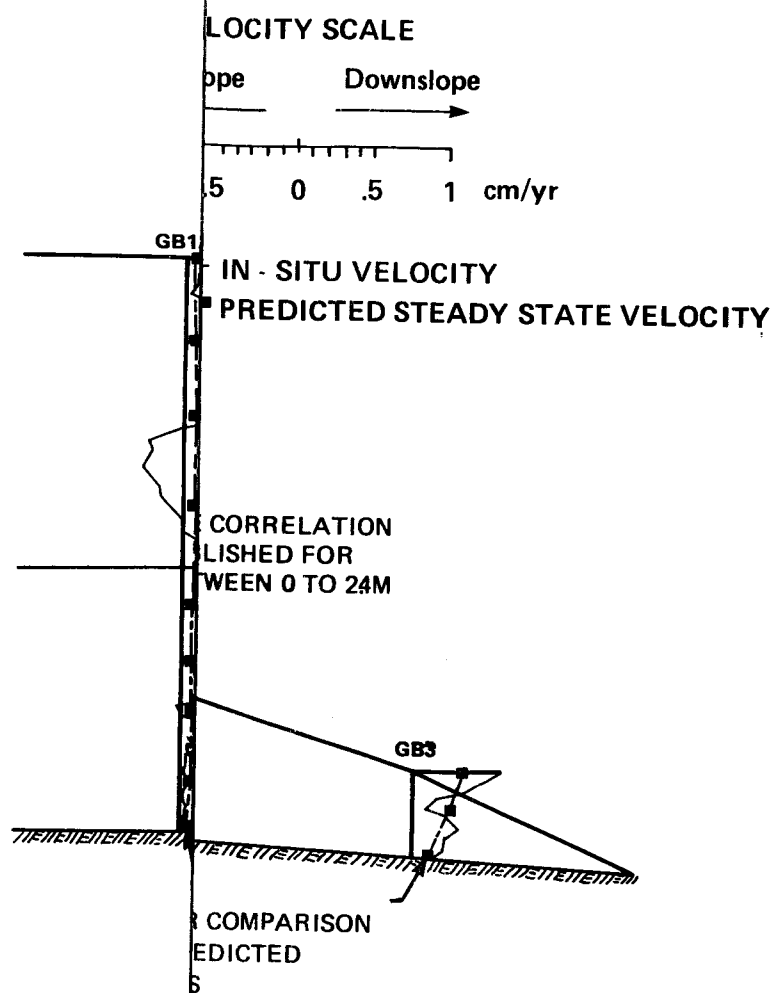


FIGURE 4.11 DEFORMED GRID FOR STEADY STATE CREEP IN SAND AND CLAY



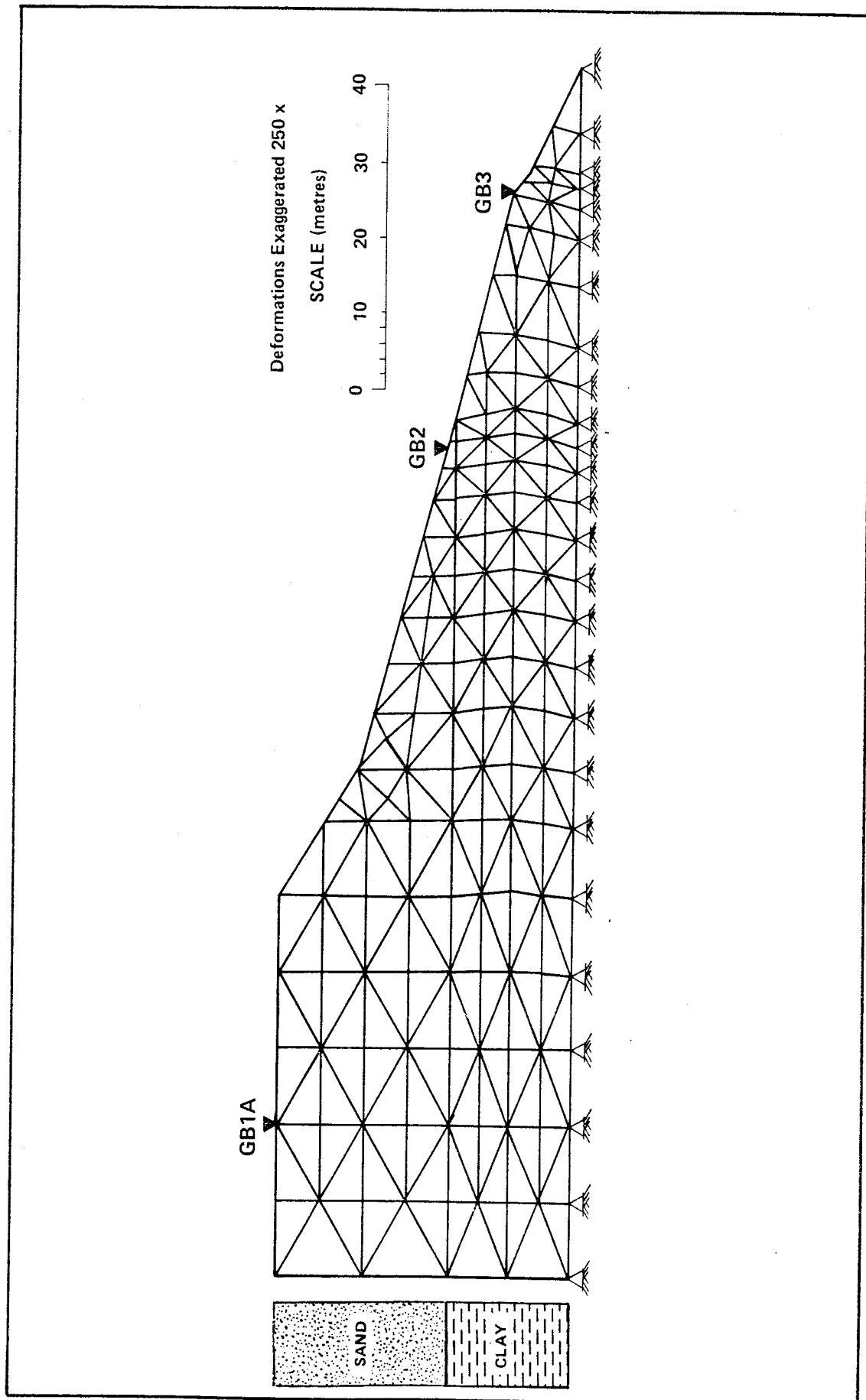


FIGURE 4.13 DEFORMED GRID FOR STEADY STATE CREEP IN CLAY ONLY

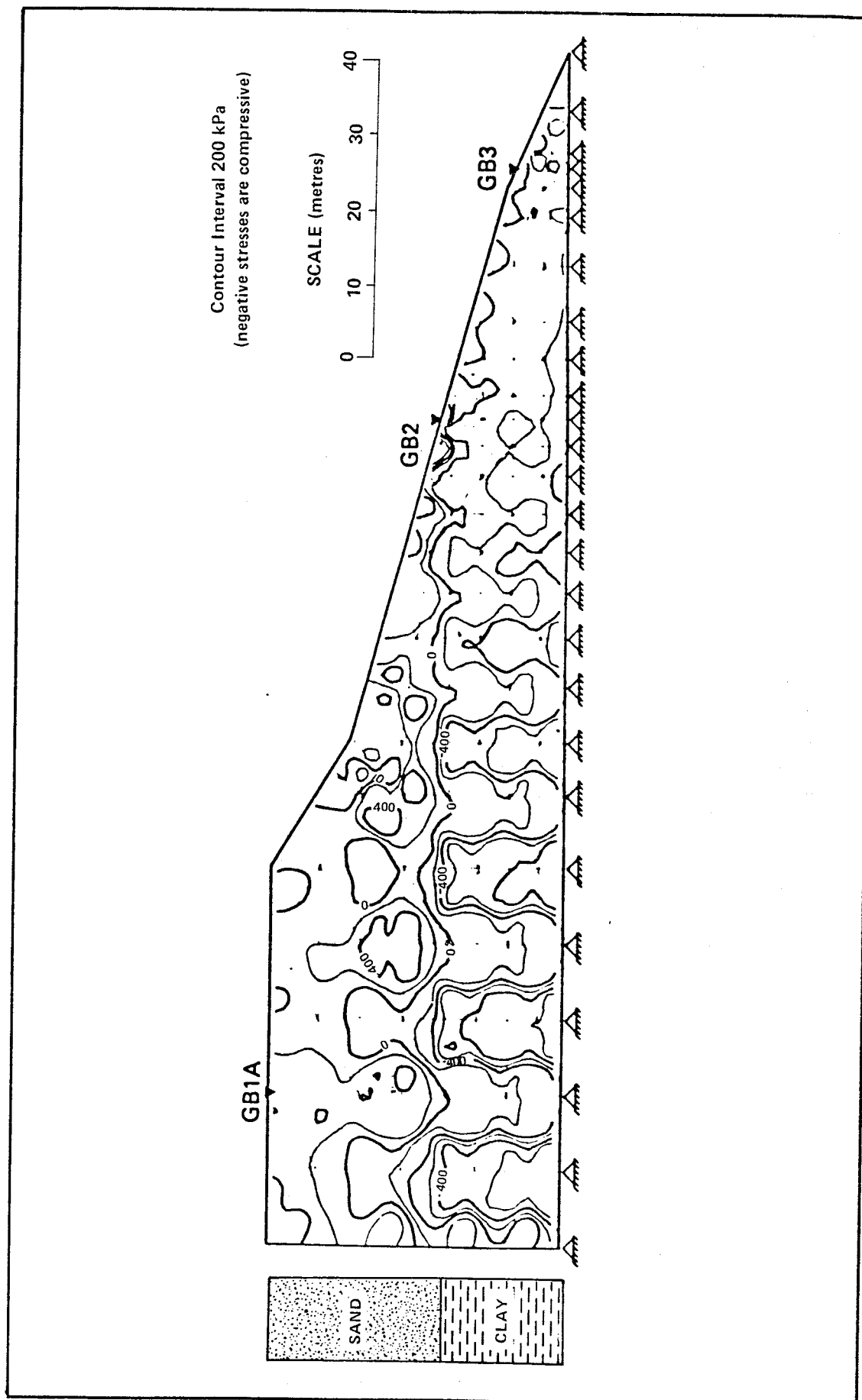


FIGURE 4.14 HORIZONTAL STRESS CONTOURS FOR STEADY STATE CREEP IN CLAY ONLY

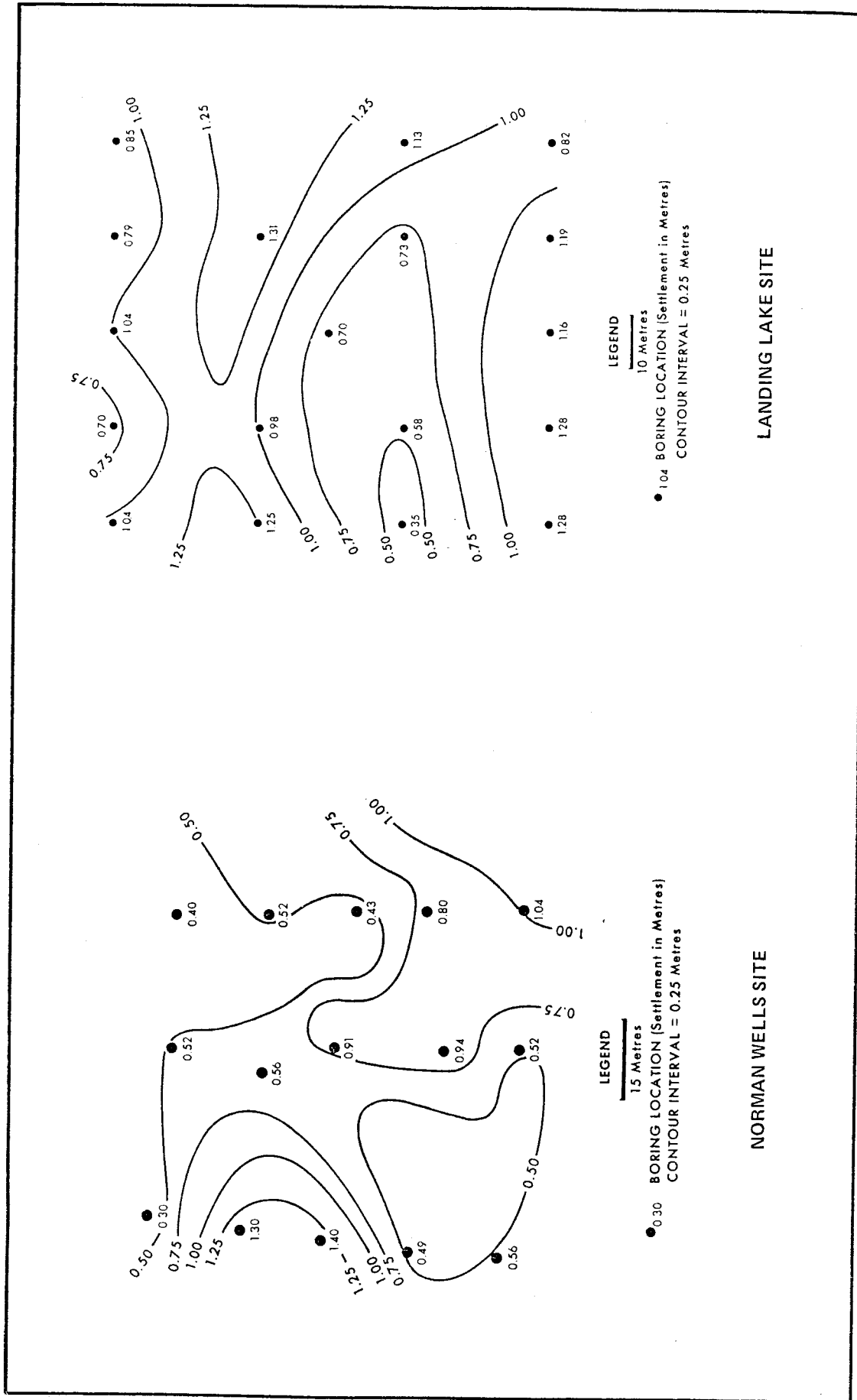


FIGURE 4.15 THAW SETTLEMENT DATA FOR NORMAN WELLS AND LANDING LAKE ICE VARIABILITY TEST SITES (Speer et. al., 1973)

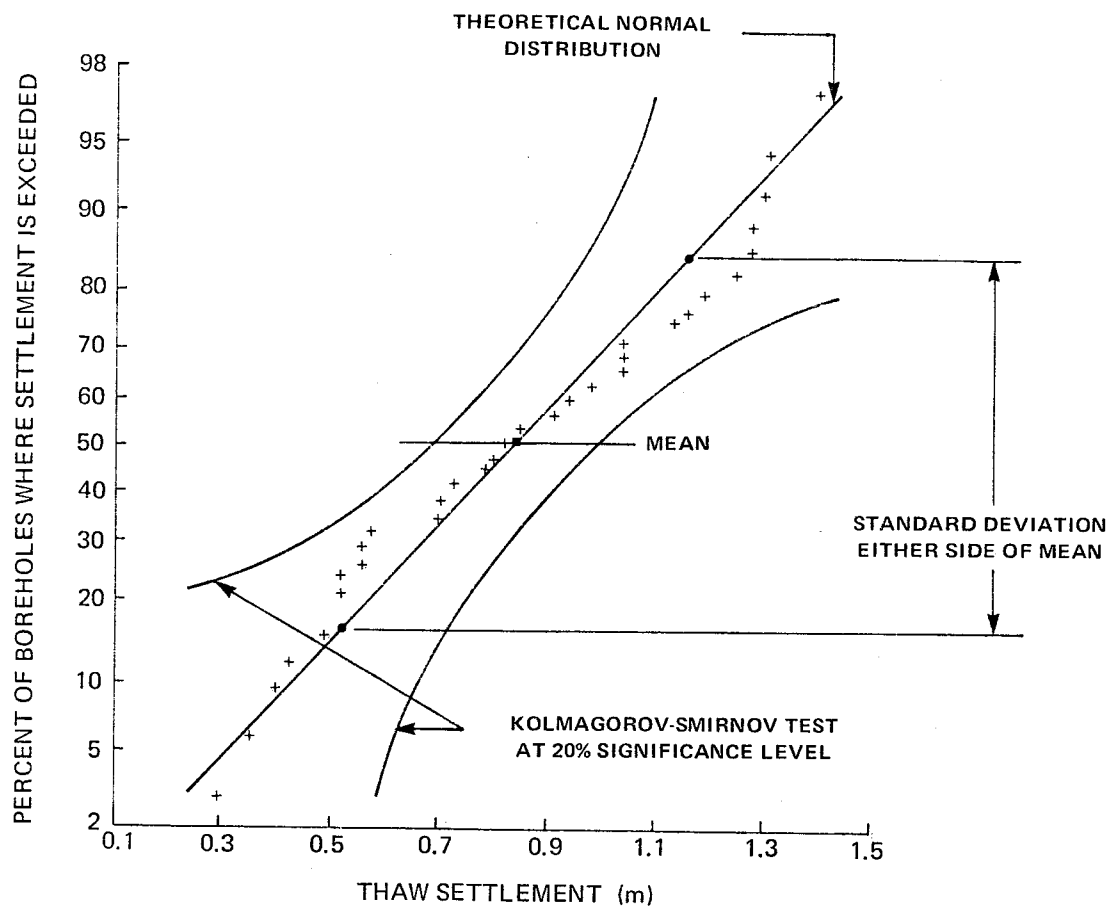


FIGURE 4.16

PROBABILITY RELATIONSHIP FOR THAW SETTLEMENT
IN ICE-RICH LACUSTRINE SILT AND CLAY (Speer et. al., 1973)

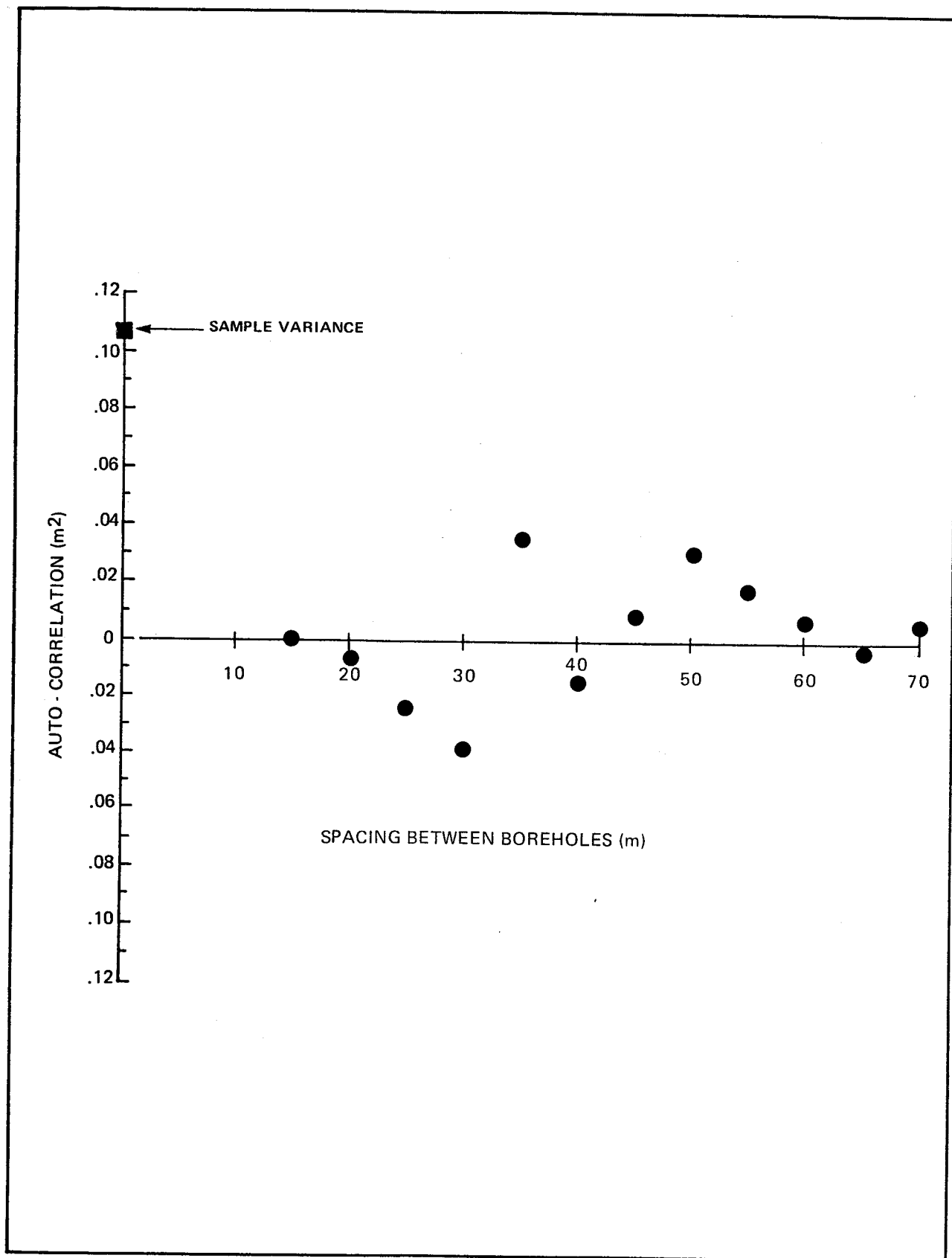


FIGURE 4.17 AUTO - CORRELATION FUNCTION FOR THAW SETTLEMENT IN ICE - RICH LACUSTRINE SILT AND CLAY

CHAPTER V

IN-SITU CREEP ANALYSIS OF FOX TUNNEL

5.1 General

The Fox Tunnel is located approximately 16 km north of Fairbanks, Alaska. The tunnel was excavated into perennally frozen silts. These silts have been of an economic interest in the Fairbanks area since the turn of the century because of the need to remove thick sections of the material to expose the underlying gold-bearing gravels. The old mining practice consisted of stripping the overburden of perennally frozen silt to expose the gravels which were subsequently worked with dredges. The Fox Tunnel excavations were part of two separate research programs from 1966 to 1969. The objectives of these research projects were to investigate methods of subsurface exploration in perennally frozen ground and evaluate subsurface openings as a shelter, storage space and site for military activities.

The first stage of this project was to excavate a tunnel 110 m long in the winters from 1963 to 1966 as shown in Figure 5.1. The excavation was carried out using a continuous mining machine and a modified blasting technique. The tunnel was excavated through ice-rich Fairbanks silt for its entire length. The tunnel portal was excavated into a near-vertical silt escarpment left from old placer mining operations. Closure measurements were undertaken following the mining operations. In November, 1966, a 1.22 m ventilation shaft was drilled from the surface and connected to the end of the tunnel. The ventilation shaft was left open the first winter to allow the heavier cold winter air to circulate through the tunnel by natural convection. The cooler air lowered the tunnel wall temperature from -1.0°C

to -6.7°C (McAnerney, 1968). The lower ground temperature substantially decreased the closure rate of the tunnel. Deformations have been negligible since the ventilation stack was drilled.

The second stage of this project involved excavating a winze from the existing tunnel to provide access to the stratigraphically lower gold bearing gravels and bedrock. This phase of the project is shown as the shaded portions in Figure 5.2. This phase was a cooperative effort between the U.S. Bureau of Mines (USBM) and the U.S. Army Cold Regions Research and Engineering Laboratory (USA CRREL). The winze terminates in a room; designated as Room B by the USA CRREL. A second room USBM Room A, was excavated south of the inclined winze. The winze and two rooms were excavated using air and conventional blasting techniques. Slabbing of the gravel occurred in regions where the thickness of gravel separating the ceiling from the silt was thin. This necessitated that this material be scaled from the roof to ensure a safe working area. Extensive deformation studies were carried out in both rooms (Thompson and Sayles, 1972; Pettibone, 1973). These will be discussed in a subsequent section.

5.2 Site Geology

The sequence of geologic sediments encountered in the main tunnel and winze excavation exposures are shown in Figure 5.3. The disintegrated schist bedrock at the tunnel site is mantled by unconsolidated, ice-cemented silts and gravels. Immediately overlying the bedrock are the gold-bearing gravel deposits. The gravel is capped by a thick re-transported silt deposit.

The floor of both rooms excavated at the end of the winze consist of Birch Creek schist. The bedrock is perennially frozen and deeply weathered

in these exposures. X-ray diffraction analyses performed on samples of the decomposed bedrock just below the gravel/bedrock interface showed a high montmorillonite peak (Sellman, 1972).

A stream deposited gravel layer lies unconformably on top of the bedrock. Stream deposition was inferred from the strong imbrication of the pebbles and cobbles. The average thickness of the gravel deposit in this region is approximately 4 m. Pockets of fine-grained materials are interbedded in the sand. The gravel is perennially frozen and bonded with ice. Pore ice is visible in the voids but the particles retain grain-to-grain contact.

A thick deposit of silt, the uppermost unit at this site, lies unconformably on top of the gravel. Although irregular, the silt-gravel contact is sharp. The silt deposit ranges in age from Illinoian to recent. The Illinoian deposits contain considerably less ground ice and organic matter than the younger units. The overlying silt of Wisconsin age is characterized by large ice wedges. Retransported silts of recent age mantle the top of the upland plain at the tunnel site. Generally, there are two types of ground ice present in the frozen silt; ice lenses and masses formed by ice segregation, and the massive ice structures of Aufeis (buried ice). Ice volumes range from 54 to 79%.

5.3 In-Situ Deformation Studies

The in-situ creep movements of the ice-rich Fairbanks silt were studied independently by Swinzow (1970), Thompson and Sayles (1972) and Pettibone (1973).

The closure rate of the main tunnel was measured by Swinzow (1970) immediately following the final excavation season. No details on the type of instrumentation or the exact time the measurements were recorded are given. Horizontal and vertical closure measurements were recorded at two different locations, Sta. 1+50 and Sta. 3+50 (45.7 m and 106.7 m, respectively, from the portal). The closure measurements recorded at the two tunnel sections are shown in Figures 5.4. The air temperature in the tunnel at the time the readings were recorded was very close to the melting point of ice (McAnerney, 1968). Measurements recorded over a 3200 hour time span showed that the deformation rate continued to attenuate with time. Very little difference was noted between horizontal and vertical closure of the tunnel at both locations. A closure of 14.6 cm was recorded at the instrumented section furthest removed from the tunnel portal. The difference in overburden between these two sections was only 2.1 m. The difference in closure between the two measuring stations is likely due to temperature since the end of the tunnel was consistently warmer during the observation period (Swinzow, 1970). The closure measurements clearly showed that artificial ventilation was inadequate to control the deformations within tolerable limits. In December, 1965 a vertical ventilation shaft was drilled at the end of the tunnel. Cold winter air created by the chimney effect lowered the temperature of the surrounding silt to -6.7°C and -10°C at distances of 0.61 m and 2.35 m from the tunnel wall. The colder tunnel wall temperature has subsequently reduced closure substantially.

McAnerney (1968) measured the vertical deformation of the tunnel at eight deformation points as shown in Figure 5.5(a). The readings plotted for one year commencing in April, 1966, are shown in Figure 5.5(b). The readings recorded by Swinzow in 1965 are superimposed on this plot. Note that the deformation is negligible after the ventilation shaft is opened and cold winter air enters the tunnel by natural convection.

Field observations of the USBM room were reported by Pettibone (1973). The USBM room was initially excavated 4.6 x 15.2 x 2.4 m high. In subsequent weeks 2.3 m was slabbed from each side to widen the room to approximately 9 m. The room was instrumented with vertical deformation points as excavation proceeded.

The instrumentation was designed to measure mass movement independent from parting. Floating points used to measure mass subsidence consisted of 1.8 m long steel rock bolts embedded in the roof in overbored holes. the parting was measured with a 152 mm wood dowel placed adjacent to the rock bolt. A 0.6 mm long rock bolt was permacreted into the bedrock floor directly underneath the rock bolt and wood dowel. Two thermocouple strings were installed in the roof and one in the wall.

The in-situ creep deformation versus time plot for one of the measuring points is shown in Figure 5.6. Deformation readings commenced in early February, 1969. Subsidence of the USBM room increased rapidly as the room was enlarged. Following excavation, the deformation rate dropped substantially. The room temperature up to mid-April was approximately -6.7°C. From mid-April to mid-July, 1969, the room temperature increased to about -3.3°C. The accompanying increase in deformation rate caused some thin slabbing near the back of the room. In early August, 1969, an insulated bulkhead was installed at Sta. 1+20. Following this the temperature increased rapidly and was maintained at approximately -2.2°C. The bulkhead was removed in early January, 1970. During the period the bulkhead was in place, the rate of floor to roof closure continually exceeded the closure between the floor and 1.8 m depth into the roof. Hydraulic props, installed as a safety precaution under portions of the roof that were scaling, continued to pick up load during this period indicating that they were picking up load from overall mass subsidence in addition to the dead load of the slab.

After the bulkhead was removed, the closure rate has decreased to approximately 1.5 cm/yr. This is consistent with the observations in the CRREL room. Colder average temperatures for the duration of observations in the USBM room resulted in less closure in this room than the CRREL room. In all, 30 cm of floor to roof deformation was recorded in 37 months with parting accounting for approximately 20% of total room closure. Several cracks were observed in the roof in April, 1970. Pettibone (1973) concludes that vertical separation and expansion of the formation account for the measured parting.

Thompson and Sayles (1972) reported the results of a finite element simulation of the closure of the USA CRREL room in the Fox Tunnel. Horizontal and vertical closure measurements were reported for one year. Laboratory samples were obtained from uniaxial creep testing.

The authors used the simple power law to predict the in-situ behavior of the room. They concluded that the form of the flow law predicted the in-situ behavior quite well. However, the back-calculated flow law predicted a creep rate 3.3 times faster than had been established in the laboratory on the same frozen silt. A few comments can be made regarding the analysis carried out by Thompson and Sayles (1972).

A two dimensional finite element analysis was used to model the room. The actual field configuration was three-dimensional. It is common practice to assume that sections within three tunnel diameters will be influenced by the end-restraint offered by the tunnel face. Both of the instrumented sections in this room are within this distance. The end effects of the room would delay the onset of in-situ steady-state creep.

There is a slight error involved in the boundary condition assumed along the gravel-bedrock interface. In their analysis, Thompson and Sayles (1972) assume the base of the gravel is free to translate in the horizontal direction due to a lean clay that is present immediately below the gravel-bedrock contact. The decomposed bedrock exposed at the base of the gravels in this room are perennially frozen and no excess segregated ice features were observed (Sellman, 1972). Thus, the clay would offer some frictional resistance which would decrease the predicted creep rate.

The form of the constitutive relationship of the placer gravels is questionable. The initial finite element simulations indicated that in-situ behavior of the room could not be predicted assuming constitutive models which exhibit continuously decreasing creep rate with time (Thompson and Sayles, 1972). The simple power law for the gravel was arrived at by a best fit to the actual field curves. The average density and moisture content of the gravel is 2.08 Mg/m^3 and 10%, respectively. Sellman (1972) described the gravel as retaining grain-to-grain contact with no massive ice forms present.

In light of the review of the constitutive behavior of frozen soils presented in Chapter 2, it would appear that the behavior of the frozen gravel will be governed by frictional response. Swinzow (1970) reported that a room excavated in cold, densely packed, permafrost till did not deform in excess of measurement error after 3 years of operation. Thus, depending on temperature, the in-situ deformation behavior of the gravel may be more accurately represented by elastic behavior or a creep law which is a function of confining pressure.

Examination of the laboratory creep data of the Fairbanks silt is necessary to establish whether a steady state condition did exist. Unconfined compression creep tests were performed on undisturbed samples of frozen Fairbanks silt obtained from the tunnel exposure. The test temperature was -1.7°C and the stress levels ranged from 245 kPa to 2100 kPa. A simple power law was fitted to the laboratory creep data with the following parameters:

$$\dot{\epsilon} = 1.16 \times 10^{-9} \sigma^{-4} \quad (5.1)$$

where the units for strain rate and stress are year^{-1} and kPa, respectively.

Thompson and Sayles (1972) reported strains in the laboratory typically in excess of 40% in 8 days. Tests performed by Savigny (1980) on ice-rich fine-grained glaciolacustrine clay showed that compatibility at the soil-ice interface would not exist at strains of this magnitude. Thus, the minimum strain rates measured in the laboratory may well exceed the steady-state rate and the exponent of 4 is considered high.

A more detailed investigation of the U.S. Army CRREL Room was undertaken to gain further insight into the in-situ creep behavior of the frozen silt.

The horizontal and vertical closure of the USA CRREL room were monitored at two instrumented test sections, Sta. 1+83 and Sta. 2+02. The horizontal and vertical room deformations at each section were recorded between four closure points installed in the roof, walls and floor. Each point consisted of a 32 mm diameter wooden dowel embedded 0.3 m into the permafrost. A steel tape in conjunction with a steel ruler with an accuracy of 1.6 mm (1/16 inch) was used to measure the closure between opposing

points in the walls, and roof and floor. In-situ ground temperatures were monitored by four strings of thermocouples installed at Sta. 1+96. The thermocouples were installed in the roof, floor and each side wall. The ground temperatures were recorded at 0.3 m intervals from the walls and roof to a total depth of 1.83 m.

The closure and temperature instrumentation could not be installed for three weeks after the room was excavated. An insulated bulkhead was installed at the entrance to the room (Sta. 1+81) approximately eight weeks after the initial closure and temperature readings were recorded. The purpose of this bulkhead was to maintain a constant room temperature thus minimizing the effect of seasonal fluctuations in air temperature on the permafrost temperature. Approximately 18 weeks later, a second bulkhead was installed by the USBM at Sta. 1+20 and subsequently removed five months after installation.

The vertical and horizontal closure measurement recorded at Sta. 1+83 and 2+02 in the USA CRREL room are shown in Figures 5.7 and 5.8, respectively. The in-situ ground temperatures measured in the roof and walls are also plotted in Figures 5.7 and 5.8. The vertical and horizontal closure rate determined from the closure versus time data (Figures 5.7 and 5.8) is shown in Figure 5.9. The closure observations were recorded for a period of 3.65 years.

The vertical and horizontal closure measurements increased rapidly during the first year of observations. This is probably due to the general warming trend in the surrounding permafrost as indicated by the temperature measurements. After removal of the USBM bulkhead at Sta. 1+20, the rate of closure decreased as a consequence of the cooler permafrost temperatures. The greatest vertical and horizontal movements were recorded at Sta. 2+02.

After 3 1/2 years of observations, 48 cm and 16.5 cm of vertical and horizontal movement, respectively, had taken place.

The temperature dependence on in-situ creep behavior is clearly illustrated by the vertical closure at Sta. 1+83. The first bulkhead was installed within .75 m of this instrumented section. The colder ambient air temperature just outside of the room would lower the permafrost temperature at Sta. 1+83. This would explain the slower vertical deformation rate at Sta. 1+83. From January, 1970 to October, 1972, the deformation readings indicate a steady-state closure rate at Sta. 1+83 at a velocity of 1 cm/yr. For the duration of this time interval the permafrost temperature has been maintained at an average value of -2.2°C .

The vertical closure at Sta. 2+02 shows much less sensitivity to changes in permafrost temperature brought about by installation and removal of the USBM bulkhead. At this section, the closure rate continued to attenuate with time for the entire period of time closure observations were recorded. During the final ten months of observations, the closure rate at Sta. 2+02 was 2.5 cm/yr.

The horizontal closure at Sta. 1+83 and 2+02 continued to accelerate for the first year movements were recorded. Following removal of the USBM bulkhead, the horizontal movements at Sta. 1+83 fluctuated seasonally with 1.3 cm of net closure being recorded for the final 2.5 years of observations. During the warm summer months and early fall, 1 to 2 cm of horizontal movement was recorded. Very little or negative closure was recorded during the cold winter months. Attenuation of the horizontal closure rate at Sta. 2+02 was observed for 1 year following removal of the USBM bulkhead. Seasonal fluctuations have subsequently taken place. One

centimetre of net horizontal closure was recorded during the final 1.7 years of observations.

A finite element simulation of the USA CRREL room was carried out using the programme CREEP. The purpose of this simulation was to assess the validity of the simple power law for ice as an upper bound solution for the in-situ creep deformations recorded in the frozen silt while the frozen placer gravels were treated as an elastic material.

The discretized continuum of the CRREL room used for this finite element simulation is shown in Figure 5.10. Only the Fairbanks silt and placer gravels were discretized. It was assumed that the competent bedrock floor would not contribute to vertical closure of the room. Only one half of the tunnel was analyzed due to symmetry. The thickness of overburden above the roof centerline was 16.9 m. The ground surface was assumed stress free. The two vertical boundaries were fixed by roller constraints which allowed movement in the vertical direction only. The base of the gravel was placed on roller constraints which allowed movement in the horizontal direction only as was done in finite element analysis carried out by Thompson and Sayles (1972). The elastic material properties of the frozen silt and gravel are given in Table 5.1.

The Young's Modulus values and unit weights were obtained from Thompson (1970). Creep properties of the silt and gravel were assigned the values of pure ice at -2°C (Morgenstern et al., 1980), i.e.:

$$\dot{\epsilon} = 2.0 \times 10^{-8} \frac{1}{\sigma^3} \quad (5.2)$$

where the units of stress and strain rate are kPa and yr^{-1} , respectively.

TABLE 5.1
ELASTIC MATERIAL PROPERTIES FOR
FROZEN SILT AND GRAVEL
FOX TUNNEL, ALASKA

PROPERTY	SILT	GRAVEL
Young's Modulus	346 MPa	718 MPa
Poissons' Ratio	.495	.450
Unit Weight	1.47 Mg/m ³	2.08 Mg/m ³

The nodal point labelled V in Figure 5.10 was chosen to represent the in-situ vertical closure points. The vertical displacement of node V was assumed to represent the total vertical closure since the closure point installed in the floor was fixed into bedrock.

The vertical steady-state velocity predicted by the programme CREEP using the simple power law given by equation 5.2 to model the creep behavior of the frozen silt was 1 mm/yr. The accumulated closure of the node point labelled H in Figure 5.10 after 2.5 years was 11 mm. The horizontal

displacement of node H was assumed to be one half of the total horizontal closure. The predicted results are at least one order of magnitude less than the in-situ measurements. The design of the instrumentation installed in the CRREL room was such that vertical separation in the ceiling could not be measured independently of total floor to ceiling closure. In the USBM room, 20% of the total room closure was due to parting in the ceiling. Thus, a portion of the total closure measured in the CRREL room is likely due to vertical separation in the ceiling.

In this case, the simple power law for ice does not represent a valid upper bound for the in-situ creep behavior of the frozen silt. However, examination of the closure rate versus time data presented in Figure 5.9 clearly shows that steady-state conditions did not exist during the first year of observations as assumed by Thompson and Sayles (1972). It is entirely possible that the overall closure measurements that were recorded consisted of both plastic and creep flow. Plastic flow in this sense refers to time-dependent failure of the frozen soil.

5.3 Assessment of In-Situ Deformation Behavior

The in-situ deformation studies of the Fox Tunnel reported by Swinzow (1970), Thompson and Sayles (1972) and Pettibone (1973) all showed that the deformations in the ice-rich Fairbanks silt became excessive when the room air temperature approached the ambient soil temperature of approximately -1.7°C . In each case, the room temperature was lowered to restrain the movements within tolerable limits.

The stress release in the ceiling of the USA CRREL room and the USBM room initiated by excavating the underlying frozen placer gravels is similar to that occurring in a soil mass located above a yielding trap door. In the

past, it has been recognized that the behavior of a soil immediately above a flexible underground structure and that of a granular mass above a yielding bottom of a silo are equivalent for all practical purposes (Ladanyi and Hoyaux, 1969). This assumption allows arching or bin theory to be used to estimate the contact pressures between the ground and an underground structure.

The general expression for the pressure acting on the roof of a tunnel derived from arching theory is given by:

$$\sigma_v = \frac{B(\gamma - c/B)(1 - e^{-K \tan \phi D/B})}{K \tan \phi} \quad (5.3)$$

where: σ_v = vertical contact pressure
 B = half width of the zone of arching
 (half width of the underground rooms in this case)
 D = depth of the roof below the surface
 γ = unit weight of soil above the roof
 $c + \phi$ = shear strength parameters of the soil above the roof
 K = ratio between vertical and horizontal stress

It is clear from the above expression that an underground opening located in a cohesionless material will not remain stable without additional ground support provided by a tunnel liner or timber lagging. For an unlined opening in a cohesive material, the following relationship must be satisfied:

$$c > B\gamma \quad (5.4)$$

If the cohesion is less than γB , the roof of the underground opening must be supported. For the USA CRREL room, values of $B = 2.44$ m and $\gamma = 1.47$ Mg/m³ require a minimum cohesion of 35.1 kPa to ensure stability of the room.

Vialov (1962) proposed that the strength of a frozen soil can be represented by a series of Mohr-Coulomb failure envelopes, each one corresponding to a specific time to failure. The relationship between maximum shear and normal stress takes the form:

$$\tau(t) = c_t + \sigma_n \tan \phi_t \quad (5.5)$$

where: $\tau(t)$ = shear strength at time = t

σ_n = maximum normal stress on shear plane

c_t, ϕ_t = cohesion and angle of internal friction, respectively
(functions of time and temperature)

The time variation of the cohesion intercept, c_t , can be described by

$$c_t = B / \ln(t_f/B) \quad (5.6)$$

where: B, B = experimentally determined material parameters

t_f = time to failure

Time and temperature have very little influence on the angle of internal friction.

The laboratory studies carried out to date indicate that the long term strength of frozen soil is frictional. The angle of internal friction can

well be approximated by the effective friction angle for the same soil in an unfrozen state. For coarse-grained soils, the friction angle is relatively independent of temperature and strain rate. Frictional resistance is hindered by the presence of ice in ice-rich frozen soils. In very ice-rich soils which have very little grain-to-grain contact, the strength will be that of the ice and in the long term will approach zero.

The general nature of the failure envelopes for frozen soils is nonlinear. The time variation of the cohesion and angle of internal friction is reflected by the mutual distribution of each envelope. The inclination of each envelope represents the time variation of the angle of internal friction. The loss of cohesion with time is represented by the cohesion intercept of each envelope.

In predicting the time-dependent settlements for foundations in frozen soils, Ladanyi (1981) recognized that creep and consolidation can occur simultaneously in frozen soils containing large amounts of unfrozen water. Although the treatment of creep and consolidation as two distinct phenomena would seem appropriate for predicting delayed settlements in frozen soils, experimental limitations necessitate that these two phenomena be lumped together. Thus, frozen soil is treated as a quasi-single-phase medium with an empirically derived creep relationship. Recognizing that under certain conditions consolidation volume change may have a significant impact on the delayed settlement of a footing, Ladanyi (1981) studied the frozen soil's response to a stress increment with the aid of the Rendulic plot. These plots conveniently enable any given stress increment in a triaxial test to be separated into its hydrostatic and deviatoric components.

Interaction between strength and deformation is illustrated by considering a triaxial test on a normally consolidated sample of frozen soil

containing three phases: mineral particles, ice and unfrozen water. The sample is initially in equilibrium at point O', Figure 5.11. A significant difference between frozen and unfrozen soil is that a total stress increment, $\Delta\sigma$, (O'A), can be applied to a frozen soil exceeding its long term strength. As the soil is allowed to deform under closed system conditions, the stresses are internally shared by the mineral particles and the pore filling. The unfrozen water can only support hydrostatic pressure. The shear stresses are supported by the mineral phase and temporarily by the ice. Since the ice phase supports a portion of the shear stress, the effective strength is mobilized only to point B'. The hydrostatic pressure generated by this straining, assumed to be equal in the ice and unfrozen water, is given by:

$$\Delta u_i = \Delta\sigma_{oct} - \Delta\sigma'_{oct} \quad (5.7)$$

This is somewhat less than for the same soil in an unfrozen state. Similarly, the shear stress, $\Delta\tau_{oct,i}$, acting in the ice phase is given by:

$$\Delta\tau_{oct,i} = \Delta\tau_{oct} - \Delta\tau'_{oct} \quad (5.8)$$

where: $\Delta\tau_{oct}$ and $\Delta\tau'_{oct}$ are the total applied shear stress increment and effective shear stress assumed by the soil, respectively.

Maintaining closed system conditions, the frozen soil will creep and the ice will gradually transfer its applied shear stress to the soil skeleton. The point B' will move along the effective stress path to mobilize full soil strength at B. Creep brings the ice closer to failure producing a loss of

strength with time. The failure surface shrinks from its initial position at $t=0$ until failure occurs at point A at $t = t_A$. Since point A lies beyond the long term strength, creep failure will inevitably occur. The stresses at failure will be $\Delta\tau'_{oct}$ and $\Delta\sigma'_{oct}$ in the soil skeleton, Δu_i in the ice and unfrozen water and $\Delta\tau_{oct,i}$ in the ice.

Opening the system at point B' will initiate consolidation following the stress path B'A'. The total deformation of the soil will consist of consolidation and creep, occurring simultaneously. Full soil strength mobilization at point A' will depend on the rate of consolidation relative to the steady state creep rate.

The foregoing example of a triaxial specimen clearly showed that at least for short term conditions, the strength of the ice matrix can be relied on to support stress increments that exceed the soil's long term strength. However, the shape and position of the long term strength envelope must be known to clearly define the long term response of a frozen soil to an arbitrary stress increment.

In using the quasi-single phase approach, it is important that the stress paths followed in the laboratory to develop the constitutive equation coincide with those expected under field conditions. The stress paths shown in Figure 5.11 are for an axially symmetric stress state which would be suitable for calculating the time dependent response of the frozen soil underneath a circular footing. Under conditions of plane strain, such as the stress state in the soil mass surrounding an underground opening, the use of a simple stress space in which the coordinates are the vertical and horizontal principal stresses would be more appropriate. This arises from the fact that the intermediate principal stress is generally unknown.

The stress paths shown in the Rendulic plot, Figure 5.11, have been replotted in a vertical and horizontal stress space as shown in Figure 5.12. In this plot, the stress paths for undrained extension tests are also shown. The deviatoric stress is given by either the vertical or horizontal distance between the point characterizing the stress state and the hydrostatic pressure line, $\sigma'_V = \sigma'_H$.

Although a clear understanding of effective stress changes in a frozen soil during shear has not been developed to date, Ladanyi's (1981) conceptualization provides a useful framework to study the response of a frozen soil to a stress increment. The interaction between strength and deformation will shed some light on the in-situ deformation behavior of the CRREL room in the Fox Tunnel.

Excavation of an underground opening will initiate deformation in the soil mass surrounding the opening. Considering the CRREL room at the Fox Tunnel, the ice-rich silt comprising the roof will deform much more readily than the frozen dense gravel forming the walls. Downward movement of the unsupported ceiling must now be resisted by shear stresses that develop at the interface between the yielding and stationary soil mass.

Before excavation of the USA CRREL room, the vertical stress along the ice-rich silt/gravel contact was equal to the overburden pressure. After excavation, the vertical stress along the ceiling of the room is zero. This stress release is accompanied by a corresponding increase in stress in the adjoining soil mass above the frozen gravel. Considering the two elements shown in Figure 5.13(a), the stress in element A will be a small fraction of what it was before excavation commenced. The vertical pressure in element B will have to increase by the same amount to maintain equilibrium. The horizontal stress these two elements exert upon each other must remain equal

and opposite. The induced shear stress at the interface between elements A and B will cause a rotation of the principal stresses above the tunnel supports.

Although the actual stress paths followed by the soil elements surrounding a rectangular opening are quite complicated, some insight into the interaction between strength and deformation of frozen soil may be gained by studying the simple stress paths shown in Figure 5.13(b). The stress release in element A is given by the stress path O'A while an equal but opposite stress increment is applied to element B, ie. path O'B. As shown in the figure, the stress release in element A exceeds the long term $\sigma_H > \sigma_V$ failure line. However, an equal stress increase in element B lies below the long term $\sigma_V > \sigma_H$ failure line.

This simplified stress path description clearly shows the importance of considering the stress path in selecting an appropriate constitutive relationship for predicting the initial time dependent response of the frozen soil surrounding an underground opening. Since the stress increase in the soil above the gravel does not exceed the soil's long term strength, the creep behavior may well continue to attenuate with time. On the other hand, the soil above the roof of the Tunnel that experiences a stress release that exceeds the soil's long term strength will creep at either a steady state or decreasing rate. The creep rate will depend on the relative position of the stress point A with respect to the service life strength, envelope for the frozen soil under consideration. Failure will inevitably occur for the purely frictional long term strength shown in Figure 5.13(b).

Evidence of the frozen silt approaching failure can be inferred from the vertical closure rate versus time plot at Sta 1+ 83. Installation of the USBM bulkhead at Sta 1 + 20 increased the permafrost temperature at Sta.

1+83. The warmer soil temperatures resulted in a 50% increase in vertical closure rate at Sta. 1+83 over a period of three months. A portion of the increase is due to the temperature dependence of creep. However, the warmer soil temperatures will increase the unfrozen water content of the frozen silt. Increasing a frozen soils unfrozen moisture content will shrink the delayed failure surface from its original position (ie. instantaneous or short term strength) towards the long term failure line. The soil temperature becomes cooler following removal of the USBM bulkhead at Sta. 1+20. As the soil temperature decreases, its failure surface expands towards the instantaneous strength. Thus, a numerical prediction of the in-situ deformation behavior of the CRREL room would have to account for temperature dependent creep properties and a temperature and time dependent failure criterion governing the strength of the frozen soil.

The vertical deformation instrumentation installed in the CRREL and USBM rooms at the Fox Tunnel was not able to measure vertical deformations in the gravel walls independently of the overall vertical closure of the room. The arching support provided by the overlying silt (due to the stress release above the roof) will increase the vertical stress in the frozen gravel walls in the vicinity of the room. The increase in vertical stress coupled with a horizontal stress release brought about by excavating the room will significantly increase the shear stress in the frozen gravel. The build up of shear stress in the frozen gravel walls will lead to time dependent deformations occurring in the gravel that may eventually lead to delayed failure.

Thompson and Sayles (1972) incorporated a yield criterion in their finite element program to account for plastic yield of the material. In their analysis, the effective stress as defined by Odqvist (1966) was used as the yield criterion. After each creep increment, the effective stress for each element was checked to determine whether it had exceeded a given stress level. Plastic deformations were allowed to take place in all elements that exceeded the given yield stress. It is interesting to note

that when the in situ creep behavior of the frozen silt and gravel was governed by primary creep, plastic yielding was initiated in the gravel walls. The elastic-plastic analysis carried out by Thompson and Sayles (1972) indicate that a catastrophic failure would occur which lead the authors to conclude that the long term strength of the frozen silt and gravel was zero and that the in situ creep behavior was more accurately represented by a steady-state creep relationship.

The use of effective stress for a failure criterion is over-simplified. Numerous studies have shown that the strength of frozen soil is governed by a Mohr-Coulomb failure law where the shape of the failure envelope depends on soil type, density, ice saturation, temperature and strain rate. The strength of the gravel will most certainly be governed by a Mohr-Coulomb failure envelope.

It is possible that the frozen gravel was approaching the failure envelope and experiencing large vertical strains during the period of time the soil temperatures increased while USBM bulkhead was in place at Sta 1+20. Thus, vertical displacements in the gravel walls may be superimposed on to vertical creep displacement in the frozen silt to give the total room closure as recorded by Thompson and Sayles (1972). The strain rate measured in the laboratory is independent of the vertical deformations, in the gravel walls. This is very likely why the creep equation which gave the best fit to the observed room closure yielded a strain rate 3.3 times greater than the strain rate measured in the laboratory.

5.5 Underground Circular Cavities in Permafrost

Major design considerations of tunnels in frozen ground are stand-up time, change in diameter of the unlined opening with time and the change in pressure on the tunnel liner with time. The first two factors will

influence the decision on whether or not to support the underground opening with a tunnel liner. The third factor will provide guidelines for dimensioning the tunnel liner. A small parametric study of shallow circular tunnels in frozen ground was carried out to study closure phenomena and the stress changes in the frozen soil surrounding the opening.

The first part of this study deals with the closure behavior of an unlined circular opening excavated at various depths in warm ice-rich permafrost. The warm ice-rich permafrost was chosen as the host soil medium as it would present an upper bound for the closure rates. A circular opening was chosen since it is more amenable to simpler analysis under certain conditions.

The height to diameter (H/D) ratios of the circular openings varied from 1.5 to 4.0. The finite element mesh for H/D=2.0 is shown in Figure 5.14. A homogeneous soil profile was assumed for all analysis. Material properties were assigned on the basis of data reported in the literature.

The flow law for frozen soil used in this study was:

$$\dot{\epsilon} = 2.0 \times 10^{-8} \sigma^3 \quad (5.9)$$

This flow law is identical to the flow law for ice at -2°C (Morgenstern et al., 1980) and represents a valid upper bound to the in-situ creep behavior of fine-grained, ice-rich frozen soils of the Mackenzie River Valley, N.W.T.

A plot of the closure velocity versus depth of overburden for the crown, springline and invert of the tunnel is shown in Figure 5.15. The depth of overburden in this figure is expressed non-dimensionally as the H/D

ratio. For all three points considered, the closure rate increases with the depth of overburden. Also, the downward movement of the crown is greater than the upward movement of the invert creating a sag in the roof. Thus, for shallow tunnels, the initial circular cross section undergoes a transition to an oval shape as the frozen soil creeps inward. At greater depths, this transition will be less pronounced because the relative change in overburden stress from crown to invert will be smaller. Based on these preliminary results, it would appear that a circular opening located 9 to 10 diameters below the surface would behave axisymmetrically. This is somewhat deeper than a circular opening in an elastic soil or rock.

Velocity vectors depicting magnitude and direction throughout the region analyzed for $H/D=2.0$ and $H/D=3.0$ are shown in Figures 5.16 and 5.17, respectively. The relative magnitude in these figures is given by the length of the line segment. For both tunnel depths, the frozen soil flows in towards the opening. The transformed cross section is shown on a grossly exaggerated scale by the dashed line joining the velocity vectors around the periphery of the opening.

Contours of effective strain rate for $H/D=2.0$ and $H/D=3.0$ are shown in Figures 5.18 and 5.19, respectively. Although the magnitude of individual strain rate components are not given by the effective strain rate, the contours serve the purpose of indicating where the high strain rate gradients exist. The greatest effective strain rate gradients directly above the crown of the opening and diminish as one sweeps around from the crown to the invert. Higher gradients exist for the tunnel located at $H/D=3.0$ (see Figure 5.19) indicating that the disturbing influence of the ground surfaces decreases with deeper tunnels. Both Figures 5.18 and 5.19 show the high strain rate gradients extend only 1.5 to 2 tunnel diameters above the crown of the opening. This zone of influence will decrease with deeper tunnels.

It was clear from the foregoing analyses that shallow circular cavities located in warm ice-rich permafrost would become unstable unless some measures were undertaken to arrest the deformations. The second part of this study dealt with the closure of a lined circular opening in permafrost. Only one opening located 2 diameters below the surface was analyzed. The same finite element mesh as shown in Figure 5.14 was used except that a narrow ring of elements was added to the periphery of the opening to simulate a 100 mm thick concrete liner. The geometry is shown in Figure 5.20.

This analysis was carried out by calculating the elastic gravity stress field surrounding the circular opening without the liner elements present. After this step, the elements representing the tunnel liner were added to the geometry and the creep solution was initiated. It was felt that this sequence would be more representative of actual construction operations.

The velocity vector diagram for the lined tunnel placed 2 diameters below the ground surface is shown in Figure 5.21. As shown in the figure, the tunnel liner plays a significant role in altering the velocity distribution in the frozen soil surrounding the opening. The overall movement of the soil mass and tunnel opening is downward due to gravity stress. In the vicinity of the tunnel, the creeping frozen soil is deflected around the tunnel since its stiffness is much greater than that of the frozen soil. The velocity profile is a function of the stiffness properties of the tunnel liner. The tunnel liner also serves to drastically reduce the downward velocity of the crown of the tunnel. In this case, the crown velocity is reduced by 3 orders of magnitude for a tunnel at the same depth with no liner present.

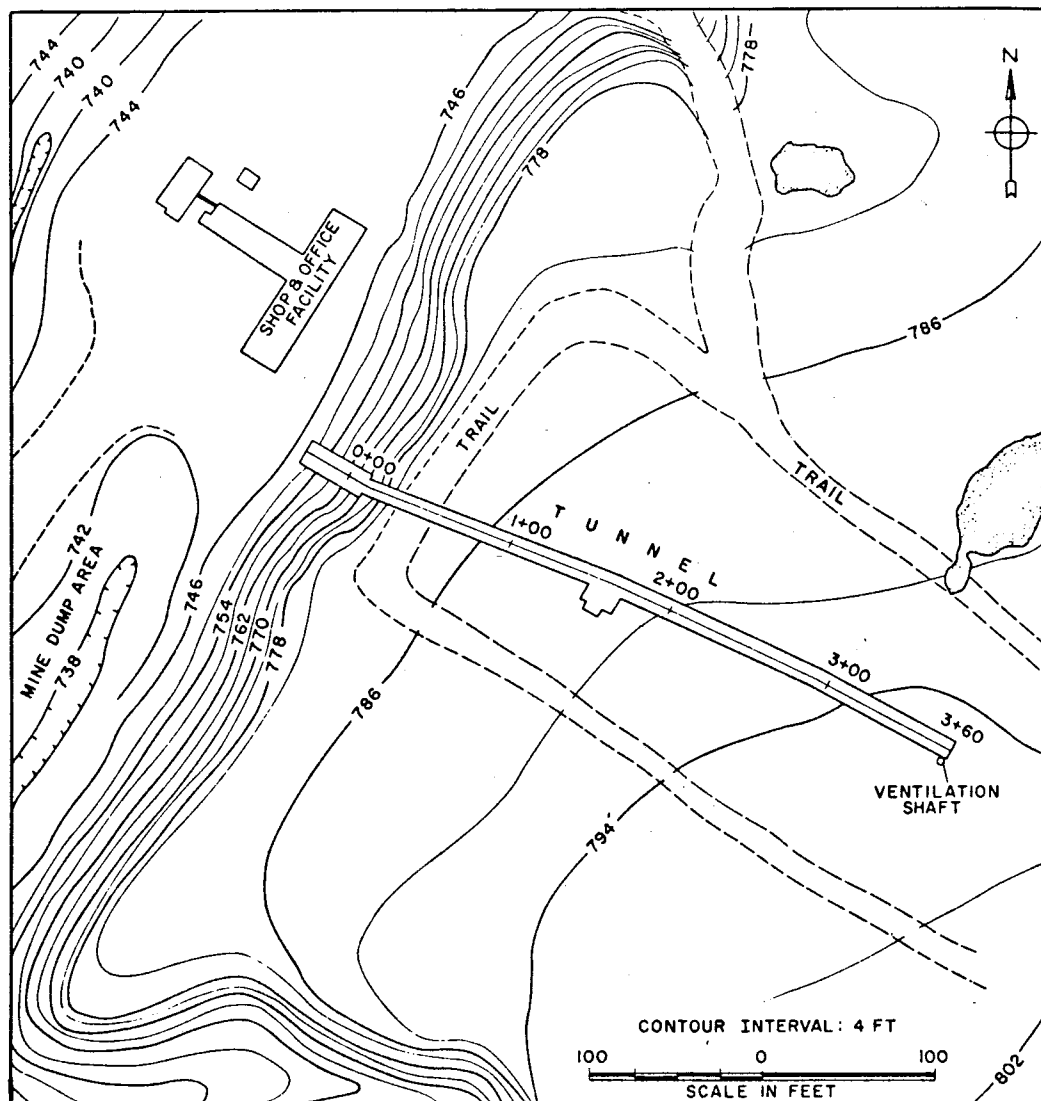


FIGURE 5.1 SITE PLAN OF FOX TUNNEL SITE,
ALASKA (Sellman, 1967)

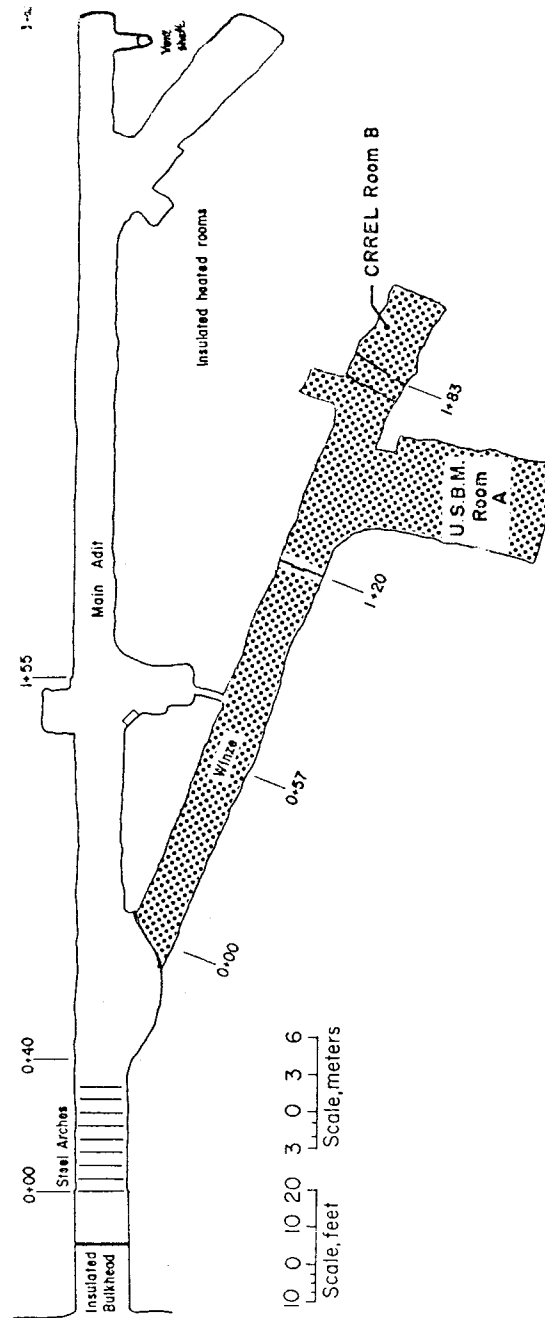


FIGURE 5.2 PLAN OF FOX TUNNEL SHOWING SECOND STAGE OF EXCAVATION (Pettibone, 1973)

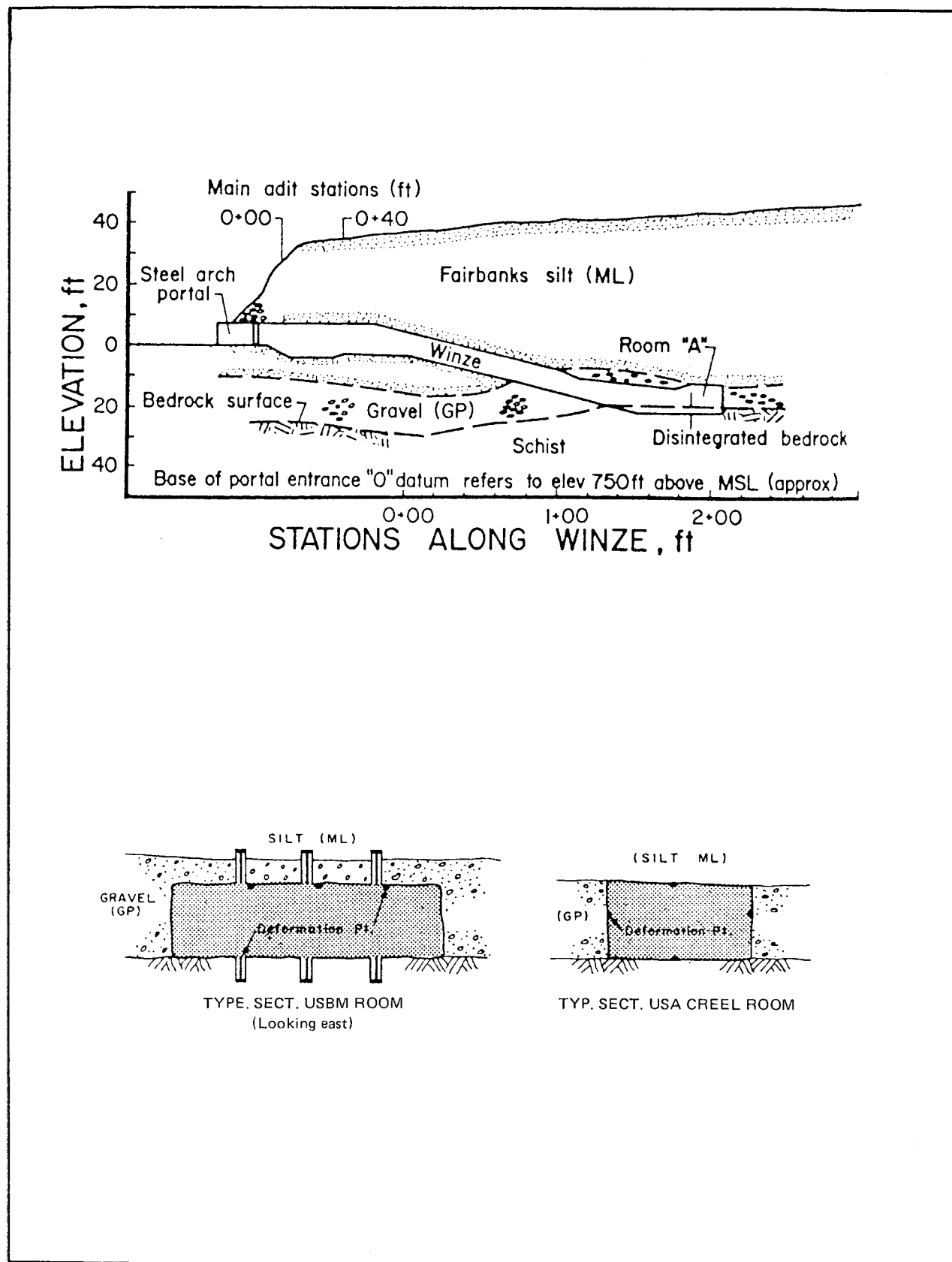


FIGURE 5.3

GEOLOGIC CROSS SECTION, FOX TUNNEL SITE
(Pettibone, 1973)

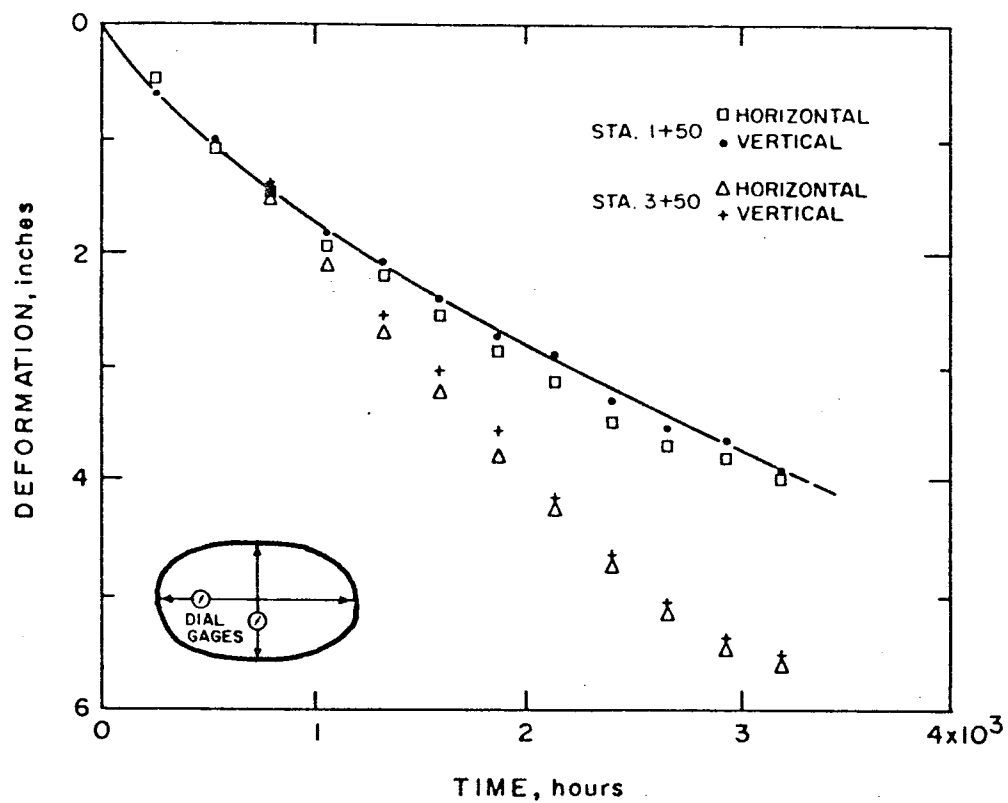


FIGURE 5.4

CLOSURE MEASUREMENTS OF MAIN TUNNEL
FOLLOWING 1965 CONSTRUCTION SEASON
(Swinzow, 1970)

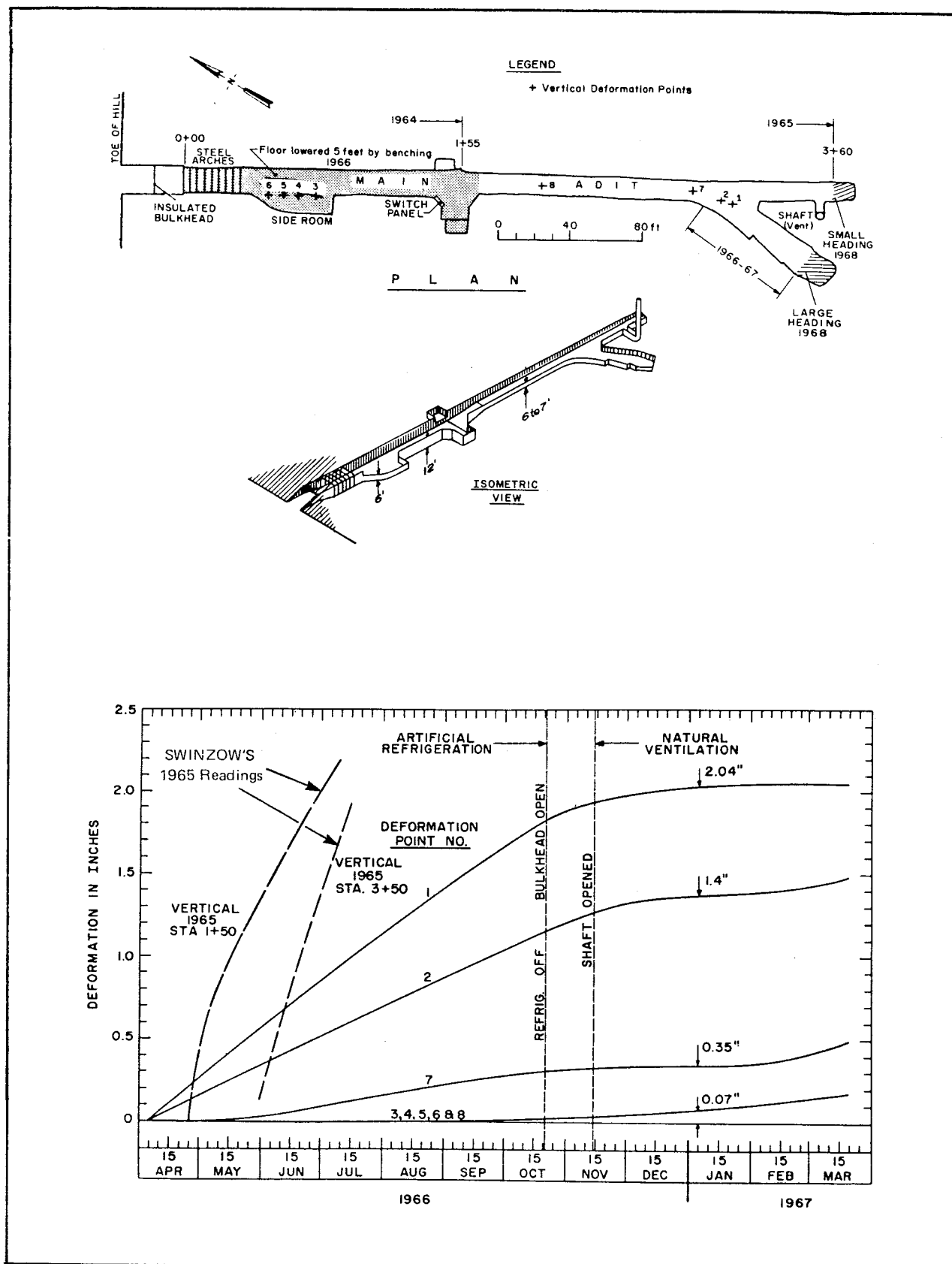


FIGURE 5.5

CLOSURE MEASUREMENTS OF MAIN TUNNEL
FOLLOWING 1965 CONSTRUCTION SEASON
(McAnerney, 1968)

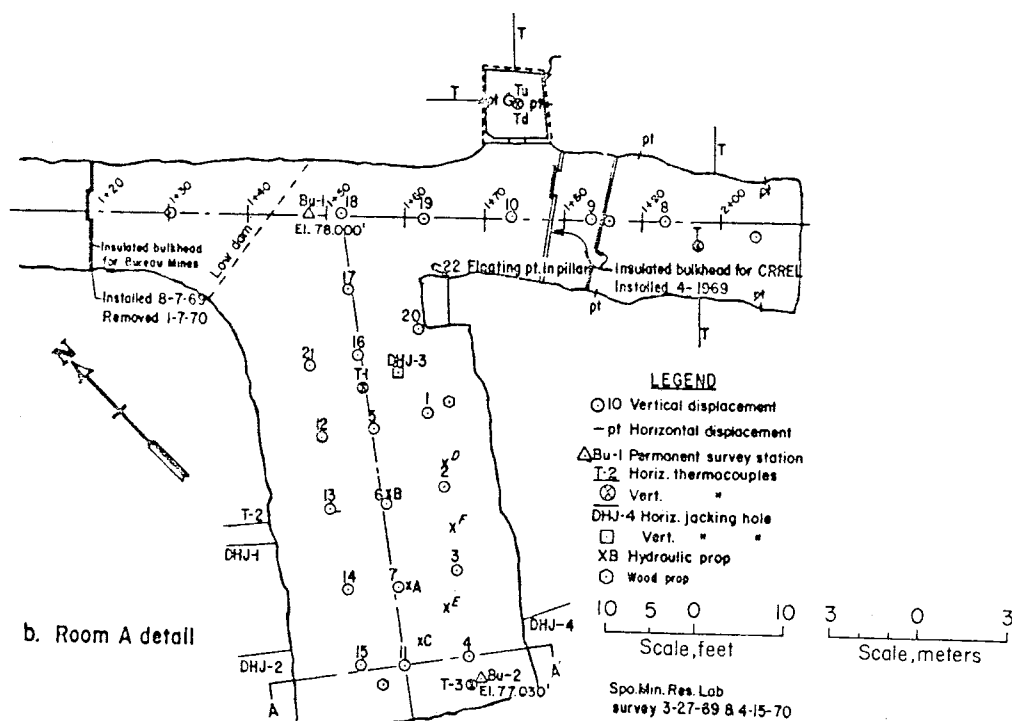


FIGURE 5.6

VERTICAL CLOSURE OF USBM ROOM IN THE FOX TUNNEL
(Pettibone, 1973)

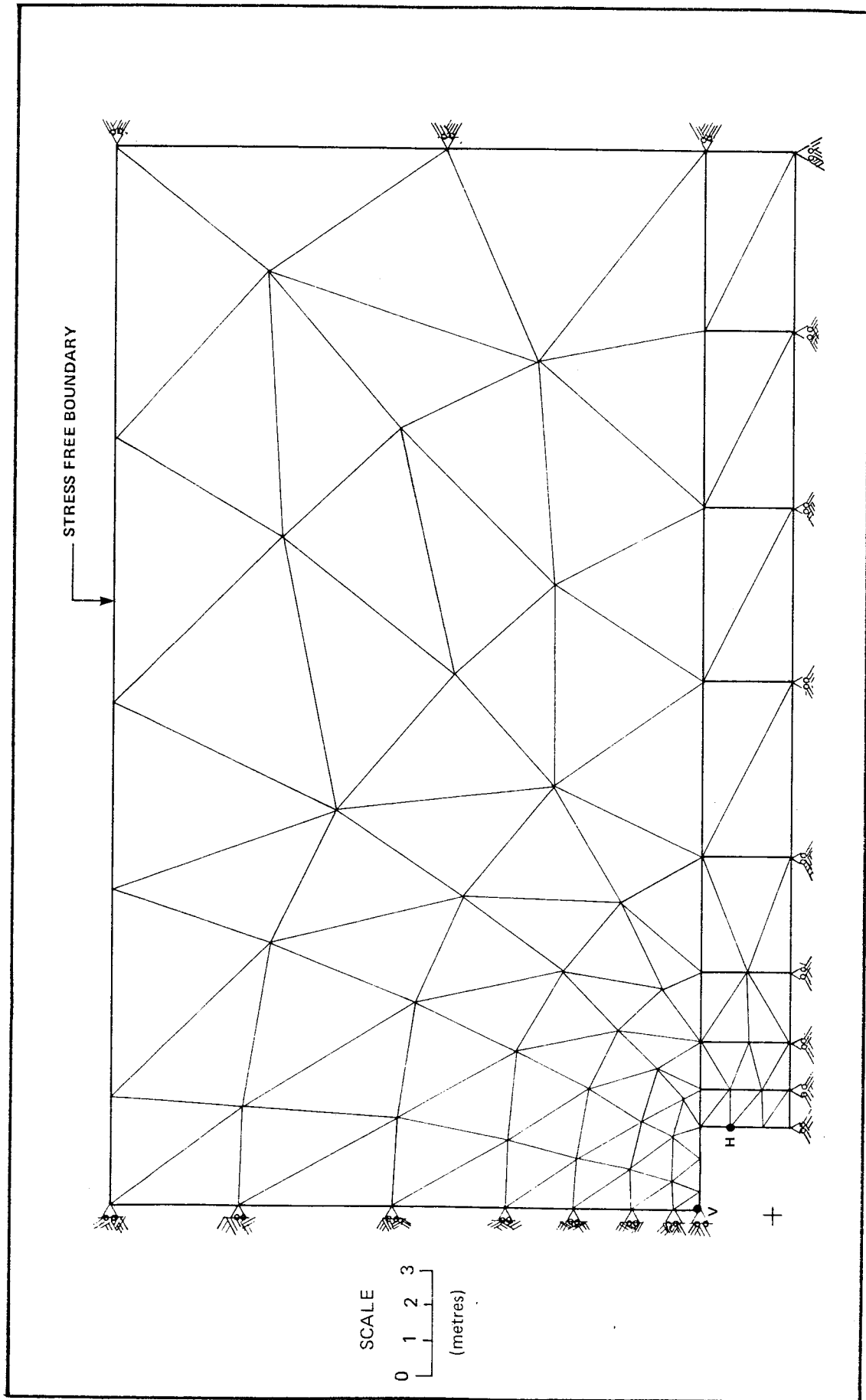


FIGURE 5.10 FINITE ELEMENT DISCRETIZATION OF USA CRREL ROOM IN THE FOX TUNNEL

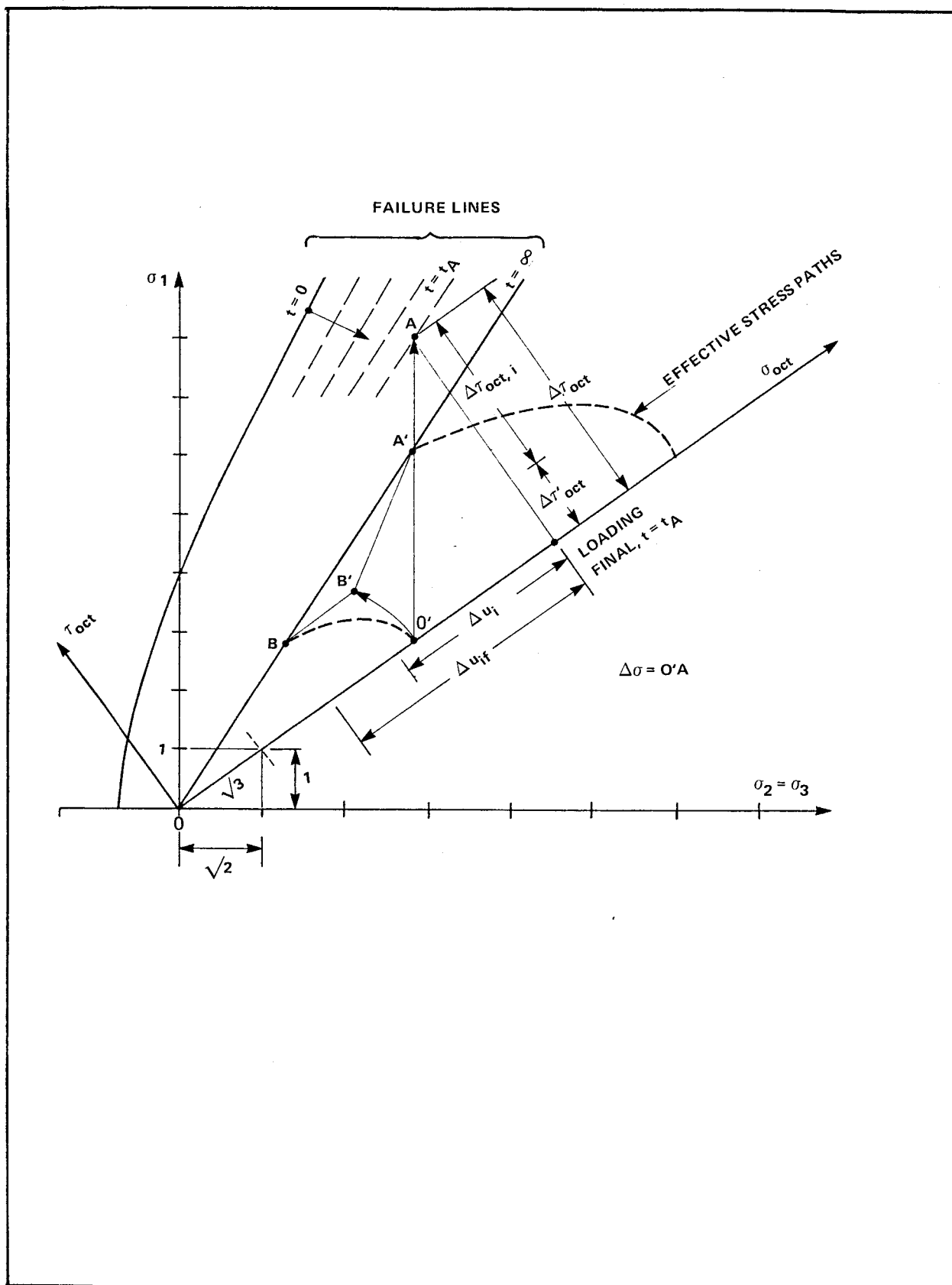
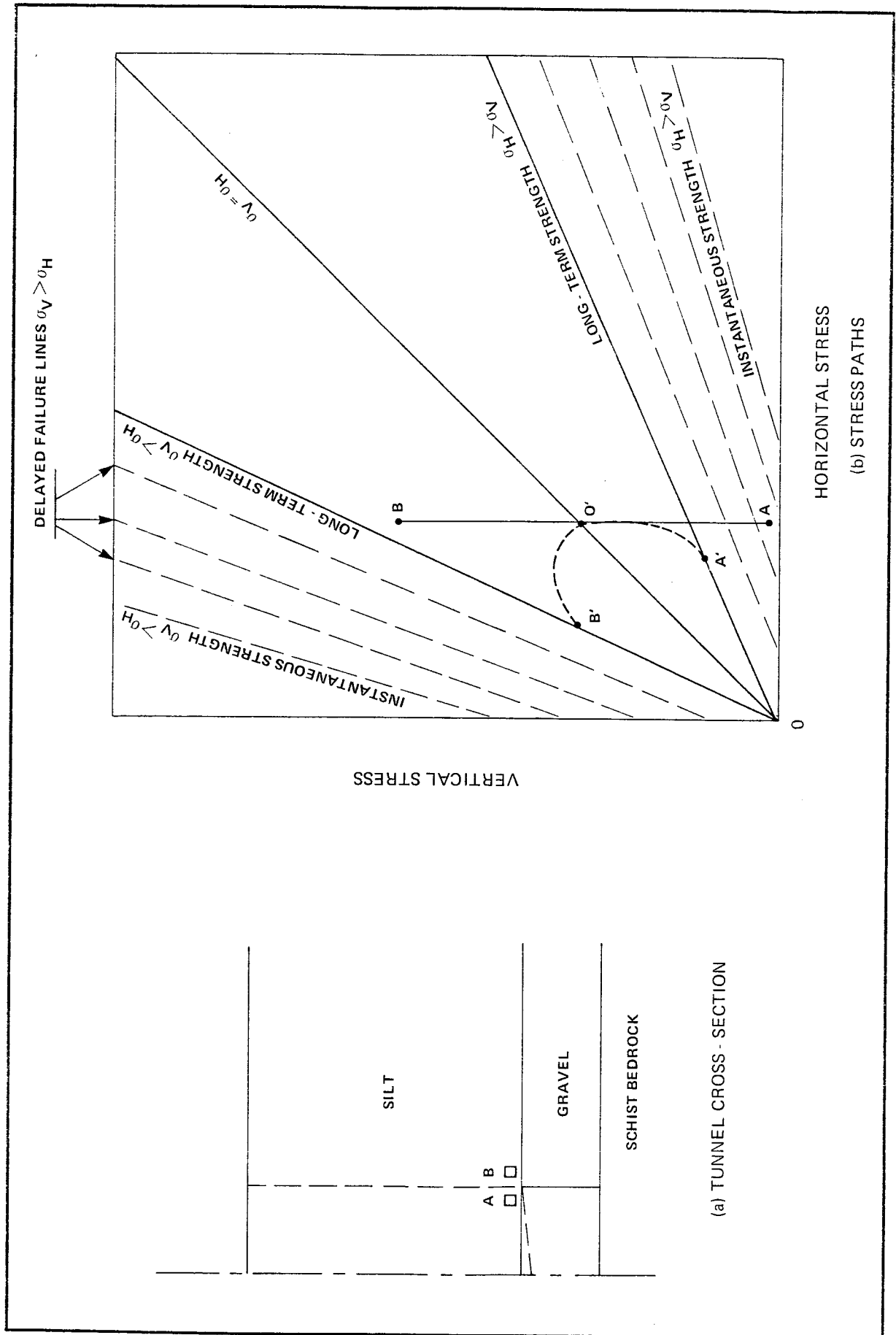


FIGURE 5.11

RENDULIC PLOT OF NORMALLY CONSOLIDATED PLASTIC FROZEN SOIL STRESSED BEYOND ITS LONG TERM STRENGTH (Ladanyi, 1981)



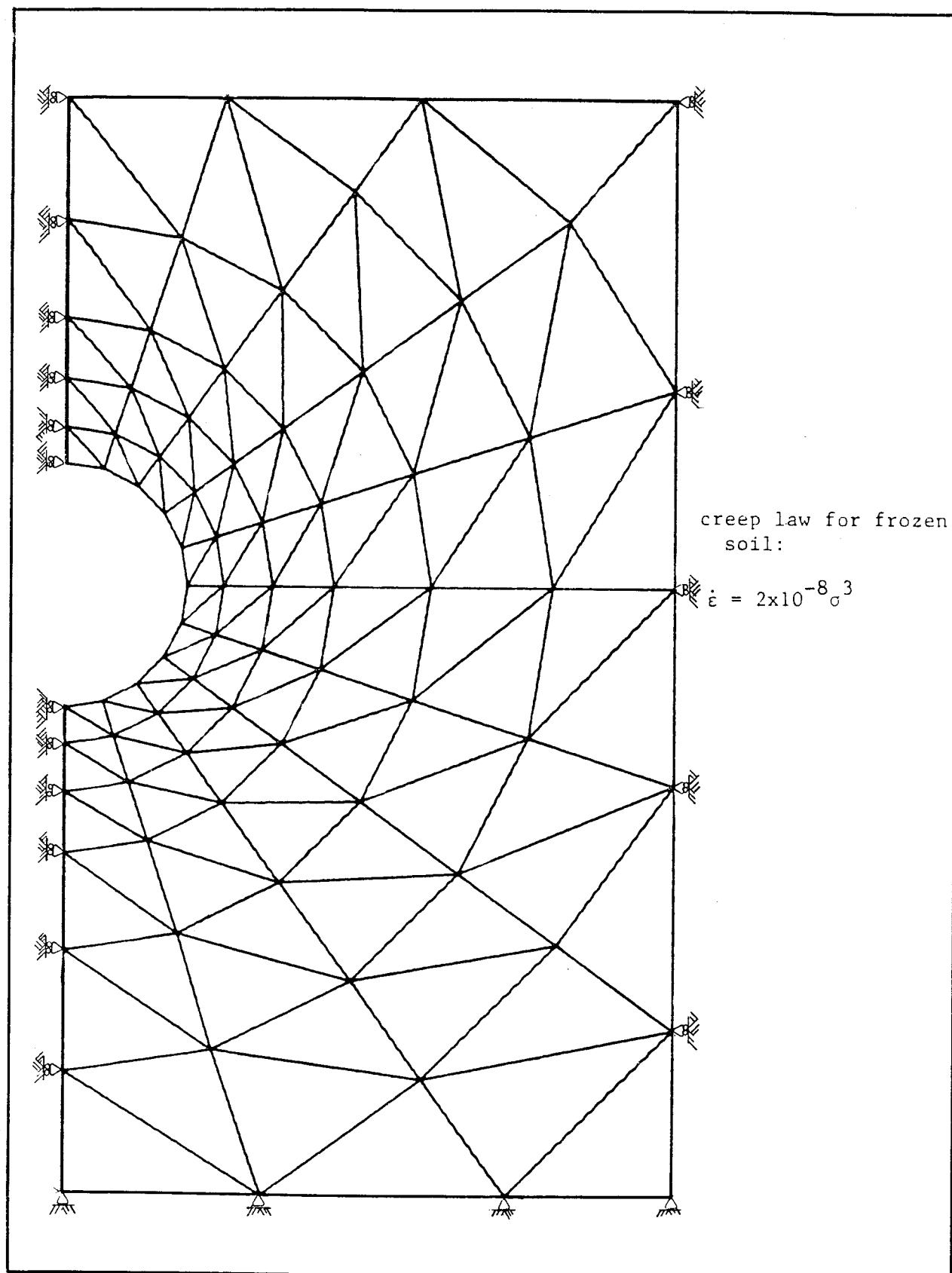


FIGURE 5.14

FINITE ELEMENT MESH FOR TUNNEL AT $H/D = 2.0$

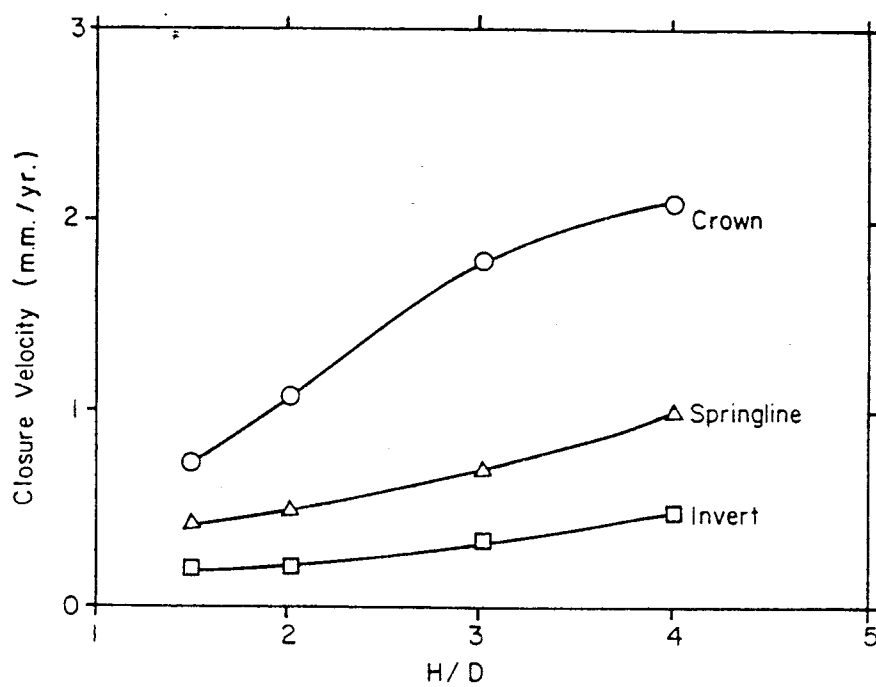


FIGURE 5.15 TUNNEL CLOSURE VELOCITY VERSUS DEPTH OF OVERBURDEN

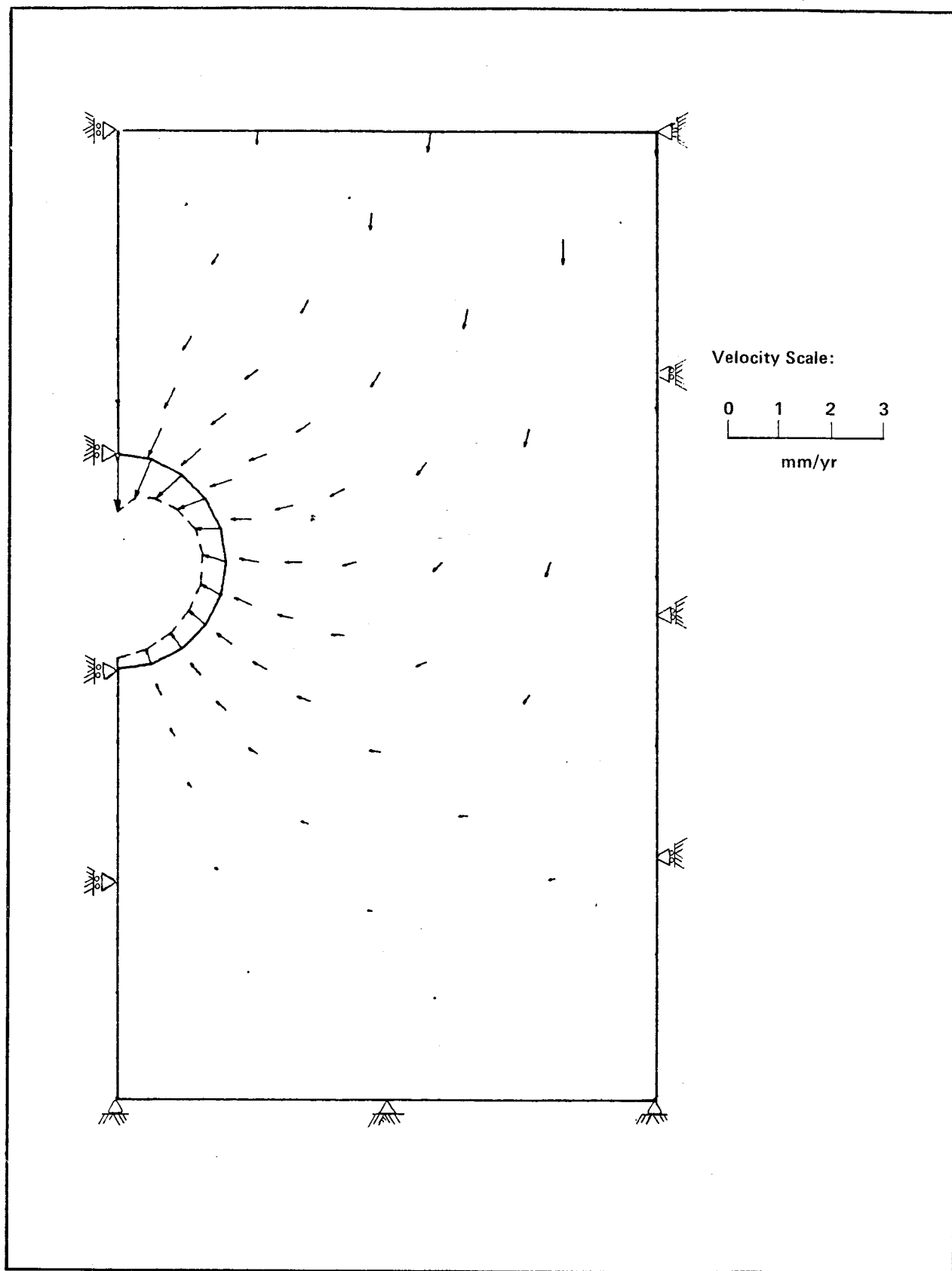


FIGURE 5.16

VELOCITY VECTOR PLOT FOR TUNNEL AT $H/D = 2.0$

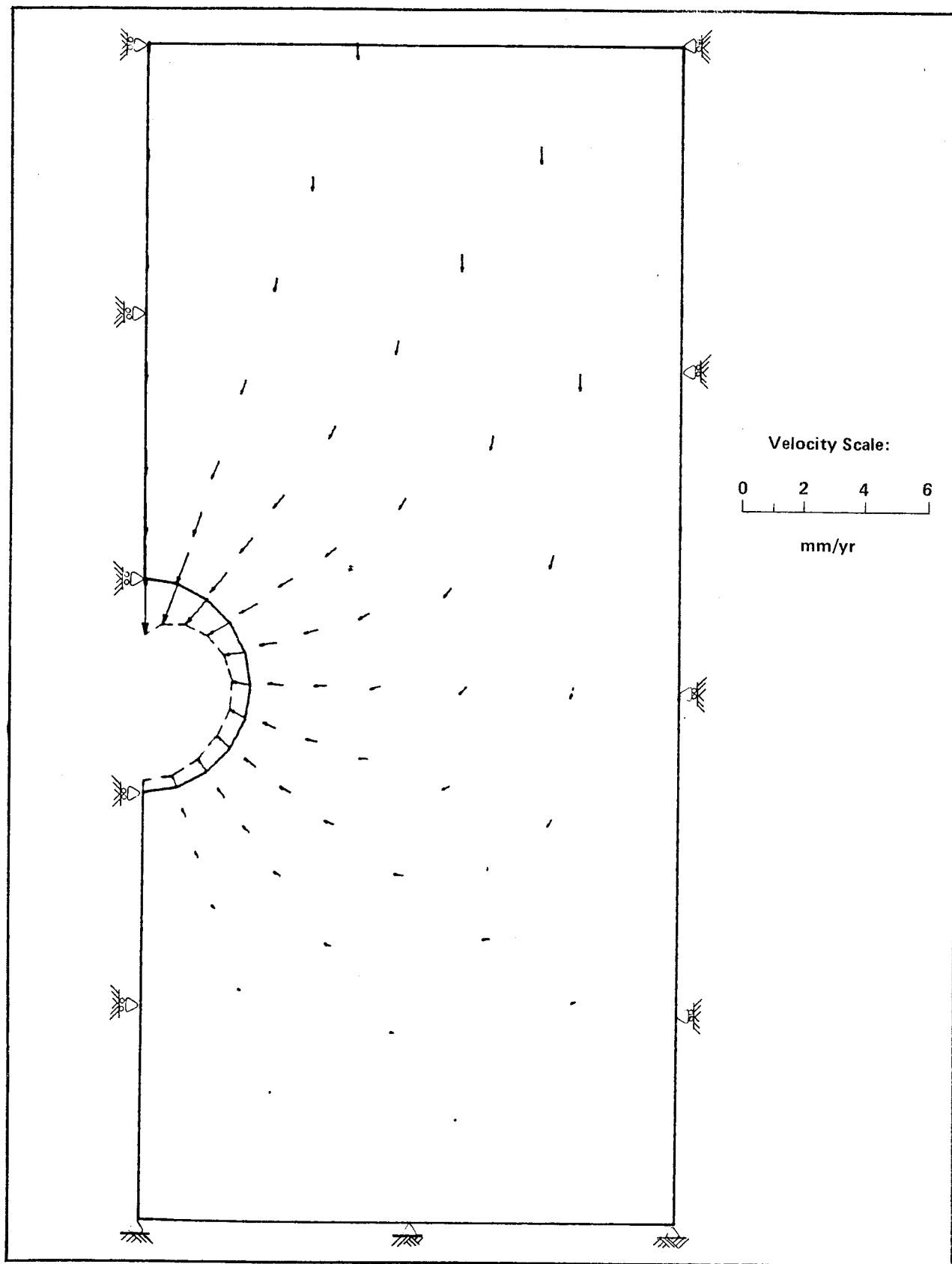


FIGURE 5.17 VELOCITY VECTOR PLOT FOR TUNNEL AT $H/D = 3.0$

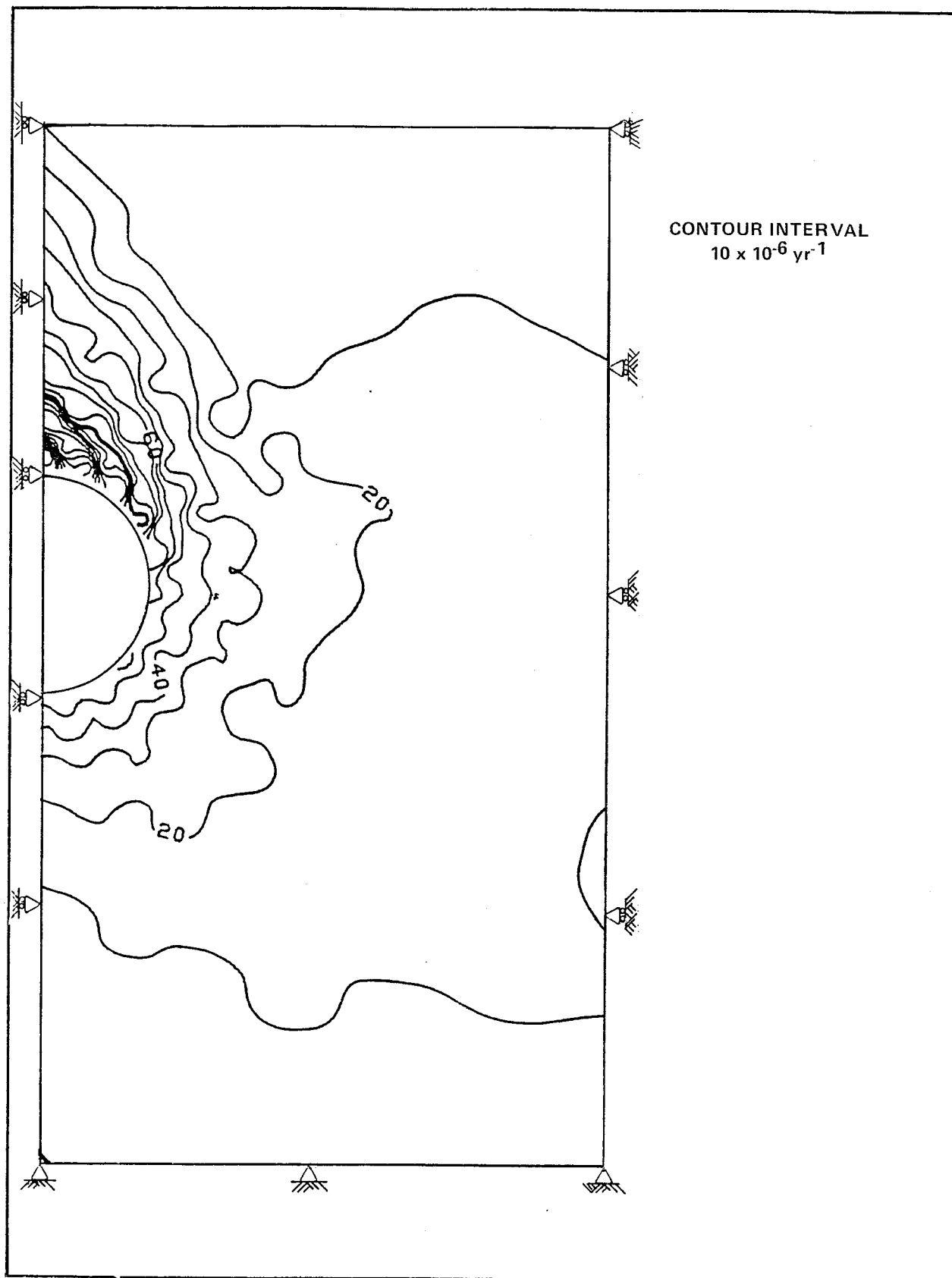


FIGURE 5.18 EFFECTIVE STRAIN RATE CONTOURS FOR TUNNEL AT $H/D = 2.0$

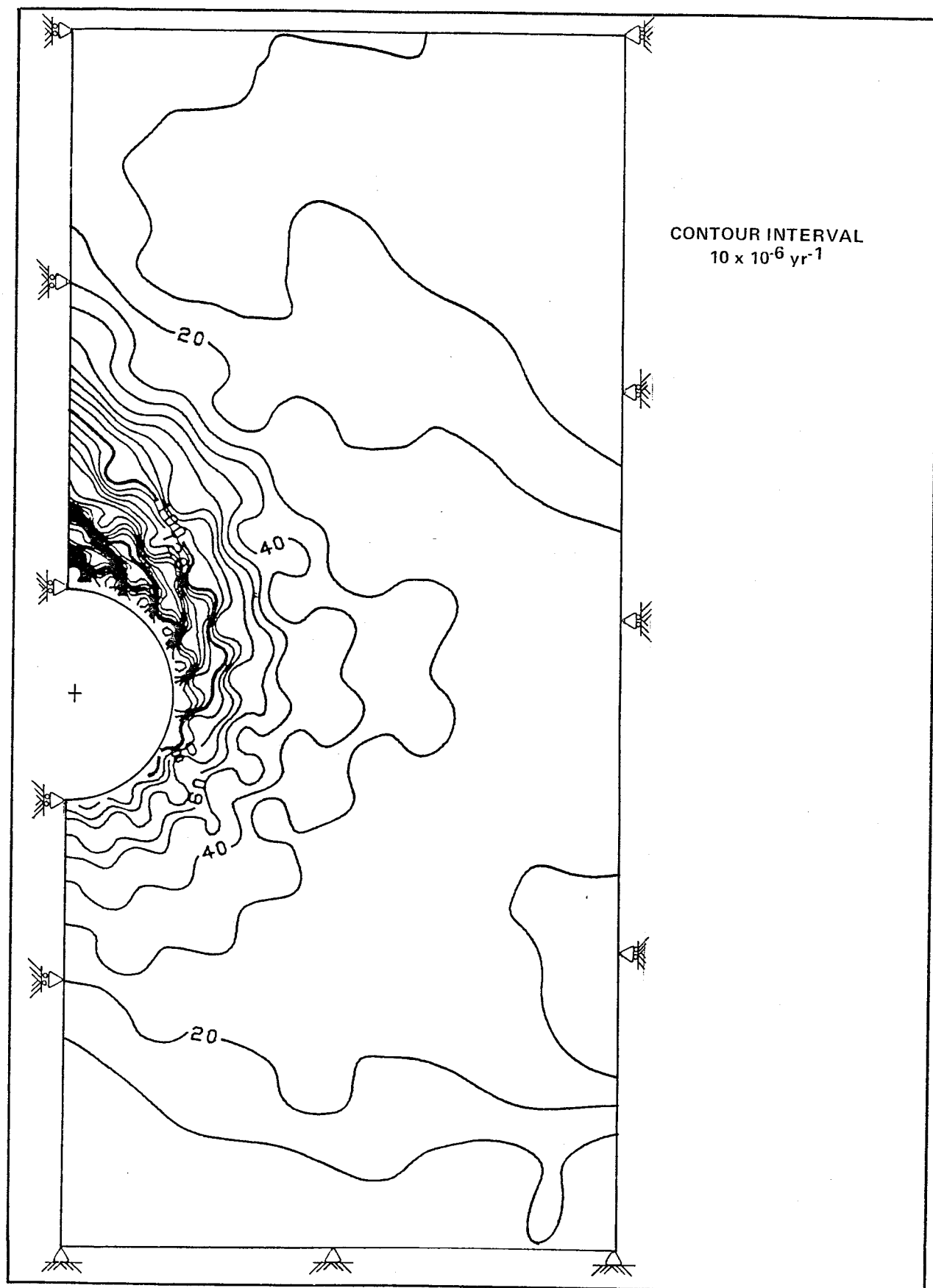


FIGURE 5.19 EFFECTIVE STRAIN RATE CONTOURS FOR TUNNEL AT $H/D = 3.0$

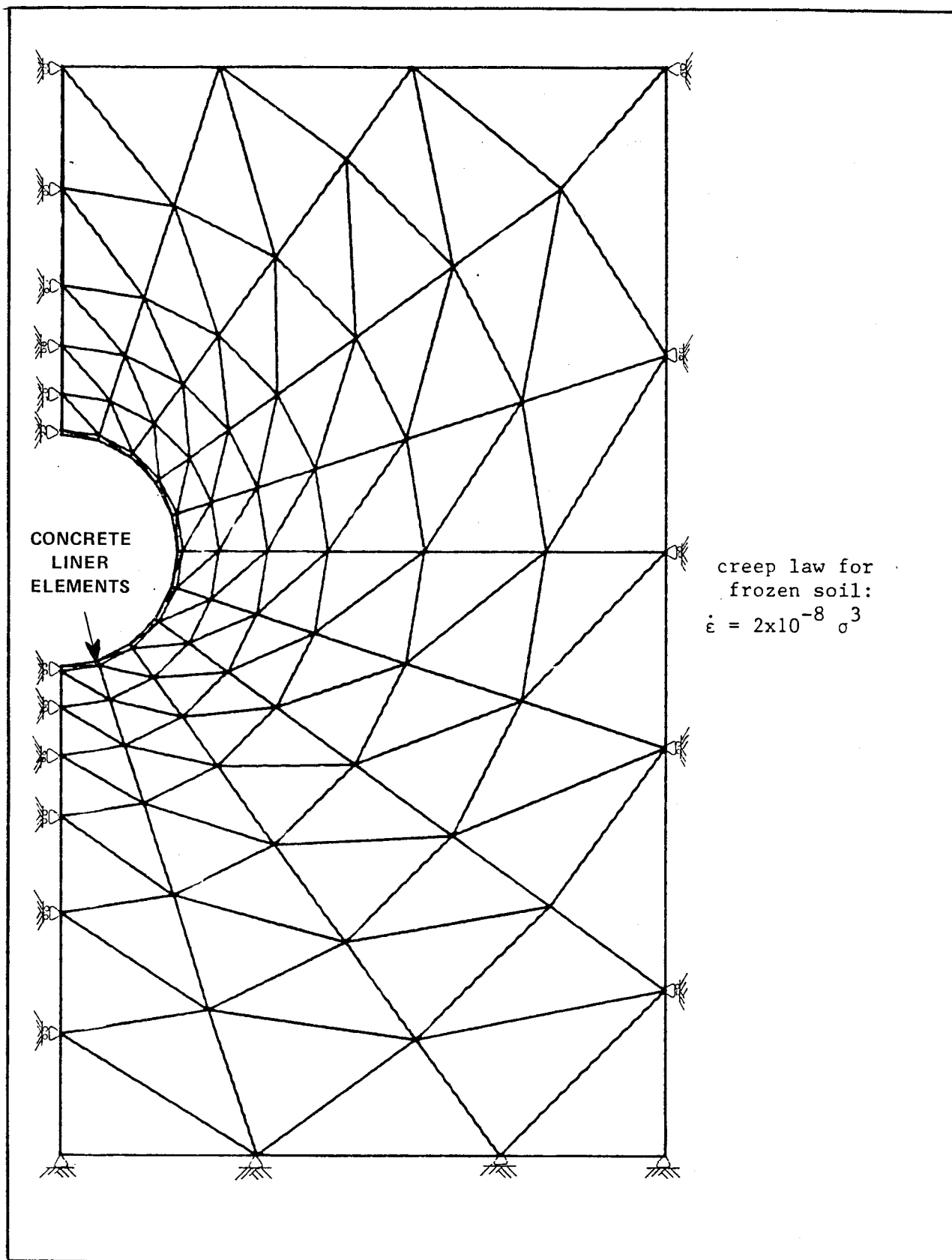


FIGURE 5.20

FINITE ELEMENT MESH FOR LINED TUNNEL AT $H/D = 2.0$

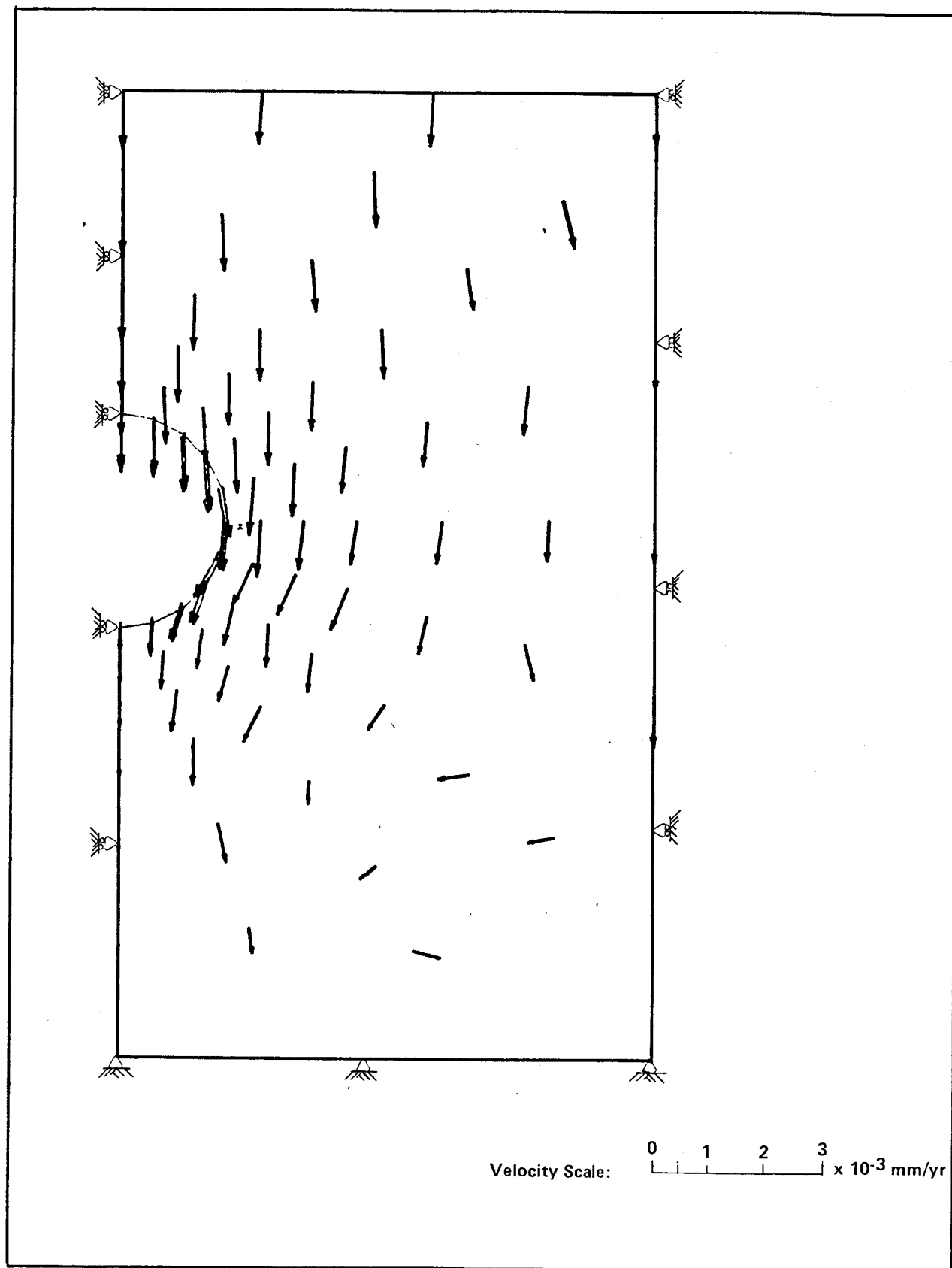


FIGURE 5.21 VELOCITY VECTOR PLOT FOR LINED TUNNEL AT $H/D = 2.0$

CHAPTER VI

CONCLUDING REMARKS

6.1 GENERAL

An analytical study of the in-situ creep behavior in ice-rich frozen soil has been reported herein. A review of the deformation behavior of frozen soil has lead to the conclusion that the steady state creep of ice-rich frozen soil is adequately described by an empirically derived simple power law which relates strain rate to stress. The flow law for steady state creep in polycrystalline ice (Morgenstern et al., 1980) represents a valid upper bound for the constitutive behavior of frozen soil under equivalent loading conditions. The importance of testing natural permafrost samples with representative segregated ice features was also noted.

The incremental initial strain finite element procedure was applied to two case histories of naturally occurring creep in ice-rich permafrost. The in-situ creep behavior of a relatively steep slope on the left bank of the Great Bear River at the proposed Arctic Gas crossing was studied by Savigny (1980). Naturally occurring creep in ice-rich Fairbanks silt at the Fox Tunnel near Fairbanks, Alaska was studied by Swinzow (1970), Thompson and Sayles (1972) and Pettibone (1973). The following sections summarize the major conclusions of the analytical creep studies at these two sites.

6.2 Conclusions

6.2.1 Great Bear River

The finite element creep analysis of the left bank of the Great Bear River at the proposed Arctic Gas crossing has shown that the in-situ creep behavior of the glaciolacustrine clay can be accurately represented by the simple power law. The flow law which gave the best fit between observed and predicted velocities had an exponent of 3.0 and a coefficient of 3.33×10^{-9} . This strain rate is six times slower than the value for polycrystalline ice at -2°C (Morgenstern et al., 1980).

The exact form of the constitutive relationship for the glaciodeltaic sand remains inconclusive. In the first creep analysis, the creep behavior of the sand was assumed to be given by the simple power law with the same parameters as those in the glaciolacustrine clay. The velocities predicted by the simple power law were much greater than the observed values. A second analysis was carried out in which the sand was treated as an elastic material. This proved to be much too serious a restriction on the creep movements in the sand and the glaciolacustrine clay where it was overlain by sand. This analysis did predict the build up of horizontal tensile stress in the sand much greater than the sand would be able to sustain under in-situ conditions.

The creep movements recorded at the GB1A and GB3 inclinometer locations at the Great Bear River showed that more movement was associated with large ice lenses when they are widely separated (Savigny, 1980). Numerical modelling of permafrost as composite soil-ice medium would require an accurate mapping of the lateral continuity of these large ice lenses. Statistical analysis of the data derived from the MVPRL ice variability

test sites at Norman Wells N.W.T. and Landing Lake N.W.T. suggest that for a uniform stratigraphic unit, the overall ground ice content of permafrost can be described as a normally distributed variable. Thus, for practical purposes, the assumption of homogeneity is valid for predicting overall time-dependent ground movements in ice-rich frozen soil.

A general conclusion derived from the study of the left bank of the Great Bank River at the proposed Arctic Gas crossing that the time-dependent deformation behavior of homogeneous slopes consisting of ice-rich fine-grained permafrost can be modelled by a simple power law. Using the creep parameters for polycrystalline ice will provide an upper bound to the creep velocities.

6.2.2 Fox Tunnel

The in-situ deformation studies of the Fox Tunnel reported by Swinzow (1970), Thompson and Sayles (1972), and Pettibone (1973) showed that underground openings excavated in warm ice-rich permafrost can become unstable when the ambient air temperature is near the soil temperature. In each case, primary creep behavior contributed substantially to the overall room closure.

The closure data of the USA CRREL room recorded over a 3 year period clearly showed attenuating creep for well over one year. The closure versus time data for the USA CRREL room and the USBM room in the Fox Tunnel both have the same general form for the first 3 years of operation. Thus, Thompson and Sayles (1972) were premature in concluding that steady state creep dominated the closure of the USA CRREL room during the first year of operation.

A finite element creep analysis of the USA CRREL room showed that the simple power law for polycrystalline ice does not represent a valid upper bound for the in-situ creep behavior of the frozen Fairbanks silt. However, analyzing the stress changes above a rectangular underground opening in terms of arching theory clearly shows that the stress paths followed by soil elements above the roof of the opening are entirely different from that normally used in constant stress or constant strain rate creep tests. The strain rate predicted by finite element computations was 3.3 times faster than the laboratory measured value (Thompson and Sayles, 1972) because the vertical room closure measurements could not distinguish between vertical compression in the gravel walls and downward creep movement of the overlying ice-rich silt. For this particular case it was not possible to isolate in-situ creep deformations by measuring the closure between opposing points in the roof, floor and walls.

6.3 Recommendations for Further Study

The analytical study of naturally occurring creep in a slope and underground cavity located in ice-rich permafrost has shown that the simple power law can be used to predict in-situ deformations provided that the stress does not approach failure.

The largest discrepancies between predicted and observed behavior appear to coincide with regions where the stresses experienced by the soil mass are unlike the stress paths used in laboratory constant stress or constant strain rate creep tests. This was illustrated at the GB2 inclinometer location at the Great Bear River. The stresses at this location are altered from the simple shear condition because the downslope creep movement in the glaciolacustrine clay is impeded by the overlying sand just upslope of this location. Downslope of the GB2 inclinometer, the glaciolacustrine clay creeps at a faster rate since there is no overlying sand to restrict movement. Thus, the clay in the vicinity of the GB2 inclinometer would experience an unloading in the horizontal direction.

The frozen silt above the roof in the USA CRREL room in the Fox Tunnel experiences an unloading due to excavation of the room. Arching behavior will transfer the vertical stress release in the roof to the frozen silt above the walls and to the frozen gravel walls. Thus, while the soil which experiences greater compressive stress may still lie below the compression failure envelope, the soil above the roof of the tunnel may experience tensile failure at some time after the room is excavated. The stress state can become quite complicated whenever two or more frozen materials are present with different creep properties. Thus, it is important to obtain an accurate assessment of the creep behavior of the materials relative to each other.

Future numerical modelling of ice-rich permafrost under complex stress conditions should be directed towards studying the stress changes experienced by the soil as deformations develop and to incorporate the concept of a limiting long term strength into the analysis. Toward this end, laboratory studies should be carried out to study the creep behavior of frozen soils under stress conditions that simulate unloading. Also, future field programs should have carefully designed and installed instrumentation to ensure that the in-situ creep deformations can be isolated from other deformation processes which may be occurring simultaneously.

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APPENDIX A

FORMULATION OF FINITE ELEMENT EQUATIONS FOR THE INCREMENTAL INITIAL STRAIN PROCEDURE

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FORMULATION OF FINITE ELEMENT EQUATIONS FOR
INCREMENTAL INITIAL STRAIN PROCEDURE

This appendix presents a brief review of how the incremental theory of creep is introduced into the finite element method using the initial strain procedure.

In the incremental theory of creep, the increment of creep strain is treated as an initial strain for any one time interval and is assumed to be constant for that time interval.

The increment of total strain, $\Delta\epsilon^T$, for any time interval can be assumed to consist of an increment of elastic strain, $\Delta\epsilon^E$, and an increment of creep strain, $\Delta\epsilon^C$. This can be expressed as:

$$\{\Delta\epsilon^T\} = \{\Delta\epsilon^E\} + \{\Delta\epsilon^C\} \quad (A1)$$

Solving equation A1 for the elastic strain gives:

$$\{\Delta\epsilon^E\} = \{\Delta\epsilon^T\} - \{\Delta\epsilon^C\} \quad (A2)$$

and applying Hooke's law:

$$\{\Delta\sigma\} = [D]\{\Delta\epsilon^T - \Delta\epsilon^C\} \quad (A3)$$

where: $[D]$ = matrix of elastic constants

For the linear elastic case, an increment of strain energy density of a body is given by:

$$\Delta U(x, y, z) = \frac{1}{2} \{\Delta \epsilon^E\}^T \{\Delta \sigma\} \quad (A4)$$

Substitution of equation A2 and A3 into A8 gives:

$$\Delta U(x, y, z) = \frac{1}{2} \{\Delta \epsilon^T - \Delta \epsilon^C\}^T [D] \{\Delta \epsilon^T - \Delta \epsilon^C\} \quad (A5)$$

An increment of potential energy of a body is expressed mathematically as:

$$\begin{aligned} \Delta \pi_p = & \int_V \Delta U(x, y, z) dV - \int_V (\bar{X} \Delta u + \bar{Y} \Delta v + \bar{Z} \Delta w) dV \\ & - \int_{S_T} (\bar{T}_x \Delta u + \bar{T}_y \Delta v + \bar{T}_z \Delta w) dS_T \end{aligned} \quad (A6)$$

where V represents the volume of the body and S_T is the surface on which surface tractions are prescribed. The last two integrals in equation A6 represent the work done by the external forces; i.e. the body forces \bar{X} , \bar{Y} , \bar{Z} , and the surface tractions, \bar{T}_x , \bar{T}_y , and \bar{T}_z . The bar denotes quantities that are prescribed. Substituting equation A5 into A6 yields:

$$\begin{aligned}
\Delta \pi_p = & \frac{1}{2} \int_V \{\Delta \varepsilon^T\}^T [D] \{\Delta \varepsilon^T\} dV - \frac{1}{2} \int_V \{\Delta \varepsilon^T\}^T [D] \{\Delta \varepsilon^C\} dV \\
& - \frac{1}{2} \int_V \{\Delta \varepsilon^C\}^T [D] \{\Delta \varepsilon^T\} dV + \frac{1}{2} \int_V \{\Delta \varepsilon^C\}^T [D] \{\Delta \varepsilon^C\} dV \quad (A7) \\
& - \frac{1}{2} \int_V \{\Delta \psi(x, y, z)\}^T \{F\} dV - \int_{S_T} \{\Delta \psi(x, y)\}^T \{T\} dS_T
\end{aligned}$$

where: $\{\Delta \psi(x, y, z)\}^T = [\Delta u \quad \Delta y \quad \Delta z]$

$$\{F\} = \begin{pmatrix} \bar{X} \\ \bar{Y} \\ \bar{Z} \end{pmatrix} \quad \{T\} = \begin{pmatrix} \bar{T}_x \\ \bar{T}_y \\ \bar{T}_z \end{pmatrix}$$

Using standard finite element calculations, one can write:

$$\{\Delta \varepsilon\} = [B] \{\Delta q\} \quad (A8)$$

$$\{\Delta \psi(x, y, z)\} = [N] \{\Delta q\} \quad (A9)$$

where $[B]$ = matrix relating element strains to the values of the displacement at its nodes.

$[N]$ = coefficient matrix for displacement interpolation model.

Substituting equation A8 and A9 into equation A7 yields:

$$\begin{aligned}
 \Delta\pi_p = & \frac{1}{2} \int_V \{\Delta q\}^T [B]^T [D] [B] \{\Delta q\} dV - \int_V \{\Delta q\}^T [B]^T [D] \{\Delta \epsilon^C\} dV \\
 & + \frac{1}{2} \int_V \{\Delta \epsilon^C\}^T [D] \{\Delta \epsilon^C\} dV - \int_V \{\Delta q\}^T [N]^T \{F\} dV \\
 & - \int_{S_T} \{\Delta q\}^T [N]^T \{T\} dS_T
 \end{aligned} \tag{A10}$$

The Theorem of Minimum Potential Energy states:

"Of all possible displacement configurations a body can assume which satisfy compatibility on the constraints or kinematic boundary conditions, the configuration satisfying equilibrium makes the potential energy assume a minimum value".

Mathematically, this is expressed as:

$$\delta(\Delta\pi_p) = 0 \tag{A11}$$

Taking the first variation of equation A10 with respect to the incremental displacements, $\{\Delta q\}$, and setting it equal to zero gives:

$$\begin{aligned}
 \{\delta(\Delta q)\}^T \int_V [B]^T [D] [B] dV \{\Delta q\} - \int_V [B]^T [D] \{\Delta \epsilon^C\} dV \\
 - \int_V [N]^T \{F\} dV - \int_{S_T} [N]^T \{T\} dS_T = 0
 \end{aligned} \tag{A12}$$

Since the variation of the displacements are arbitrary, the quantity enclosed in the brackets must vanish. The variation of the third integral in equation A10 is zero because it is not a function of the nodal displacements. The set of equilibrium equations for each time interval is:

$$\{K\}_s \{\Delta q\}_s = \{P\}_s + \{\Delta F_c\}_s \quad (A13)$$

where:

$$\{K\}_s = \sum_{m=1}^M \{K\}_m^{el} = \sum_{m=1}^M \left(\int_V [B]^T [D] [B] dV \right)_m$$

$$\{P\}_s = \sum_{m=1}^M \{P\}_m^{el} = \sum_{m=1}^M \left(\int_V [N]^T \{F\} dV + \int_{S_T} [N]^T \{T\} dS_T \right)_m$$

$$\{\Delta F_c\}_s = \sum_{m=1}^M \{\Delta F_c\}_m^{el} = \sum_{m=1}^M \left(\int_V [B]^T [D] \{\Delta \epsilon_c\} dV \right)_m$$

The summation sign is carried over the total number of elements, M , to obtain the equilibrium equations for the entire assemblage.

The vector $\{F_c\}_s$ represents the creep strain nodal load vector. Thus, it can be seen that for each interval of time, creep strains are allowed to take place. Equivalent nodal forces are then calculated which would be necessary to cause elastic strains of the same magnitude. These fictitious nodal forces are then added to the load vector and new set of nodal displacements are calculated during that time interval with equation A13.

APPENDIX B

CALCULATION OF FICTITIOUS NODAL FORCES

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CALCULATION OF FICTITIOUS NODAL FORCES

This appendix presents the derivation of the fictitious nodal creep forces used in the initial strain approach to solving creep problems.

In each time interval, the increment of creep strain is given by the stress-strain-rate law, i.e.:

$$\Delta \dot{\epsilon}_{ij} = \frac{3}{2} (\Delta \dot{\epsilon}_e / \sigma_e) s_{ij} \quad (B1)$$

Using Hooke's law, the increments of creep strain are converted to incremental creep stresses:

$$\{\sigma^C\} = [D] \{\epsilon^C\} \quad (B2)$$

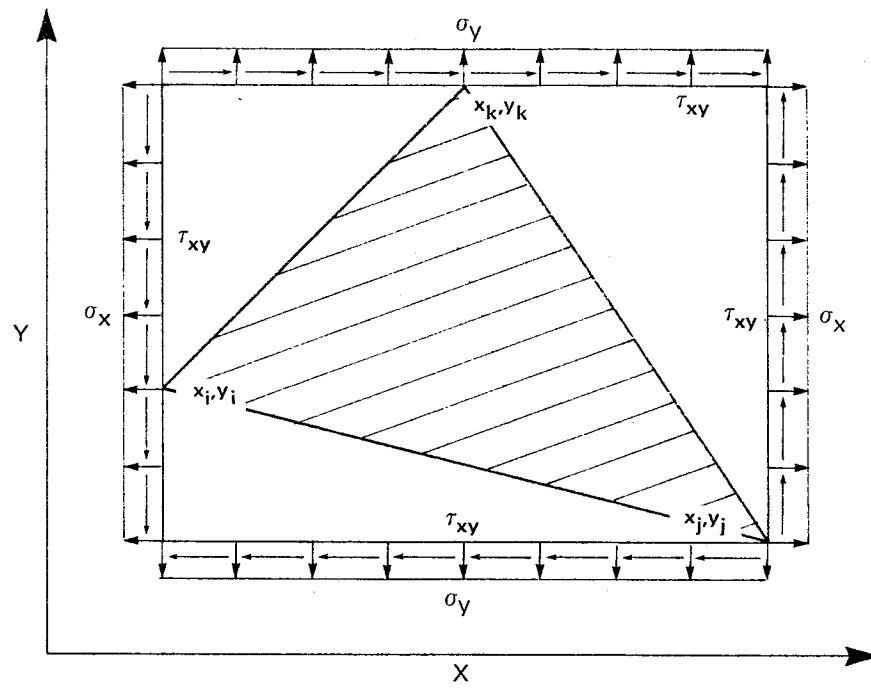
The constant stress state for the constant strain triangular finite element is represented by Figure B-1(a).

The stresses can be replaced by a set of statically equivalent forces acting at the node points of the constant strain triangle. Elements with higher order displacement interpolation functions require that the nodal forces be work-equivalent loads. However, for the constant strain triangle, work-equivalent and statically-equivalent loads are equal.

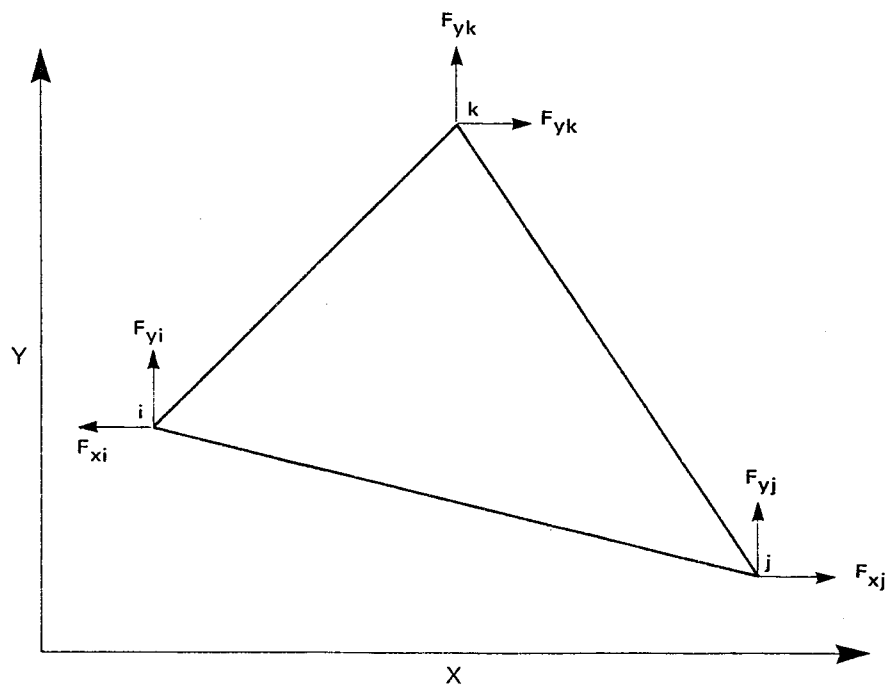
The incremental stresses are related to the nodal forces for each element as follows:

$$\begin{Bmatrix} F_{xi} \\ F_{xj} \\ F_{xk} \\ F_{yi} \\ F_{yj} \\ F_{yk} \end{Bmatrix} = \frac{1}{2} \begin{bmatrix} y_k - y_i & 0 & x_j - x_k \\ y_i - y_j & 0 & x_k - x_i \\ y_j - y_k & 0 & x_i - x_j \\ 0 & x_j - x_k & y_k - y_i \\ 0 & x_k - x_i & y_i - y_j \\ 0 & x_i - x_j & y_j - y_k \end{bmatrix} \begin{Bmatrix} \sigma_x^C \\ \sigma_y^C \\ \tau_{xy}^C \end{Bmatrix}_{el} \quad (B3)$$

The fictitious nodal load vector is computed for each element in turn and then summed to obtain the incremental load vector for the entire assemblage. The incremental load vector is then used to calculate a new set of incremental displacements for the current time step.



a) CONSTANT STRESS DISTRIBUTION



b) STATICALLY-EQUIVALENT NODAL FORCES

FIGURE B.1

STATICALLY EQUIVALENT NODAL FORCES FOR
CONSTANT STRAIN TRIANGLE

APPENDIX C

USER'S MANUAL FOR CREEP

INPUT DATA FORMAT FOR PROGRAM: CREEP

 CARD TYPE NO. 1 - TITLE OF THE PROBLEM [Format(20A4)]

Columns	Variable	Entry
*****	*****	*****
1-80	TITLE	Enter the heading information that is to be printed with the output

-any alphanumeric characters can be entered

 CARD TYPE NO. 2 - PROBLEM SOLUTION CONTROL [Format(2A8)]

Columns *****	Variable *****	Entry *****
1-8	CODE1	START - For an initial run of the program, all data is input on cards RSTART - For a continuation of a previous problem, that was terminated early
9-16	CODE2	RUN - For a normal run of the program CONTIN - Write the results of the final time step on to tape for restarting the problem at a future date

-if CODE1 = RSTART, card sets 1, 2, 3, 11, 12, 14 and 15
 are required

-to restart a previous problem, the results of the last increment
 of the terminated run are read from unit ITAPE

 CARD TYPE NO. 3 - INPUT AND OUTPUT DESIGNATION [Format(2I6)]

Columns *****	Variable *****	Entry *****
1-6	ITAPE	Unit number from which the results of a terminated run will be read
7-12	JTAPE	Unit number onto which the results for the final time step will be written for a run that will be terminated early

-for an initial run of the program, all the input data is
assumed to be read from cards

-to continue a problem at a future date, a unit number
for JTAPE must be specified; otherwise solution of the
problem will terminate

 CARD TYPE NO. 4 - MASTER CONTROL CARD [Format(7I6)]

Columns *****	Variable *****	Entry *****
1-6	NN	Total number of nodes (maximum of 200)
7-12	NE	Total number of elements (maximum of 300)
13-18	NUE	Total number of materials for the elastic solution (maximum of 20)
19-24	NUC	Total number of materials for the creep solution (maximum of 20)
25-30	NGI	Gravity loading condition NGI=0 no gravity load NGI=1 gravity load considered
31-36	NLC	External load condition NLC=0 no external loads NLC=1 external loads considered
37-42	KV	Creep condition KV=0 only an elastic solution is performed KV=1 creep solution is required

```
*****
CARD TYPE NO. 5 - NODAL POINT DATA [Format(3I5,2F10.3)]
*****
```

Columns *****	Variable *****	Entry *****
1-5 6-10	I ND(I,1)	Node X-direction boundary condition code for node I ND(I,1)=0 fixed in X-direction ND(I,1)=1 free
11-15	ND(I,2)	Y-direction boundary condition code for node I ND(I,2)=0 fixed in Y-direction ND(I,2)=1 free
16-25	X(I)	X coordinate for node number I
26-35	Y(I)	Y coordinate for node number I

```
*****
```

-one node per card

```

*****
CARD TYPE NO. 6 - ELASTIC MATERIAL PROPERTIES
                  [Format(F12.0,2F6.3)]
*****

```

Columns *****	Variable *****	Entry *****
1-12	EI(I)	Young's Modulus of material number I
13-18	UI(I)	Poisson's Ratio of material number I
19-24	UNI(I)	Unit weight of material number I

```

*****
- there should be one card for each material up to a maximum of
  20

```

```

*****
CARD TYPE NO. 7 - ELEMENT INCIDENCES [Format(6X,5I6)]
*****

```

Columns *****	Variable *****	Entry *****
7-12	NI	Corner node number I
13-18	NJ	Corner node number J
19-24	NK	Corner node number K
25-30	MET	Elastic material number indicator for each element
31-36	MMET(I)	Creep material indicator for each element

```

*****
- one element per card
- node numbers I, J, K are entered in a counter-clockwise
  direction; the starting node is arbitrary.

```

```
*****
CARD TYPE NO. 8 - EXTERNAL LOADING [Format(I6)]
*****
```

Columns *****	Variable *****	Entry *****
1-6	NNL	Number of noads loaded

```
*****
```

-if NLC=0, card type no. 7 as well as card type no. 8 must not appear.

```
*****
CARD TYPE NO. 9 - EXTERNAL LOADS [Format(I6,2F12.3)]
*****
```

Columns *****	Variable *****	Entry *****
1-6	JNU	Node number of loaded node
7-18	FF(1)	Load in X-direction
19-30	FF(2)	Load in Y-direction

```
*****
```

-one node per card

-this card set must not appear if NLC=0

```

*****
CARD TYPE NO. 10 - NODAL PRINT OUT CONTROL [Format(12I6)]
*****

```

Columns *****	Variable *****	Entry *****
1-6	NNPR(I)	Nodal print out control indicator for node number I NNPR(I)= 0 print out and no rotation NNPR(I)=-1 do not print NNPR(I)= k print out and rotate an angle of k degrees. (k=an integer)
7-12	NNPR(I+1)	Nodal print out control indicator for node number I+1
13-18	NNPR(I+2)	Nodal print out control indicator for node number I+2
.	.	.
.	.	.
.	.	.
.	.	.
67-72	NNPR(I+5)	Nodal print out control indicator for node number I+5

```

*****

```

-twelve nodes per card

```

*****
CARD TYPE NO. 11 - ELEMENT PRINT OUT CONTROL [(Format (12I6))]
*****

```

Columns *****	Variable *****	Entry *****
1-6	NEPR(I)	Element print out control indicator for element number I NEPR(I)= 0 Print out and no rotation NEPR(I)=-1 do not print NEPR(I)= k print out and rotate an angle of k degrees (k=an integer)
7-12	NEPR(I+1)	Element print out control indicator for element number I+1
13-18	NEPR(I+2)	Element print out control indicator for element number I+2
.	.	.
.	.	.
.	.	.
.	.	.
67-72	NEPR(I+5)	Element print out control indicator for element number I+5

```

*****

```

-twelve elements per card

-if KV = 0 (i.e., an elastic solution only) no more data is required

-if a creep solution is desired, the following card sets must be included.


```

*****
CARD TYPE NO. 12 - CREEP SOLUTION CONTROL CARD
                  [Format(F12.0,3I6)]
*****

```

Columns *****	Variable *****	Entry *****
1-12	TDM	Maximum time allowed for creep solution
13-18	MNI	Maximum number of creep increments
19-24	NCPR	Creep solution print control NCPR=0 print all output NCPR=1 print selective output
25-30	NSPO	Total number of creep increments to be printed

```

*****

```

```

*****
CARD TYPE NO. 13 - CREEP SOLUTION PRINT CONTROL [Format (12I6)]
*****

```

Columns *****	Variable *****	Entry *****
1-6	IPRIN(1)	The increment number for which the first set of creep results are to be printed
.	.	.
.	.	.
.	.	.
.	.	.
.	.	.
.	IPRIN(IMAX)	The increment number for which the Ith set of creep results are to be printed (IMAX = NSPO)

```

*****

```

```

*****
CARD TYPE NO. 14 - CREEP MATERIAL PROPERTIES
                  [Format(2(E15.4,F5.2),F5.2)]
*****

```

Columns *****	Variable *****	Entry *****
1-15	COEF1(I)	Coefficient in the creep law
16-20	EXP1(I)	Stress exponent in the creep law
21-35	COEF2(I)	Coefficient in the creep law
36-40	EXP2(I)	Stress exponent in the creep law
41-45	EXP3(I)	Time exponent in the creep law

```

*****

```

-one card for each material

-the general form of the power law creep relationship is:

$$\dot{\epsilon}_e = [A_1 \sigma_e^{n_1} + A_2 \sigma_e^{n_2}] t^m$$

where $\dot{\epsilon}_e$ = effective strain rate

σ_e = effective strain

t = time

A_1, A_2 = coefficients

n_1, n_2 = creep exponents

m = time exponent

 CARD TYPE NO. 15 - CREEP SOLUTION CONVERGENCE PARAMETERS
 [Format(3f10.3)]

Columns *****	Variable *****	Entry *****
1-10	ETA0	Maximum ratio of effective elastic strain to increment of effective creep strain for first time interval
11-20	ETA1	Maximum ratio of effective elastic strain to increment of effective creep strain for succeeding time intervals
21-30	OMEGA	Maximum fractional change in effective stress allowed per creep increment

-to ensure convergence, ETA0, ETA1 and OMEGA should lie within the following ranges:

$$\begin{aligned}
 10 &\leq \text{ETA0} \leq 25 \\
 1 &\leq \text{ETA1} \leq 10 \\
 0.03 &\leq \text{OMEGA} \leq 0.10
 \end{aligned}$$

-default values have been set equal to:

$$\begin{aligned}
 \text{ETA0} &= 25 \\
 \text{ETA1} &= 10 \\
 \text{OMEGA} &= 0.03
 \end{aligned}$$

-these default values will produce the minimum time increment

```

*****
CARD TYPE NO. 16 - NODE PRINT OUT CONTROL (CREEP SOLUTION ONLY)
                  [Format(12I6)]
*****

```

Columns *****	Variable *****	Entry *****
1-6	NNPR(I)	Node print out control indicator for node number I NNPR(I)= 0 print out and no rotation NNPR(I)=-1 do not print NNPR(I)= k print out and rotate an angle of k degrees (k=an integer)
7-12	NNPR(I+1)	Node print out control indicator for node number I+1
13-18	NNPR(I+2)	Node print out control indicator for node number I+2
.	.	.
.	.	.
.	.	.
.	.	.
67-72	NNPR(I+5)	Node print out control indicator for node number I+5

```

*****

```

-twelve nodes per card

 CARD TYPE NO. 17 - ELEMENT PRINT OUT CONTROL (CREEP SOLUTION
 ONLY) [Format (12I6)]

Columns *****	Variable *****	Entry *****
1-6	NEPR(I)	Element print out control indicator for element number I NEPR(I)= 0 Print out and no rotation NEPR(I)=-1 do not print NEPR(I)= k print out and rotate an angle of k degrees (k=an integer)
7-12	NEPR(I+1)	Element print out control indicator for element number I+1
13-18	NEPR(I+2)	Element print out control indicator for element number I+2
.	.	.
.	.	.
.	.	.
.	.	.
67-72	NEPR(I+5)	Element print out control indicator for element number I+5

-twelve elements per card

APPENDIX D

PROGRAMME LISTING FOR CREEP

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:29	PAGE 0002
56.000	0015	C	COMMON /BLK15/	ETAO, ETA1, OMEGA, RHO, BETA		
57.000	0016	C	COMMON /BLK16/	MMET(300), COEF1(20), COEF2(20), EXP1(20), EXP2(20), EXP3(20)		
58.000						
59.000						
60.000		C	COMMON /BLK17/	CODE1, CODE2		
61.000	0017	C				
62.000		C				
63.000		C				
64.000		C				
65.000		C				
66.000	0018		READ(5,50)((TITLE(I),I=1,20)			
67.000	0019		WRITE(6,51)((TITLE(I),I=1,20)			
68.000	0020		READ(5,49)CODE1, CODE2			
69.000	0021		READ(5,3)ITAPE, JTAPE			
70.000	0022		IF(CODE1.EQ.RSTART)GOTO 68			
71.000	0023		READ(5,3)NN, NE, NUC, NGI, NLC, KV			
72.000	0024		WRITE(6,4)NN, NE, NUC, NLC			
73.000	0025		WRITE(6,12)			
74.000	0026		IF(NGI.EQ.0.AND.NLC.EQ.1)WRITE(6,13)			
75.000	0027		IF(NGI.EQ.1.AND.NLC.EQ.0)WRITE(6,14)			
76.000	0028		IF(NGI.EQ.1.AND.NLC.EQ.1)WRITE(6,15)			
77.000		C				
78.000		C	-----NN = NUMBER OF NODES			
79.000		C	-----NE = NUMBER OF ELEMENTS			
80.000		C	-----NUE = NUMBER OF DIFFERENT MATERIALS FOR THE ELASTIC			
81.000		C	-----NUC = SOLUTION (MAXIMUM OF 20)			
82.000		C	-----NLC = NUMBER OF DIFFERENT MATERIALS FOR THE CREEP			
83.000		C	-----NGI = SOLUTION (MAXIMUM OF 20)			
84.000		C	-----KV = GRAVITY LOADING CONDITION			
85.000		C	NGI = 0 NO GRAVITY LOAD			
86.000		C	NGI = 1 GRAVITY LOAD CONSIDERED			
87.000		C	-----NLC = EXTERNAL LOAD CONDITION			
88.000		C	NLC = 0 NO EXTERNAL LOADS APPLIED			
89.000		C	NLC = 1 EXTERNAL LOADS CONSIDERED			
90.000		C	-----KV = CREEP CONDITION			
91.000		C	KV = 0 ONLY AN ELASTIC SOLUTION IS PERFORMED			
92.000		C	KV = 1 CREEP CONSIDERED			
93.000		C				
94.000	0029		READ(5,5)((ND(I,J),J=1,2),X(I),Y(I),I=1,NN)			
95.000		C				
96.000		C	-----ND(I,J) FIXITY FACTORS FOR EACH NODE			
97.000		C	ND(I,J) = 0 FIXED			
98.000		C	ND(I,J) = 1 FREE			
99.000		C	-----X(I) = X-COORDINATE FOR NODE I			
100.000		C	-----Y(I) = Y-COORDINATE FOR NODE I			
101.000		C				
102.000	0030		WRITE(6,35)			
103.000	0031		WRITE(6,37)((I,ND(I,J),J=1,2),X(I),Y(I),I=1,NN)			
104.000		C				
105.000	0032		REWIND 1			
106.000	0033		REWIND 2			
107.000	0034		NN2=2*NN			
108.000	0035		NN=1			
109.000	0036		DO 1 I=1,NN			
110.000	0037		DO 1 J=1,2			

MTS	FORTAN IV G COMPILER(21.8) WITH SDS SUPPORT	MAIN	04-21-80	11:35:29	PAGE 0003
111.000	0038	IF(ND(II,JJ).EQ.O) GO TO 1			
112.000	0039	ND(II,JJ)=NU			
113.000	0040	NU=NU+1			
114.000	0041	1			
115.000	0042	CONTINUE			
116.000		NU=NU-1			
117.000		C			
118.000	0043	DO 102 I=1,NUE			
119.000	0044	READ(5,6)E(I),UI(I),UNI(I)			
120.000	0045	102 CONTINUE			
121.000		C			
122.000		C-----EI = ELASTIC MODULUS			
123.000		C-----UI = POISSON'S RATIO			
124.000		C-----UNI = UNIT WEIGHT			
125.000	0046	SARR=0.0			
126.000	0047	IKK=0			
127.000	0048	IFP=0			
128.000	0049	NS=0			
129.000	0050	UC=0			
130.000		C			
131.000	0051	WRITE(6,23)			
132.000	0052	DO 104 I=1,NE			
133.000	0053	READ(5,8)(NODES(I,J),J=1,3),MET(I),MMET(I)			
134.000	0054	IF(MET(I).EQ.O)MET(I)=1			
135.000	0055	IF(MMET(I).EQ.O)MMET(I)=1			
136.000		C			
137.000		C-----NODES(I,J) = ELEMENT CORNER NODES (ENTERED COUNTERCLOCKWISE)			
138.000		C-----MET(I) = ELASTIC MATERIAL NUMBER INDICATOR (UP TO 20)			
139.000		C-----MMET(I) = CREEP MATERIAL NUMBER INDICATOR (UP TO 20)			
140.000		C			
141.000	0056	WRITE(6,36)I,(NODES(I,J),J=1,3),MET(I),MMET(I)			
142.000	0057	104 CONTINUE			
143.000		C			
144.000	0058	DO 105 J=1,NE			
145.000	0059	NI=NODES(J,1)			
146.000	0060	NJ=NODES(J,2)			
147.000	0061	NK=NODES(J,3)			
148.000		C			
149.000		C-----DETERMINE THE ELEMENT AREA AND THE TRANSPOSE MATRIX P			
150.000		C			
151.000	0062	CALL GEOM			
152.000		C			
153.000	0063	MNE=MET(J)			
154.000	0064	UN=UNI(MNE)			
155.000	0065	ARW=(UN*ARR)/3.			
156.000	0066	SARR=SARR+ARR			
157.000		C			
158.000		C-----DETERMINE THE CODE NUMBERS, BAND WIDTH, AND THE NODAL			
159.000		C GRAVITY LOADING FOR EACH ELEMENT			
160.000		C			
161.000	0067	CALL CNODE			
162.000		C			
163.000		C-----CALCULATE THE ELEMENT STIFFNESS MATRIX S			
164.000		C			
165.000	0068	CALL FSM			

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:29	PAGE 0004
166.000	C		C-----WRITE ELEMENT DATA ON TAPE			
167.000	C					
168.000						
169.000	0069		WRITE(1)(NP(1),I=1,6),(PI(K),K=1,9),ARR			
170.000	0070		WRITE(2)(NP(1),I=1,6),((S(MM,MN),MN=1,6),MM=1,6)			
171.000						
172.000	0071		C			
173.000			C			
174.000	0072		C			
175.000	0073		C			
176.000			C			
177.000			C			
178.000			C			
179.000			C			
180.000	0074		C			
181.000	0075		C			
182.000	0076		C			
183.000	0077		C			
184.000	0078		C			
185.000	0079		C			
186.000	0080		C			
187.000			C			
188.000			C			
189.000			C			
190.000	0081		C			
191.000	0082		C			
192.000	0083		C			
193.000	0084		C			
194.000	0085		C			
195.000	0086		C			
196.000	0087		C			
197.000	0088		C			
198.000	0089		C			
199.000	0090		C			
200.000	0091		C			
201.000	0092		C			
202.000	0093		C			
203.000	0094		C			
204.000	0095		C			
205.000	0096		C			
206.000	0097		C			
207.000	0098		C			
208.000	0099		C			
209.000			C			
210.000			C			
211.000			C			
212.000			C			
213.000			C			
214.000			C			
215.000			C			
216.000			C			
217.000			C			
218.000			C			
219.000	0100		C			
220.000			C			

MTS FORTRAN IV G COMPILER(21.B) WITH SDS SUPPORT				MAIN	04-21-80	11:55:29	PAGE 0005
221.000	0101		CALL NLOAD				
222.000		C					
223.000	0102	22	CONTINUE				
224.000	0103		DO 124 JJ=1,NU				
225.000	0104		F(JJ)=0.0				
226.000	0105		F(JJ)=F(JJ)+FG(JJ)				
227.000	0106		FN(JJ)=FG(JJ)				
228.000	0107	124	CONTINUE				
229.000		C					
230.000		C	-----PRINT OUT NODAL LOADS				
231.000		C					
232.000	0108		WRITE(6,28)				
233.000	0109		WRITE(6,18)				
234.000	0110		DO 16 I=1,NU				
235.000	0111	16	WRITE(6,19)I,EI(I),UI(I),UNI(I)				
236.000	0112		WRITE(6,80)				
237.000		C					
238.000	0113		CALL CNLP				
239.000		C					
240.000		C	-----SOLVE FOR THE ELASTIC DISPLACEMENTS				
241.000		C					
242.000	0114		CALL BAND(1)				
243.000		C					
244.000		C	-----PRINT OUT THE ELASTIC DISPLACEMENTS				
245.000		C					
246.000	0115		WRITE(6,30)				
247.000		C					
248.000	0116		CALL CNDI				
249.000		C					
250.000		C	-----READ IN THE NODAL AND ELEMENT PRINT OUT CONTROL VECTORS				
251.000		C					
252.000	0117		READ(5,3)(NNPR(II),II=1,NN)				
253.000		C					
254.000		C	-----NNPR(II) = NODAL PRINT OUT CONTROL				
255.000		C	NNPR(II) = 0 PRINT OUT AND NO ROTATION				
256.000		C	NNPR(II) = -1 DO NOT PRINT				
257.000		C	NNPR(II) = K PRINT OUT AND ROTATE AN ANGLE OF K				
258.000		C	DEGREES				
259.000		C					
260.000	0118		READ(5,3)(NEPR(II),II=1,NE)				
261.000		C					
262.000		C	-----NEPR(II) = ELEMENT PRINT OUT CONTROL				
263.000		C	NEPR(II) = 0 PRINT OUT AND NO ROTATION				
264.000		C	NEPR(II) = -1 DO NOT PRINT				
265.000		C	NEPR(II) = K PRINT OUT AND ROTATE AN ANGLE OF K				
266.000		C	DEGREES				
267.000		C					
268.000		C	-----PRINT OUT THE ELASTIC ELEMENT STRAINS AND STRESSES				
269.000		C					
270.000	0119		WRITE(6,33)				
271.000	0120		WRITE(6,46)				
272.000		C					
273.000	0121		CALL CESS				
274.000		C					
275.000		C	-----CALCULATE AND PRINT OUT THE AVERAGE NODAL STRESSES				

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				MAIN	04-21-80	11:35:29	PAGE 0006
276.000	C						
277.000	0122	WRITE(6,34)					
278.000	0123	WRITE(6,47)					
279.000	C						
280.000	0124	CALL CNAVE					
281.000	C						
282.000	0125	IF(KV.EQ.O) STOP					
283.000	C						
284.000	C	*****					
285.000	C	*					
286.000	C	* START CREEP SOLUTION *					
287.000	C	*					
288.000	C	*****					
289.000	C						
290.000	0126	WRITE(6,900)					
291.000	0127	WRITE(6,42)					
292.000	C						
293.000	C	-----READ ALL THE NECESSARY DATA OFF OF TAPE IF THE CURRENT RUN					
294.000	C	IS A CONTINUATION OF A PREVIOUS PROBLEM THAT WAS TERMINATED					
295.000	C	EARLY					
296.000	C						
297.000	0128	68 IF(CODE1.EQ.RSTART)CALL INPUT					
298.000	C						
299.000	C	-----CONTROL AND TYPE OF CREEP SOLUTION REQUIRED.					
300.000	C						
301.000	0129	READ(5,57)TOM,MNI,NCPR,NSPO					
302.000	C						
303.000	C	-----TDM = MAXIMUM TIME ALLOWED FOR CREEP SOLUTION					
304.000	C	-----MNI = MAXIMUM NUMBER OF CREEP INCREMENTS					
305.000	C	-----NCPR = PRINT CONTROL					
306.000	C	NCPR = 0 PRINT ALL OUTPUT					
307.000	C	NCPR = 1 PRINT SELECTIVE OUTPUT					
308.000	C	-----NSPO = TOTAL NUMBER OF CREEP INCREMENTS TO BE PRINTED					
309.000	C						
310.000	0130	READ(5,3)((IPRIN(I),I=1,NSPO)					
311.000	C						
312.000	C	-----IPRIN(I) = VECTOR CONTAINING THE INCREMENT NUMBERS FOR WHICH					
313.000	C	THE CREEP RESULTS ARE TO BE PRINTED					
314.000	C						
315.000	0131	IF(CODE1.EQ.RSTART)GOTO 70					
316.000	C						
317.000	C	-----READ CREEP MATERIAL PROPERTIES					
318.000	C						
319.000	0132	WRITE(6,65)					
320.000	0133	DO 69 I=1,NUC					
321.000	0134	READ(5,62)COEF1(I),EXP1(I),COEF2(I),EXP2(I),EXP3(I)					
322.000	0135	IF(EXP3(I).EQ.O)EXP3(I)=1.0					
323.000	C						
324.000	C	-----COEF1(I),COEF2(I) = COEFFICIENTS IN THE CREEP LAW					
325.000	C	-----EXP1(I),EXP2(I) = STRESS EXPONENTS IN THE CREEP LAW					
326.000	C	-----EXP3(I) = TIME EXPONENT IN THE CREEP LAW					
327.000	C						
328.000	0136	WRITE(6,66))COEF1(I),COEF2(I),EXP1(I),EXP2(I),EXP3(I)					
329.000	0137	CONTINUE					
330.000	C						

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:29	PAGE 0007
331.000	0138		READ(5,21)ETA0,ETA1,OMEGA			
332.000		C				
333.000		C	C-----ETA0,ETA1,OMEGA = CREEP SOLUTION CONVERGENCE PARAMETERS			
334.000		C				
335.000	0139		IF(ETA0.EQ.O.O)ETA0=25.			
336.000	0140		IF(ETA1.EQ.O.O)ETA1=1.			
337.000	0141		IF(OMEGA.EQ.O.O)OMEGA=0.1			
338.000	0142		WRITE(6,24)ETA0,ETA1,OMEGA			
339.000		C				
340.000	0143	70	IF(NCPR.EQ.O.O)GOTO 71			
341.000	0144		READ(5,3)(NNPR(II),II=1,NN)			
342.000	0145		READ(5,3)(NEPR(II),II=1,NE)			
343.000	0146		GOTO 74			
344.000	0147	71	CONTINUE			
345.000	0148		DO 72 II=1,NN			
346.000	0149	72	NNPR(II)=O			
347.000	0150		DO 73 II=1,NE			
348.000	0151	73	NEPR(II)=O			
349.000		C				
350.000	0152	74	IPR=1			
351.000	0153		IF(CODE1.EQ.START)TIM=0.0			
352.000		C				
353.000		C	C-----THIS SECTION PROVIDES THE NON-LINEAR CREEP SOLUTION.			
354.000		C				
355.000	0154	501	JC=JC+1			
356.000	0155		IKK=IKK+1			
357.000	0156		IPP=1			
358.000	0157		IF(JC.EQ.IPRIN(IPR))IPP=0			
359.000	0158		IF(IPP.EQ.O)IPR=IPR+1			
360.000		C				
361.000		C	C-----GENERATE THE TIME INTERVAL FOR THE INCREMENT OF CREEP			
362.000		C				
363.000	0159		CALL TIMGEN			
364.000		C				
365.000	0160		TIM=TIM+TAV			
366.000	0161		IF(IPP.EQ.O)GOTO 480			
367.000	0162		WRITE(6,53)JC,TAV,TIM			
368.000	0163		GO TO 1544			
369.000	0164	480	IF(JC.GT.1)WRITE(6,900)			
370.000	0165		WRITE(6,53)JC,TAV,TIM			
371.000	0166	1544	CONTINUE			
372.000		C				
373.000		C	C-----COMPUTE THE CREEP STRAIN INCREMENTS AND THE INCREMENTAL			
374.000		C	FICTICIOUS LOAD VECTOR			
375.000		C				
376.000	0167		CALL TSME			
377.000		C				
378.000		C	C-----SUM THE FORCES TO OBTAIN TOTAL APPLIED LOAD			
379.000		C				
380.000	0168		DO 505 JU=1,NU			
381.000	0169		F(JU)=F(JU)+FN(JU)			
382.000	0170	505	CONTINUE			
383.000		C				
384.000		C	C-----PRINT THE INCREMENTAL AND TOTAL NODAL LOADS			
385.000		C				

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				MAIN	04-21-80	11:35:29	PAGE 0008
386.000	0171		IF(IPP.EQ.1)GOTO 81				
387.000	0172		WRITE(6,54)JC				
388.000		C					
389.000	0173		CALL CNLP				
390.000		C					
391.000		C	C-----SOLVE FOR THE NEW INCREMENTAL NODAL DISPLACEMENTS				
392.000		C					
393.000	0174	81	CALL BAND(2)				
394.000		C					
395.000		C	C-----PRINT THE INCREMENTAL AND TOTAL NODAL DISPLACEMENTS				
396.000		C					
397.000	0175		IF(IPP.EQ.1)GOTO 1545				
398.000	0176		WRITE(6,55)JC				
399.000	0177	1545	CONTINUE				
400.000		C					
401.000	0178		CALL GNDI				
402.000		C					
403.000		C	C-----PRINT OUT THE TOTAL AND INCREMENTAL ELEMENT STRAINS AND				
404.000		C	STRESSES				
405.000		C					
406.000	0179		IF(IPP.EQ.1)GOTO 1546				
407.000	0180		WRITE(6,56)JC				
408.000	0181		WRITE(6,46)				
409.000	0182	1546	CONTINUE				
410.000		C					
411.000	0183		CALL CESS				
412.000		C					
413.000		C	C-----COMPUTE AND PRINT OUT THE TOTAL AND INCREMENTAL AVERAGE				
414.000		C	NODAL STRESSES				
415.000		C					
416.000	0184		IF(IPP.EQ.1)GOTO 1547				
417.000	0185		WRITE(6,58)JC				
418.000	0186		WRITE(6,47)				
419.000	0187	1547	CONTINUE				
420.000		C					
421.000	0188		CALL CNAVE				
422.000		C					
423.000		C	C-----CHECK TO SEE IF A STEADY STATE CONDITION HAS BEEN ACHIEVED				
424.000		C					
425.000	0189		CALL STEADY				
426.000		C					
427.000		C	C-----CHECK TO SEE IF THE MAXIMUM TIME OR MAXIMUM NUMBER OF				
428.000		C	INCREMENTS HAVE BEEN EXCEEDED				
429.000		C					
430.000	0190		IF(TIM.LT.TDM.AND.JC.LT.MNI)GOTO 501				
431.000	0191		IF(TIM.GE.TDM) WRITE(6,60)TIM.TDM				
432.000	0192		IF(JC.GE.MNI)WRITE(6,61)JC.MNI				
433.000		C					
434.000		C	C-----WRITE THE COMPLETE SET OF RESULTS FOR THE FINAL CREEP				
435.000		C	INCREMENT AND WRITE THE NECESSARY DATA ON TAPE IF A				
436.000		C	CONTINUATION IS DESIRED AT A LATER DATE				
437.000		C					
438.000	0193		CALL FINAL				
439.000		C					
440.000		C	C-----FORMAT STATEMENTS				

MTS	FORTAN IV G COMPILER(21.B)	WITH SDS SUPPORT	MAIN	04-21-80	11:35:29	PAGE 0009
441.000	C					
442.000	0194	3	FORMAT(1216)			
443.000	0195	4	FORMAT(1216)			
444.000		8	'DESCRIBING THE FINITE ELEMENTS'//2X,'**',43X,'**//2X,'** INPUT DATA '			
445.000		8	'ELEMENTS'//5X,'NUMBER OF NODES'//2X,'**',43X,'**//2X,'**'//5X,'NUMBER OF '			
446.000		8	'ELEMENTS'//5X,'NUMBER OF MATERIALS FOR ELASTIC '			
447.000		8	'SOLUTION'//5X,'NUMBER OF MATERIALS FOR CREEP '			
448.000		8	'SOLUTION'//5X,'NUMBER OF MATERIALS FOR CREEP '			
449.000	0196	5	FORMAT(3(4X,213,2F7.3))			
450.000	0197	6	FORMAT(F12.0,2F6.3)			
451.000	0198	8	FORMAT(6X,516)			
452.000	0199	12	FORMAT(11/5X,'LOADING CASES CONSIDERED:')			
453.000	0200	13	FORMAT(11/5X,'EXTERNAL LOADING ONLY')			
454.000	0201	14	FORMAT(11/5X,'GRAVITY LOADING ONLY')			
455.000	0202	15	FORMAT(11/5X,'EXTERNAL + GRAVITY LOADING')			
456.000	0203	17	FORMAT(1X,'SIZE OF THE PROBLEM IS ',16,' PROGRAM CAPACITY ',			
457.000		8	'EXCEEDED.')			
458.000	0204	18	FORMAT(11/5X,'MATERIAL PROPERTIES:')			
459.000	0205	19	FORMAT(11/5X,'MATERIAL NO. ',13,3X,'YOUNG'S MODULUS = ',1PE15.4,'/23X,			
460.000		8	'POISSON'S RATIO = ',0PF15.3,'/23X,' UNIT WEIGHT = ',			
461.000		8	'1PE15.4')			
462.000	0206	21	FORMAT(3F10.3)			
463.000	0207	23	FORMAT(11/5X,'ELEMENT INCIDENCE LIST (NODES ENTERED COUNTER-',			
464.000		8	'CLOCKWISE')//5X,'ELEMENT I-NODE J-NODE K-NODE,			
465.000		8	'MATERIAL NUMBER CREEP LAW INDICATOR')//			
466.000	0208	24	FORMAT(11/5X,'CREEP SOLUTION CONVERGENCE PARAMETERS: ETA0 = ',			
467.000		8	'F6.2/45X,'ETA1 = ',F6.2/45X,'OMEGA = ',F6.2)			
468.000	0209	28	FORMAT(11/5X,'RESULTS 4//2X,'**',26X,'**//2X,'** ELASTIC SOLUTION',			
469.000		8	'RESULTS 4//2X,'**',26X,'**//2X,'**')//			
470.000	0210	30	FORMAT(11/5X,'NODAL DISPLACEMENTS',//5X,'NODAL			
471.000		8	'Y-DISPLACEMENT',//			
472.000	0211	33	FORMAT(11/5X,'ELEMENT STRAINS AND STRESSES')			
473.000	0212	34	FORMAT(11/5X,'AVERAGE NODAL STRESSES')			
474.000	0213	35	FORMAT(11/5X,'NODAL POINT DATA',//2(5X,'NODAL			
475.000		8	'POINT COORDINATES',5X)/2(5X,'NUMBER',7X,			
476.000		8	'X Y',12X,'X-COORD Y-COORD',5X)/)			
477.000	0214	36	FORMAT(5X,15,1X,3110.11X,12,18X,12)			
478.000	0215	37	FORMAT(2(5X,14,6X,214,7X,2F12.3,5X))			
479.000	0216	42	FORMAT(11/5X,'CREEP SOLUTION RESULTS *'			
480.000		8	'2X,'**',24X,'**//2X,'**')//			
481.000	0217	46	FORMAT(11/5X,'ELEMENT X-STRAIN Y-STRAIN XY-STRAIN X-STRESS			
482.000		8	'Y-STRESS XY-STRESS MAX-STRESS MIN-STRESS X-SHEAR			
483.000		8	'DIRECTION',//			
484.000	0218	47	FORMAT(11/5X,'NODAL X-STRESS Y-STRESS XY-STRESS MAX STRESS',			
485.000		8	'MIN STRESS MAX SHEAR DIRECTION')//			
486.000	0219	49	FORMAT(20A8)			
487.000	0220	50	FORMAT(20A4)			
488.000	0221	51	FORMAT(11/5X,'TITLE OF PROBLEM : ',20A4)			
489.000	0222	53	FORMAT(11/5X,'CREEP INCREMENT NUMBER',15,5X,'INCREMENT OF TIME = ',			
490.000		8	'1PE13.3/2X,21/5X,'TOTAL ELAPSED TIME',1PE13.3)			
491.000	0223	54	FORMAT(11/5X,'NODAL LOADS DUE TO CREEP INCREMENT',14//14X,			
492.000		8	'INCREMENT OF LOAD TOTAL APPLIED LOAD',5X,'NODAL',			
493.000		8	'21/ X-LOAD Y-LOAD')//			
494.000	0224	55	FORMAT(11/5X,'NODAL DISPLACEMENTS AND DISPLACEMENT RATE DUE TO ',			
495.000		8	'CREEP INCREMENT',15//18X,'INCREMENT OF DISPLACEMENT',8X,			

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:29	PAGE 0010
496.000			&	'TOTAL ACCUMULATED DISPLACEMENT', 14X, 'DISPLACEMENT RATE', /		
497.000			&	5X, 'NODE', 1X, 2, 'X-DISPLACEMENT', 'Y-DISPLACEMENT', /		
498.000			&	8X, 'X-VELOCITY', 8X, 'Y-VELOCITY', /		
499.000	0225		&	FORMAT(///5X, 'ELEMENT STRAINS AND STRESSES DUE TO CREEP',		
500.000			&	'INCREMENT', 14/7X, '(FIRST LINE GIVES TOTAL, SECOND LINE',		
501.000			&	'GIVES INCREMENT')',		
502.000	0226		&	FORMAT(F12.0, 516)		
503.000	0227		&	FORMAT(///5X, 'AVERAGE NODAL STRESSES FOR CREEP INCREMENT', 14/7X,		
504.000			&	'(FIRST LINE GIVES TOTAL, SECOND LINE GIVES INCREMENT)')',		
505.000	0228		&	FORMAT(///1X, F12.0, 'MINUTES REQUESTED, SOLUTION TERMINATED')',		
506.000			&	'OF', 1F12.0, 'MINUTES REQUESTED, SOLUTION TERMINATED')',		
507.000	0229		&	FORMAT(///2X, 'CREEP INCREMENT #', 15, 'EXCEEDS THE MAXIMUM OF', 15,		
508.000			&	'REQUESTED, SOLUTION TERMINATED')',		
509.000	0230		&	FORMAT(2(E15.4, F5.2), F5.2)		
510.000	0231		&	FORMAT(///5X, 'CREEP LAW PARAMETERS:')',		
511.000	0232		&	FORMAT(///5X, 'CREEP LAW NO.', 13, 3X, 2, 'COEFFICIENT =', 1PE11.3, 5X),		
512.000			&	/24X, 2, 'EXPONENT =', 0PF11.3, 5X/35X, 'TIME EXPONENT =',		
513.000			&	1PE10.3)		
514.000	0233		&	FORMAT(1B14)		
515.000	0234		&	FORMAT(///5X, 'NODAL LOADS', /5X, 'NODE', 'X-LOAD', 'Y-LOAD', /)		
516.000	0235		&	FORMAT(1,)		
517.000	0236		&	FORMAT(///2X, 'HALF BANDWIDTH =', 15)		
518.000	0237			STOP		
519.000	0238			END		

MTS	FORTAN IV G COMPILER(21.B) WITH SDS SUPPORT	GEOM	04-21-80	11:35:34	PAGE 0002
575.000	0013	ARR=((XI-XJ)*(VI-YK)-(XI-XK)*(VI-YJ))/2.			
576.000	0014	IF(ARR.GT.0.0)GOTO 1			
577.000	0015	WRITE(6,10)J			
578.000	0016	STOP			
579.000		C			
580.000		C-----COMPUTE COEFFICIENTS OF THE SHAPE FUNCTIONS			
581.000		C			
582.000	0017	PT(1)=XJ*YK-XK*YJ			
583.000	0018	PT(2)=XK*VI-XI*YK			
584.000	0019	PT(3)=XI*YJ-XJ*YI			
585.000	0020	PT(4)=YJ-YK			
586.000	0021	PT(5)=YK-YI			
587.000	0022	PT(6)=VI-YJ			
588.000	0023	PT(7)=XK-XJ			
589.000	0024	PT(8)=XI-XK			
590.000	0025	PT(9)=XJ-XI			
591.000		C			
592.000		C-----FORMAT STATEMENTS			
593.000		C			
594.000	0026	10 FORMAT(///5X,'ERROR: ELEMENT NUMBER',I4,' HAS A NEGATIVE AREA. ',			
595.000		8'SOLUTION TERMINATED')			
596.000	0027	RETURN			
597.000	0028	END			

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				MAIN	04-21-80	11:35:34	PAGE 0001
598.000	C	*****	C	*****	*****	*****	*****
599.000	C	*****	C	*****	*****	*****	*****
600.000	C	*****	C	*****	*****	*****	*****
601.000	C	*****	C	*****	*****	*****	*****
602.000	C	*****	C	*****	*****	*****	*****
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616.000	C	*****	C	*****	*****	*****	*****
617.000	C	*****	C	*****	*****	*****	*****
618.000	C	*****	C	*****	*****	*****	*****
619.000	C	*****	C	*****	*****	*****	*****
620.000	C	*****	C	*****	*****	*****	*****
621.000	C	*****	C	*****	*****	*****	*****
622.000	C	*****	C	*****	*****	*****	*****
623.000	C	*****	C	*****	*****	*****	*****
624.000	C	*****	C	*****	*****	*****	*****
625.000	C	*****	C	*****	*****	*****	*****
626.000	C	*****	C	*****	*****	*****	*****
627.000	C	*****	C	*****	*****	*****	*****
628.000	C	*****	C	*****	*****	*****	*****
629.000	C	*****	C	*****	*****	*****	*****
630.000	C	*****	C	*****	*****	*****	*****
631.000	C	*****	C	*****	*****	*****	*****
632.000	C	*****	C	*****	*****	*****	*****
633.000	C	*****	C	*****	*****	*****	*****
634.000	C	*****	C	*****	*****	*****	*****
635.000	C	*****	C	*****	*****	*****	*****
636.000	C	*****	C	*****	*****	*****	*****
637.000	C	*****	C	*****	*****	*****	*****
638.000	C	*****	C	*****	*****	*****	*****
639.000	C	*****	C	*****	*****	*****	*****
640.000	C	*****	C	*****	*****	*****	*****
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643.000	C	*****	C	*****	*****	*****	*****
644.000	C	*****	C	*****	*****	*****	*****
645.000	C	*****	C	*****	*****	*****	*****
646.000	C	*****	C	*****	*****	*****	*****
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649.000	C	*****	C	*****	*****	*****	*****
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651.000	C	*****	C	*****	*****	*****	*****
652.000	C	*****	C	*****	*****	*****	*****

MTS FORTRAN IV G COMPILER(21.8) WITH SOS SUPPORT			CNODE	04-21-80	11:35:34	PAGE 0002
659.000	0010	NB=O				
654.000	0011	DO 12 IJ=1,NU				
655.000	0012	FG(IJ)=O.O				
656.000	0013	12 CONTINUE				
657.000	0014	13 CONTINUE				
658.000		C				
659.000		C-----SET UP DEGREE OF FREEDOM CODE MATRIX				
660.000		C				
661.000	0015	NP(1)=ND(NI,1)				
662.000	0016	NP(2)=ND(NJ,1)				
663.000	0017	NP(3)=ND(NK,1)				
664.000	0018	NP(4)=ND(NI,2)				
665.000	0019	NP(5)=ND(NJ,2)				
666.000	0020	NP(6)=ND(NK,2)				
667.000		C				
668.000		C-----COMPUTE NODAL LOADS DUE TO GRAVITY.				
669.000		C				
670.000	0021	IF(ND(NI,2).EQ.O)GOTO 10				
671.000	0022	IK=ND(NI,2)				
672.000	0023	FG(IK)=EG(IK)-ARW				
673.000	0024	CONTINUE	10			
674.000	0025	IF(ND(NJ,2).EQ.O)GOTO 11				
675.000	0026	IK=ND(NJ,2)				
676.000	0027	FG(IK)=FG(IK)-ARW				
677.000	0028	CONTINUE	11			
678.000	0029	IF(ND(NK,2).EQ.O)GOTO 25				
679.000	0030	IK=ND(NK,2)				
680.000	0031	FG(IK)=FG(IK)-ARW				
681.000	0032	CONTINUE	25			
682.000		C				
683.000		C-----COMPUTE THE MAXIMUM HALF BAND WIDTH OF CURRENT PROBLEM				
684.000		C				
685.000	0033	MAX=O				
686.000	0034	MIN=3000				
687.000	0035	DO 5 KK=1,6				
688.000	0036	IF(NP(KK).EQ.O)GOTO 5				
689.000	0037	IF(NP(KK).MAX)6,6,7				
690.000	0038	MAX=NP(KK)	7			
691.000	0039	IF(NP(KK)-MIN)8,5,5	6			
692.000	0040	MIN=NP(KK)	8			
693.000	0041	CONTINUE	5			
694.000	0042	NB1=MAX-MIN				
695.000	0043	IF(NB1.GT.NB) NB=NB1				
696.000	0044	RETURN				
697.000	0045	END				

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:35	PAGE 0001
698.000	C	*****				
699.000	C	*				
700.000	C	*				
701.000	C	*	FSM			
702.000	C	*				
703.000	C	*				
704.000	C	*				
705.000	C	*				
706.000	C					
707.000	C		SUBROUTINE FSM			
708.000	C					
709.000	C	CC				
710.000	C					
711.000	C		SUBROUTINE FSM - DETERMINES THE INDIVIDUAL ELEMENT STIFFNESS			
712.000	C		MATRICIES			
713.000	C					
714.000	C		DATA REQUIRED FROM COMMON BLOCKS			
715.000	C					
716.000	C		PT = MATRIX CONTAINING THE COEFFICIENTS OF THE SHAPE FUNCTIONS			
717.000	C		E = YOUNG'S MODULUS			
718.000	C		U = POISSON'S RATIO			
719.000	C		ARR = AREA OF ELEMENT			
720.000	C					
721.000	C		OUTPUT PLACED IN BLK9			
722.000	C					
723.000	C		S = ELEMENT STIFFNESS MATRIX			
724.000	C					
725.000	C	CC				
726.000	C					
727.000	C					
728.000	C		C-----COMMON AND DIMENSION STATEMENTS			
729.000	C					
730.000	C					
731.000	C		REAL*8 PT,S,ARR,EI,UI,E,U,EU,R(6,6),P(6,6)			
732.000	C		INTEGER*2 MET,NP			
733.000	C					
734.000	C		COMMON /BLK2/ MET(300),EI(20),UI(20)			
735.000	C					
736.000	C		COMMON /BLK9/ PT(9),S(6,6),NP(6)			
737.000	C					
738.000	C		COMMON /BLK10/ J,NJ,NK,ARR			
739.000	C					
740.000	C		C-----INITIALIZATION			
741.000	C					
742.000	C		DO 1 II=1,6			
743.000	C		DO 1 JJ=1,6			
744.000	C		S(II,JJ)=0.0			
745.000	C		P(II,JJ)=0.0			
746.000	C		CONTINUE			
747.000	C					
748.000	C		C-----SELECT THE COEFFICIENTS OF THE STRAIN-DISPLACEMENT MATRIX			
749.000	C					
750.000	C		DO 110 II=1,3			
751.000	C		DO 110 JJ=1,3			
752.000	C		KK=3*(II-1)+JJ			

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			FSM	O4-21-80	11:35:35	PAGE 0002
753.000	0015	P(II,JJ)=PT(KK)				
754.000	0016	MJJ=II+3				
755.000	0017	NJJ=JJ+3				
756.000	0018	P(MJJ,NJJ)=P(II,JJ)				
757.000	0019	110 CONTINUE				
758.000		C				
759.000		C-----SET UP CONSTITUTIVE MATRIX				
760.000		C				
761.000	0020	MNE=MET(J)				
762.000	0021	E=E1(MNE)				
763.000	0022	U=UI(MNE)				
764.000	0023	EU = E/((1.+U)*(1.-2.*U))				
765.000	0024	S(2,2)=EU*(1.-U)				
766.000	0025	S(2,6)=EU*U				
767.000	0026	S(3,3)=EU*(1.-2.*U)/2.				
768.000	0027	S(3,5)=S(3,3)				
769.000	0028	S(5,3)=S(3,3)				
770.000	0029	S(5,5)=S(3,3)				
771.000	0030	S(6,2)=S(2,6)				
772.000	0031	S(6,6)=S(2,2)				
773.000		C				
774.000		C-----CALCULATE ELEMENT STIFFNESS MATRIX				
775.000		C				
776.000	0032	DO 4 JJ=1,6				
777.000	0033	DO 4 II=1,6				
778.000	0034	R(JJ,II)=0.0				
779.000	0035	DO 4 KK=1,6				
780.000	0036	R(JJ,II)=R(JJ,II)+S(JJ,KK)*P(KK,II)				
781.000	0037	CONTINUE				
782.000	0038	DO 5 JJ=1,6				
783.000	0039	DO 5 II=1,6				
784.000	0040	S(JJ,II)=0.0				
785.000	0041	DO 5 KK=1,6				
786.000	0042	S(JJ,II)=S(JJ,II)+P(KK,JJ)*R(KK,II)/(4.*ARR)				
787.000	0043	CONTINUE				
788.000	0044	RETURN				
789.000	0045	END				

MIS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:36	PAGE 0001
790.000	C	*****				
791.000	C	*****				
792.000	C	*				
793.000	C	*	NLOAD			
794.000	C	*				
795.000	C	*****				
796.000	C	*****				
797.000	C	*****				
798.000	C	*****				
799.000	C	*****				
800.000	C	*****	SUBROUTINE NLOAD			
801.000	C	*****				
802.000	C	*****				
803.000	C	*****				
804.000	C	*****				
805.000	C	*****				
806.000	C	*****				
807.000	C	*****				
808.000	C	*****				
809.000	C	*****				
810.000	C	*****				
811.000	C	*****				
812.000	C	*****				
813.000	C	*****				
814.000	C	*****				
815.000	C	*****				
816.000	C	*****				
817.000	C	*****				
818.000	C	*****				
819.000	C	*****				
820.000	C	*****				
821.000	C	*****				
822.000	C	*****				
823.000	C	*****				
824.000	C	*****				
825.000	C	*****				
826.000	C	*****				
827.000	C	*****				
828.000	C	*****				
829.000	C	*****				
830.000	C	*****				
831.000	C	*****				
832.000	C	*****				
833.000	C	*****				
834.000	C	*****				
835.000	C	*****				
836.000	C	*****				
837.000	C	*****				
838.000	C	*****				
839.000	C	*****				
840.000	C	*****				
841.000	C	*****				
842.000	C	*****				
843.000	C	*****				
844.000	C	*****				

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:37	PAGE 0001
867.000	C	*****				
868.000	C	*****				
869.000	C	*				*
870.000	C	*	CNLP			*
871.000	C	*				*
872.000	C	*****				*
873.000	C	*****				*
874.000	C	*****				*
875.000	0001		SUBROUTINE CNLP			
876.000	C					
877.000	C	CC				
878.000	C	CC				
879.000	C					
880.000	C	SUBROUTINE CNLP - PRINTS OUT INCREMENTAL AND TOTAL NODAL				
881.000	C	LOADS				
882.000	C					
883.000	C	DATA REQUIRED FROM COMMON BLOCKS				
884.000	C					
885.000	C	JC = CREEP INCREMENT NUMBER				
886.000	C	NN = TOTAL NUMBER OF NODES				
887.000	C	ND = RESTRAINED DEGREE OF FREEDOM CODE MATRIX				
888.000	C	F = ACCUMULATED TOTAL LOAD VECTOR				
889.000	C	FN = INCREMENTAL FICTICIOUS LOAD VECTOR				
890.000	C					
891.000	C	CC				
892.000	C					
893.000	C					
894.000	C	C-----DIMENSION STATEMENTS				
895.000	C					
896.000	0002	REAL*8 FN,F,TAV				
897.000	0003	INTEGER*2 ND				
898.000						
899.000	0004	COMMON /BLK1/ ND(200,2)				
900.000						
901.000	0005	COMMON /BLK3/ FN(400),F(400)				
902.000						
903.000	0006	COMMON /BLK11/ NN,NE,NUE,NUC,NGI,NLC,KV,NJ,NB,NB1				
904.000						
905.000	0007	COMMON /BLK12/ JC,TIM,TAV,IPP				
906.000						
907.000		C-----INITIALIZATION				
908.000						
909.000	0008	SUMX1=0.0				
910.000	0009	SUMY1=0.0				
911.000	0010	SUMX2=0.0				
912.000	0011	SUMY2=0.0				
913.000	0012	DO 1 JJ=1,NN				
914.000						
915.000		C-----SELECT INCREMENTAL AND TOTAL X-FORCES FROM FN AND F				
916.000						
917.000	0013	IF(ND(JJ,1).NE.0)GOTO 4				
918.000	0014	FX1=0.0				
919.000	0015	FX2=0.0				
920.000	0016	GOTO 5				
921.000	0017	4 IK=ND(JJ,1)				

MIS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				CNLP	04-21-80	11:35:37	PAGE 0002
922.000	0018		FX1=FN(IK)				
923.000	0019		FX2=F(IK)				
924.000	0020	5	CONTINUE				
925.000		C					
926.000		C	-----SELECT INCREMENTAL AND TOTAL Y-FORCES FROM FN AND F				
927.000		C					
928.000	0021		IF(ND(JJ,2).NE.O)GOTO 6				
929.000	0022		FY1=O.O				
930.000	0023		FY2=O.O				
931.000	0024		GOTO 7				
932.000	0025	6	IK=ND(JJ,2)				
933.000	0026		FY1=FN(IK)				
934.000	0027		FY2=F(IK)				
935.000	0028	7	CONTINUE				
936.000		C					
937.000		C	-----SUM FORCES TO OBTAIN TOTAL APPLIED LOAD				
938.000		C					
939.000	0029		SUMX1=SUMX1+FX1				
940.000	0030		SUMY1=SUMY1+FY1				
941.000	0031		SUMX2=SUMX2+FX2				
942.000	0032		SUMY2=SUMY2+FY2				
943.000		C					
944.000		C	-----PRINT OUT INCREMENTAL AND TOTAL LOADS				
945.000		C					
946.000	0033		IF(JC.EQ.O)GOTO 10				
947.000	0034		WRITE(6,8) JJ,FX1,FY1,FX2,FY2				
948.000	0035		GOTO 1				
949.000	0036	10	WRITE(6,9)JJ,FX1,FY1				
950.000	0037	1	CONTINUE				
951.000	0038		IF(JC.EQ.O)GOTO 15				
952.000	0039		WRITE(6,11)SUMX1,SUMY1,SUMX2,SUMY2				
953.000	0040		GOTO 20				
954.000	0041	15	WRITE(6,12)SUMX1,SUMY1				
955.000		C					
956.000		C	-----FORMAT STATEMENTS				
957.000		C					
958.000	0042	8	FORMAT(SX,IX,IX,4F11.2)				
959.000	0043	9	FORMAT(SX,IX,IX,2F11.2)				
960.000	0044	11	FORMAT(/T7,'SUM OF INCREMENTAL FORCES: X-DIRECTION = ',E10.4,				
961.000		&	/T35,'Y-DIRECTION = ',E10.4//T7,'SUM OF TOTAL FORCES: ',				
962.000		&	'X-DIRECTION = ',E10.4//T29,'Y-DIRECTION = ',E10.4//				
963.000	0045	12	FORMAT(/T7,'SUM OF APPLIED LOADING: X-DIRECTION = ',E10.4,				
964.000		&	/T32,'Y-DIRECTION = ',E10.4//				
965.000	0046	20	RETURN				
966.000	0047		END				

MTS FORTRAN IV G COMPILER(21.B) WITH SDS SUPPORT			MAIN	04-21-80	11:35:38	PAGE 0001
967.000	C				
968.000	C				
969.000	C				
970.000	C				
971.000	C	BAND			
972.000	C				
973.000	C				
974.000	C				
975.000	C				
976.000	C	SUBROUTINE BAND(LT)			
977.000	C				
978.000	C				
979.000	C				
980.000	C				
981.000	C				
982.000	C				
983.000	C	INPUT			
984.000	C				
985.000	C	LT = EXECUTION PARAMETER			
986.000	C	LT=1: DECOMPOSE STIFFNESS MATRIX AND BACK SUBSTITUTE			
987.000	C	LT=2: BACK SUBSTITUTE ONLY			
988.000	C				
989.000	C	DATA REQUIRED FROM COMMON BLOCKS			
990.000	C				
991.000	C	A = ASSEMBLED STRUCTURE STIFFNESS MATRIX			
992.000	C	B = LOAD VECTOR			
993.000	C	N = TOTAL NUMBER OF UNKNOWN			
994.000	C	M = MAXIMUM HALF BAND WIDTH			
995.000	C				
996.000	C	OUTPUT PLACED IN BLK3			
997.000	C				
998.000	C	B = NODAL X AND Y DISPLACEMENTS			
999.000	C				
1000.000	C				
1001.000	C				
1002.000	C				
1003.000	C				
1004.000	C				
1005.000	C				
1006.000	C				
1007.000	C	COMMON /BLK3/ B(400)			
1008.000	C				
1009.000	C	COMMON /BLK11/ NP,NE,NUE,NUC,NGI,NLC,KV,N,NB,M			
1010.000	C				
1011.000	C	COMMON /BLK13/ A(5000)			
1012.000	C				
1013.000	C	DET=1.E-8			
1014.000	C	MM=M-1			
1015.000	C	NM=N*M			
1016.000	C	NM1=N*M-NM			
1017.000	C	IF (LT.NE.1) GO TO 55			
1018.000	C	MP=M+1			
1019.000	C	KK=2			
1020.000	C	FAC=DET			
1021.000	C	A(1)=1./DSORT(A(1))			

MTS FORTRAN IV G COMPILER(21.B) WITH SDS SUPPORT			BAND	04-21-80	11:35:38	PAGE 0002
1022.000	0015	BIGL=A(1)				
1023.000	0016	SML=A(1)				
1024.000	0017	A(2)=A(2)*A(1)				
1025.000	0018	A(MP)=1./DSORT(A(MP)-A(2)*A(2))				
1026.000	0019	IF(A(MP).GT.BIGL)BIGL=A(MP)				
1027.000	0020	IF(A(MP).LT.SML)SML=A(MP)				
1028.000	0021	MP=MP+M				
1029.000	0022	DO 62 J=MP,NM1,M				
1030.000	0023	JP=J-MM				
1031.000	0024	MZC=O				
1032.000	0025	IF(KK.GE.M) GO TO 1				
1033.000	0026	KK=KK+1				
1034.000	0027	II=1				
1035.000	0028	JC=1				
1036.000	0029	GO TO 2				
1037.000	0030	KK=KK+M	1			
1038.000	0031	II=KK-MM				
1039.000	0032	JC=KK-MM				
1040.000	0033	DO 65 I=KK,JP,MM	2			
1041.000	0034	IF(A(I).EQ.O.)GO TO 64				
1042.000	0035	GO TO 66				
1043.000	0036	JC=JC+M	64			
1044.000	0037	MZC=MZC+1	65			
1045.000	0038	ASUM1=O.				
1046.000	0039	GO TO 61				
1047.000	0040	MMZC=MM+MZC	66			
1048.000	0041	II=II+MZC				
1049.000	0042	KM=KK+MMZC				
1050.000	0043	A(KM)=A(KM)*A(JC)				
1051.000	0044	IF(KM.GE.JP)GO TO 6				
1052.000	0045	KJ=KM+MM				
1053.000	0046	DO 5 I=KJ,JP,MM				
1054.000	0047	ASUM2=O.				
1055.000	0048	IM=I-MM				
1056.000	0049	II=II+I				
1057.000	0050	KI=II+MMZC				
1058.000	0051	DO 7 K=KM,IM,MM				
1059.000	0052	ASUM2=ASUM2+A(KI)*A(K)				
1060.000	0053	KI=KI+MM	7			
1061.000	0054	A(I)=(A(I)-ASUM2)*A(KI)	5			
1062.000	0055	CONTINUE	6			
1063.000	0056	ASUM1=O.				
1064.000	0057	DO 4 K=KM,JP,MM				
1065.000	0058	ASUM1=ASUM1+A(K)*A(K)	4			
1066.000	0059	S=A(J)-ASUM1	61			
1067.000	0060	IF(S.LT.O.)DET=S				
1068.000	0061	IF(S.EQ.O.)DET=O				
1069.000	0062	IF(S.GT.O.)GO TO 63				
1070.000	0063	NROW=(J+MM)/M				
1071.000	0064	WRITE(6,99) NROW				
1072.000	0065	FORMAT(35HOERROR CONDITION ENCOUNTERED IN ROW,16)	99			
1073.000	0066	GOTO 56				
1074.000	0067	A(J)=1./DSORT(S)	63			
1075.000	0068	IF(A(J).GT.BIGL)BIGL=A(J)				
1076.000	0069	IF(A(J).LT.SML)SML=A(J)				

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			BAND	04-21-80	11:35:38	PAGE 0003
1077.000	0070	62	CONTINUE			
1078.000	0071		IF(SML.E.FAC*BIGL)GO TO 54			
1079.000	0072		GO TO 53			
1080.000	0073	54	DET=0.			
1081.000	0074	56	IF(DET)70,71,72			
1082.000	0075	70	WRITE(6,73)DET			
1083.000	0076	73	FORMAT(IX,'MATRIX SM IS NOT POSITIVE DEFINITE. DET=',E15.7)			
1084.000	0077		STOP			
1085.000	0078	71	WRITE(6,74)DET			
1086.000	0079	74	FORMAT(IX,'DETERMINANT IS ZERO. DET=',E15.7)			
1087.000	0080		STOP			
1088.000	0081	72	CONTINUE			
1089.000	0082	53	DET=SML/BIGL			
1090.000	0083	55	B(1)=B(1)*A(1)			
1091.000	0084		KK=1			
1092.000	0085		K1=1			
1093.000	0086		J=1			
1094.000	0087		DO 8 L=2,N			
1095.000	0088		BSUM1=0.			
1096.000	0089		LM=L-1			
1097.000	0090		J=J+M			
1098.000	0091		IF(KK.GE.M)GO TO 12			
1099.000	0092		KK=KK+1			
1100.000	0093		GO TO 13			
1101.000	0094	12	KK=KK+M			
1102.000	0095		K1=K1+1			
1103.000	0096	13	JK=KK			
1104.000	0097		DO 9 K=K1,LM			
1105.000	0098		BSUM1=BSUM1+A(JK)*B(K)			
1106.000	0099		JK=JK+MM			
1107.000	0100	9	CONTINUE			
1108.000	0101	8	B(L)=(B(L)-BSUM1)*A(J)			
1109.000	0102		B(N)=B(N)*A(NM1)			
1110.000	0103		NM=N-1			
1111.000	0104		NN=N-1.			
1112.000	0105		ND=N			
1113.000	0106		DO 10 L=1,NN			
1114.000	0107		BSUM2=0.			
1115.000	0108		NL=N-L			
1116.000	0109		NL1=N-L+1			
1117.000	0110		NM1=N-1			
1118.000	0111		NJ1=N-1			
1119.000	0112		IF(L.GE.M)ND=ND-1			
1120.000	0113		DO 11 K=NL1,ND			
1121.000	0114		NJ1=NJ1+1			
1122.000	0115		BSUM2=BSUM2+A(NJ1)*B(K)			
1123.000	0116	11	CONTINUE			
1124.000	0117	10	B(NL)=(B(NL)-BSUM2)*A(NM)			
1125.000	0118		RETURN			
1126.000	0119		END			

MTS FORTRAN IV G COMPILER(21.B) WITH SDS SUPPORT				MAIN	04-21-80	11:35:41	PAGE 0001
1127.000	C						
1128.000	C						
1129.000	C	*					
1130.000	C	*					
1131.000	C	*	CNDI				
1132.000	C	*					
1133.000	C	*					
1134.000	C						
1135.000	0001		SUBROUTINE CNDI				
1136.000	C						
1137.000	C						
1138.000	C						
1139.000	C						
1140.000	C		SUBROUTINE CNDI - PRINTS OUT INCREMENTAL AND TOTAL				
1141.000	C		DISPLACEMENTS				
1142.000	C						
1143.000	C		DATA REQUIRED FROM COMMON BLOCKS				
1144.000	C						
1145.000	C		NN = TOTAL NUMBER OF NODES				
1146.000	C		ND = RESTRAINED DEGREE OF FREEDOM CODE MATRIX				
1147.000	C		FN = NODAL X AND Y DISPLACEMENTS				
1148.000	C		IPP = INCREMENT PRINT OUT CONTROL FLAG				
1149.000	C		JC = CREEP INCREMENT NUMBER				
1150.000	C						
1151.000	C		OUTPUT PLACED IN BLK4				
1152.000	C						
1153.000	C		DX = TOTAL NODAL X-DISPLACEMENTS				
1154.000	C		DY = TOTAL NODAL Y-DISPLACEMENTS				
1155.000	C		DXI = INCREMENTAL X-DISPLACEMENT				
1156.000	C		DYI = INCREMENTAL Y-DISPLACEMENT				
1157.000	C		VELX = CREEP VELOCITY IN X-DIRECTION FOR CURRENT INCREMENT				
1158.000	C		VELY = CREEP VELOCITY IN Y-DIRECTION FOR CURRENT INCREMENT				
1159.000	C						
1160.000	C						
1161.000	C						
1162.000	C						
1163.000	C		-----COMMON AND DIMENSION STATEMENTS				
1164.000	C						
1165.000	0002		REAL*8 FN,TAV				
1166.000	0003		INTEGER*2 ND				
1167.000	C						
1168.000	0004		COMMON /BLK1/ ND(200,2)				
1169.000	C						
1170.000	0005		COMMON /BLK3/ FN(400)				
1171.000	C						
1172.000	0006		COMMON /BLK4/ DX(200),DY(200),DXI(200),DYI(200),				
1173.000			VELX(200),VELY(200)				
1174.000	C						
1175.000	0007		COMMON /BLK11/ NN,NE,NUE,NJC,NGI,NLC,KV,NU,NE,NB1				
1176.000	C						
1177.000	0008		COMMON /BLK12/ JC,TIM,TAV,IPP				
1178.000	C						
1179.000	0009		IF(JC.NE.O)GOTO 3				
1180.000	0010		DO 2 I=1,NN				
1181.000	0011		DX(I)=O.O				

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				CNDI	04-21-80	11:35:41	PAGE 0002
1182.000	0012		DY(I)=0.0				
1183.000	0013	2	CONTINUE				
1184.000	0014	3	CONTINUE				
1185.000	0015		DO 1 JJJ=1,NN				
1186.000		C					
1187.000		C	-----SELECT X-DISPLACEMENT FROM FN				
1188.000		C					
1189.000	0016		IF(ND(JJ,1).NE.0)GOTO 5				
1190.000	0017		DXI(JJ)=0.0				
1191.000	0018		GOTO 4				
1192.000	0019	5	IK=ND(JJ,1)				
1193.000	0020		DXI(JJ)=FN(IK)				
1194.000	0021	4	CONTINUE				
1195.000		C					
1196.000		C	-----SELECT Y-DISPLACEMENT FROM FN				
1197.000		C					
1198.000	0022		IF(ND(JJ,2).NE.0)GOTO 6				
1199.000	0023		DYI(JJ)=0.0				
1200.000	0024		GOTO 9				
1201.000	0025	6	IK=ND(JJ,2)				
1202.000	0026		DYI(JJ)=FN(IK)				
1203.000	0027	9	CONTINUE				
1204.000		C					
1205.000		C	-----CALCULATE TOTAL DISPLACEMENTS				
1206.000		C					
1207.000	0028		DX(JJ)=DX(JJ)+DXI(JJ)				
1208.000	0029		DY(JJ)=DY(JJ)+DYI(JJ)				
1209.000		C					
1210.000		C	-----CALCULATE CREEP VELOCITY				
1211.000		C					
1212.000	0030		IF(JC.EQ.0)GOTO 10				
1213.000	0031		VELX(JJ)=DXI(JJ)/TAV				
1214.000	0032		VELY(JJ)=DYI(JJ)/TAV				
1215.000		C					
1216.000		C	-----PRINT OUT INCREMENTAL AND TOTAL DISPLACEMENTS				
1217.000		C					
1218.000	0033		IF(IPP.EQ.1)GOTO 1				
1219.000	0034		IF(JC.EQ.0)GOTO 10				
1220.000	0035		WRITE(6.8)JJ,DXI(JJ),DYI(JJ),DX(JJ),DY(JJ),VELX(JJ),VELY(JJ)				
1221.000	0036		GOTO 1				
1222.000	0037	10	WRITE(6.11)JJ,DXI(JJ),DYI(JJ)				
1223.000	0038	1	CONTINUE				
1224.000		C					
1225.000		C	-----FORMAT STATEMENTS				
1226.000		C					
1227.000	0039		FORMAT(5X,I4,1X,6(1PE18.4))				
1228.000	0040	11	FORMAT(5X,I3,2(1PE18.4))				
1229.000	0041		RETURN				
1230.000	0042		END				

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:42	PAGE 0001
1231.000	C	*****				
1232.000	C	*****				
1233.000	C	*				*
1234.000	C	*				*
1235.000	C	*	CESS			*
1236.000	C	*	*****			*
1237.000	C	*****				*****
1238.000	C					
1239.000	C					
1240.000	C	0001	SUBROUTINE CESS			
1241.000	C					
1242.000	C					
1243.000	C	CC				
1244.000	C					
1245.000	C		SUBROUTINE CESS - CALCULATES AND PRINTS OUT THE ELEMENT STRAINS			C
1246.000	C		AND STRESSES FOR THE ELASTIC SOLUTION AND THE			C
1247.000	C		INCREMENTAL AND TOTAL ELEMENT STRAINS AND			C
1248.000	C		STRESSES FOR THE CREEP SOLUTION			C
1249.000	C					
1250.000	C		DATA REQUIRED FROM COMMON BLOCKS			C
1251.000	C					
1252.000	C		NN = NUMBER OF NODES			C
1253.000	C		NE = NUMBER OF ELEMENTS			C
1254.000	C		KV = CREEP CONDITION INDICATOR			C
1255.000	C		JC = CREEP INCREMENT NUMBER			C
1256.000	C		FN = NODAL X AND Y DISPLACEMENTS			C
1257.000	C		EXL = INCREMENT OF CREEP STRAIN IN X-DIRECTION			C
1258.000	C		EYL = INCREMENT OF CREEP STRAIN IN Y-DIRECTION			C
1259.000	C		EZL = INCREMENT OF CREEP STRAIN IN Z-DIRECTION			C
1260.000	C		EXYL = INCREMENT OF CREEP SHEAR STRAIN			C
1261.000	C		UI = POISSON'S RATIO FOR EACH MATERIAL			C
1262.000	C		EI = ELASTIC MODULUS FOR EACH MATERIAL			C
1263.000	C		MET = ELASTIC MATERIAL NUMBER INDICATOR			C
1264.000	C		NODES = ARRAY CONTAINING THE ELEMENT INCIDENCE LIST			C
1265.000	C		NIPT = PRINT OUT CONTROL INDICATOR FOR EACH ELEMENT			C
1266.000	C		IPP = INCREMENT PRINT OUT CONTROL FLAG			C
1267.000	C					
1268.000	C		OUTPUT PLACED IN BLK6, BLK7 AND BLK8			C
1269.000	C					
1270.000	C		EX = ACCUMULATED TOTAL ELEMENT STRAINS IN X-DIRECTION			C
1271.000	C		EY = ACCUMULATED TOTAL ELEMENT STRAINS IN Y-DIRECTION			C
1272.000	C		EXY = ACCUMULATED TOTAL ELEMENT SHEAR STRAINS			C
1273.000	C		SEX = ACCUMULATED ELASTIC ELEMENT SIGMA-X STRESSES			C
1274.000	C		SEY = ACCUMULATED ELASTIC ELEMENT SIGMA-Y STRESSES			C
1275.000	C		SEZ = ACCUMULATED ELASTIC ELEMENT SIGMA-Z STRESSES			C
1276.000	C		SEXY = ACCUMULATED ELASTIC ELEMENT SHEAR STRESSES			C
1277.000	C		SEXYL = CHANGE IN ELASTIC EFFECTIVE STRESS FOR INCREMENT #JC			C
1278.000	C		SEYL = ELASTIC EFFECTIVE STRESSES FOR INCREMENT #JC			C
1279.000	C		FS1 = ACCUMULATED ELASTIC ELEMENT STRAINS IN X-DIRECTION			C
1280.000	C		FS2 = ACCUMULATED ELASTIC ELEMENT STRAINS IN Y-DIRECTION			C
1281.000	C		FS3 = ACCUMULATED ELASTIC ELEMENT SHEAR STRAINS			C
1282.000	C		EEFFC = EFFECTIVE CREEP STRAIN			C
1283.000	C		CXI = SUM OF ADJOINING ELEMENT SIGMA-X STRESSES FOR EACH			C
1284.000	C		NODE			C
1285.000	C		CYI = SUM OF ADJOINING ELEMENT SIGMA-Y STRESSES FOR EACH			C

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				MAIN	04-21-80	11:35:42	PAGE 0002
1286.000	C	NODE					
1287.000	C	CXVI = SUM OF ADJOINING ELEMENT SHEAR STRESSES FOR EACH NODE					C
1288.000	C	CNC = TOTAL NUMBER OF ELEMENTS CONNECTED TO EACH NODE					C
1289.000	C						C
1290.000	C	CC					C
1291.000	C						
1292.000	C						
1293.000	C	-----DIMENSION STATEMENTS					
1294.000	C						
1295.000	0002	REAL*8 FN,EX,EY,EXY,SEX,SEY,SEZ,SEXY,EYL,EZL,EXYL, & SEXYL,SEYL,XS,YS,XYS,FS1E,FS2E,FS3E,DISP(6),TAV,PT(9), & SIX,SIZ,SIXY,EI,UI,EU,COMM,EXC,EYC,EZC,EXYC,EEFFC & INTEGER*2 NNPR,NPR,NP(6),MET,NODES					
1296.000							
1297.000	0003						
1298.000							
1299.000	C	COMMON /BLK2/ MET(300),EI(20),UI(20)					
1300.000							
1301.000	C	COMMON /BLK3/ FN(400)					
1302.000							
1303.000	C	COMMON /BLK5/ NNPR(300),NEPR(300)					
1304.000	0006						
1305.000	C	COMMON /BLK6/ EX(300),EY(300),EXY(300), & EXL(300),EYL(300),EZL(300),EXYL(300), & EXC(300),EYC(300),EZC(300),EXYC(300), & FS1(300),FS2(300),FS3(300)					
1306.000	0007						
1307.000							
1308.000							
1309.000							
1310.000	C	COMMON /BLK7/ SEX(300),SEY(300),SEZ(300),SEXY(300), & CXI(200),CYI(200),CXII(200),CNC(200), & SNX(200),SNY(200),SNXY(200)					
1311.000	0008						
1312.000							
1313.000							
1314.000	C	COMMON /BLK8/ SEYL(300),SEXYL(300),EEFFC(300)					
1315.000	0009						
1316.000	C	COMMON /BLK11/ NN,NE,NUE,NUC,NGI,NLC,KV,NU,NB,NB1					
1317.000	0010						
1318.000	C	COMMON /BLK12/ JC,TIM,TAV,IPP					
1319.000	0011						
1320.000	C	COMMON /BLK14/ NODES(300.3)					
1321.000	0012						
1322.000	C						
1323.000	C	-----INITIALIZATION					
1324.000	C						
1325.000	C						
1326.000	0013	D0 128 JJ=1,NN					
1327.000	0014	CXI(JJ)=0.0					
1328.000	0015	CYI(JJ)=0.0					
1329.000	0016	CXYI(JJ)=0.0					
1330.000	0017	CNC(JJ)=0.0					
1331.000	0018	CONTINUE					
1332.000	0019	SIGMAX=0.0					
1333.000	0020	IF(JC,NE.0)GOTO 140					
1334.000	0021	D0 130 J=1,NE					
1335.000	0022	EX(J)=0.0					
1336.000	0023	EY(J)=0.0					
1337.000	0024	EXY(J)=0.0					
1338.000	0025	FS1(J)=0.0					
1339.000	0026	FS2(J)=0.0					
1340.000	0027	FS3(J)=0.0					

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				CESS	04-21-80	11:35:42	PAGE 0003
1341.000	0028	SEX(J)=0.0					
1342.000	0029	SEY(J)=0.0					
1343.000	0030	SEZ(J)=0.0					
1344.000	0031	SEXY(J)=0.0					
1345.000	0032	EXL(J)=0.0					
1346.000	0033	EYL(J)=0.0					
1347.000	0034	EZL(J)=0.0					
1348.000	0035	EXYL(J)=0.0					
1349.000	0036	EXC(J)=0.0					
1350.000	0037	EXC(J)=0.0					
1351.000	0038	EZC(J)=0.0					
1352.000	0039	EXYC(J)=0.0					
1353.000	0040	CONTINUE	130				
1354.000	0041	CONTINUE	140				
1355.000		C	C				
1356.000		C-----CALCULATE STRAIN AND STRESS FOR EACH ELEMENT	C				
1357.000		DO 1 J=1,NE					
1358.000	0042	MNE=MET(J)					
1359.000	0043	E=ET(MNE)					
1360.000	0044	U=UI(MNE)					
1361.000	0045		C				
1362.000		C-----READ ELEMENT DATA	C				
1363.000		READ(1)(NP(I),I=1,6),(PT(K),K=1,9),ARR	C				
1364.000	0046		C				
1365.000		C-----SELECT X AND Y DISPLACEMENTS OF EACH NODE IN ELEMENT #J	C				
1366.000		DO 2 K=1,6					
1367.000		DISP(K)=0.0					
1368.000		LL=NP(K)					
1369.000	0047	IF(LL.EQ.0) GO TO 2					
1370.000	0048	DISP(K)=FN(LL)					
1371.000	0049		2				
1372.000	0050	CONTINUE	C				
1373.000	0051		C				
1374.000	0052	C-----CALCULATE INCREMENTAL TOTAL ELEMENT STRAIN	C				
1375.000		XS=(DISP(1)*PT(4)+DISP(2)*PT(5)+DISP(3)*PT(6))/(2.*ARR)					
1376.000		YS=(DISP(1)*PT(7)+DISP(2)*PT(8)+DISP(3)*PT(9)+DISP(4)*PT(4)					
1377.000		& +DISP(5)*PT(5)+DISP(6)*PT(6))/(2.*ARR)					
1378.000	0053	YS=(DISP(4)*PT(7)+DISP(5)*PT(8)+DISP(6)*PT(9))/(2.*ARR)					
1379.000	0054		C				
1380.000	0055	C-----ACCUMULATE TOTAL ELEMENT STRAIN	C				
1381.000		EX(J)=EX(J)+XS					
1382.000		EY(J)=EY(J)+YS					
1383.000		EXY(J)=EXY(J)+XYS					
1384.000	0056		C				
1385.000	0057	C-----CALCULATE INCREMENTAL ELASTIC ELEMENT STRAIN	C				
1386.000	0058	FS1E=XS-EXL(J)					
1387.000		FS2E=YS-EYL(J)					
1388.000		FS3E=XYS-EXYL(J)					
1389.000			C				
1390.000	0059		C				
1391.000	0060						
1392.000	0061		C				
1393.000		C-----ACCUMULATE ELASTIC ELEMENT STRAIN	C				
1394.000							
1395.000							

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				CESS	04-21-80	11:35:42	PAGE 0004
1396.000	C						
1397.000	0062	FS1(J)=FS1(J)+FS1E					
1398.000	0063	FS2(J)=FS2(J)+FS2E					
1399.000	0064	FS3(J)=FS3(J)+FS3E					
1400.000	C						
1401.000	C						
1402.000	C						
1403.000	0065	COMM=E/((1.+U)*(1.-2.*U))					
1404.000	0066	SIX=COMM*((1.-U)*FS1E+U*FS2E)-COMM*U*EZL(J)					
1405.000	0067	SIY=COMM*(U*FS1E+(1.-U)*FS2E)-COMM*U*EZL(J)					
1406.000	0068	SIZ=U*(SIX+SIY)-E*EZL(J)					
1407.000	0069	SIXY=COMM*((1.-2.*U)/2.)*FS3E					
1408.000	0070	SEXYL(J)=DSORT(.5*(SIX-SIY)**2+(SIY-SIZ)**2+(SIZ-SIX)**2 +6.*SIY**2))					
1409.000	8						
1410.000	C						
1411.000	C						
1412.000	C						
1413.000	0071	SEX(J)=SEX(J)+SIX					
1414.000	0072	SEY(J)=SEY(J)+SIY					
1415.000	0073	SEXY(J)=SEXY(J)+SIYX					
1416.000	0074	SEZ(J)=U*(SEX(J)+SEY(J))-E*EZL(J)					
1417.000	0075	SEYL(J)=DSORT(.5*((SEX(J)-SEY(J))**2+(SEY(J)-SEZ(J))**2 +(SEZ(J)-SEXY(J))**2+6.*(SEXY(J)**2)))					
1418.000	8						
1419.000	0076	IF(SEYL(J).LT.SIGMAX)GOTO 108					
1420.000	0077	SIGMAX=SEYL(J)					
1421.000	0078	IELM=J					
1422.000	0079	108 IF(JC.EQ.O)GOTO 109					
1423.000	C						
1424.000	C						
1425.000	C						
1426.000	0080	EXC(J)=EXC(J)+EXL(J)					
1427.000	0081	EYC(J)=EYC(J)+EYL(J)					
1428.000	0082	EZC(J)=EZC(J)+EZL(J)					
1429.000	0083	EXYC(J)=EXYC(J)+EXYL(J)					
1430.000	0084	EEFFC(J)=DSORT((4./3.)*(EXC(J)**2+EXC(J)*EYC(J) +EYC(J)**2+(EXYC(J)**2)/4.))					
1431.000	8						
1432.000	C						
1433.000	C						
1434.000	C						
1435.000	0085	IF(IPP.EQ.1.OR.NEPR(J).LT.O)GOTO 112					
1436.000	0086	CALL SMAX(SEX(J),SEY(J),SEXY(J),XMT,XM2,XM3,XM4)					
1437.000	0087	WRITE(6,5)J,EX(J),EY(J),EXY(J),SEX(J),SEY(J),SEXY(J), XM1,XM2,XM3,XM4					
1438.000	8						
1439.000	0088	IF(NEPR(J).EQ.O)GOTO 110					
1440.000	0089	CALL ROTATE(EX(J),EY(J),EXY(J),NEPR(J),C1,C2,C3)					
1441.000	0090	CALL ROTATE(SEX(J),SEY(J),SEXY(J),NEPR(J),D1,D2,D3)					
1442.000	0091	WRITE(6,7)C1,C2,C3,D1,D2,D3					
1443.000	0092	CONTINUE					
1444.000	C						
1445.000	C						
1446.000	C						
1447.000	0093	IF(JC.EQ.O)GOTO 112					
1448.000	0094	WRITE(6,6)XS,YS,XYS,SIX,SIY,SIXY					
1449.000	0095	WRITE(6,605)FS1E,FS2E,FS3E					
1450.000	0096	WRITE(6,606)FS1(J),FS2(J),FS3(J)					

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MTS  FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT      CESS      04-21-80      11:35:42      PAGE 0005

1451.000 0097      WRITE(6,609)EXL(J),EVL(J),EZL(J),EXYL(J)
1452.000 0098      WRITE(6,608)EXC(J),EYC(J),EYC(J),EYC(J)
1453.000 0099      WRITE(6,607)SEYL(J),SEZ(J)
1454.000 0100      CONTINUE
1455.000 C
1456.000 C-----SUM THE ELEMENT STRESSES THAT ARE CONNECTED TO EACH NODE
1457.000 C
1458.000 DO 77 K=1,3
1459.000 L=NODES(J,K)
1460.000 CXI(L)=CXI(L)+SIX
1461.000 CXY(L)=CXY(L)+SIX
1462.000 CXYI(L)=CXYI(L)+SIXY
1463.000 CMC(L)=CMC(L)+1.
1464.000 C
1465.000 0107 CONTINUE
1466.000 0108 1 CONTINUE
1467.000 C
1468.000 0109      WRITE(6,610)SIGMAX,IELM
1469.000 C
1470.000 C-----FORMAT STATEMENTS
1471.000 C
1472.000 0110 5      FORMAT(5X,I5,2X,9(IPE11.3),OPF10.2)
1473.000 0111      FORMAT(12X,6(IPE11.3))
1474.000 0112      FORMAT(12X,6(IPE11.3), (ROTATED STRAINS AND STRESSES))
1475.000 0113      FORMAT(12X,3(IPE11.3), (ELASTIC STRAIN INCREMENTS))
1476.000 0114      FORMAT(12X,3(IPE11.3), (TOTAL ELASTIC STRAIN))
1477.000 0115      FORMAT(14X,'EFFECTIVE STRESS = ',IPE11.3,14X,'SIGMA-Z = ',
1478.000      & IPE11.3/)
1479.000 0116      FORMAT(12X,4(IPE11.3), (TOTAL ACCUMULATED CREEP STRAIN,
1480.000      & 'X, Y, Z & XY COMPONENTS'))
1481.000 0117      FORMAT(12X,4(IPE11.3), (CREEP STRAIN INCREMENTS,
1482.000      & 'X, Y, Z & XY COMPONENTS'))
1483.000 0118      FORMAT(15X,'MAXIMUM EFFECTIVE STRESS = ',IPE12.3,' OCCURS IN ',
1484.000      & ELEMENT ',I5)
1485.000 0119      REWIND.1
1486.000 0120      RETURN
1487.000 0121      END

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MTS	FORTAN IV G COMPILER(21.8) WITH SDS SUPPORT	MAIN	04-21-80	11:35:47	PAGE 0001
1488.000	C	*****			
1489.000	C	*****			
1490.000	C	*****			
1491.000	C	*****			
1492.000	C	*****			
1493.000	C	*****			
1494.000	C	*****			
1495.000	C	*****			
1496.000	C	*****			
1497.000	C	*****			
1498.000	C	*****			
1499.000	C	*****			
1500.000	C	*****			
1501.000	C	*****			
1502.000	C	*****			
1503.000	C	*****			
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1536.000	C	*****			
1537.000	C	*****			
1538.000	C	*****			
1539.000	C	*****			
1540.000	C	*****			
1541.000	C	*****			
1542.000	C	*****			

MTS FORTRAN IV G COMPILER(21.B) WITH SDS SUPPORT			CNAVE	04-21-80	11:35:47	PAGE 0002
1543.000	0008	DO 1,J=1,N				
1544.000	0009	CXI(J)=CXI(J)/CNC(J)				
1545.000	0010	CYI(J)=CYI(J)/CNC(J)				
1546.000	0011	CXI(J)=CXI(J)/CNC(J)				
1547.000		C				
1548.000		C-----ACCUMULATE AVERAGE NODAL STRESSES				
1549.000		C				
1550.000	0012	SNX(J)=SNX(J)+CXI(J)				
1551.000	0013	SNY(J)=SNY(J)+CYI(J)				
1552.000	0014	SNXY(J)=SNXY(J)+CXI(J)				
1553.000		C				
1554.000		C-----PRINT OUT ACCUMULATED AVERAGE NODAL STRESSES				
1555.000		C				
1556.000	0015	IF(IPP.EQ.1.OR.NNPR(J).LT.O)GOTO 1				
1557.000	0016	CALL SMAX(SNX(J),SNY(J),SNXY(J),XM1,XM2,XM3,XM4)				
1558.000	0017	WRITE(6,4)J,SNX(J),SNY(J),SNXY(J),XM1,XM2,XM3,XM4				
1559.000	0018	IF(NNPR(J).EQ.O)GOTO 112				
1560.000	0019	CALL ROTATE(SNX(J),SNY(J),SNXY(J),S1,S2,S3)				
1561.000	0020	WRITE(6,7) S1,S2,S3				
1562.000	0021	112 CONTINUE				
1563.000		C				
1564.000		C-----PRINT OUT INCREMENTAL AVERAGE NODAL STRESSES				
1565.000		C				
1566.000	0022	IF(JC.EQ.O)GOTO 1				
1567.000	0023	WRITE(6,5)CXI(J),CYI(J),CXI(J)				
1568.000	0024	1 CONTINUE				
1569.000		C				
1570.000		C-----FORMAT STATEMENTS				
1571.000		C				
1572.000	0025	4 FORMAT(5X,14,6(1PE12.3),OPF11.2)				
1573.000	0026	5 FORMAT(9X,3(1PE12.3))				
1574.000	0027	7 FORMAT(9X,3(1PE12.3), (ROTATED STRESSES))				
1575.000	0028	RETURN				
1576.000	0029	END				

MYS FORTRAN IV G COMPILER(21.B) WITH SDS SUPPORT			MAIN	04-21-80	11:35:48	PAGE 0001
1577.000	C	*****				
1578.000	C	*****				
1579.000	C	*****				
1580.000	C	*****				
1581.000	C	*****	SMAX			
1582.000	C	*****				
1583.000	C	*****				
1584.000	C	*****				
1585.000	C	*****				
1586.000	C	*****				
1587.000	C	*****				
1588.000	C	*****				
1589.000	C	*****				
1590.000	C	*****				
1591.000	C	*****				
1592.000	C	*****				
1593.000	C	*****				
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1596.000	C	*****				
1597.000	C	*****				
1598.000	C	*****				
1599.000	C	*****				
1600.000	C	*****				
1601.000	C	*****				
1602.000	C	*****				
1603.000	C	*****				
1604.000	C	*****				
1605.000	C	*****				
1606.000	C	*****				
1607.000	C	*****				
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1613.000	C	*****				
1614.000	C	*****				
1615.000	C	*****				
1616.000	C	*****				
1617.000	C	*****				
1618.000	C	*****				
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1621.000	C	*****				
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1623.000	C	*****				
1624.000	C	*****				
1625.000	C	*****				
1626.000	C	*****				
1627.000	C	*****				
1628.000	C	*****				
1629.000	C	*****				
1630.000	C	*****				
1631.000	C	*****				

[illegible]

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:48	PAGE 0001
1637.000	C	*****				
1638.000	C	*				
1639.000	C	*				
1640.000	C	*	ROTATE			
1641.000	C	*				
1642.000	C	*				
1643.000	C	*				
1644.000	C					
1645.000	0001		SUBROUTINE ROTATE(A,B,C,KK,RSEX,RSEXY)			
1646.000	C					
1647.000	C	CC				
1648.000	C					
1649.000	C					
1650.000	C		SUBROUTINE ROTATE - DETERMINES THE STRESSES IN ANY DESIRED			
1651.000	C		DIRECTION			
1652.000	C					
1653.000	C		INPUT			
1654.000	C					
1655.000	C		A = SIGMA-X STRESS			
1656.000	C		B = SIGMA-Y STRESS			
1657.000	C		C = SHEAR STRESS			
1658.000	C		KK = ANGLE THE STRESSES ARE TO ROTATED			
1659.000	C					
1660.000	C		OUTPUT			
1661.000	C					
1662.000	C		RSEX = ROTATED SIGMA-X STRESS			
1663.000	C		RSEY = ROTATED SIGMA-Y STRESS			
1664.000	C		RSEXY = ROTATED SHEAR STRESS			
1665.000	C	CC				
1666.000	C					
1667.000	C					
1668.000	C					
1669.000	C					
1670.000	C		C-----ROTATE THE STRESSES THROUGH AN ANGLE OF KK DEGREES			
1671.000	C					
1672.000	0002		DEL1= FLOAT(KK)			
1673.000	0003		DEL=DEL1*(3.141593/180.)			
1674.000	0004		RSEX=A*(COS(DEL)**2)+2.*C*COS(DEL)*SIN(DEL)+B*(SIN(DEL)**2)			
1675.000	0005		RSEY=A*(SIN(DEL)**2)-2.*C*COS(DEL)*SIN(DEL)+B*(COS(DEL)**2)			
1676.000	0006		RSEXY=(B-A)*SIN(DEL)*COS(DEL)+C*(COS(DEL)**2-SIN(DEL)**2)			
1677.000	0007		RETURN			
1678.000	0008		END			

MIS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				MAIN	04-21-80	11:35:49	PAGE 0001
1679.000	C	*****					
1680.000	C	*					
1681.000	C	*					
1682.000	C	*					
1683.000	C	*					
1684.000	C	*					
1685.000	C	*					
1686.000	C	*					
1687.000	C	*					
1688.000	C	*					
1689.000	C	*					
1690.000	C	*					
1691.000	C	*					
1692.000	C	*					
1693.000	C	*					
1694.000	C	*					
1695.000	C	*					
1696.000	C	*					
1697.000	C	*					
1698.000	C	*					
1699.000	C	*					
1700.000	C	*					
1701.000	C	*					
1702.000	C	*					
1703.000	C	*					
1704.000	C	*					
1705.000	C	*					
1706.000	C	*					
1707.000	C	*					
1708.000	C	*					
1709.000	C	*					
1710.000	C	*					
1711.000	C	*					
1712.000	C	*					
1713.000	C	*					
1714.000	C	*					
1715.000	C	*					
1716.000	C	*					
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1718.000	C	*					
1719.000	C	*					
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1721.000	C	*					
1722.000	C	*					
1723.000	C	*					
1724.000	C	*					
1725.000	C	*					
1726.000	C	*					
1727.000	C	*					
1728.000	C	*					
1729.000	C	*					
1730.000	C	*					
1731.000	C	*					
1732.000	C	*					
1733.000	C	*					

MIS FORTRAN IV G COMPILER(21.B) WITH SDS SUPPORT				TIMING	04-21-80	11:35:49	PAGE 0002
1734.000							
1735.000							
1736.000							
1737.000							
1738.000	0010						
1739.000							
1740.000							
1741.000							
1742.000	0011						
1743.000	0012						
1744.000	0013						
1745.000	0014						
1746.000	0015						
1747.000	0016						
1748.000	0017						
1749.000	0018						
1750.000	0019						
1751.000	0020						
1752.000	0021						
1753.000	0022						
1754.000	0023						
1755.000							
1756.000	0024						
1757.000	0025						
1758.000							
1759.000	0026						
1760.000	0027						
1761.000	0028						
1762.000							
1763.000	0029						
1764.000	0030						
1765.000	0031						
1766.000							
1767.000							
1768.000							
1769.000	0032						
1770.000	0033						
1771.000	0034						
1772.000	0035						
1773.000	0036						
1774.000	0037						
1775.000	0038						
1776.000	0039						
1777.000	0040						
1778.000							
1779.000	0041						
1780.000							
1781.000	0042						
1782.000	0043						

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:51	PAGE 0001
1783.000	C	*****				
1784.000	C	*****				
1785.000	C	*****				
1786.000	C	*****				
1787.000	C	*****	TSME			
1788.000	C	*****				
1789.000	C	*****				
1790.000	C	*****				
1791.000	C	*****				
1792.000	C	*****	SUBROUTINE TSME			
1793.000	C	*****				
1794.000	C	*****				
1795.000	C	*****				
1796.000	C	*****				
1797.000	C	*****				
1798.000	C	*****				
1799.000	C	*****				
1800.000	C	*****				
1801.000	C	*****				
1802.000	C	*****				
1803.000	C	*****				
1804.000	C	*****				
1805.000	C	*****				
1806.000	C	*****				
1807.000	C	*****				
1808.000	C	*****				
1809.000	C	*****				
1810.000	C	*****				
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1827.000	C	*****				
1828.000	C	*****				
1829.000	C	*****				
1830.000	C	*****				
1831.000	C	*****				
1832.000	C	*****				
1833.000	C	*****				
1834.000	C	*****				
1835.000	C	*****				
1836.000	C	*****				
1837.000	C	*****				

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				TSME	04-21-80	11:35:51	PAGE 0002
1838.000	0003		INTEGER*2 MET,MMET,NP(6)				
1839.000		C					
1840.000	0004		COMMON /BLK2/ MET(300),EI(20)				
1841.000		C					
1842.000	0005		COMMON /BLK3/ FN(400)				
1843.000		C					
1844.000	0006		COMMON /BLK6/ EX(300),EY(300),EXY(300),				
1845.000		C	EXL(300),EYL(300),EZL(300),EXYL(300)				
1846.000							
1847.000	0007		COMMON /BLK7/ SEX(300),SEZ(300),SEXY(300)				
1848.000		C					
1849.000	0008		COMMON /BLK8/ SEYL(300),SEXYL(300),EEFFC(300)				
1850.000		C					
1851.000	0009		COMMON /BLK11/ NN,NE,NUE,NUC,NGI,KV,NU,NB,NB1				
1852.000		C					
1853.000	0010		COMMON /BLK12/ JC,TIM,TAV,IPP				
1854.000		C					
1855.000	0011		COMMON /BLK16/ MMET(300),COEF1(20),COEF2(20),EXP1(20),EXP2(20),				
1856.000		C	EXP3(20)				
1857.000		C					
1858.000		C	-----INITIALIZATION				
1859.000		C					
1860.000	0012		DO 4 I=1,NU				
1861.000	0013		FN(I)=0.0				
1862.000	0014	10	CONTINUE				
1863.000		C					
1864.000	0015		DO 4 J=1,NE				
1865.000	0016		MNE=MET(I)				
1866.000	0017		MNC=MMET(I)				
1867.000	0018		E=E1(MNE)				
1868.000	0019		B1=COEF1(MNC)				
1869.000	0020		B2=COEF2(MNC)				
1870.000	0021		CN1=EXP1(MNC)				
1871.000	0022		CN2=EXP2(MNC)				
1872.000	0023		CM=EXP3(MNC)				
1873.000	0024		CM1=1./CM				
1874.000	0025		A1=B1*(SEYL(I)**CN1)				
1875.000	0026		A2=B2*(SEYL(I)**CN2)				
1876.000		C					
1877.000	0027		IF(JC.EQ.1)GOTO 8				
1878.000	0028		TF=EEFFC(I)/(A1+A2)**CM1				
1879.000	0029		GOTO 9				
1880.000	0030	8	TF=0.0				
1881.000		C					
1882.000		C	-----CALCULATE THE INCREMENTS OF CREEP STRAIN IN EACH DIRECTION				
1883.000		C					
1884.000	0031	9	CRE=(A1+A2)*((TF+TAV)**CM-TF**CM)				
1885.000	0032		EXL(I)=(CRE*(2.*SEX(I)-SEY(I)-SEZ(I)))/(2.*SEYL(I))				
1886.000	0033		EYL(I)=(CRE*(2.*SEX(I)-SEY(I)-SEZ(I)))/(2.*SEYL(I))				
1887.000	0034		EZL(I)=(CRE*(2.*SEX(I)-SEY(I)-SEZ(I)))/(2.*SEYL(I))				
1888.000	0035		EXYL(I)=(3.*CRE*SEXY(I))/SEYL(I)				
1889.000		C					
1890.000		C	-----READ ELEMENT DATA				
1891.000	0036		READ(I)(NP(K),K=1,6),(PT(K),K=1,9),ARR				
1892.000							

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MTS  FORTRAN IV G COMPILER(21.8) WITH SOS SUPPORT      TSME      04-21-80      11:35:51      PAGE 0003

1893.000      C
1894.000      C-----CALCULATE THE FICTICIOUS CREEP FORCES
1895.000      C
1896.000      0037      COMM=E/(1+.5)
1897.000      0038      FCX=COMM*EXL(1)
1898.000      0039      FCY=COMM*EYL(1)
1899.000      0040      FCXY=COMM*EXYL(1)/2
1900.000      0041      FL(1)=(PT(4)*FCX+PT(7)*FCXY)/2.
1901.000      0042      FL(2)=(PT(5)*FCX+PT(8)*FCXY)/2.
1902.000      0043      FL(3)=(PT(6)*FCX+PT(9)*FCXY)/2.
1903.000      0044      FL(4)=(PT(4)*FCXY+PT(7)*FCY)/2.
1904.000      0045      FL(5)=(PT(5)*FCXY+PT(8)*FCY)/2.
1905.000      0046      FL(6)=(PT(6)*FCXY+PT(9)*FCY)/2.
1906.000      C
1907.000      C-----ADD THE FICTICIOUS CREEP FORCES INTO THE LOAD VECTOR
1908.000      C
1909.000      0047      DO 7 J=1,6
1910.000      0048      IF (NP(J).EQ.0)GOTO 7
1911.000      0049      IJ=NP(J)
1912.000      0050      FN(IJ)=FN(IJ)+FL(J)
1913.000      0051      7 CONTINUE
1914.000      0052      4 CONTINUE
1915.000      0053      REWIND 1
1916.000      0054      RETURN
1917.000      0055      END

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MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:53	PAGE 0001
1918.000	C					
1919.000	C					
1920.000	C					
1921.000	C					
1922.000	C					
1923.000	C					
1924.000	C					
1925.000	C					
1926.000	C					
1927.000	C					
1928.000	C					
1929.000	C					
1930.000	C					
1931.000	C					
1932.000	C					
1933.000	C					
1934.000	C					
1935.000	C					
1936.000	C					
1937.000	C					
1938.000	C					
1939.000	C					
1940.000	C					
1941.000	C					
1942.000	C					
1943.000	C					
1944.000	C					
1945.000	C					
1946.000	C					
1947.000	C					
1948.000	C					
1949.000	C					
1950.000	C					
1951.000	C					
1952.000	C					
1953.000	C					
1954.000	C					
1955.000	C					
1956.000	C					
1957.000	C					
1958.000	C					
1959.000	C					
1960.000	C					
1961.000	C					
1962.000	C					
1963.000	C					
1964.000	C					
1965.000	C					
1966.000	C					
1967.000	C					
1968.000	C					
1969.000	C					
1970.000	C					
1971.000	C					
1972.000	C					
1973.000	C					

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				MAIN	04-21-80	11:35:55	PAGE 0001
1990.000	C						
1991.000	C						
1992.000	C	*					
1993.000	C	*					
1994.000	C	*	FINAL				
1995.000	C	*					
1996.000	C	*					
1997.000	C	*					
1998.000	C						
1999.000	C						
0001	C		SUBROUTINE FINAL				
2000.000	C						
2001.000	C						
2002.000	C						
2003.000	C						
2004.000	C						
2005.000	C						
2006.000	C						
2007.000	C						
2008.000	C						
2009.000	C						
2010.000	C						
2011.000	C						
2012.000	C						
2013.000	C						
2014.000	C						
2015.000	C						
2016.000	C						
2017.000	C						
2018.000	C						
2019.000	C						
2020.000	C						
2021.000	C						
2022.000	C						
2023.000	C						
2024.000	C						
2025.000	C						
2026.000	C						
2027.000	C						
2028.000	C						
2029.000	C						
2030.000	C						
2031.000	C						
2032.000	C						
2033.000	C						
2034.000	C						
2035.000	C						
2036.000	C						
2037.000	C						
2038.000	C						
2039.000	C						
2040.000	C						
2041.000	C						
2042.000	C						
2043.000	C						
2044.000	C						

MIS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				FINAL	04-21-80	11:35:55	PAGE 0002
2045.000							
2046.000							
2047.000	0010	C		FS1(300),FS2(300),FS3(300)			
2048.000				COMMON /BLK7/ SEX(300),SEY(300),SEZ(300),SEXY(300),			
2049.000				CXI(200),CXI(200),CXI(200),CNC(200),			
2050.000				SNX(200),SNV(200),SNXY(200)			
2051.000	0011	C		COMMON /BLK8/ SEYL(300),SEXYL(300),EEFFC(300)			
2052.000							
2053.000	0012	C		COMMON /BLK11/ NN,NE,NUE,NUC,NGI,NLC,KV,NU,NB,NB1			
2054.000							
2055.000	0013	C		COMMON /BLK12/ JC,TIM,TAV,IPP,ITAPE,UTAPE			
2056.000							
2057.000	0014	C		COMMON /BLK13/ SM(5000)			
2058.000							
2059.000	0015	C		COMMON /BLK14/ NODES(300,3)			
2060.000							
2061.000	0016	C		COMMON /BLK15/ ETAO,ETA1,OMEGA,RHO,BETA			
2062.000							
2063.000	0017	C		COMMON /BLK16/ MMET(300),COEF1(20),COEF2(20),EXP1(20),EXP2(20),			
2064.000				EXP3(20)			
2065.000							
2066.000	0018	C		COMMON /BLK17/ CODE1,CODE2			
2067.000							
2068.000				C-----PRINT OUT THE NODAL DISPLACEMENTS, ELEMENT STRESSES AND STRAINS			
2069.000				C AND THE NODAL STRESSES FOR THE LAST TIME STEP			
2070.000							
2071.000	0019	C		WRITE(6,600)			
2072.000	0020			WRITE(6,605)JC,TAV,TIM			
2073.000							
2074.000	0021	C		WRITE(6,610)JC			
2075.000	0022			DO 5 I=1,NN			
2076.000	0023			WRITE(6,615)I,DXI(I),DVI(I),DX(I),DV(I),VELX(I),VELY(I)			
2077.000	0024	5		CONTINUE			
2078.000							
2079.000	0025	C		WRITE(6,620)JC			
2080.000	0026			DO 10 J=1,NE			
2081.000	0027			CALL SMAX(SEX(J),SEY(J),SEXY(J),XM1,XM2,XM3,XM4)			
2082.000	0028			WRITE(6,621)J,EX(J),EY(J),EXY(J),SEX(J),SEY(J),SEXY(J),			
2083.000				XM1,XM2,XM3,XM4			
2084.000	0029			WRITE(6,622)FS1(J),FS2(J),FS3(J)			
2085.000	0030			WRITE(6,623)EXC(J),EYC(J),EZC(J),EXYC(J)			
2086.000	0031			WRITE(6,624)SEYL(J),SEZ(J)			
2087.000	0032	10		CONTINUE			
2088.000							
2089.000	0033	C		WRITE(6,630)JC			
2090.000	0034			DO 20 J=1,NN			
2091.000	0035			CALL SMAX(SNX(J),SNV(J),SNXY(J),XM1,XM2,XM3,XM4)			
2092.000	0036			WRITE(6,635)J,SNX(J),SNV(J),SNXY(J),XM1,XM2,XM3,XM4			
2093.000	0037	20		CONTINUE			
2094.000							
2095.000				C-----WRITE ALL DATA ON TAPE REQUIRED FOR A CONTINUATION RUN			
2096.000				C AT A LATER DATE			
2097.000							
2098.000	0038	C		IF(CODE2.EQ.RUN)RETURN			
2099.000	0039			REWIND JTAPE			

[illegible]

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT			MAIN	04-21-80	11:35:59	PAGE 0001
2157.000	C					
2158.000	C					
2159.000	C					
2160.000	C					
2161.000	C					
2162.000	C					
2163.000	C					
2164.000	C					
2165.000	C					
2166.000	C					
2167.000	C					
2168.000	C					
2169.000	C					
2170.000	C					
2171.000	C					
2172.000	C					
2173.000	C					
2174.000	C					
2175.000	C					
2176.000	C					
2177.000	C					
2178.000	C					
2179.000	C					
2180.000	C					
2181.000	C					
2182.000	C					
2183.000	C					
2184.000	C					
2185.000	C					
2186.000	C					
2187.000	C					
2188.000	C					
2189.000	C					
2190.000	C					
2191.000	C					
2192.000	C					
2193.000	C					
2194.000	C					
2195.000	C					
2196.000	C					
2197.000	C					
2198.000	C					
2199.000	C					
2200.000	C					
2201.000	C					
2202.000	C					
2203.000	C					
2204.000	C					
2205.000	C					
2206.000	C					
2207.000	C					
2208.000	C					
2209.000	C					
2210.000	C					
2211.000	C					

MTS FORTRAN IV G COMPILER(21.8) WITH SDS SUPPORT				INPUT	04-21-80	11:35:59	PAGE 0002
2212.000	C						
2213.000	0010	COMMON /BLK8/ SEYL(300), SEXYL(300), EFFFC(300)					
2214.000	C						
2215.000	0011	COMMON /BLK11/ NN, NE, NUC, NUC, NLC, KV, NU, NB, NB1					
2216.000	C						
2217.000	0012	COMMON /BLK12/ JC, TIM, TAV, IPP, ITAPE, JTAPE					
2218.000	C						
2219.000	0013	COMMON /BLK13/ SM(5000)					
2220.000	C						
2221.000	0014	COMMON /BLK14/ NODES(300,3)					
2222.000	C						
2223.000	0015	COMMON /BLK15/ ETAO, ETA1, OMEGA, RHO, BETA					
2224.000	C						
2225.000	0016	COMMON /BLK16/ MMET(300), COEF1(20), COEF2(20), EXP1(20), EXP2(20),					
2226.000	&	EXP3(20)					
2227.000	C						
2228.000	0017	REWIND ITAPE					
2229.000	0018	READ(ITAPE)NN, NE, NUC, NUC, KV, NU, NB, NB1					
2230.000	0019	READ(ITAPE)((ND(1,J), J=1,2), I=1, NN)					
2231.000	0020	READ(ITAPE)((NODES(1,J), J=1,3), MET(1), MMET(1), I=1, NE)					
2232.000	0021	READ(ITAPE)(ET(1), UT(1), I=1, NUC)					
2233.000	0022	READ(ITAPE)(COEF1(1), COEF2(1), EXP1(1), EXP2(1), EXP3(1), I=1, NUC)					
2234.000	0023	READ(ITAPE)(F(1), I=1, NU)					
2235.000	0024	NV=NB1*NU					
2236.000	0025	READ(ITAPE)(SM(1), I=1, NV)					
2237.000	0026	READ(ITAPE)JC, TIM, TAV					
2238.000	0027	READ(ITAPE)ETAO, ETA1, OMEGA, RHO, BETA					
2239.000	0028	READ(ITAPE)(DX(1), DY(1), I=1, NN)					
2240.000	0029	READ(ITAPE)(EX(1), EV(1), EXY(1), I=1, NE)					
2241.000	0030	READ(ITAPE)(EXC(1), EYC(1), E2C(1), EXYC(1), I=1, NE)					
2242.000	0031	READ(ITAPE)(FS1(1), FS2(1), FS3(1), I=1, NE)					
2243.000	0032	READ(ITAPE)(SEX(1), SEXY(1), SEXY(1), I=1, NE)					
2244.000	0033	READ(ITAPE)(SNX(1), SNY(1), SNXY(1), I=1, NN)					
2245.000	0034	READ(ITAPE)(SEYL(1), SEXYL(1), SEFFFC(1), I=1, NE)					
2246.000	0035	DO 5 I=1, NE					
2247.000	0036	READ(ITAPE)(NP(K), K=1, 6), (PT(K), K=1, 9), ARR					
2248.000	0037	WRITE(1)(NP(K), K=1, 6), (PT(K), K=1, 9), ARR					
2249.000	0038	CONTINUE					
2250.000	0039	REWIND 1					
2251.000	0040	REWIND ITAPE					
2252.000	C						
2253.000	0041	WRITE(6, 600)JC, ITAPE, TIM					
2254.000	0042	WRITE(6, 610)JC					
2255.000	C						
2256.000	0043	WRITE(6, 615)					
2257.000	0044	DO 10 I=1, NN					
2258.000	0045	WRITE(6, 616)1, DX(1), DY(1)					
2259.000	C						
2260.000	0046	WRITE(6, 620)					
2261.000	0047	DO 20 I=1, NE					
2262.000	0048	WRITE(6, 621)1, EX(1), EV(1), EXY(1)					
2263.000	0049	WRITE(6, 622)FS1(1), FS2(1), FS3(1)					
2264.000	0050	WRITE(6, 623)EXC(1), EYC(1), E2C(1), EXYC(1)					
2265.000	0051	WRITE(6, 624)SEX(1), SEXY(1), SEXY(1), SEXY(1)					
2266.000	0052	WRITE(6, 625)SEYL(1), SEFFFC(1)					

APPENDIX E

THICK WALL CYLINDER CLOSED FORM SOLUTION

APPENDIX E

THICK WALL CYLINDER CLOSED FORM SOLUTION

This appendix presents a comparison of the finite element results predicted by the programme CREEP and the closed form steady state creep of a thick wall cylinder subjected to a uniform internal pressure.

The finite element grid is shown in Figure E-1. The mesh consists of 90 elements and 61 nodes. The ratio of outer to inner radius is equal to 2.0. The vertical and horizontal boundaries are placed on roller supports which allow translation in the radial direction only.

The inner radius is subjected to a uniform pressure of 100 kPa. The elastic material properties are:

$$E = 10 \times 10^6 \text{ kPa}$$

$$\nu = 0.495$$

The flow law of the material is;

$$\dot{\epsilon} = (1.0 \times 10^{-14}) \sigma^{4.5} \quad (\text{E-1})$$

where the units are $\text{kPa}^{-4.5}$ and hr^{-1} .

The steady state creep velocity for the inner radius is shown in Figure E-2. The closed form solution neglects elastic response and thus maintains a constant value. After approximately 0.75 hr of creep simulation, the finite element solution predicts a constant velocity 2.2% lower than the closed form solution.

The closed form and finite element elastic and steady state effective stresses are compared in Figure E-3. As shown in the figure, there is very good agreement for both elastic and steady state creep stresses.

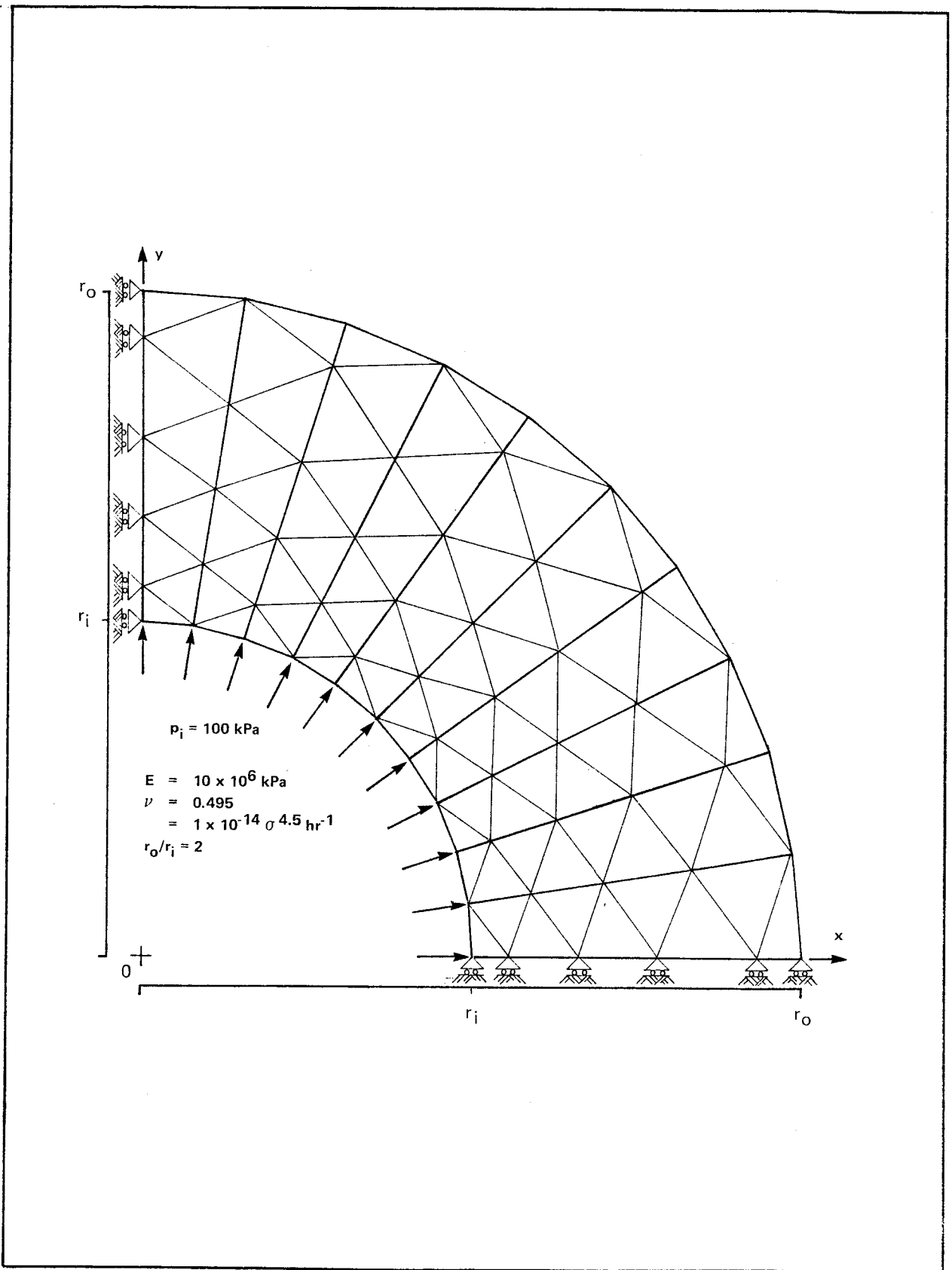


FIGURE E-1

THICK WALL CYLINDER FINITE ELEMENT MESH

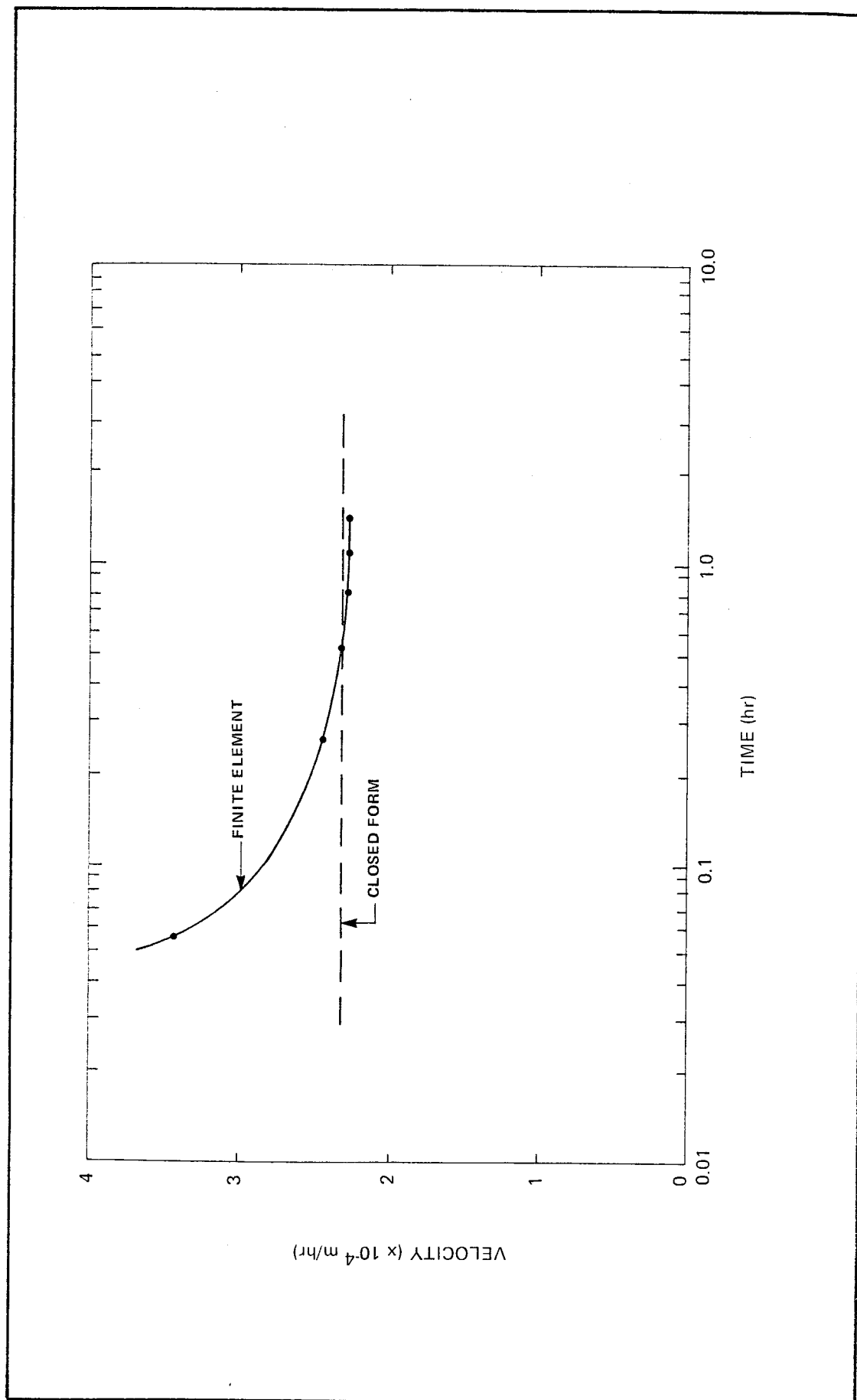


FIGURE E-2 STEADY - STATE CREEP VELOCITY