

University of Alberta

Strength Mobilization in Stiff clays and Weak Rocks

by

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ABSTRACT

This thesis focuses on establishing a framework to identify the operational strength of quasi-frictional materials and the distribution of strength parameters at low confining pressure with plastic strain. In the first part of the thesis, two pre-peak strength thresholds are identified as promising candidates to quantify the operational strength of geomaterials instead of peak and residual strengths. In the second part of the thesis, a comprehensive experimental investigation on a locked sand Formation is presented to reveal the upper and lower strength thresholds by using local strain measurements. In the third part of the thesis, a simplified approach is suggested to partition shear strength parameters with plastic strain. Finally, numerical simulation results are presented to reveal the merits of strain-dependent strength parameters in capturing progressive failure of slopes in stiff clays.

Based on the current investigation, it was shown that the onset of dilation (onset of volumetric strain reversal) may be considered as the upper strength threshold for interlocked/bonded geomaterials to avoid localization and excessive deformation. On the other hand, to avoid initiation of interlocking/bonding degradation, the onset of lateral strain nonlinearity may be considered as the lower strength threshold. The high uniaxial compressive strength of locked sand is due to its interlocked fabric in the absence of virtual bonding, as revealed from microstructural observations. The failure of locked sand under uniaxial

compressive stress condition is achieved by the creation of dilation bands, forming columns/walls that buckle to form a persistent shear band. The interlocking and frictional resistance parameters are not mobilized simultaneously in locked sand. For stiff clays, the cohesion and friction are not mobilized simultaneously as implied in Mohr-Coulomb criterion. The similarity in the shearing response of various bonded geomaterials (the existence of a pre-peak plateau that marks the initiation of bond degradation) and the transition between frictional and bonded materials can be sufficiently captured by using the strain-dependent strength parameters. The concept of strain-dependent strength parameters is a powerful approach to capture the progressive failure of slopes in stiff clays and to interpret the deviation of the operational strength from the laboratory peak strength at the instant of failure.

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DEDICATION

I am dedicating this thesis to the people who influenced my life, personality, and career through their love.

*To my parents (Said and Nadia), my late grandmother
(Safia), and my lovely wife Marwa*

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CHAPTER 1: INTRODUCTION

1.1 General

In the field of geotechnical engineering, there is a wide spectrum of geomaterials ranging from soils to hard rocks. Geotechnical engineers have the responsibility of assuring the short- and long-term performance of civil engineering systems, e.g., underground excavations, foundations, earth-retaining structures, and slopes constructed on these geomaterials. As most civil engineering activities occur at relatively shallow depths, the geomechanical response of geomaterials at low stress levels is significant. Traditionally, the design of engineering systems has been based on the laboratory peak strength envelope with a suitable factor of safety to limit deformation to the pre-peak regime. However, back analyses of failures in stiff clays and weak and hard rocks have revealed an operational strength between the peak and residual strengths. Hence, a design based on the peak strength may be unconservative even with the application of a conventional factor of safety. On the other hand, accounting for residual strength may lead to a conservative and expensive design. Therefore, the main question is what operational strength, based on laboratory results, can be accounted for at low confining stresses to prevent short- and long-term failures? (see Figure 1.1). Answering this question will require an understanding of the mechanical behaviour at low confining stress to assign such a strength threshold.

This chapter provides a brief description of the problem, followed by an explanation of the thesis' objectives, methodology, and organization.

1.2 Classification of geomaterials and shear response

Figure 1.2 presents a simplified classification of geomaterials as well as their inherent shearing response and failure envelopes. For simplification, weak and hard rocks (i.e., bonded stiff clays, clayshales, locked sands, and sandstones) are considered as competent materials formed through the introduction of bonding or interlocking during diagenesis of soils. "Diagenesis" is defined as the process through which sediments are transformed into rocks. It is the sum of the physical, inorganic, chemical, or biochemical changes in a sedimentary deposit after initial accumulation, including mechanical compaction, addition of new material, removal of materials and transformation by change in the mineral phase, and the replacement of the mineral phase by another one (Friedman and Sanders 1978). Granular and cohesive soils are thought to be particulate in nature, and, hence, their strength is governed by the frictional resistance between particles with no virtual bonding. Cohesion in cohesive soils is understood to be the result of electrostatic, electromagnetic attraction, primary valence, and adhesion in contrast to true bonding in rocks partly due to the precipitation of cementation in pores (Mitchell 1993). The frictional component in all geomaterials is a function of the basic friction angle dictated by the mineral characteristics and dilation to allow for grains' arrangement at low stress levels (Taylor 1948; Thurairajah 1961; Rowe 1962; Bolton 1986). For rocks, strength is

derived from both friction (basic friction and dilation) and is dominated by diagenetic effects (interlocking and/or bonding).

The stress-strain curves for dense sands, overconsolidated (OC) clays, and rocks at low confining pressure show a reduction in load-carrying capacity after attaining the peak strength (the post-peak strain weakening), and the post-peak strain weakening is accompanied by positive volumetric strain (dilation). The extra component of the strength at peak is dominated by the dilation and bonding/interlocking for soils and rocks, respectively. On the other hand, loose sands and normally consolidated (NC) clays show a constant resistance that is dictated by the basic frictional resistance of particles with contractive volumetric strain and no dilation component (see Figure 1.2). For intermediate materials (quasi-frictional materials, i.e., stiff clays and weak rock), the extra component cannot be attributed completely either to bonding/interlocking or dilation as in soils or rocks.

1.3 Source of peak strength in stiff clays

The significance of cohesion and its contribution to peak strength for stiff clays as quasi-frictional materials constituting the boundary between soil and rocks remains controversial. However, no confusion should result when the cohesive component in clays is due to electrostatic, electromagnetic attraction, primary valance, and adhesion, in contrast to true bonding due to the precipitation of cementation in pores (Mitchell 1993). This issue has been clarified by comparing

the mechanical response of natural stiff or bonded clays with that of reconstituted clays (clays with the same mineralogical composition at water contents close to the liquid limit). Leroueil and Vaughn (1990), Burland et al. (1996), Cotecchia and Chandler (1997), Amorosi and Rampello (1998), and Cotecchia and Chandler (2004), attributed the behavioral deviation between reconstituted clays (particulate frictional material) and natural clays at peak due to the existence of bonds acquired during diagenetic processes.

Schofield (1998, 2006) argued that the extra resistance in stiff/overconsolidated reconstituted clays is an interlocking component (due to the dilatancy rate) and cannot be attributed to the cohesive effect at peak strength. Hence, a design should be based on critical state friction angle (the state at which no change in volumetric strains and shear stresses and can be approximated by the basic friction angle). This argument is in agreement with Rowe et al.'s (1963) analysis of two reconstituted clays, by using Rowe's (1962) correction for dilation. In contrast, Rampello et al. (1994) compared the failure envelopes of natural bonded clay (Vallerccia clay) under the reconstituted and natural state. Accounting for the dilation rate (by using Rowe's stress dilatancy relation) and the void ratio at failure, resulted in zero cohesion for reconstituted clay; however, a cohesive component was evident for the natural clay. Hence, it is recognized that bonding represents a main source of strength in natural stiff clays at peak strength regardless of its value. Figure 1.3 presents a schematic illustration for the corrected envelopes, accounting for dilation for both reconstituted and natural

clays. As this figure reveals, bonding is an intrinsic source of strength for natural clays, in contrast to unbonded clays (reconstituted clays), in which only basic frictional resistance can be accounted for.

1.4 Laboratory stress-strain curves and failure modes of geomaterials

The stress-strain response and the failure modes of tested cylindrical geomaterials under compression are dependent on the investigated range of stresses. Novello and Johnston (1995) showed that soils and weak and hard rocks show identical mechanical behaviour, and that their behaviour is solely a function of the confining pressure.

Figure 1.4 shows a schematic diagram of typical failure modes for geomaterials as a function of the confining stresses and the magnitude of bonding/interlocking. The increase in confining pressure or reduction of bonding between particles results in a transition from rock to soil. The first mode (A) corresponds to pure localized extension in the direction of the major principal stress, a failure mode common in highly cemented clays and rocks. Mode (C) corresponds to the localized shear mode of failure encountered at medium confining pressures. Mode (B) is an intermediate stage where geomaterials fail under mixed modes of localized extension and shear failure when tested at low to medium confining pressures. Mode (D) corresponds to distributed multiple shear cataclasis resulting in bulging of the tested geomaterials and occurs at high confining pressures, at which diagenetic effects are alleviated due to breakage of bonding

or interlocking. For soils, Mode (C) is recognized for both dense sands and OC clays. However, Mode (D) is the failure mode associated with loose sands and normally consolidated clays. The reduction of diagenetic effects with the increase of confining pressure results in a transition from purely bonded/interlocked geomaterials (Mode (A)) to purely frictional material (Mode (D)).

Figure 1.5 presents the stress-strain curves associated with such deformation modes. These curves can be divided into three stages that are dependent on the range of confining pressures as well as the diagenetic effects. These stages are purely elastic responses followed by strain hardening and/or weakening. For modes of deformation A through C, a significant reduction in strength occurs due to the brittleness of interlocking or bonding as well as the localization of the deformation into a distinct band. However, the reduction in strength and its rate tend to be less abrupt with an increase in confining pressure, partly due to reduction in interlocking/bonding effect. In case (D), no reduction in strength occurs and geomaterials experience infinite plastic strain and ductile behaviour with an indistinct localized mode of deformation. The general effect of the confining pressure is to increase the plastic (permanent) strain experienced by geomaterial within the failure surface. For case A, the loss of strength is sudden, and the strength along the failure surface is due to bonding/interlocking with almost negligible mobilized frictional resistance. For cases B and C, the strength on the failure surface is governed by both the friction and bonding/interlocking resistance. However, the contribution of the bonding/interlocking resistance is

more significant in case B than in case C. In case D, the resistance is governed by friction as most of the contribution due to bonding/interlocking is lost due to the application of high confining pressures.

The failure modes as well as the source of strength are understood in case A (strength is governed by bonding/interlocking), in case C (strength is governed by basic frictional resistance and dilation), and in case D (strength is governed by basic frictional resistance). However, the source of strength in case B and the mutual contribution of friction (its basic and dilation contributions) and bonding/interlocking resistance are not fully understood as the failure mode is the combined contribution of extension and shear, which is encountered mainly in quasi-frictional materials (e.g., stiff clays and weak rocks).

1.5 Mobilized field strength in field problems

The geotechnical literature provides extensive case histories of excessive movement and failures in stiff clays and rocks, and attributes these failures to the post-peak strain weakening and progressive failure due to bond degradation (e.g., Skempton, 1964, 1970; Conlon 1966; Bjerrum 1967; Pelli et al. 1991; Martin and Chandler 1994; Cooper et al. 1998; Olalla and Cueallar 2001). The back-analysis of failures usually indicates an operational strength somewhere between the peak and residual strengths. In this case, a design based on the peak strength can be unconservative. On the other hand, a design based on the residual strength can be utterly conservative and cost-ineffective.

Figure 1.6 shows the relationship between the mobilized strength normalized to the peak strength and plasticity index, based on the back analysis of several case histories involving stiff clays and weak rocks for clarification purpose (Mesri and Abdel-Ghaffar 1993). The brittleness index of such geomaterials, compiled by the author according to the available data, is imposed on this figure. The brittleness of stiff clays and weak rocks was calculated by using the brittleness index (IB) introduced by Bishop (1967) to describe the reduction of the load-carrying capacity from the peak (τ_p) to the residual strength (τ_r) as follows:

$$[1.1] \quad IB = \frac{\tau_p - \tau_r}{\tau_p}.$$

Figure 1.6 reveals that the increase in the brittleness index (the increase in the post-peak strength reduction) results in a reduction in the mobilized strength ratio from 100% to 50%. Up to a brittleness index of 35%, the mobilized strength is equivalent to the peak strength. For brittleness indices greater than 35 %, a significant reduction in the mobilized strength ratio occurs. For most of the collected points, the mobilized strength is almost 50-75% of the peak resistance. For hard rocks, Martin (1997) found that the mobilized strength in underground excavations could be as low as 30% of the laboratory peak strength. This level of strength corresponded to the level at which bond degradation was initiated in the pre-peak regime. Martin and Chandler (1994) argued that the strain hardening to attain peak resistance is a temporary effect that cannot be accounted for or even mobilized in the field (see case B in Figure 1.4). Conlon (1966) argued that at the

instant of failure in slope stability problems in stiff natural clays, the strength is governed by the bonding resistance and that the full frictional resistance cannot be accounted for. This argument is in agreement with that of Martin and Chandler (1994) for hard rocks.

Leroueil and Vaughn (1990) revealed the similarity in strength and volume change characteristics for natural soils and weak rocks due to bonding despite the difference in the bonds' formation condition. These researchers stated that the bond degradation is a progressive process resulting in the release of locked-in elastic energy that is best described as a function of the strain or strain energy. In the author's opinion, many bonded geomaterials including natural stiff clays, weak rocks, and hard rocks share the same characteristics, as nature does not differentiate among many types of bonded geomaterials. Bonded geomaterials possess a component of strength (the bonds between particles) that is initially responsible for extra resistance, and in spite of the bond strength, are susceptible to degradation during the shearing process.

1.6 Simultaneous mobilization of strength parameters

In practice, engineers traditionally use the peak strength parameters, for intact geomaterials with no weakness surfaces, with a suitable factor of safety based on experience in an attempt to ensure stability and to avoid excessive deformation by limiting the strength to the pre-peak regime. This goal is customarily achieved for geomaterials by using the empirical Mohr-Coulomb

hypothesis with the cohesion intercept and friction components of the strength. This approach adopts the assumption that friction and cohesion mobilize simultaneously throughout the deformation process, and that they are constant and independent of strain. However, many researchers have cast doubt on the validity of this approach, and have postulated the conceptual mechanism of the non-simultaneous mobilization of the frictional and bonding strength for natural clays, weak rocks, and hard rocks (e.g., Schmertmann and Osterberg 1960; Lambe 1960; Conlon 1966; Mencl 1965; Trollope 1967; Lajtai 1974; Martin and Chandler 1994).

If the assumption of simultaneous mobilization of the strength parameters is valid, then the parameters used despite applying the factor of safety, will be in error, for they will not correspond to the mobilized strength parameters at the working strain level, particularly where the bonding resistance degrades well before the peak strength (De Mello 1988), see Figure 1.7. These observations lead to the following questions:

1. What is the distribution of the strength parameters with the strain?;
2. What is the contribution of the frictional and bonding/interlocking resistance to the peak resistance at different confining pressures?, and, more importantly;
3. What defines the onset of failure in terms of stress and strain?

To prevent excessive movements and to ensure the short- and long-term stability of geotechnical structures, the shearing behavior of geomaterials in the pre-peak regime must be assessed to establish the threshold at which failure commences in favor of the peak and residual strengths, which are the signatures and consequences of failure. Observations of stress-strain curves showed that geomaterials exhibit significant change in their behaviour in the pre-peak regime, as will be shown in Chapter 2. Once this threshold is determined, it can be assigned as a safety boundary beyond which failure is imminent. Consequently, more attention should be paid to the pre-peak regime, in which failure is initiated, instead of focusing on the peak and post-peak regimes.

1.6.1 Laboratory and field stress-strain perspective for quasi-frictional materials

Figure 1.8 shows a typical stress strain curve for a quasi-frictional material at low confining pressure that is similar to case B in Figure 1.5. The peak resistance can be attributed to both frictional resistance and bonding/interlocking resistance, a combination of case A and C. The diagenetic resistance (interlocked or bonding resistance) tends to degrade in the pre-peak regime (point 1) followed by the mobilization of the basic frictional resistance and dilation at the peak (point 2). Hence, the peak strength is the by-product of dilation, basic friction, and the residual diagenetic effect. However, the residual strength is governed by the basic frictional resistance and residual bonding/interlocking, if it persists.

Figure 1.9 shows a schematic diagram for two materials, (E) and (I), derived from the same particulate source (O). The two configurations, E and I, differ in the way the particles are constrained even though macroscopically they are assumed to have equivalent strength. For configuration E, the particles are constrained externally (confining stresses, σ_3) and for configuration I, the particles are internally constrained (interlocking/bonding, σ_i). Although both materials are equally constrained, each displays different behaviour upon shearing. Configuration E shows higher strength when tested under higher confining pressure ($\sigma_3^1 > \sigma_3^2$), due to mainly its frictional nature (Figure 1.10). However, configuration I, although sheared at a confining pressure of σ_3^1 , can display a reduction in strength that is corresponding to a reduction in the applied confining stress (from σ_3^1 to σ_3^2) for configuration E. This result can be attributed to the reduction in local confinement between particles due to the bonding/interlocking degradation. As a result, the frictional resistance will not be fully mobilized, unlike dense frictional material with no diagenetic effect (E). This difference implies that the frictional resistance is not fully mobilized and that the strength is governed by the residual bonding/interlocking component and the partially mobilized frictional resistance (basic frictional resistance). Several investigators have found that the dilation tend to be suppressed prior to the bond degradation by comparing the mechanical behaviour of natural and reconstituted geomaterials (e.g., Coop and Atkinson 1993; Cotecchia and Chandler 1997; Cuccovillo and Coops 1999; Lo and Wardani, 2002,2003; Cresswell and Barton, 2003). These researchers

suggested that the dissipated energy prior to bond degradation should be accounted for along with the energy dissipated in friction. Fedá (2004) came to similar mechanistic conclusions for strength mobilization in bonded geomaterials.

In order to allow for the full mobilization of frictional resistance (basic friction and dilation), the loss of internal confinement should be compensated for by an extra confinement or even by allowing materials to attain higher strain levels after failure initiation, as shown in Figure 1.10. In a laboratory, this reduction in confinement can be compensated for by automatically adjusting the confinement, but it cannot be adjusted in the field, where the confining stresses cannot be controlled. In addition, most failures are sudden, but they will not be so if geomaterials actually deform the same way as they do in laboratory to reach full frictional resistance (Crawford 1963). Hence, a deviation should be expected between the laboratory results and the field mobilized strength.

1.7 Laboratory local and global strain measurements

In the previous sections, most of the discussion was related to stress and variation in geomaterials' strength. In this section, the discussion focuses on the measurement of strain in the laboratory and the significance of local strain measurements. Roscoe (1970) pointed out that "stress is a philosophical concept ... deformation is a physical reality". In engineering practice, problems are

customarily expressed in terms of stresses because geomaterials experience little strain before failure and because stresses can be indirectly calculated or measured (Watterson 1999). However, if the aim is to integrate stresses and strains into a unified framework, accurate measurement of strain is indispensable. For soils, Roscoe (1970) stressed the value of understanding failure's precursory strains, because eventual failure surfaces and their orientation are determined well prior to the peak strength. Desrues and Chambon (2002) also highlighted the importance of appropriately simulating the pre-peak response of soils up to the onset of localization instead of matching the peak strength, which is merely a sign of established failure.

Currently, it is recognized that local axial and lateral strain measurements in soils, weak and hard rocks are mandatory to reveal behavioral response in both the pre-peak and post-peak regime. (e.g., Shibuya et al. 1994; Martin and Chandler 1994; Scholey et al. 1995; Sheng et al. 1997; Fujii et al. 1998; Hayano et al. 1999; Bésuelle et al. 2000; Yimsiri et al. 2005; Sulem and Ouffroukh 2006). It is generally agreed that external axial strain measurements overestimate axial strains at any strain level due to bedding effects (irregularities at specimen extremes). As well, local strain measurements can reveal additional information regarding the onset of localization (the loss of strain uniformity) and the onset of bond degradation. Martin and Chandler (1994) investigated the behaviour of Lac Du Bonnet granite (strong bonded geomaterial) by using a triaxial apparatus while monitoring the local axial and lateral strains. These researchers revealed

additional strength thresholds in pre-peak regime through variation in local axial and lateral strains. These strength thresholds were true material properties independent of the scale and loading rate. Furthermore, Martin and Chandler (1994) argued that the peak strength was related to the testing boundary conditions, i.e., the length to diameter ratio, sample size and loading rate.

1.7.1 Testing conditions and strain measurements

Testing of cylindrical samples for both soils and rocks under compression is one of the common tests in geotechnical laboratories. It is implicitly assumed that the stress and strains are homogeneous within the tested specimens, or that specimens deform as a right cylinder. Hence, arbitrary measurement of the deformation within specimens is acceptable. In geotechnical laboratories, axial strains are traditionally measured externally between the top and bottom platens. Lateral strains are customarily inferred from volume change measurements assuming uniformity of strains. However, this assumption is invalid when the end effects interfere with the experimental results, leading to data misinterpretation. The existence of rough platens results in the development of local confinement due to the mobilized shear stresses at the sample ends. As a consequence, samples become more restrained laterally at the ends compared to elsewhere. Hence, global measurements of the strain can be in error and reflect a system response rather than a behavioral response. An important factor also affecting the measured axial strains is the bedding errors due to irregular surfaces at the specimen/platen interface

(Figure 1.11). The bedding errors result in an increase in the measured global axial strains due to the breakage of asperities at the specimen/platen interface for soils and rocks (e.g., Hawkes and Mellor 1970; Brady and Brown 1985; Scholey et al. 1995; Hayano et al. 1999; Korinets and Alehossein 2002).

Figure 1.11 shows a schematic diagram of a typical 2:1 cylindrical soil specimen in a triaxial test with conventional upper and lower platens. The same effects recognized for a soil specimen are recognized for a rock specimen due to end effects and bedding errors. Along the specimen axis, the distributions of axial, lateral, and volumetric strains accounting for bedding errors and end restraints are shown. Accordingly, the specimen can be divided into three zones along its axis (Rowe and Barden 1964). The top and bottom zones are called “dead zones”. The middle zone (the “active zone”) represents the most homogeneous zone, which is not influenced by platen effects where deformation is not impeded. Within the dead zones, the deformation is highly restrained due to the end effects imposed by platens, causing shear stresses and frictional resistance at sample’s extremes. The distribution of shear stresses on the platens was studied by Lee (1978). Lee (1978) showed that for soils, the shear stresses tend to diminish within the middle of the specimen. Such shear stresses result in non-uniform deformation within the specimen, and, thereby, inherent strain non-homogeneity. An additional result of end effects is the development of a predominant shear band in soils at failure (Rowe and Barden 1964; Bishop

and Green 1965; Kirkpatrick and Belshaw 1968; Lade 1982). In the absence of shear stresses (perfect end conditions), the specimen should deform homogeneously, as in the middle zone, and multiple shear bands due to the uniformity of the strain should be rendered (Raju et al. 1972; Lade 1982).

Frost and Yang (2003) investigated the effect of end restraints and microstructure evolution in sand specimens, tested under the triaxial compression condition, at a confining pressure of 50 kPa. The specimen with lubricated ends showed the most uniform axial strains, lateral strains, and void ratios during the course of shearing along the specimen's vertical axis. The inclusion of the lubricated ends resulted in a slight reduction in the peak strength and the initial secant modulus, in contrast to the results with non-lubricated ends with the same slenderness ratio due to a reduction in the shear stresses at sample/platen interface. A direct effect of end restraints is the reduction in the expansive lateral strains and void ratio (dilation) at sample extremes, in contrast to the effects at the middle third (see Figure 1.10). Frost and Lang (2000) provided a quantitative measure of the local void ratio distribution for fine uniform quartz sand during triaxial compression test by using optical image analysis of images for resin impregnated specimens. The test results showed that the local void ratio was the largest near the center of the tested specimens (see Figure 1.10). The local void ratio first increased at the center of specimens during shearing and migrated towards the ends as the axial strain increased due to the end effects for lubricated

specimen regardless of the slenderness ratio. Frost and Lang (2000) stressed on the importance of the localized measurements in preference to global measurements to ensure true material response.

As the end effects are due to the mismatch between the stiffness of the platen and samples, it is more significant in soils than in rocks. However, the bedding errors due to the irregularity of samples ends are generally unavoidable in both soils and rocks. Although these effects are completely understood, local strain measurements are still for research purposes, especially for soils.

1.8 Locked sands

In the intermediate zone within the spectrum of geomaterials, a class of geomaterials displayed both soil and rock characteristics because of mild diagenesis. These geomaterials can be classified into bond-dominant geomaterials (bonding is responsible for the material's coherent mass) and fabric-dominant geomaterials (interlocking is the binding agent between particles). Bond-dominant geomaterials have been frequently studied in the literature (e.g., Vaughn and Leroueil (1990) have presented a comprehensive data compilation and analysis for stiff clays and weak rocks), where as the exploration of fabric-dominant geomaterials' behavior has been comparatively neglected in geotechnical studies (Barton 1993). An example of fabric-dominant geomaterials is locked sands where mild diagenesis is reflected in the material's

unusual particles' contact characteristics resulting in an interlocked fabric with no virtual bonding (Dusseault and Morgenstern 1979).

Locked sands are classified as *arenaceous* (predominately quartz sand particles) weak sedimentary rocks that constitute the boundary between soils and rocks. In dense sands, the strength is primarily a function of density, which yields a weak interlocking between particles with relative densities less than 100%. In sandstone, the strength is primarily a function of the amount of interstitial cement between sand particles. Despite the lack of cementation in locked sands pores, locked sands possess strength properties much higher than those of dense sands and much lower than those of sandstones. In addition, in contrast to dense sands, locked sands exhibit relative densities greater than 100% (Dusseault and Morgenstern 1979). The deviation from dense sand strength properties stems from mild diagenetic processes that reduce sand mass porosity with negligible or no cementation (Barton, 1993).

In north-eastern Alberta and particularly within 115 km of Fort McMurray, Dusseault (1977) recorded the existence of oil-sand and oil-free sand natural steep slopes. Dusseault (1977) found no interrelationship between slopes' heights and inclination, and the oil content. Dusseault and Morgenstern (1978) showed experimentally that the strength of oil sands was merely a reflection of their interlocked fabric, and, hence, was independent of the oil content. Protulipac (1991) recommended oil-free locked sands as a reasonable substitute

for oil sands to avoid sample disturbance due to gas expulsion where strength and deformation characteristics are a concern. Entire experimental investigations of locked sands from the McMurray area and other localities have typically focused on ultimate strength and volume change characteristics by using external strain measurements (Dusseault and Morgenstern 1979; Barton et al. 1986; Richards and Barton 1999; Wong 1999; Dittes and Labuz 2002; Creswell and Barton 2003).

Recently, Creswell and Powrie (2004) investigated the shearing response of locked sand using local strain measurements. However, no attention was paid to the significance of the local lateral strains' variation during shearing. Creswell and Powrie (2004) argued that due to the large difference between the peak and critical state strengths of locked sands, a geotechnical design based solely on the critical state friction angle would be unduly conservative even with a factor of safety of 1. Alternatively, these researchers recommended a design based on a limiting stress envelope corresponding to the onset of dilation (the point at which the reversal of the volumetric strain occurs) with a no-tension cut-off and a factor of safety on the peak stress ratios to avoid progressive failure. However, Creswell and Powrie (2004) also recommended further investigation into this design approach.

1.9 Objective of the thesis

Despite the wealth of information on the strength of geomaterials, no generalized framework exists for identifying the operational strength, particularly in the intermediate zone between soils and rocks (quasi-frictional materials). The main challenge in determining the operational strength for such materials lies partly in capturing the relative contribution of the interlocking/bonding and frictional resistance at working stress and strain levels, or in the pre-peak regime.

The main objectives of this thesis are summarized as follows:

1. To examine high-quality laboratory results for various geomaterials to understand failure precursors in the pre-peak regime, and to capture the onset of failure strength thresholds for soils and rocks.
2. To conduct a comprehensive experimental study on a McMurray fabric-dominant weak rock (locked sand) to reveal the following:
 - a. The source of the exceptional uniaxial strength of such a sand formation in the absence of bonding between particles and to confirm this formation's fitting into locked sands category;
 - b. The significance of local strain, compared to our routine external strain measurements for McMurray sand formations, in terms of failure mechanism, deformation modulus, and post-peak response;
 - c. The pre-peak damage thresholds for locked sand to capture the damage evolution based on the variation of local axial and lateral strains;

- d. The contribution of the frictional and the interlocking strength components with the plastic shear strain;
 - e. The relationship between the onset of localization and the onset of dilation as a candidate for the strength threshold;
 - f. The mechanism and kinematics of failure of locked sand;
 - g. The relationship between the onset of dilation and the long-term strength (onset of creep) of locked sand as a possible candidate for long-term strength; and
 - h. The upper and lower strength thresholds for locked sand.
3. To develop a simplified approach for partitioning the shear strength parameters (cohesion and friction angle) with plastic strain for stiff clays, and to show the significance of non-simultaneous mobilization of strength parameters in revealing:
- a. The resemblance in bonded geomaterials' shearing responses at low confining pressure irrespective of these geomaterials' origin; and
 - b. The transition between soils (frictional materials) and rocks (bonded materials) based on the mutual contribution of frictional and bonding resistance to overall resistance.
4. To highlight the significance of the concept of non-simultaneous mobilization of strength parameters in capturing the progressive failure of slopes and to interpret the deviation of the operational strength of stiff clays from the peak strength with such conceptual framework.

1.10 Methodology

The objective was achieved through the compilation and reanalysis of high-quality experimental results using local strain measurements and modern techniques to reveal the relationship between the onset of the shear-induced dilation and the onset of localization. The compiled data included examples for dense sands, overconsolidated clays, sandstones, locked sands, and granite. To extend the value of the shear-induced dilation threshold, examples were given for different geomaterials to show the relationship between the onset of shear-induced dilation and long-term strength. For bonded geomaterials (where strength is governed by the bonding resistance), another threshold (extension-induced dilation) was shown to be the lower strength threshold where bonding resistance is in effect.

The second objective was accomplished by testing specimens cut from block samples of oil-free McMurray sand formation from the High-Hill River banks. Preliminary tests served to reveal the sand's fitting into locked sand category. These tests consisted of microstructure observations of SEM images, blue-dyed thin section, and mineralogical compositions using the energy dispersion spectrum. Uniaxial compressive strength tests were conducted to reveal this formation's intrinsic strength by using local axial and lateral strain measurements. These tests highlighted the significance of local strain measurements of McMurray sand formations. The experimental results revealed the pre-peak strength thresholds, the failure kinematics based on the variation of the local

axial and lateral strains, the relationship between the onset of dilation and localization, and the contribution of frictional and interlocking strength parameters at different strain levels. To further investigate the long-term strength of such a formation and the relationship between the long-term strength and the onset of dilation, multistage creep test was conducted. Finally, direct shear box (DSB) tests were carried out to show the relationship between the maximum mobilized strength in DSB and the pre-peak strength thresholds estimated from the uniaxial compression test.

In order to meet the third objective, stress strain curves, based on triaxial compressive strength results, for two natural bonded stiff clays were investigated as an example to show the mutual contribution of cohesion/bonding and frictional resistance at different plastic strain levels. The analysis highlighted the effect of the mobilization and degradation of strength parameters on stress-strain curves and post-peak brittleness at different confining pressures. Parametric study within such a framework revealed the concept of non-simultaneous mobilization of strength's significance by providing evidence from experimental results for other bonded geomaterials with different origins.

The fourth objective was achieved by carrying out a numerical simulation using the peak strength parameters and strain-dependent parameters using the finite difference code FLAC. The analyses also included a comparison between the position of slip surface obtained by using the strain-dependent strength

parameters and those associated with peak strength parameters, estimated from limit equilibrium analysis. The analyses highlighted the significance of the concept of non-simultaneous mobilization of strength parameters in capturing the progressive failure of slopes in stiff clays and in interpreting the deviation of the operational strength from the peak strength.

1.11 Organization of the thesis

Chapter 1 explains the problem's background and the thesis' objectives, methodology, and organization.

In Chapter 2, two pre-peak strength thresholds (the onset of shear-induced dilation and the extension-induced dilation thresholds) are discussed. The relationship between the onset of shear-induced dilation and localization is established by using compiled high-quality test results for dense sands, sandstones, locked sand, stiff clay, granite, and rock salt. The data are based on local strain measurements, acoustic emissions, and shear velocity tests. The significance of the onset of shear-induced dilation for all classes of geomaterials will be revealed by showing the relationship between the onset of dilation, the onset of localization, and the long-term strength. As well, the extension-induced dilation will be shown to be an important lower strength threshold where bonding is a major source of strength.

Chapter 3 presents the details of the experimental set-up and preliminary tests conducted on samples from the High-Hill River sand formation, procured from north-eastern Alberta, as well as a summary of locked sands' characteristics. The tests include index tests, microstructure observations using SEM, analysis of thin sections under a petrographic microscope, and qualitative mineralogical composition using the energy dispersive X-ray spectrum to reveal the fit of this formation into the locked sand category. The remainder of the chapter provides details on the testing procedures, including the calibration of miniature LVDTs for local and axial strain measurements, proof testing, stability tests, and temperature effects. The details of the modification for an existing mechanical configuration are also presented. Finally, the significance of local strain measurements is presented by contrasting local and external axial strain results.

Chapter 4 presents the results of the uniaxial compressive strength results from testing locked sand using local strain measurements. The first series of analyses reveals the relationship between the onset of dilation and localization, the damage thresholds and failure evolution and kinematics, the contribution of interlocking resistance and frictional resistance at different strain levels, and the mobilized strength parameters at different strain levels. The second series of analyses shows an attempt to rule out the heterogeneity associated with boundary conditions (end effects, strain rate, and specimen slenderness ratio) allowing for propagation of pre-peak columnar structures in the axial direction. The third series of analyses shows the relationship between the onset of dilation

and the long-term strength by conducting a multistage uniaxial creep test. Finally, the maximum mobilized strength and its relationship to the damage initiation threshold will be shown by using direct shear box test results.

Chapter 5 describes the methodology used to partition the cohesion/bonding and frictional resistance with the plastic strain level in Mohr-Coulomb and q-p spaces to show the mutual contribution of frictional and bonding resistances at different strain levels. Their contribution was illustrated by using two triaxial compressive strength results for bonded stiff clays. Once the parameters have been partitioned, a parametric study was conducted to reveal their contribution at different strain levels and confining pressures. The non-simultaneous mobilization of the strength parameters was used to capture the peak strength envelope's nonlinearity and the transition between soils and rocks based on their mutual contribution at different confining pressures. The parametric study also revealed the source of the initial "plateau" in the pre-peak regime at low confining pressures. This plateau is a characteristic behavior of various bonded geomaterials, as will be shown by using examples from the literature for weak rocks, sensitive clays, and artificially bonded soils.

Chapter 6 describes a numerical simulation to reveal the significance of strain-dependent parameters in the assessment of slope stability. To achieve that purpose, a homogeneous slope, with a height of 20 m and a side slope of 45°, was used for illustration. The peak strength parameters and the strain dependent

strength parameters estimated for a stiff clay (Nanticoke clay) in chapter 5 was used in the analyses. The analyses comprised a comparison between the slip surfaces and factors of safety, associated with peak strength parameters, obtained from the limit equilibrium analysis and the finite difference analysis by using the strength reduction technique. The rest of the chapter was devoted to show numerically the merits of the non-simultaneous mobilization of strength parameters in capturing the progressive failure and to interpret the deviation of the operational strength from the peak strength.

Chapter 7 summarizes the conclusions reached in the research and their significance, and provides recommendations for further investigations.

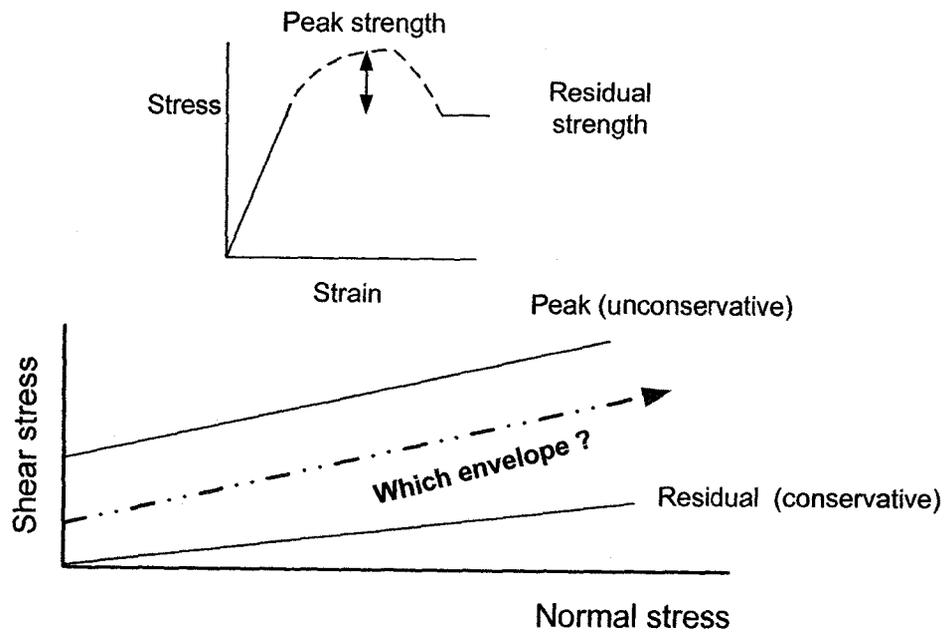


Figure 1.1 Peak and residual strength envelopes for geomaterials.

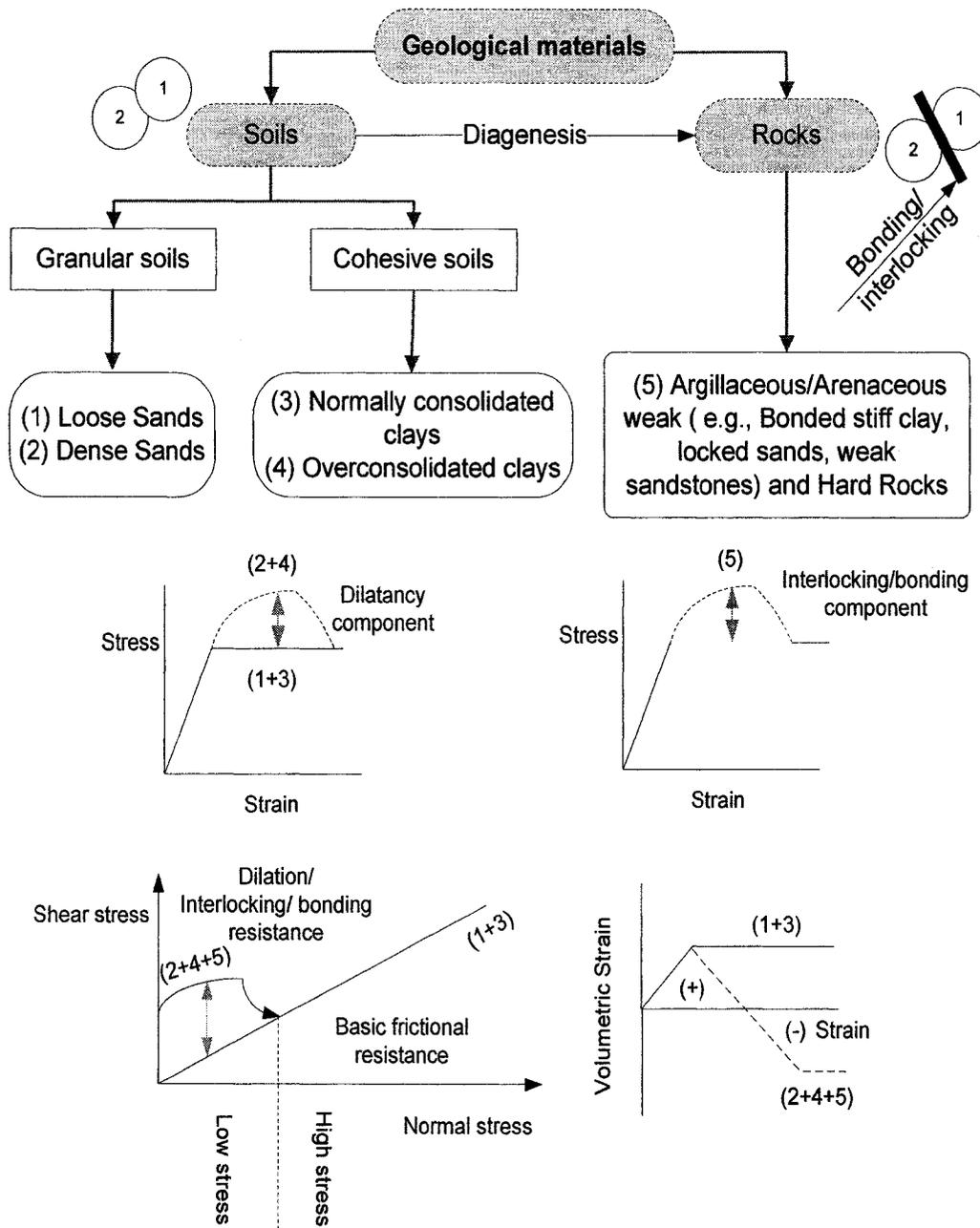


Figure 1.2 Classification of geological materials based on frictional and bonding/interlocking contribution to strength.

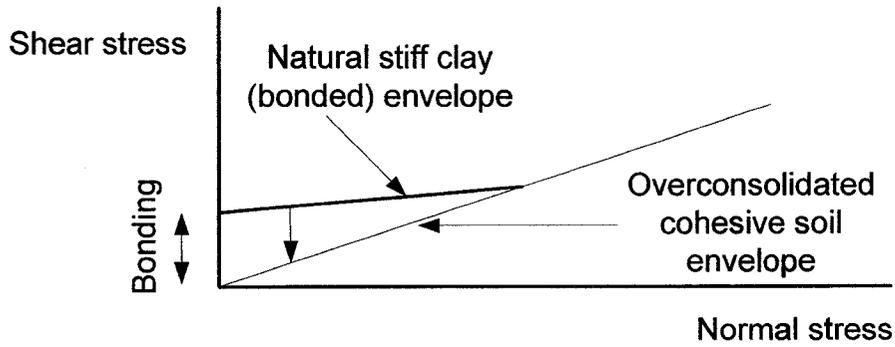


Figure 1.3 Peak strength envelopes for cohesive and bonded clays.

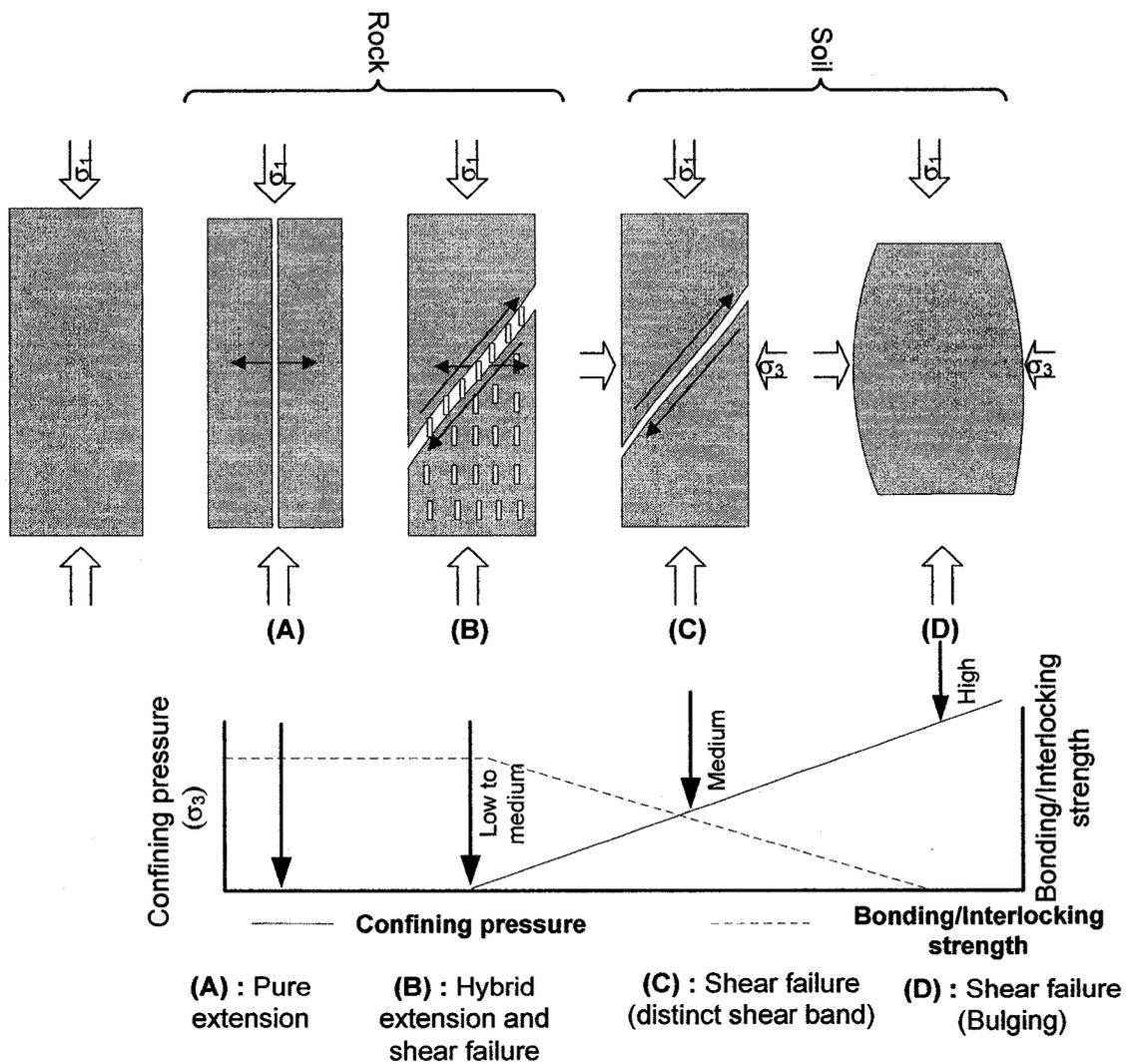


Figure 1.4 Failure modes of cylindrical specimens based on confining pressures and bonding/interlocking between particles.

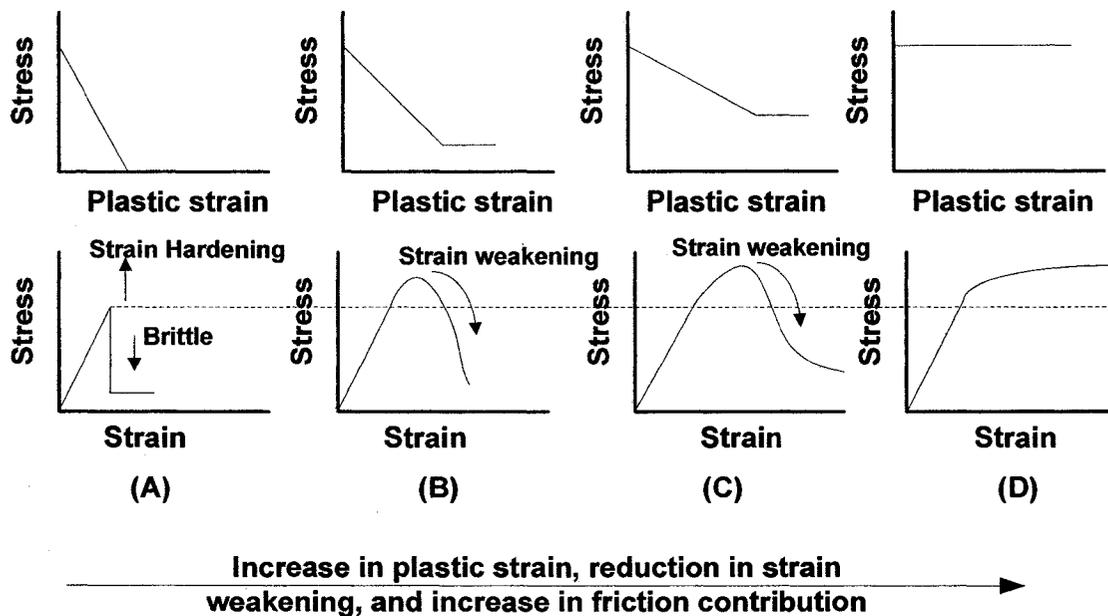


Figure 1.5 Typical stress strain curves at different confining pressure and plastic strain along failure surfaces.

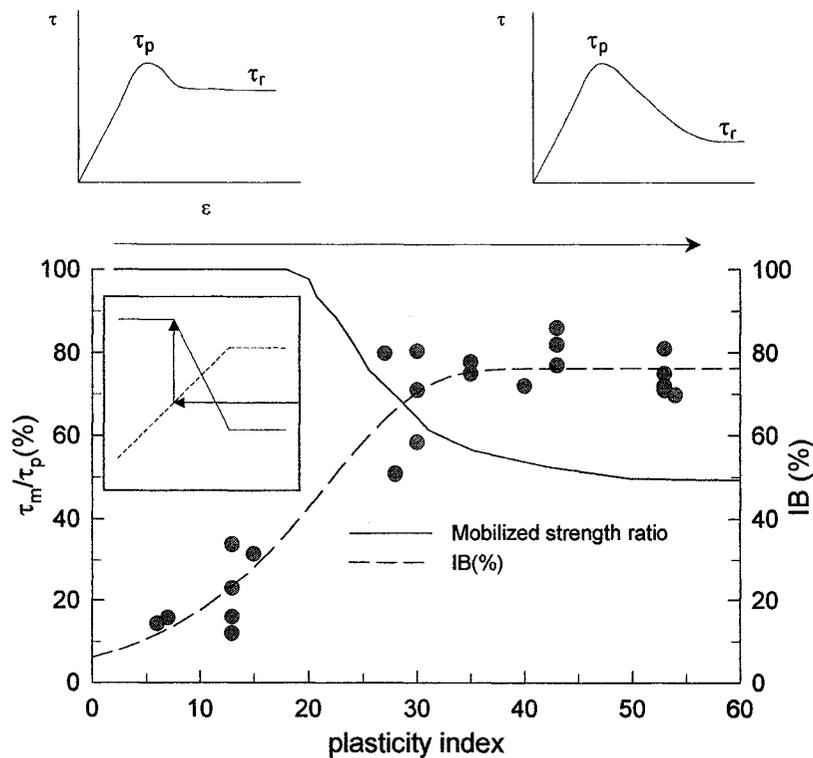


Figure 1.6 Mobilized strength ratio (τ_m/τ_p) and brittleness indices (IB) for stiff clays and weak rocks.

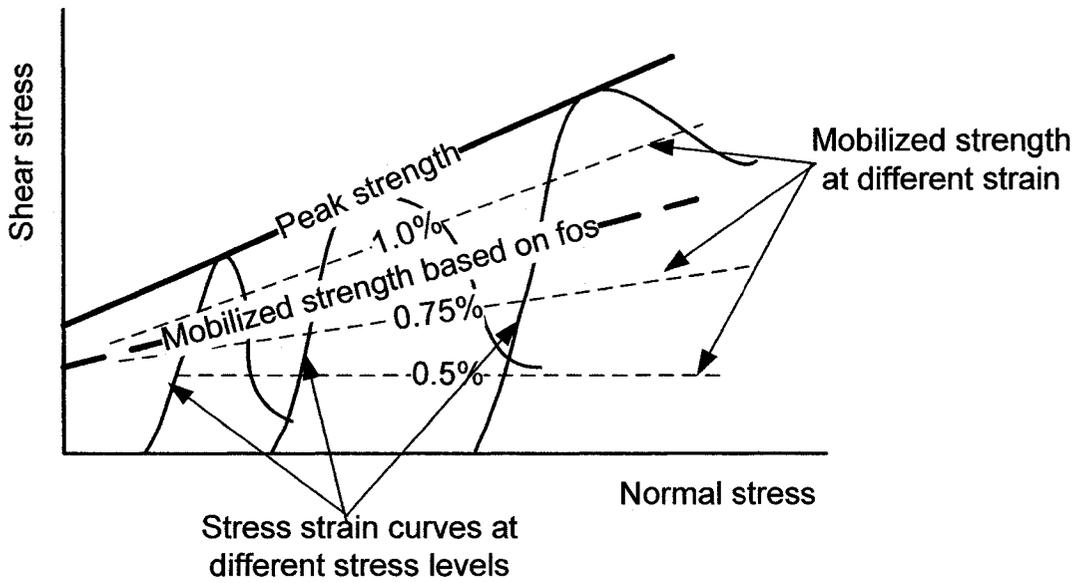


Figure 1.7 Variation of mobilized strength at different strain levels and mobilized strength based on conventional factor of safety (modified from De Mello 1988).

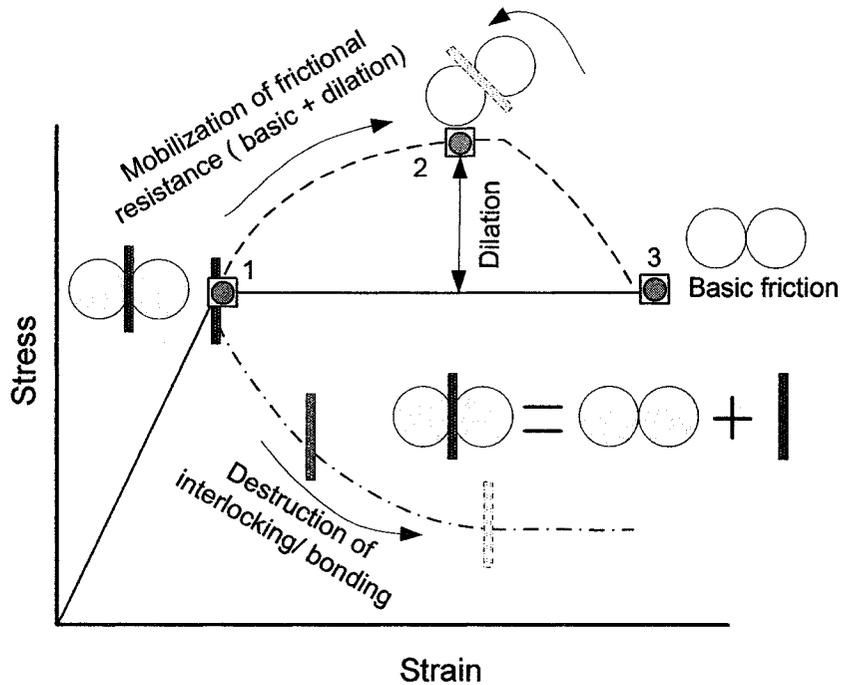


Figure 1.8 Stress strain perspective for strength mobilization in bonded/interlocked geomaterials.

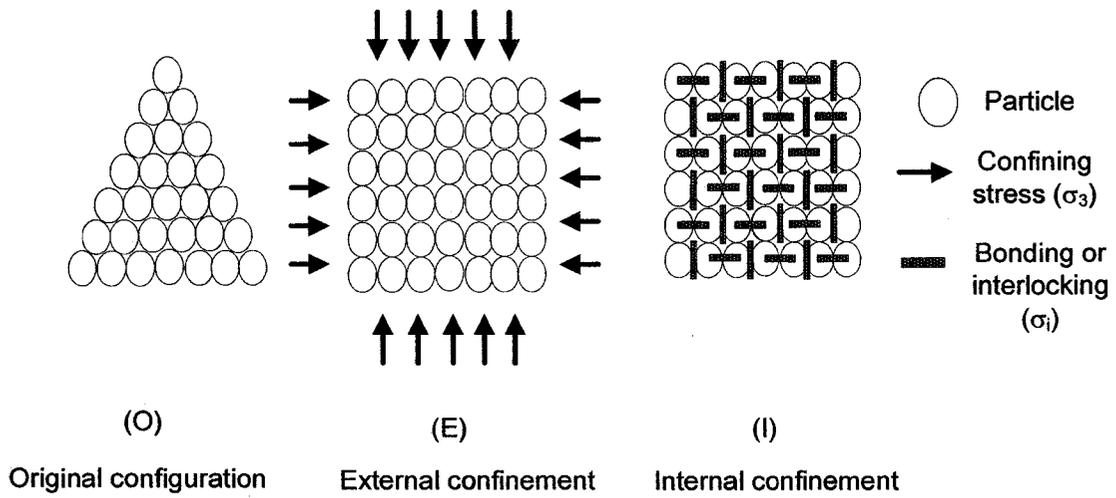


Figure 1.9 A schematic diagram of a frictional and bonded/interlocked geomaterial with equivalent constraint level.

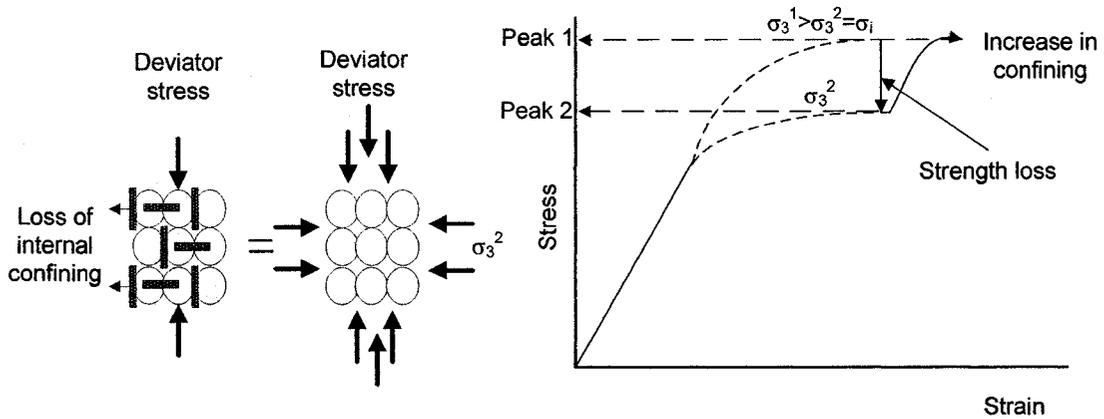


Figure 1.10 Loss of internal confining for bonded/interlocked geomaterial due to internal constraint destruction and its effect of stress strain curve compared to a frictional geomaterial.

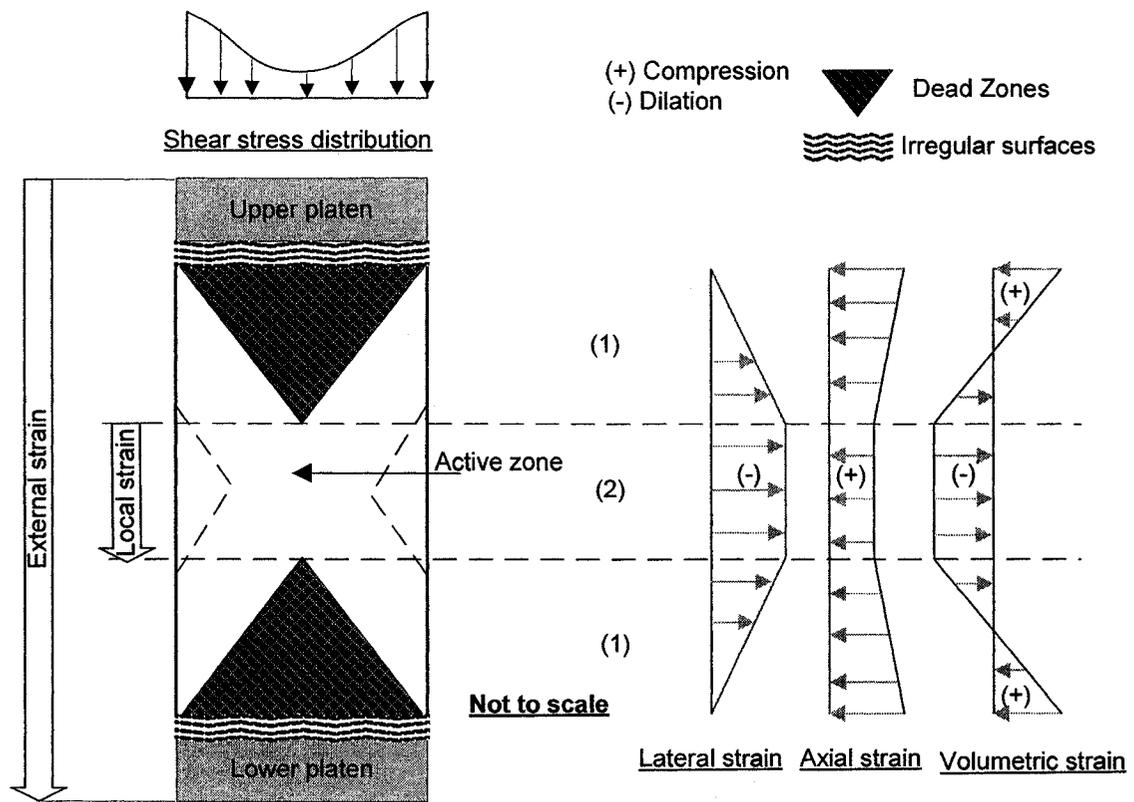


Figure 1.11 A schematic diagram to illustrate the interference between testing conditions and measured global strains.

CHAPTER 2: ONSET OF DILATION AND LOCALIZATION THRESHOLDS IN SOILS AND ROCKS

2.1 Introduction

The design of geotechnical structures necessitates a comprehensive understanding of the stress-strain curves of geomaterials to ensure the safety, stability and reliability of our structures under working stresses. By far, uniaxial and triaxial compression stress paths are the most common testing conditions used in soil and rock mechanics laboratories to determine strength and deformation characteristics. Under low confinement, or in the conditions in which most civil engineering applications are situated, most geomaterials exhibit strain weakening (strength reduction from peak to residual strength) accompanied by localized deformation in a single band or multiple bands, and also by volumetric strain reversal (dilation). As illustrated in the introductory chapter, experience gained through failures showed that a design based on the peak strength might result in an un-conservative design. On the contrary, a design based on residual strength can be over-conservative and cost-ineffective. In order to solve this dilemma, studies of microstructure changes during shearing are necessary to visualize the onset of damage and initial yielding. In pre-peak regime, geomaterials undergo changes in their microstructure, which reflect other strength thresholds rather than the peak and residual strengths as will be shown in this chapter. It should be noted that all the terms used in this chapter are in terms of effective stresses.

Although many studies of various geomaterials have been published, it remains unclear at what stress level localization is initiated and eventually results in strain weakening in the post-peak regime, and how it is related to the onset of shear-induced dilation (the onset of volumetric strain reversal). To clarify such points, a meticulous examination of high-quality experimental data, based on localized measurements, was carried out to draw attention to two pre-peak strength thresholds rather than peak and residual strengths, which are merely markers for failure. As Roscoe (1970) stated, understanding failure's precursory strains and structures is essential, because failure surfaces and their orientation are determined well prior to the peak strength.

The compiled strength and deformation data include examples for soils, weak rocks, and hard rocks to reveal the relationship between the onset of localization and the onset of volumetric strain reversal (the onset of shear-induced dilation). The bulk of this chapter discusses the relationship between the onset of shear-induced dilation and the long-term strength by using the reported creep test results for some geomaterials. In addition to exhibiting shear-induced dilation, brittle hard rocks also exhibit extension-induced dilation during compression testing due to the initiation of bond degradation, which is suggested to be the lower bound strength for all bonded geomaterials.

2.2 Definition of shear-induced dilation and localization

The phenomenon of shear-induced dilation was used by Reynolds (1885) to describe the increase in volume when dense granular materials are sheared. Reynolds (1885) attributed this distinctive feature to geometrical constraints imposed by particle packing (fabric) for dense granular materials. He believed that the unique condition for dilation was that boundary grains held interior grains so as to prevent the commencement of rearrangement. Today, it is recognized that shear-induced dilation (volumetric strain reversal) is not restricted to dense granular geomaterials but also occurs in most geomaterials at low confining pressures. Consequently, shear-induced dilation can be considered as a direct result of overcoming the different constraints (fabric and/or bonding constraints) that act as internal constraints, and reflects a transition in the behaviour of geomaterials.

The formation of a shear band (an inclined deformation band), which is a direct result of the localization of deformation and divergence from homogeneous deformation is associated with shear-induced dilation. Once localization is initiated, irreversible plastic deformation takes place, and particles move with respect to each other along thin surfaces that appear as thin lines in the sample's body. Localization takes place when deformation bifurcates from a homogeneous mode to a localized mode on one or many surfaces that coalesce to form what is called a "dominant shear band". After formation of the shear band, most

deformation occurs through the rigid body motion of one part of the sample relative to the others (Saada et al. 1999).

Volume change (compression and dilation) in compression laboratory tests, and, in particular, triaxial tests is measured by using three techniques (Alva-Hurtado and Selig 1981; Leong et al. 2004). The first technique is by monitoring the change in the volume of fluids within specimen pores, which represents a global volume change. In the second technique, the volume of fluid entering or leaving the confining pressure cell to compensate for the volume change of the soil specimen is measured. In the first two techniques, lateral strains are calculated based on the assumption of right cylinder. The third technique depends on measuring, independently, axial (ε_1) and lateral strains (ε_3) from contact or non-contact elements. The volumetric strain of a cylindrical specimen (ε_v), which is tested under triaxial test conditions, is typically calculated by using the following equation:

$$\mathbf{[2.1] \quad \varepsilon_v = \varepsilon_1 + 2\varepsilon_3.}$$

2.3 Shear-induced dilation and localization thresholds

Advances in laboratory testing techniques enabled several researchers to investigate the deformation mechanism in compression tests for soils, weak rocks, and hard rocks. These researchers used local and global measurements,

acoustic emissions, and computerized tomography to monitor the deformation process. The results from their work are presented, reviewed, and discussed.

2.3.1 Coarse-grained soils and weak rocks

2.3.1.1 Dense sand

Finno and Rechenmacher (2003) investigated the deformation of medium-to-dense concrete sand in a drained plane-strain compression test. Samples were prepared by using the vibratory method to achieve an initial void ratio of 0.543. Figure 2.1a shows the deviator-stress versus the axial-strain curve at a confining pressure of 98 kPa.

The onset of localization was identified via local lateral strain measurements by using 4 horizontally mounted LVDTs as shown in the insert in Figure 2.1c. The divergence of local lateral strains with shearing progression at point A shows the onset of localization (Figure 2.1c), which is coincident with the onset of dilation (volumetric strain reversal) in Figure 2.1b. This result is also supported by Figure 2.1d. At peak strength (point B), the strain localization results in a partially formed shear band, and subsequent shearing results in the formation of a continuous band at large strains (point C) reflected in the constant value of one of the lateral strains reflecting the rigid body motion of un-failed element. The shear band appears to be developed at point C (the critical state), where the volumetric strain and deviator stress remains constant.

Figure 2.1d shows the evolution of the dilation angle with axial strain. At point A, the dilatancy angle is equal to zero where localization commences. At point B, the dilation angle reaches its maximum value followed by a reduction in the post-peak regime. Finno and Rechenmacher (2003) stated that the peak strength is a direct result of the localization and loss of homogeneity in sheared specimens of dense sand. In these researchers' paper, point D marked the onset of localization according to their interpretation, which resulted from the subjectivity in selecting of divergence point. However, even if this interpretation is correct, the onset of dilation can be considered as a very good indicator of the onset of localization.

The results suggests that for dense sands tested under plane strain conditions, the onset of localization can be determined from the global volumetric strains and that the volumetric strain reversal is representative of the onset of shear-induced dilation, which also corresponds to the onset of shear localization. In addition, strength reduction in the post-peak regime is a direct consequence of the localization process that commences in the pre-peak regime.

2.3.1.2 *Locked sand*

Locked sands' behaviour was first investigated by Dusseault and Morgenstern (1978), who concluded that the sands' fabric, and not the bonding due to the cementation, was responsible for the high strength and

dilation angles at low confinement. Cresswell and Powrie (2004) carried out triaxial compression tests on locked sand to investigate its stress-strain characteristics. The sand used in their tests had a tangential index (the ratio of tangential contacts to long straight contacts) of 26%, a relative density index of 136%, and a void ratio of 0.47. Axial deformation was monitored by using 4 axial LVDTs, and the radial deformation was monitored with a radial strain caliper. Figure 2.2a shows the stress-ratio (q/p) versus the axial strain curve at a confining pressure of 100 kPa and Figure 2.2b depicts the volumetric strain as well as the local axial strains measured by using internal LVDTs. According to Cresswell and Powrie (2004), the divergence of the local axial LVDTs in Figure 2.2b (point A) indicates the onset of localization where the strain field becomes discontinuous. This point also coincides with the point of the shear-induced dilation (the onset of the volumetric strain reversal), as shown in Figure 2.2b. The progression of deformation from point A (the onset of localization) to B (the partial formation of shear band) to C results in the complete formation of a shear band, as was indicated previously for dense sands. Cresswell and Powrie (2004) indicated that onset of shear induced dilation represents the stress level at which sand starts to unlock and loses its integrity. They also recommended that this threshold should be the limiting strength envelope for locked sands. The ratio of the onset of dilation to peak strength is plotted in Figure 2.3, which reveals the ratio independency of the confining pressure and that the onset of dilation occurs at almost 85% of the peak strength.

2.3.1.3 Sandstone

Besuelle et al. (2000) carried out triaxial compression tests on Vosges sandstone. This sandstone is a poorly cemented sedimentary soft rock, and most of its characteristics are derived from particle contact details. It is a pink quartz sandstone (quartz = 93%) with a small percentage of feldspar and white mica, and a porosity of 22%. Axial and lateral strains were measured by using three axial and four lateral transducers based on the concept of a flexible strip with attached strain gauges. The localization was detected by comparing the three axial strain measurements. If the difference between the mean axial strain (the average of three measurements) and each of the three axial measurements remains constant, the strain field can be assumed to be homogeneous with no localization. Figure 2.4a shows the stress-strain curve at a confining pressure of 20MPa. The figure reveals a strength reduction in the post-peak regime due to localization. Figure 2.4c shows the relationship between the differences of the axial strain and mean axial strain. Besuelle et al. (2000) suggested that the first divergence (represented at point A) between the three axial strain measurements was due to bedding errors in the experiment. Once seating had occurred, the differences among the three axial strains remained nearly constant up to point C where the difference of the three transducers increased dramatically. Besuelle et al. (2000) selected this point (point C) as the onset of localization based on the axial strains' differences. By using the axial strains' differences as well as lateral strains' differences, the onset of localization can be shown to commence at point B.

The volumetric strain was also determined by using the axial and lateral strain measurements as shown in Figure 2.4b. Note that the onset of shear-induced dilation (the volumetric strain reversal) was again close to and almost coincident with the onset of localization (point B in Figure 2b). Besuelle et al. (2000) also showed that the onset of localization generally occurs at a progressively lower normalized deviator stress as the confining pressure is increased. Normalization of the onset of dilation with respect to the peak also reveals its independency of the confining pressure (see Figure 2.5). Lockner et al. (1992) also revealed the coincidence of the shear-induced dilation and the onset of localization for Berea sandstone with a porosity of 18% and tested in triaxial test at a confining pressure of 50 MPa, by using acoustic emissions.

Recently, Sulem and Ouffroukh (2005) investigated the behaviour of Fontainebleau sandstone under drained and undrained triaxial compressive stress paths at various confining pressures. The samples had a porosity of 21% and were composed of 98% quartz. Axial and lateral deformations were monitored by using two local vertical LVDTs and four diametrically opposite lateral LVDTs, respectively. The onset of shear banding was detected through the difference between the individual axial strains and/or the radial strain difference between the two opposite points. Theoretically, as long as the deformation was homogeneous, the difference between the readings would be negligible unless localization commenced. Figure 2.6 shows the results of

a drained compression test under a confinement pressure of 28 MPa. Figure 2.6b shows the volumetric strain axial strain relationship. The initial stages of the testing are dominated by compressive volumetric strains up to point A. Beyond point A, volumetric reversal commences, signifying the onset of shear-induced dilation. This point corresponds to the increase in nonlinearity (see Figure 2.6a). Figure 2.6c shows the relationship between the axial strains' differences, lateral strains' differences and average axial strains. In terms of the axial strains' differences, the difference between axial strain readings is negligible as long as the strains are homogeneous. Once localization commences, significant differences in the strains occur. For the lateral strain difference, the difference increases slightly and is followed by a reversal in the value. The point at which reversal occurred was taken as the onset of localization and coincided with the recorded localization strain level as manifested by the vertical strains. This point (B) lies very close to the onset of dilation threshold and occurs in the pre-peak regime. However, this result is caused by the subjectivity in selecting the points at which strain difference occurs. In undrained conditions, the transition from contractant to dilatant behaviour occurs much earlier than in drained tests and increases with confining pressures due to the effect of pore-pressure generation. The relationship between the onset of localization and onset of shear-induced dilation normalized to the peak strength at different confining pressures under drained conditions is illustrated in Figure 2.7, which reveals that both thresholds are almost coincident and occur at 82 % of the peak strength.

2.3.2 Fine grained soils and weak rocks

2.3.2.1 Overconsolidated silty clay

Su et al. (2004) studied the deformation mechanism of undisturbed silty clay samples under undrained triaxial compression test conditions by means of computerized tomography. Scanning was recorded at different strain levels and for different slices during shearing. The CT images were quantified by using the CT number, which is a representative number for the sample's density and has a linear relationship with it. The change in density can be observed from the changes in the brightness between two cross-sectional images at the same location. The white areas in the images are considered to represent damage and localization. Figure 2.8a shows the deviator-stress versus the axial-strain relationship at a confining pressure of 300 kPa. The sample has a void ratio of 1.043, plasticity index of 14.31, and water content of 37.72 %. The letters on the curve represent the strain levels at which CT images were taken. Figure 2.8b gives the relationship between the CT number and the axial strain obtained from the middle slice as it represents the most homogeneous location in a specimen (a location away from boundary effects). Based on the variation in the CT numbers, the stress-strain curve is divided into four portions: (a) a portion where micro-defects close up (O-A); (b) a portion where new microcracks open and develop (bond-degradation stress level, A-B); (c) a portion where local deformation and damage starts (point B); and (d) a portion where softening and a reduction in load-carrying capacity occur. According to Su et al.'s (2004) experiments, localization starts

to occur at point B where local defects and damage commence. The X-ray portion shown at this axial strain level reveals separate locations of damage sites (indicated by the light grey points) that coalesce after this stress level to form the major shear band. At peak (point C), the shear band is partially developed, and the clay loses its homogeneity. Beyond point C, a significant reduction in the CT number occurs because of the reduction in the density due to localization. Hence, the peak strength (point C) is a direct result of localization, and one can argue that point B represents the onset of shear-induced dilation and localization in these tests. Normalization of the onset of dilation (as predicted from their results) to the peak at different confining pressures is also illustrated in Figure 2.9, which shows that the ratio is independent of the confining pressure and that the onset of dilation occurs at 84 % of the peak strength.

2.3.2.2 *Bonded clay*

Vallericca Clay is a bonded overconsolidated clay and can be considered as a weak sedimentary rock. It has a plasticity index of 31.6%, water content of 28.6% and a carbonate content of 30%. In order to investigate the formation of shear bands in this weak rock, Viggiani et al. (1994) carried out a plane-strain compression test using the same technique as that used by Finno and Rechenmacher (2003). Viggiani et al. (1994) also showed that the onset of localization took place pre-peak and was coincident with the onset of shear-induced dilation for this kind of geomaterial.

2.3.3 Hard crystalline rock (granite)

Lockner et al. (1992) studied the deformation mechanism of western granite by using acoustic emissions. These researchers conducted triaxial compression tests on granite samples with a porosity of 0.01 at a confining pressure of 50 MPa. The cell was provided with six piezoelectric transducers to monitor deformation mechanism by using acoustic source-locating techniques. Figure 2.10a shows the deviator-stress versus the axial-strain curve for a granite sample as well as the acoustic emission events recorded during deformation. The letters (A to G) on the deviator stress versus axial strain curve represent the strain levels at which acoustic emissions were recorded. Figure 2.10b gives the volumetric strains-axial strain relationship. According to Lockner et al.'s (1992) acoustic emission observations (see the insert in Figure 2.10), localization starts at point A where several distributed damage sites start to coalesce. This point is also coincident with the volumetric strain reversal based on the researchers' volume change measurements. At peak strength (point B), the shear band is partially formed, and the progression of deformation beyond this point results in the complete formation of a shear band, as shown by the transition from point B to G, accompanied by a reduction in strength as a direct result of localization. Eberhardt et al. (1998), using acoustic emissions, also reported the coincidence of the onset of localization and the onset shear-induced dilation for granite.

Martin and Chandler (1994) carried out a series of damage-control tests on Lac du Bonnet Granite by using strain gauges to monitor local axial and lateral strains. The granite was medium- to coarse-grained and composed of 30% feldspar, 30% plagioclase, 30% quartz, and 10% biotite minerals. Martin and Chandler's (1994) experimental results revealed a threshold after which unstable cracks propagated. This threshold was found to be coincident with the onset of the volumetric strain reversal, which marked the onset of localization due to the coalescence of microcracks. This threshold was shown to be independent of the strain rates and the sample size and was suggested to be the long-term strength instead of the peak strength. Martin and Chandler (1994) showed the independence of this stress on the confining pressure when normalized to the peak strength (see Figure 2.11) and that the onset of dilation occurred at almost 75% of the peak strength.

Takemura and Oda (2005) carried out triaxial compression tests at different confining pressures on Indiana granite from Japan. Indiana granite is composed of 37% quartz, 33% plagioclase, 24% orthoclase, and 6% biotite and clay minerals. After loading and unloading from different stress levels relative to the peak, Takemura and Oda (2005) recorded a major reduction in transversal wave velocities at 75% of the peak strength (onset of dilation).

2.4 Significance of onset of dilation threshold

The laboratory tests reviewed clearly reveal that the failure process in a wide spectrum of geomaterials is dominated by the formation of shear bands and volumetric strain reversal. The initiation of these bands is indicated in the volumetric strain curve as it changes from compression to dilation. The onset of shear-induced dilation, i.e., volumetric strain reversal, appears to be an excellent indicator of the initiation of localization. The peak strength, however, appears to be related to the strain hardening associated with the progressive development of the shear band throughout the sample once localization onsets, and, hence, to be only a signature of the failure process in laboratory tests. Therefore, the onset of dilation is considered of significant importance compared to the peak strength. The onset of dilation signifies the occurrence of non-homogeneous deformation, which should be prevented, and, hence, it should be regarded as an upper bound of strength to avoid the onset of shear-induced failure.

The stress magnitudes associated with the onset of shear-induced dilation ($q_{dil.}$) in the laboratory tests presented have been normalized to the peak strength (q_p) and expressed as follows:

$$[2.2] \quad NDR = \frac{q_{dil.}}{q_p}.$$

Figure 2.12 shows that for the geomaterials examined, this ratio ranges from 0.75 to 0.86, regardless of whether the geomaterial is soil or rock. Once localization is initiated, failure appears to be imminent regardless of the geomaterial's bonding or fabric characteristics. This ratio has also been found to be unique for each of the investigated geomaterials and independent of the confining pressures, as shown in the previous figures. In terms of the absolute values of the normalized dilation ratio and the geomaterials' type (soil or rock), no relation exists between the void ratio and the normalized dilation ratio.

2.5 Onset of shear-induced dilation and long-term strength

The relationship between the onset of shear-induced dilation and the long-term strength of three different geomaterials is illustrated below. If the onset of shear-induced dilation can be verified to be a good indicator of the long-term strength, its significance as a limiting strength threshold is assured. The presented data are taken from recent creep tests for sand, granite, and rock salt. Most of the extensive published laboratory data for creep and long-term strength of various geomaterials do not relate the onset of dilation to the long-term strength for the tested geomaterials. However, Rusch (1959, 1960) suggested the volumetric strain reversal as the long-term strength for concrete and Beiniawiski (1967a) recommended this also for hard rocks.

2.5.1 Sand

Kuwano and Jardine (2002) investigated the creep potential of granular materials under triaxial test conditions by using local axial and lateral strain measurements. Figure 2.13 shows the undrained creep axial strain versus time for Ham River sand at various deviator stress levels. Ham River sand is a uniform sub-angular-shaped quartz sand with a mean particle size of 0.27 mm. A specimen was tested under a mean stress of 200 kPa and had a void ratio of 0.76 and a relative density of 30% (loose state). As the figure shows, the creep potential increased with the increase in the creep stress level. However, the creep potential became more significant after a deviator stress level of 200 kPa, which marked the phase transformation point (the reversal in pore pressure value from a positive value to a negative value) that is equivalent to the shear-induced dilation points in the drained tests. As a result, the onset of shear-induced dilation can be regarded as a safe boundary for excessive deformation in this geomaterial.

2.5.2 Granite

Lau and Chandler (2004) carried out damage controlled tests to determine the damage thresholds as well as the long-term strength of Lac Du Bonnet granite. The axial and lateral strains were measured by using strain gauges glued onto the specimen's surface along with three vertical LVDTs for measuring the axial strains, and a radial extensometer for the circumferential strains. Lau and Chandler (2004) showed that the state of stresses close to

the peak can be sustained for a short period prior to failure due to the temporary strain hardening after the onset of dilation. Hence, the peak strength cannot be relied on for long-term stability, as previously suggested by Schmidtke and Lajtai (1985). To investigate this concept, creep tests were conducted at different confining pressures before and after attaining the onset of dilation stress level. In creep tests, the targeted stress levels (deviator stress) remained constant while the deformation was monitored for at least 24 hrs. These experiments revealed that samples tested below this threshold cracked in a stable manner with no signs of creep failure or rupture. Only samples tested above this threshold tended to fail and showed excessive deformation. It was also observed that a significant change in lateral deformation took place before failure, when the axial strains remained almost constant or negligibly varied. Only when failure was initiated did the axial strain accelerate with the lateral strains, and, hence the failure process was dominated by creation, extension, and coalescence of vertical cracks. Creep stress levels (σ_{cp}) were assigned as a percentage of the onset of dilation stress level (σ_{cd}). The ratio between them was represented by the stress ratio parameter (β) and was expressed as follows:

$$[2.3] \quad \beta = \frac{\sigma_{cp}}{\sigma_{cd}}.$$

Figure 2.14 shows the relationship between the stress ratio (β) and the time to failure in hours at different confining pressures. Increasing confining pressures at the same stress ratio resulted in an increase in the time required to attain creep failure. Within the investigated range of confining pressures, the increase in time to failure was proportional to the increase in stress ratio. In other words, as long as the stresses remained closer to the onset of dilation stress level, the time to failure was delayed. Consequently, the onset of the dilation stress level appeared to be a plausible approximation of the creep stress level at any confining pressure. These observations were also confirmed by the experimental investigation conducted by Lin et al. (2004) by using acoustic emissions on granite samples.

2.5.3 Soft crystalline rock (rock salt)

Rock salts are usually used as the geological host rock for caverns and considered for the disposal of radioactive waste due to these materials' low permeability and ductility (Schulze et al. 2001). The most critical requirement for such application is the prevention of material from cracking to avoid the increase in permeability and hence ensuring the tightness of such rocks. To avoid such effects, Cristescu and Hunsche (1998) recommended that the dilation boundary (the onset of dilation threshold) be considered as the long-term strength. This boundary basically marked the culmination of cracking, damage, increase in acoustic emissions, increase in permeability, and promotion of creep failure. Schulze et al. (2001) conducted triaxial

compression tests on a rock salt formation in Germany to investigate the concept of the dilation boundary. These researchers used ultrasonic wave velocity to detect the onset of dilation, and, hence, damage accumulation. Figure 2.15 shows the relationship between the normalized ultrasonic wave velocities, deviator stress and axial strain for a sample tested under a mean effective stress of 5 MPa. Schulze et al. (2001) identified the onset of dilation boundary by using the reversal of the shear wave velocity. As soon as microcracks commenced, the shear wave velocity exhibited a reduction in its value, reflecting the increase in axial cracks propagating in the vertical direction and the onset of localization. No failure was recorded beyond this boundary of up to 20% strain. This was basically a case of uniform localization as opposed to what occurred in the previously discussed geomaterials where localized dilation was the mechanism of failure.

Hunsche and Hampel (1999) reported Hunsche's (1990) results of true triaxial compression results, to identify the dilation boundary and its relationship to the long-term strength. The normalization of the onset of dilation with respect to the peak strength is plotted in Figure 2.16, which shows that the ratio is independent of the confining pressure, but much lower than the recorded values for all previous geomaterials. Figure 2.17 presents an example of one of the creep tests conducted at a mean stress of 10 MPa. The creep stages and the change in the volumetric strains at each stage of testing are plotted in the figure. According to Hunsche and Hampel's (1999) results, a slow

decrease or a constant volume was observed as long as the stresses were below the onset of dilation (the dilation boundary). In the dilatant regime (where this is an increase in volumetric strain), creep rupture was attained after a certain time or deformation along with a remarkable increase in acoustic emission rate and a decrease in seismic wave velocity. According to their observations, humidity induced creep also occurred above the onset of dilation boundary. These observations suggested that the onset of dilation can be considered as a safe boundary for such geomaterials to prevent long-term instability.

Based on the presented examples, it can be inferred that the onset of dilation can be considered as an excellent candidate for long-term strength. In order to avoid excessive long-term deformation, stresses should be kept below this threshold.

2.6 Extension-induced dilation in bonded geomaterials

In the testing of hard rocks, surface-mounted strain gauges are normally used to monitor the stress-strain response and frequently combined with acoustic emissions (Brace 1964; Brace et al. 1966; Bieniawski 1967*b* ; Wawersik and Fairhurst 1970; Lajtai and Lajtai 1974; Martin and Chandler 1994; Eberhardt et al. 1998; Chang and Lee 2004; Diederichs et al. 2004). Brace et al. (1966) and Lajtai et al. (1990), among others, found based on their observations during the middle stages of compression test conducted on Lac De Bonnet granite, that only

lateral dilation of the cylindrical sample was recorded, with no axial shortening. This suggested that damage initiation and propagation in Lac du Bonnet granite was caused by extension-induced cracking in the vertical direction. The extension-induced dilation can be simply defined as the uniaxial dilation that occurs in the direction of the minor principal stress direction creating microcracks align in the major principal stress direction. These microcracks tend to coalesce, leading to shear failure and the formation of macro-fractures (Lockner et al. 1990).

Martin and Chandler (1994) used damage-controlled triaxial tests on Lac du Bonnet granite to identify the damage stress thresholds, and, more importantly the crack initiation and damage thresholds (see Figure 2.18). The crack initiation stress was caused by extension-induced dilation (30-40% of peak resistance) marked by the initial nonlinearity in lateral stiffness and volumetric strains. In zone (III), microcracks propagated in the vertical direction and formed columnar structures separated by discrete microfractures. At 70-80% of the peak resistance (zone (IV)), the onset of shear-induced dilation, these columns started to buckle, and microcracks coalesced to form major macrofractures (Sammis and Ashby 1986), resulting in the reversal of volumetric strains and localization. Martin and Chandler (1994) argued that strain hardening observed from the onset of shear-induced dilation to the peak strength resulted from the testing methodology and was not a material property. Martin (1997) has shown that the operational strength around the Mine by test tunnel, constructed at the

Underground Research Laboratory (Winnipeg, Manitoba, Canada), was coincident with the deviator-stress required to cause extension-induced dilation (crack initiation stress) in laboratory tests.

Lajtai et al. (1990) indicated that up to the onset of dilation, the formation of vertical cracks propagating in the axial direction dominated the failure mechanism. Figure 2.19 illustrates the results of a cyclic uniaxial compression test on Lac Du Bonnet granite sample. Below the onset of shear-induced dilation threshold, the axial strains were predominantly elastic, and the permanent axial strains remained unchanged; however, a significant change in permanent lateral strains took place. The observation of such vertical cracks and their propagation in the major principal direction prior to the onset of dilation in triaxial testing conditions was also recorded for Oshima granite from Japan with similar mineralogical composition to that of Lac Du Bonnet granite (Schubnel et al. 2003). By using ultrasonic wave velocities, Schubnel et al. (2003) showed that the transversal wave velocities decreased at the onset of shear-induced dilation, along with an increase in acoustic emissions, signifying localization, while the vertical velocity remained constant. In other words, formation of vertical cracks parallel to the axial stress dominated the deformation mechanism.

Vallejo (1987) carried out direct shear tests on brittle kaolinite clay with preexisting cracks. He observed that the failure in the samples started in the form of secondary cracks developed from the tips of the original cracks. These

secondary cracks were developed in the region of the tensile stresses in the samples due to shear stresses. Secondary cracks tended to extend in a direction perpendicular to maximum tensile stresses regardless of the direction of original cracks. Vallejo (1989) also studied the effect of fissures on the strength of brittle kaolinite. He showed that under compressive stresses, local tensile stresses developed at the crack tips. The extension of cracks in the direction of the major principal stress dominated the fracture process of the specimens. The original fissures were acting as stress concentrators from which cracks propagated in the tensile stress regimes so that bridges between the artificial fissures dictated the final strength.

In bonded geomaterials (regardless of the degree of bonding), another characteristic strength threshold exists after which bonds start to degrade, i.e., the extension-induced dilation threshold. The degradation of these bonds under compression is due to the geomaterials' heterogeneity where local tensile stresses develop. The cracks open laterally and extend in the direction of the major principal stress under the compression stress path in a stable fashion, forming vertical columns separated by discrete vertical microcracks. At a certain stress level, (the onset of shear-induced dilation, i.e., the volumetric reversal stress level) these separated columns start to interact to form a shear band, and strain becomes localized in a band composed of these microcracks (Figure 2.20). Directly capturing the extension-induced dilation threshold would be difficult, if not impossible, to achieve from global volume change measurements.

Determining the extension-induced dilation requires local lateral strain measurements so that the nonlinearity of lateral strains can be accurately captured.

2.7 Summary and conclusions

In this chapter, high-quality laboratory data for various geomaterials were compiled and reviewed to reveal the salient pre-peak strength thresholds observed in these data sets. The first part revealed the relationship between the onsets of shear-induced dilation (the volumetric strain reversal stress level) and localization (loss of strain uniformity) for various geomaterials. In the second part, three examples were given to clarify the investigation of the relationship between the onset of shear-induced dilation and the long-term strength (it is defined here as the excessive creep induced deformation that ultimately can lead to rupture). In the third part of the chapter, another threshold was introduced for bonded geomaterials (extension-induced dilation, i.e., the initiation of bond degradation). Based on these investigations, the following conclusions can be drawn.

1. Geomaterials display a marked change in their behaviour prior to reaching the laboratory peak strength, and, hence, pre-peak characteristics may be very important to understand failure and its evolution.
2. The onset of shear-induced dilation (volumetric strain reversal) is an excellent indicator of the onset of localization for a wide spectrum of

geomaterials regardless of their origin and can be directly estimated from global measurements.

3. The peak and residual strengths are direct result of localization and can be considered as signs of failure. The peak strength is due to temporary strain hardening resulting from the localization and propagation of failure.
4. The ratio of the stress level required for initiating shear-induced dilation normalized to the peak strength, for the investigated materials except for the rock salt, ranges from 0.76 to 0.86. This ratio is independent of fabric and/or bonding as well as the confining pressures for the investigated geomaterials.
5. For the investigated geomaterials, the onset of shear-induced dilation may also be an excellent indicator of the creep limit. Beyond this threshold, creep potential becomes significant, and, hence, the onset of shear-induced dilation can be considered suitable for avoiding excessive deformation.
6. The onset of shear-induced dilation may be considered as upper strength boundary threshold for many geomaterials.
7. The onset of extension-induced dilation marks the initiation of bond degradation for bonded geomaterials regardless of the bonding strength. The extension-induced dilation can be estimated only by interpreting the variation in local lateral strains and simply marks the commencement of bond degradation in the pre-peak regime.

8. The pre-peak strength thresholds (the onset of shear-induced dilation and the onset of extension-induced dilation) revealed in this chapter seem promising for determining the upper and lower bounds for geomaterials instead of using the peak and residual strengths.

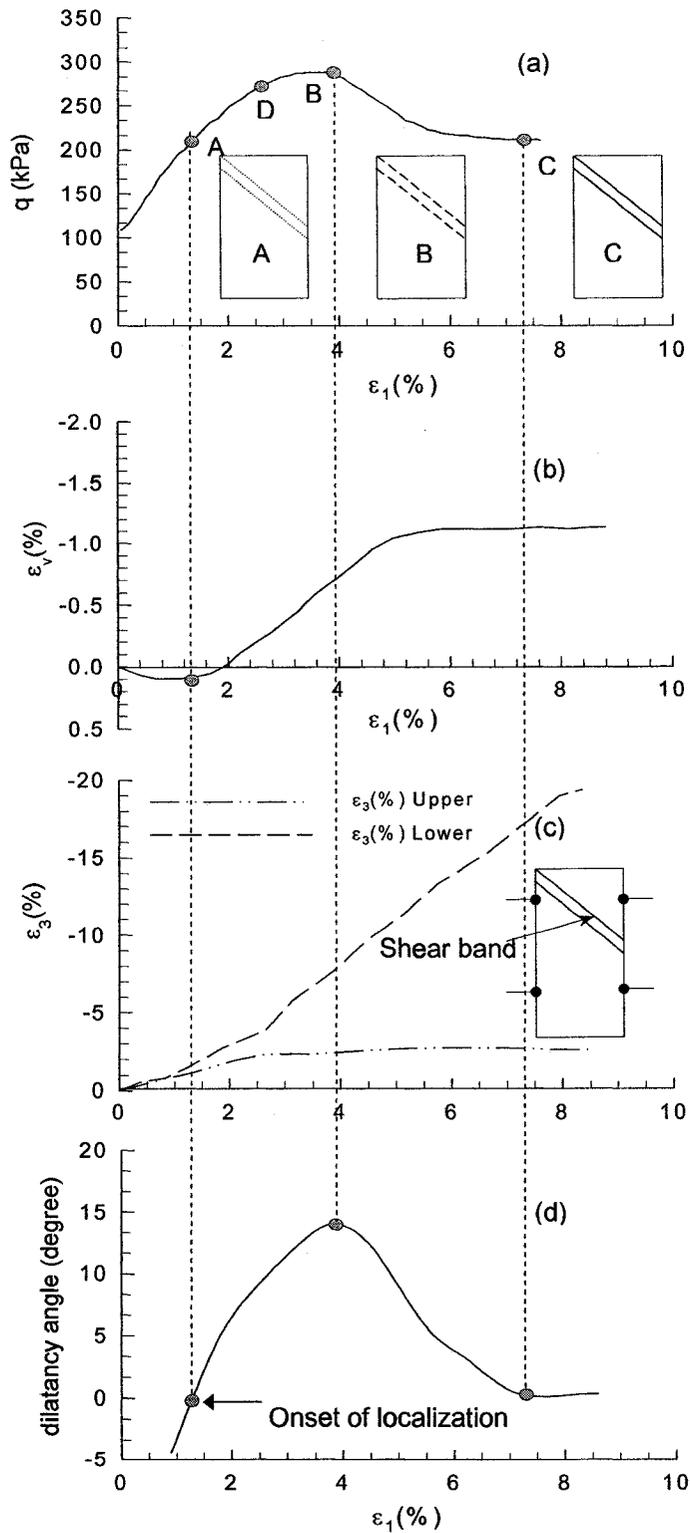


Figure 2.1 Stress-strain and volumetric strain curves for dense sand (modified from Finno and Rechenmacher 2003): (a) Deviator stress-strain curve; (b) Volumetric strain curve; (c) lateral strain curves; and (d) dilatancy angle.

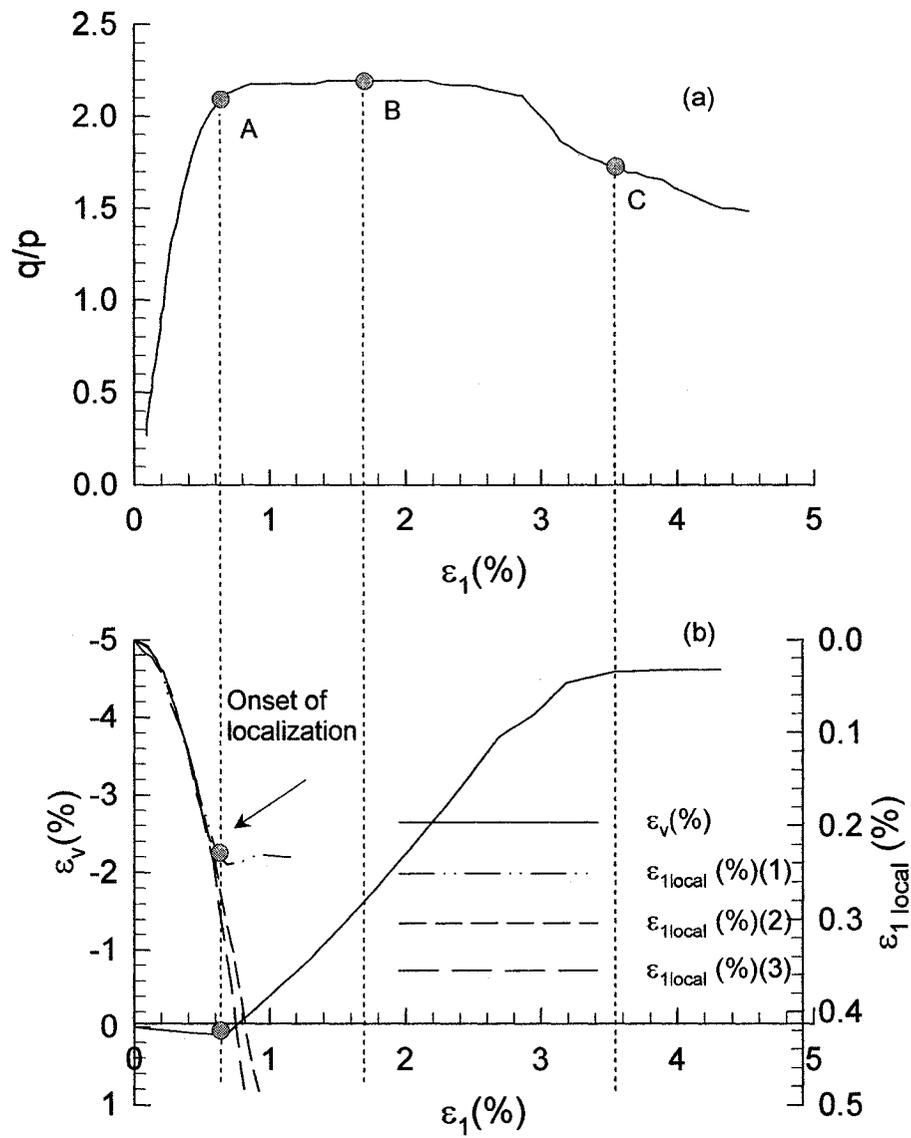


Figure 2.2 Stress ratio-strain and volumetric strain curves for interlocked sand (modified from Cresswell and Powrie 2004): (a) Stress ratio-strain curve, and (b) Volumetric and local strain curves.

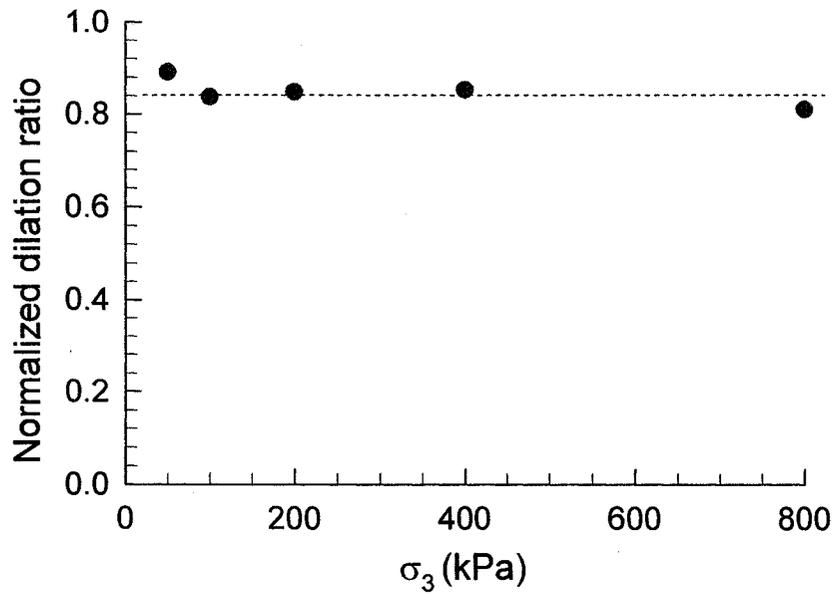


Figure 2.3 The ratio of onset of dilation normalized to the peak strength for locked sand at various confining pressures.

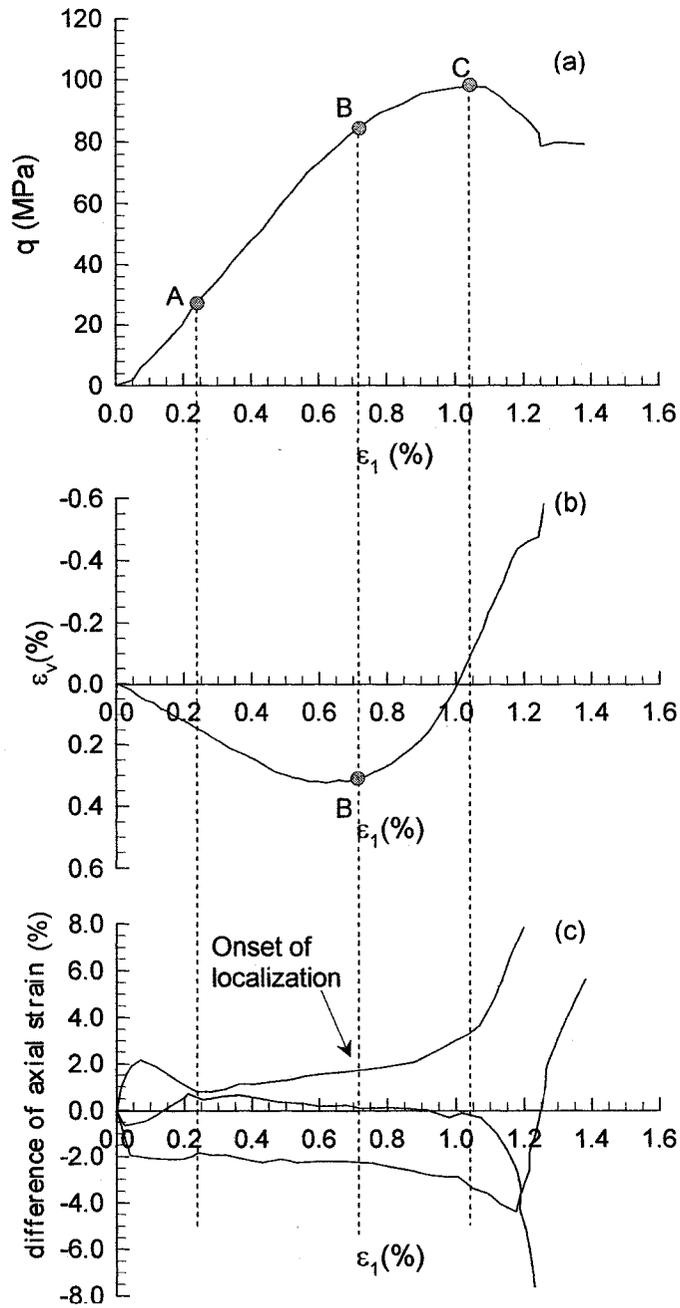


Figure 2.4 Stress strain and volumetric strains curves for Vosges sandstone (modified from Besuelle et al. 2000): (a) Deviator stress- axial strain curve; (b) Volumetric strain curve; and (c) axial strain difference curves.

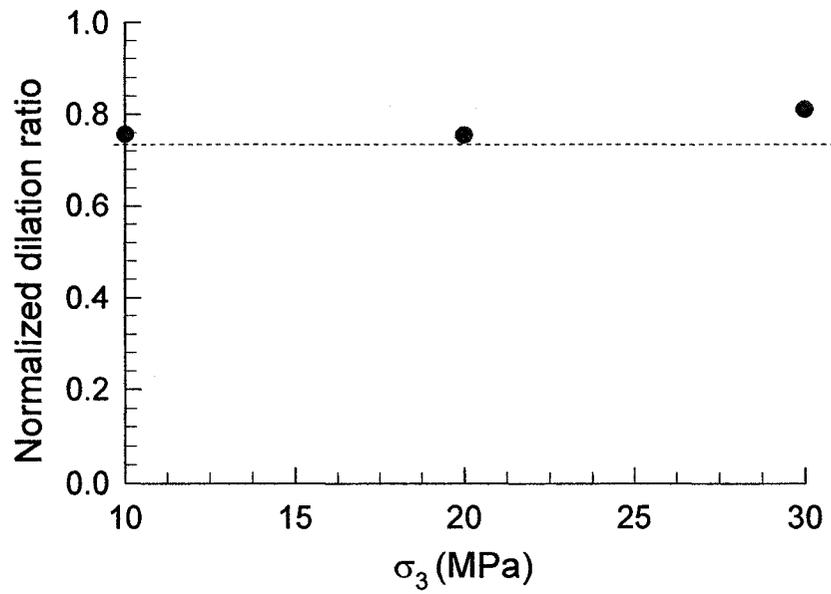


Figure 2.5 The ratio of onset of dilation normalized to the peak strength for Vosges sandstone at various confining pressures.

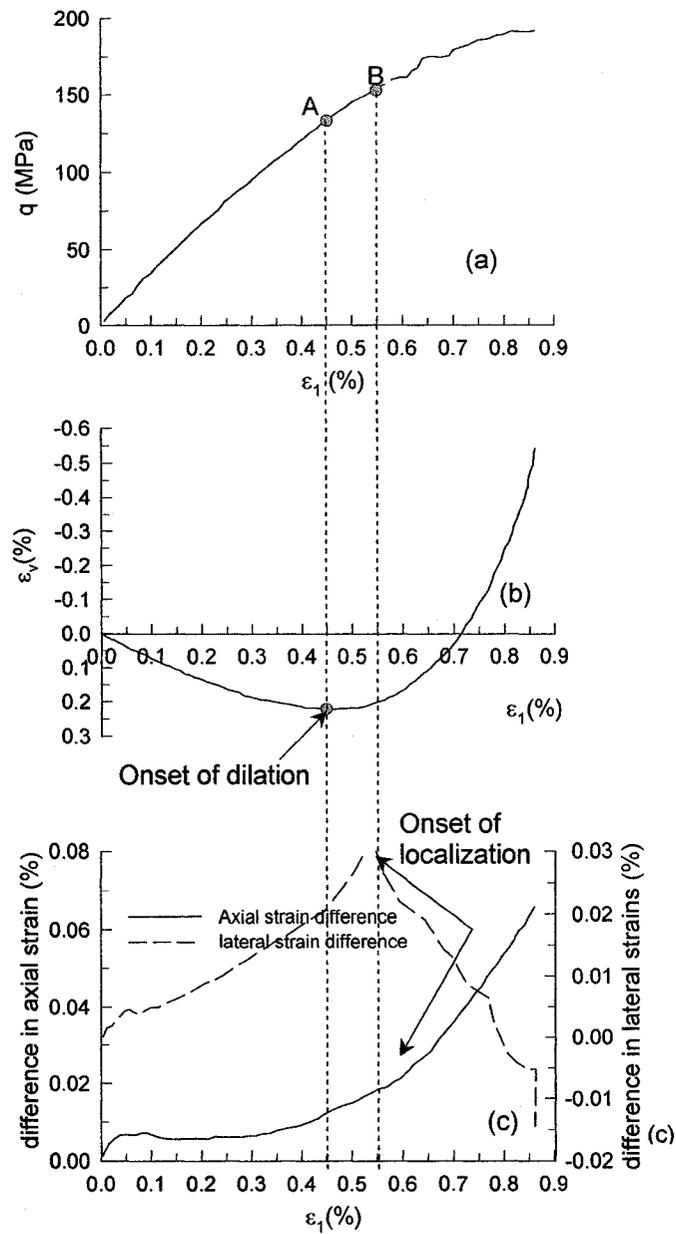


Figure 2.6 Stress strain and volumetric strains curves for Fontainebleau sandstone (modified from Sulem and Ouffroukh 2005): (a) Deviator stress-axial strain curve; (b) Volumetric strain curve; and (c) axial and lateral strain differences curves.

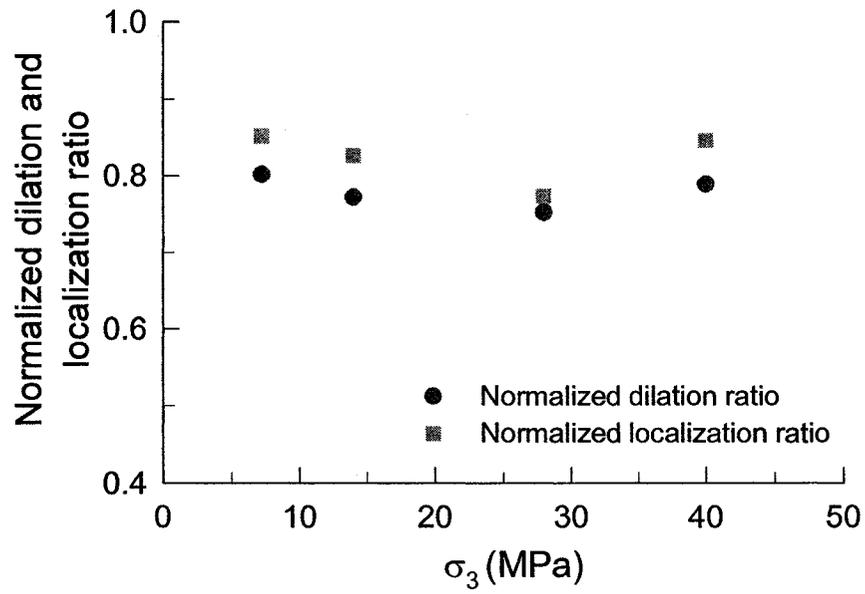


Figure 2.7 The ratio of onsets of dilation and localization normalized to the peak strength for Fontainebleau sandstone at various confining pressures.

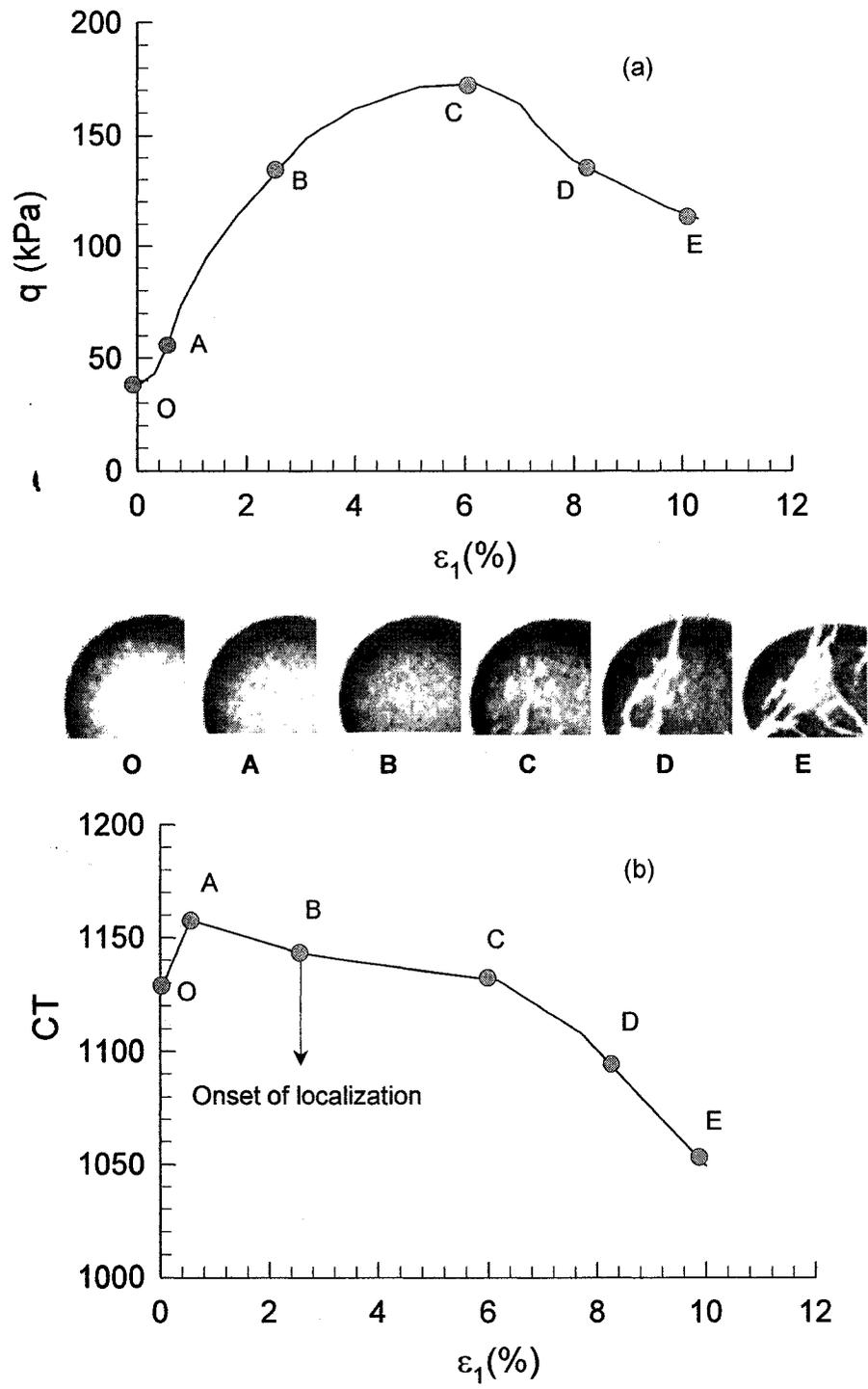


Figure 2.8 Stress-strain and variation of CT number with strain (modified from Su et al. 2004): (a) Deviator-strain curve; and (b) variation of CT number with strain.

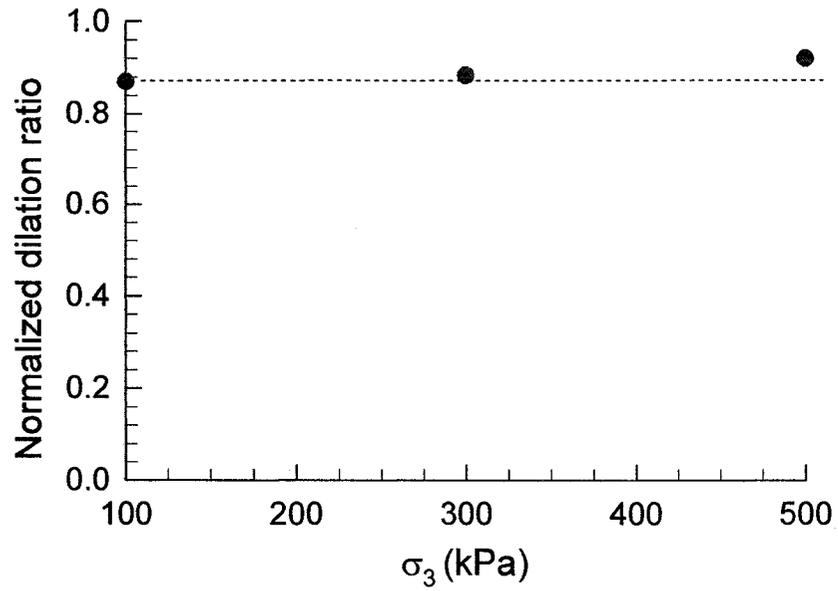


Figure 2.9 The ratio of onset of dilation normalized to the peak strength for an OC clay at various confining pressures.

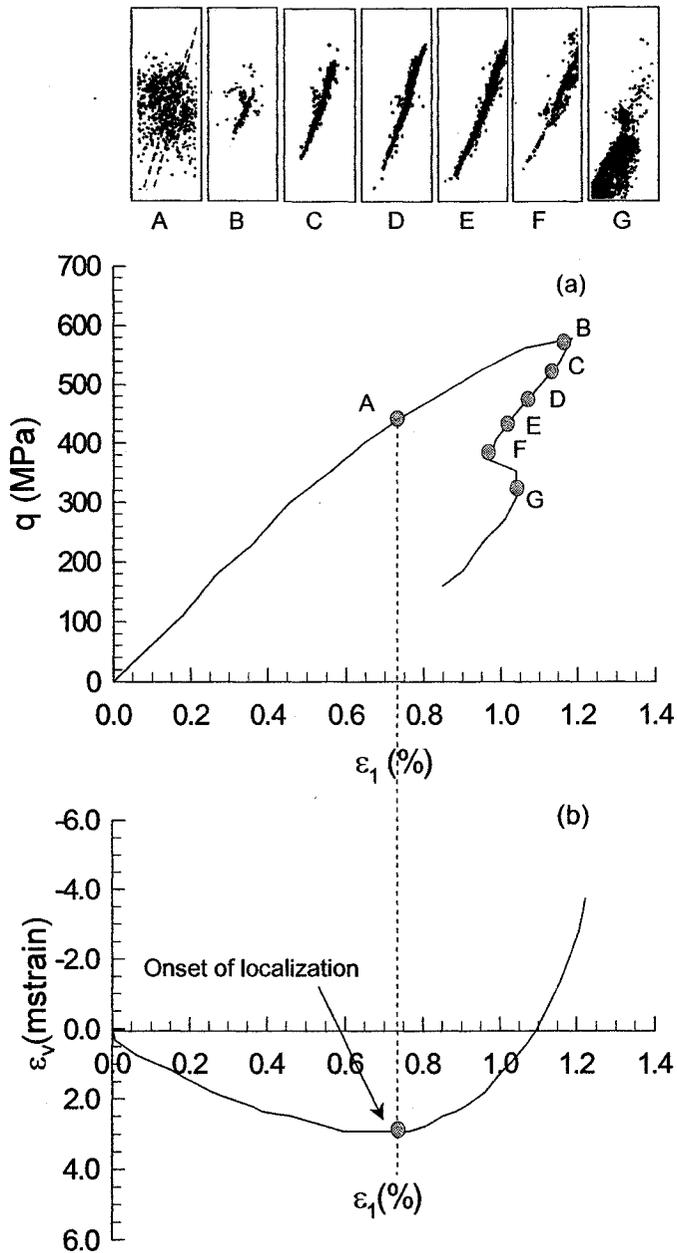


Figure 2.10 Stress-strain and volumetric strain for granite (modified from Lockner et al. 1992): (a) Deviator stress -strain curve; and (b) Volumetric strain curve.

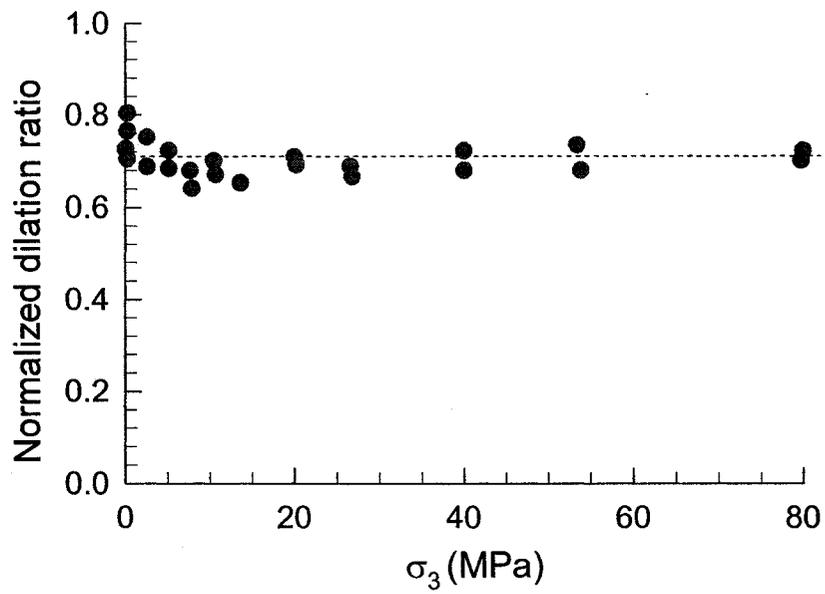


Figure 2.11 The ratio of onset of dilation normalized to the peak strength for La Du Bonnet granite at various confining pressures (modified from Martin and Chandler 1994).

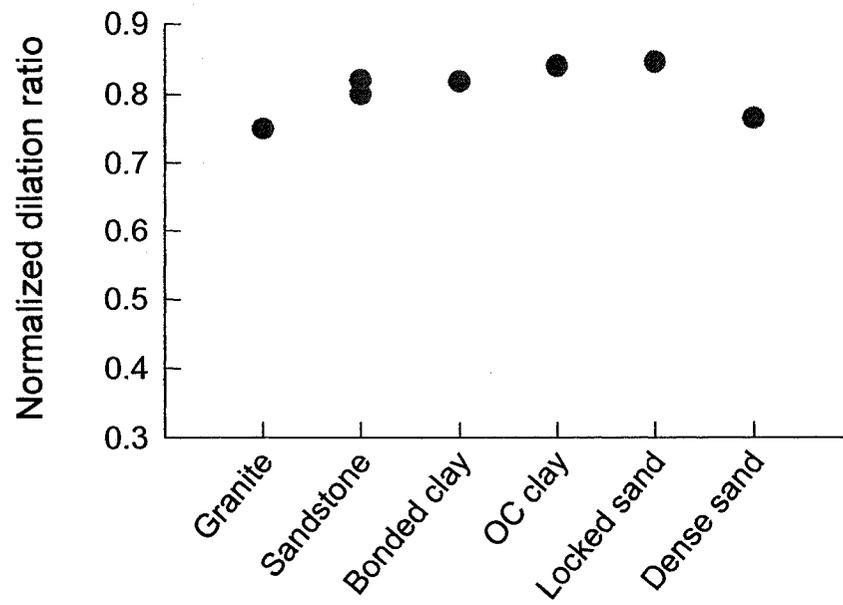


Figure 2.12 Independence of dilation ratio of geomaterials' nature.

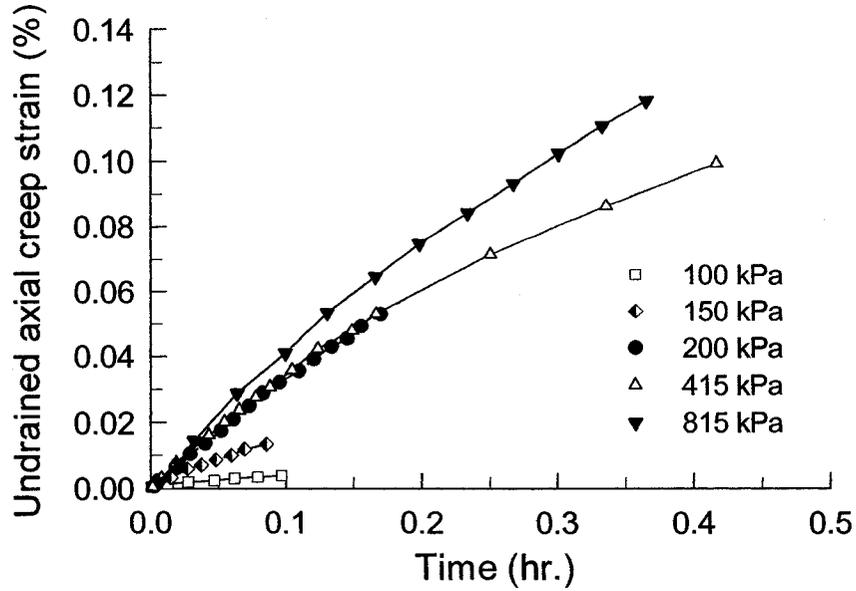


Figure 2.13 Undrained axial creep strain with time under different stress levels for Ham River Sand (modified from Kuwano and Jardine 2002).

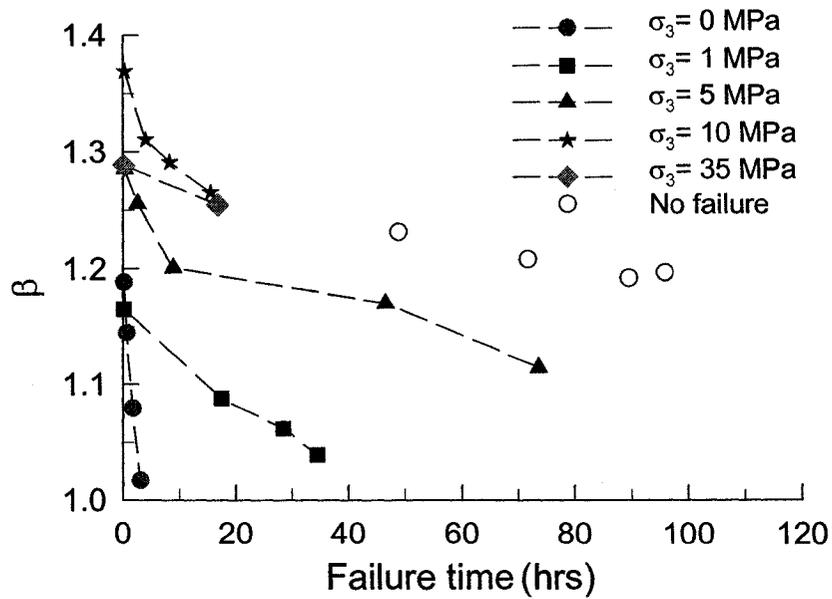


Figure 2.14 Normalized creep stress ratio versus time to failure at different confining pressures (modified from Lau and Chandler 2004).

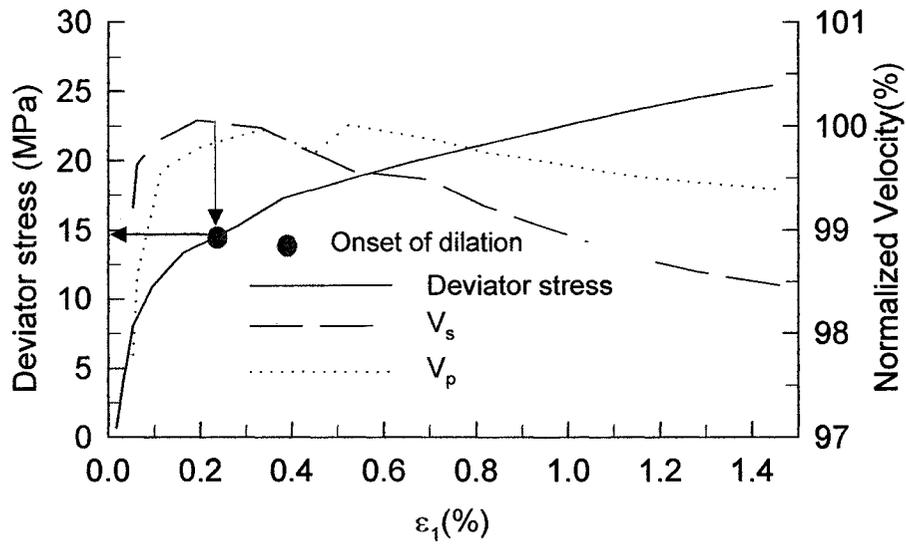


Figure 2.15 Normalized ultrasonic velocities and deviator stress vs. axial strain for rock salt (modified from Schulze et al. 2001).

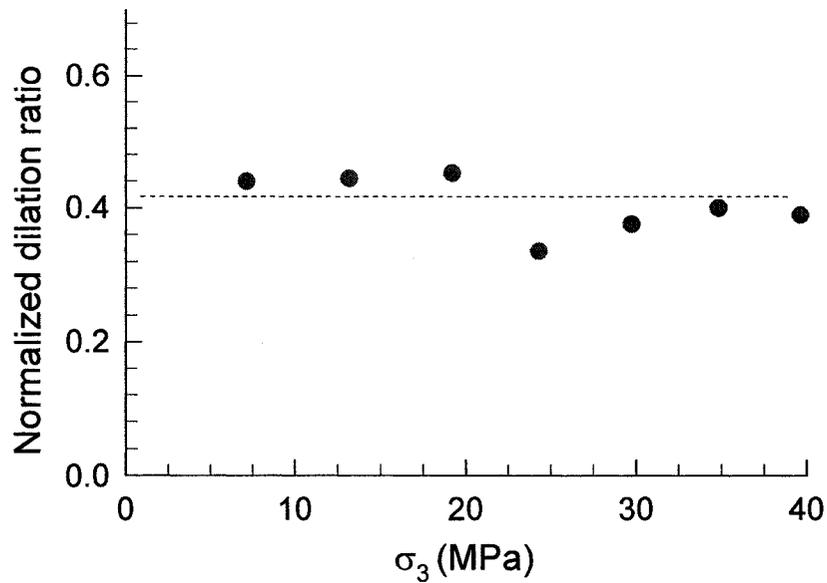


Figure 2.16 Normalized dilation ratio of rock salt (modified from Hunsche and Hampel 1999).

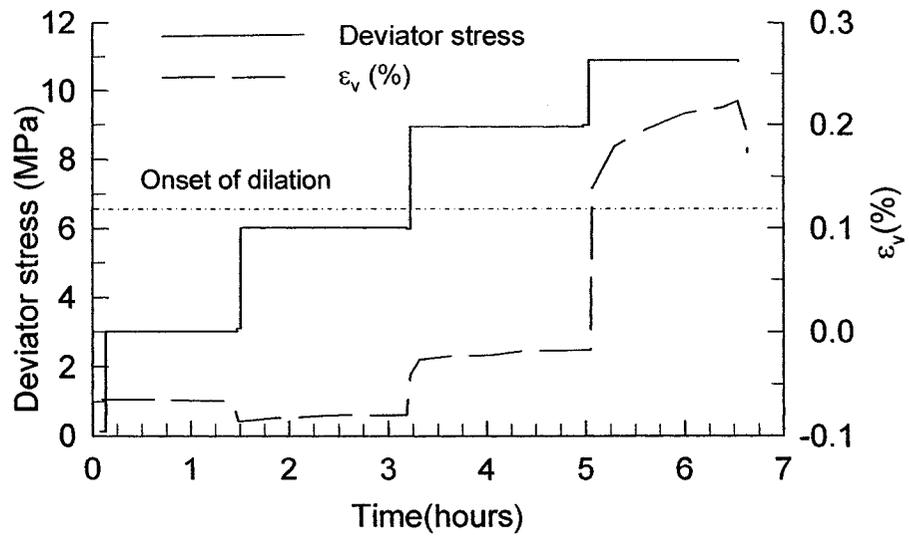


Figure 2.17 Results of a true triaxial creep test for rock salt at constant mean stress at four shear stress increments prior to and past to onset of dilation (modified from Hunsche and Hampel 1999).

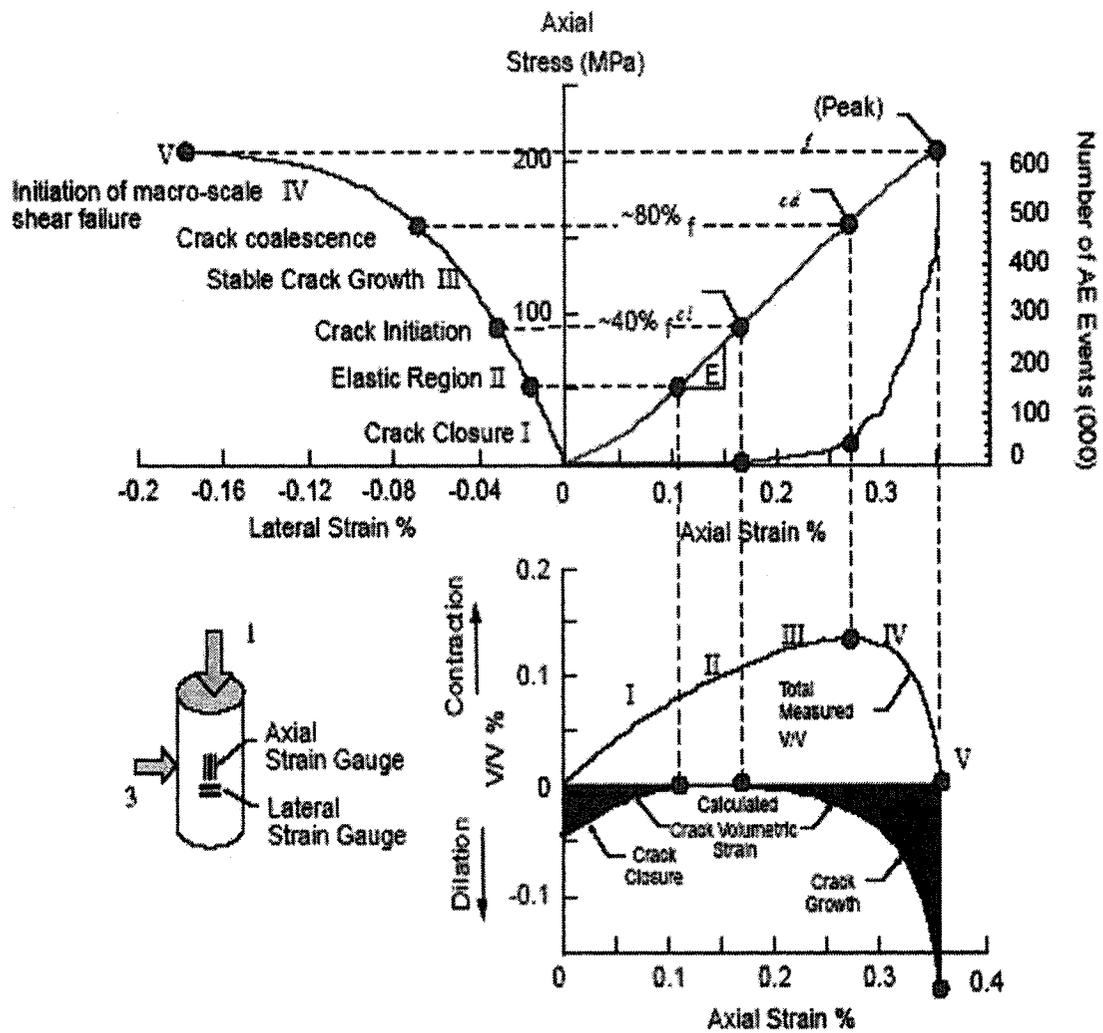


Figure 2.18 Stages of crack development for Lac Bu Bonnet granite using damage tests (modified from Martin and Chandler 1994).

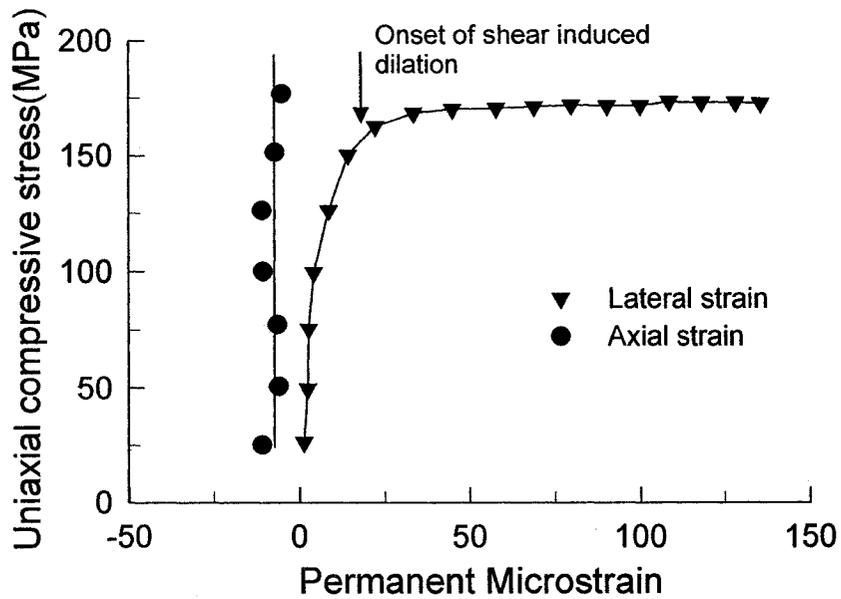


Figure 2.19 Results of a cyclic uniaxial test indicating the variation of permanent axial and lateral strains with evolution of deformation below onset of dilation threshold (modified from Lajtai et al.1990).

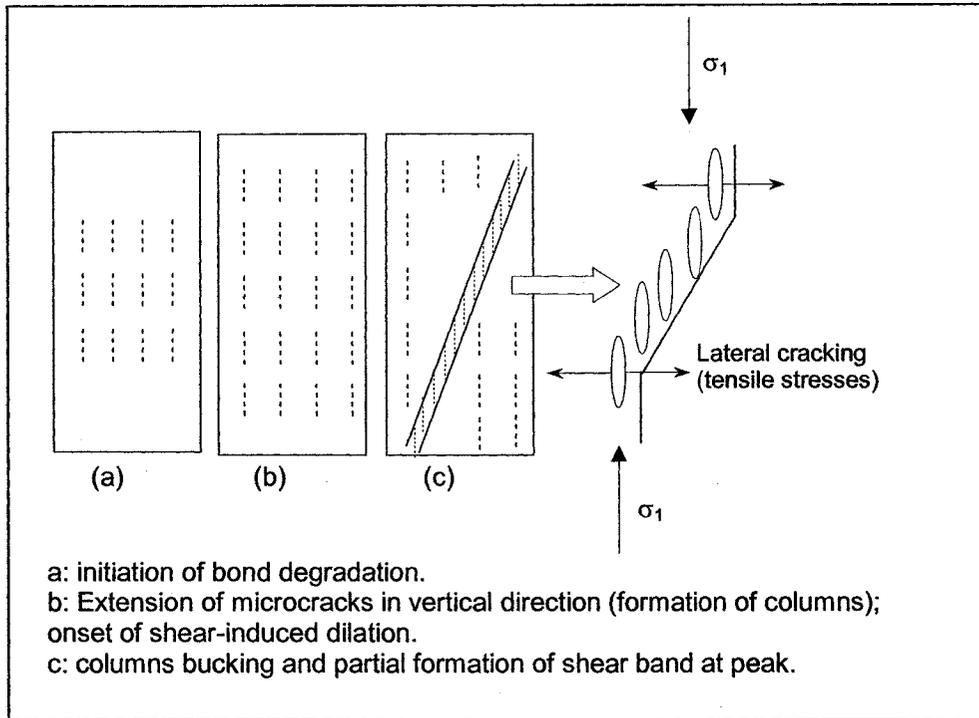


Figure 2.20 Analogue for deformation mechanism in bonded geomaterials.

CHAPTER 3: HIGH-HILL RIVER FORMATION GENERAL CHARACTERISTICS AND UNIAXIAL TEST SET-UP

3.1 Introduction

Nature provides a wide spectrum of geomaterials ranging from soils to hard rocks. In the middle range, a class of geomaterials exists that exhibits the characteristics of both soils and rocks. This zone's geomaterials can be classified into bond-dominant geomaterials (where bonding is the binding agent between particles) and fabric-dominant geomaterials (where interlocking is the binding agent between particles) or *argillaceous* (clay-based) and *arenaceous* (sand-based) intermediate geomaterials. Only recently, the mechanical behaviour of these bond-dominant geomaterials has received considerable attention (e.g., Leroueil and Vaughn 1990; Burland et al. 1996; Cotecchia and Chandler 1997; Amorosi and Rampello 1998; Cotecchia and Chandler 2004). In contrast to bond-dominant geomaterials, fabric-dominant geomaterials have been relatively neglected in geotechnical studies. Our understanding of the mechanical behaviour of this class of geomaterials behaviour has been limited due to the difficulty of sampling, the difficulty in sample preparation due to friability, and the scarcity of such geomaterials, where strength is derived from the grain surface characteristics (fabric) rather than bonding (Barton 1993).

The oil reserves in Alberta's oil sands are a huge source of reliable energy and income for Canada (20.4 billion \$CDN in 2004, according to Statistics Canada). Canada's oil reserves are estimated to be 179 billion recoverable oil barrels (175

billion are located in oil sand) and rank second only to the reserves of Saudi Arabia (259 Billion barrels), see Figure 3.1. The extraction of the oil from the oil-sand formations has resulted in many studies particularly of the Athabasca oil sands as the richest oil reservoir (approximately $210 \times 10^9 \text{ m}^3$ of bitumen). Figure 3.2 shows the major oil locations in Alberta. Nearly $142 \times 10^9 \text{ m}^3$ of the recoverable reserves is located in the lower cretaceous McMurray sand Formation and the Wabisca member of the Clearwater Formation. These reserves have an aerial extent of 46800 km^2 and an average thickness of 34m (Wightman et al. 1989; ERCB 1990). In these types of Formations, particles are interlocked with each other and have heavy bitumen and gas in their pores.

Dusseault and Morgenstern (1978) revealed that factors like gas-bitumen-water minerals, bitumen's viscosity, mineral cementation, and negative pore pressures make no contribution to locked sand strength. Protulipac (1991) showed that oil-free locked sands could be employed to investigate the behaviour of oil sands to avoid the sample disturbance caused by the expansion associated with gas evolution during the sampling of oil sands (McKay 1989), and, hence, disruption of the interlocked fabric (Agar 1984). Experimental investigations of oil sands have revealed that a small increase in the porosity may result in a significant reduction in strength (Agar et al. 1987). Oil-free sand was recovered from cliffs 70 m high along the High-Hill river approximately 60 km east of the McMurray area. This sand is regarded as a perfect candidate for studying the behaviour of fabric-dominant geomaterials and oil sands. In Alberta, a better understanding of

the oil-free locked sand's mechanical behaviour results in a better understanding of oil sands as a geomaterial with interlocked fabric. This understanding may result in enhancing the recovery processes.

This chapter provides a synopsis of the general characteristics of the locked sands according to the previous research on such formations. Microstructural observations and mineralogical composition analyses are then used to verify that the investigated sand formation fits into the locked sands category. Next, the experimental set-up for uniaxial compression test is described in detail, including the LVDTs' calibration, stability and proof testing, and the testing mechanical configuration and modifications. Finally, the significance of local strain measurements in investigating the behaviour of locked sand behaviour is shown by contrasting local axial and external strains' results.

3.2 General characteristics of locked sands

3.2.1 *Locked sands among geotechnical materials*

Locked sands are unique formations that constitute the sands/sandstones borderline and retain properties that are not entirely analogous to those of either soil properties or rocks. Figure 3.3 schematically illustrates a classification of geomaterials in terms of diagenetic effects and uniaxial strength. This figure shows that locked and cohesive sands are on the borderline between soils and hard rocks and belong to the arenaceous family.

The overlap between the boundaries is intentional, for demarcating them is very difficult. Sand and sandstone behaviors have been explored in soil mechanics and rock mechanics, respectively, in terms of compressibility and strength. In dense sand, the strength is a function of the degree of densification and yields a weak interlocking between particles with relative densities of less than 100%. In sandstone, strength is primarily a function of the amount of interstitial cement. Locked sand, despite the lack of authigenic cementation in its pores, possesses strength properties that are much higher than those of dense sand and much less than those of sandstone. Moreover, locked sands can exist at relative densities greater than 100%, in contrast to those of dense sands (Dusseault and Morgenstern 1979). The peculiar strength properties and the deviation from dense sands' characteristics stems from mild diagenetic processes that reduce the original sand mass's porosity after sedimentation to a compact mass with negligible or no cementation (Barton, 1993).

3.2.2 *Diagnostic aids for locked sands identification*

Dusseault and Morgenstern (1979) coined the expression "locked sands" to distinguish the characteristics of geologically old sands from recently deposited sands, compacted sands, and sandstones. Table 3.1 lists these deposits' characteristics as they have been identified in previous investigations. These characteristics are summarized as follows:

1. **Quartzose sands:** these formations are older than Quaternary in age, poorly cemented, and can be reduced to their grains component through rubbing (they are friable). They belong to the orthoquartzite family, which is characterized by the existence of at least 95% quartz-rich detrital grains cemented with authigenic quartz (Pottjohn et al. 1972). These formations show a tendency to disintegrate when immersed in water (Barton et al. 1986). However, this tendency is dependent on the amount of fine-grains in the matrix. Frequent cross beddings with traces of fine grains (such as kaolinite, illite, and iron oxide) have been reported in various investigations (Dusseault 1977; Barton et al. 1986; Richards and Barton 1999).

2. **Microfabric characteristics:** grain-to-grain contacts are large with interpenetrative fabric and rugose (ridged or wrinkled) texture (see Figure 3.4 for classification of the contacts). The contacts' characteristics are usually attributed to pressure solution at the particle contacts and/or post-depositional quartz overgrowth. Interlocking is measured through the tangential index (TI, the percentile of tangential contact to the total contact). Richards and Barton (1999) showed that no relation exists between the TI and the porosity of such sands.

3. **Packing:** porosities are lower than those of dense sands and cannot be reproduced in a laboratory, and relative densities can be in excess of 100%.

Therefore, locked sand exhibits low compressibility and can be used to support foundations where small deformation is expected.

4. *Shear strength characteristics:* locked sand is characterized by brittle behavior, small strain at failure, high friction angles of 57° - 70° , extreme dilation, a higher residual friction angle of 30° - 35° compared to that of the usual Quartzose sands due to rough surface textures, and a reduction in strength and dilation rates due to wetting. The high TI (less interlocking) results in a negligible cohesion intercept (Richards and Barton 1999). Fabric cohesion (the cohesion intercept due to interlocking) is observed only in fine- to medium-grained sands due to interlocking. In contrast, fabric cohesion is negligible in coarse-grained locked sands. Furthermore, the cement content, if it exists, and the fabric cohesion for such materials have no obvious relationship.

3.3 Investigated area and sampling

3.3.1 *Oil-free slopes and sampling area*

Dusseault (1977) carried out an extensive laboratory and field investigation of the lithology and stratigraphy of immature (created during glaciations) oil-sand outcrops in the Athabasca Region. All oil-sand and oil-free sand slopes were found to be within 115 km of Fort McMurray, and its formation was attributed to the erosive force created via the down-cutting action of young vigorous streams (the shaded area in Figure 3.5). As glacial influence

operated a short time ago, these slopes are considered to be young in age, and they follow river streams in this regime.

Within this regime, several exposed oil-free sand cliffs up to 70 m high, along High Hill River and Cottonwood Creek exist in the McMurray area. Block samples were procured from the outcrops along the High-Hill River banks. The block samples retrieved from the cliff exposure were located 60 km east of McMurray and one kilometer northeast of the junction of the High Hill and Clearwater Rivers (see Figure 3.5). The oil-free block sample was completely dry in situ, presumably due to a shale cap at the top of the oil sand formation. The blocks had dimensions of 22 cm by 22 cm by 15 cm and were cut from the exposed faces by using an electrical wire saw. These blocks were placed in wooden boxes encased with expanding foam for transportation and storage.

3.3.2 Geological settings of McMurray Formation

Figure 3.6 shows a geological survey map, depicting the bedrock in Alberta, produced by The Alberta Energy and Utilities Board and Alberta Geological Survey agencies (EUB and AGS 2005). According to the geological survey and the identified geological elements, the investigated formation belongs to the McMurray formation of the Mannville group (the group included the base of the McMurray formation to the Grand Rapids Formation), which was deposited in the lower Cretaceous period and includes the McMurray

Formation, the Wabiskaw member, the Clearwater Formation, and the Grand Rapids Formation. As most of the bitumen in the Athabasca oil sands is within the McMurray formation (ERCB 1990), several researchers have endeavored to identify the stratigraphy, geological settings and depositional history by using sedimentological, paleontological, and ichnological data sets (e.g., Carrigy 1959; Flach and Mossop 1985; Ranger and Pemberton 1992; Wightman and Pemberton 1997; Hein et al. 2000).

3.3.2.1 Stratigraphy and depositional history of McMurray Formation

The word “McMurray” was first proposed by F.H. McLean in 1917 in order to describe strata containing tar sands. Carrigy (1959) provided the first informal tripartite stratigraphic subdivision of the McMurray Formation’s members as the lower, middle, and upper McMurray Formation, for a schematic depiction see Figure 3.7a. The lower McMurray Formation unconformably overlies a succession of Devonian evaporates and carbonates (Carrigy 1973a; Wightman and Pemberton 1997). In the Devonian period, dissolution of evaporites caused the overlying carbonates of the Waterways Formation to collapse and subside, creating the basin where the sediments were accumulated. Large paleovalleys formed by the salt dissolution collapsed, and less resistant carbonates developed on the unconformity surface, creating a localized central trunk drainage system (Ranger and Pemberton 1992; Wightman and Pemberton 1997) and hosted the McMurray Formation, which was originally exposed to extreme weathering before deposition. The

McMurray Formation was deposited in the early Cretaceous period marked by continental subsidence as well as sea transgression. Sedimentation continued until the late Cretaceous period, and approximately 760-1220 m of sediments were deposited over the McMurray Formation after its deposition termination due to a relative fall in the sea level. Subsequently, erosion began in late Cretaceous period for over 70 Ma until the onset of continental glaciations in North America where large quantities of McMurray sands were removed by glacial scour.

The McMurray Formation accumulated in incised valleys formed by fluvial processes and subsequently transgressed by marginal-marine environments during an early Cretaceous sea-level rise. Thus, the McMurray Formation displays a continuum of sedimentary environments, from fluvial in the lower parts, to estuarine in the middle, to marine shoreface near the top. The Lower McMurray Formation is generally medium- to coarse-grained and massive-appearing to crudely cross-bedded, and contains no ichnofossils. The formation's beds are most commonly interpreted as having a fluvial nature. The Middle McMurray deposits are dominated by inclined heterolithic stratification, which is interpreted to represent deposition on tidally influenced point bars and which forms the primary reservoir unit in the Athabasca Oil sands. Notably, the Middle McMurray Formation contains a brackish-water trace fossil assemblage and, accordingly, is interpreted as representing a fossil estuary (Pemberton et al. 1982). The Upper McMurray deposits are

variable, but generally contain a comparatively open marine signal. These deposits are normally interpreted as shallow and low-energy shore-face deposits, and small deltaic complexes (Gingras and Rokosh 2004).

3.3.2.2 *Lithology of McMurray members*

Dusseault (1977) summarized the related lithological features of McMurray oil sand formation. According to his description, this formation is heterogeneous with respect to its grain size but homogeneous in terms of its mineralogy. The formation is characterized by well sorted un-cemented sand consisting of over 94% quartz (belonging to the orthoquartzite family) and clay minerals which are mostly illite and kaolinite. Figure 3.7b shows the grain-size distributions of all the McMurray members, marked as Group band I (Lower member), band II (Middle member), and band III (Upper member). The same figure shows the grain-size distribution of our samples, which will be discussed subsequently. Table 3.2 summarizes the depositional environment as well as the lithology of McMurray Formation sand members.

3.4 **Physical properties**

The sampled blocks appear as white to light gray in color, characterized by thin cross-black beddings and a rough texture. Under an optical microscope, the majority of the detrital grains are clear, white, or translucent. A light finger tip's pressure, along with continuous rubbing, results in a degradation of the interlocked structure, reflecting either extremely weak or no cementation.

Table 3.3 presents a summary of the pertinent physical properties of such a formation based on grain-size analysis and maximum and minimum void ratio test results in accordance with ASTM Standards. According to the ASTM Classification of Soils for Engineering Purposes (Unified Soil Classification System) (D 2847), this deposit is classified as poorly graded fine to medium sand (SP). A calculated relative density of 125% reflects the existence of exceptional grain packing, which cannot be reproduced in the laboratory and confirms that this deposit fits into the locked sands category, described in Table 3.1. Based on the locked sands' classification introduced by Barnes and Dusseault (1982), describing the degree of diagenesis and its reflection of the contact nature based on the difference between the minimum and the actual porosities, this sand belongs to the second category (which has a porosity reduction of 2.7 %). This category is characterized by small fabric cohesion (an interlocked fabric) and mild diagenesis reflected in tangential, long, and occasional interpenetrative contacts. This deposit is composed mainly of well-sorted sand particles with a small portion of fine particles and a high relative density due to mild diagenesis.

3.5 Micro-fabric analyses and mineralogical composition

The visual inspection of the blocks showed their extreme friability, interlocking of particles, and existence of crystals with different colors. In order to investigate such observations, scanning electron microscopy (SEM) observations, examination of colored thin sections by using an optical microscope, and energy

dispersion X-ray analysis were carried out on these samples. Before interpreting the results, the source of diagenesis in sedimentary rocks is discussed.

3.5.1 Source of interlocking due to diagenesis

Diagenesis processes of predominantly sand-size particles have been extensively studied in Sedimentology and Petrology for decades to reveal the source of pore reduction of detrital grains. "Diagenesis" is defined as the process through which sediments are transformed into rocks. It is the sum of physical, inorganic, chemical, or biochemical changes in a sedimentary deposit after initial accumulation including mechanical compaction, addition of new material, removal of materials and transformation by change in the mineral phase, and replacement of the mineral phase by another one (Friedman and Sanders 1978). The mechanical compaction stage of sediments is achieved by rotation to a stable position under overburden. During this stage, fractures can be easily formed as a result of the huge pressure on the particles, forming floating and tangential contacts (Taylor 1950). The second stage is composed of the alteration of sedimentary deposits through chemical and/or physical processes. In this stage, other kinds of contact such as concavo-convex (interpenetrative), long and short contacts, and sutured contacts are formed (Taylor 1950; Siever 1959). The transformation of sand to sandstone (lithification) is achieved through pressure solution and quartz overgrowth. Pressure solution occurs at particle contacts in aqueous solution where dissolution at contacts takes place due to

unequal stress distribution at them. This process results in a wavy interlocked fabric with liberation of silica in solution. The increase in the thickness of the fine-grains at particle contacts results in an increase in pressure solution due to the increase in diffusivity (Heald and Larese 1974). Quartz overgrowth (Authigenic silica or Syntaxial rim) is a direct result of silica introduced into the solution either by pressure solution or silica-rich streams, the infiltration of silica, or the diagenesis of shales and the transformation of smectite to illite with an increase in soluble silica (Sibley and Blatt 1976). The overgrowth occurs on detrital grains, giving a euhedral shape (well-formed crystal with a sharp, easily recognized faces) as a secondary growth in selective directions based on the density of the fine-grained coatings. Fine-grained coating can result in uneven growths of detrital grains (Heald and Larese 1974). This growth is responsible for the grain's unusual shapes of grains and can produce high interlocking by pushing adjacent grains or even by overlapping two growths of two adjacent grains (Pittman 1972). Unfortunately, differentiating between the two mechanisms by using simple optical analysis is very difficult, due to the same crystallographic and optical orientation of the overgrowth and the underlying nucleating detrital grains, unless a dust ring is detectable on particle surfaces (Pittman 1972). However, this mechanism can be identified by using the Cathode luminescence petrology introduced by Sippel (1968), which is beyond the scope of this investigation. Authigenic clays exist in such formations in the form of pore lining, pore filling,, or even fracture filling (Wilson and Pittman 1977). Heald and Larese (1974) found that

argillaceous coating can prevent quartz overgrowth from forming strong bonds, with uneven growths or throat nature with neighboring grains, and, hence, can result in high friability through failure of the coating.

3.5.2 SEM and Energy Dispersive X-ray Analyses

An examination of the locked sand samples was carried out by utilizing a JEOL 6301F (Field Emission Scanning Electron Microscope) to provide high-resolution digital images at magnifications ranging from 20x to 250,000x, as well as an energy-dispersive X-ray analysis of the identified elements. The undisturbed specimens were trimmed, and broken off in the middle to give a clean, fresh face for observations. The base of each specimen was attached to a metallic scanning electron microscope stub by using low-resistance contact glue. After the special silver glue had been cured for 24 hours, the top fracture surface was blown gently with air to disperse any possible debris. Dried SEM specimens were coated with a conducting gold film to prevent surface charging by the electron beam. Throughout the SEM process, the specimens were kept under room temperature and subjected to 4×10^{-7} torr of vacuum. Figure 3.8 shows selected SEM images for illustration.

The particles shown in Figure 3.8 are characterized by high angularity and irregular shapes. Fine materials acting as coating and pore filling are also evident. Most of the particle contacts are tangential, long, and short straight with a few concavo-convex and sutured contacts. Fractures within detrital

grains are also clear in most images, which can be attributed to the initial stage of diagenesis (mechanical compaction). Figure 3.8 also reveals crystal overgrowth resulting in intimacy between neighboring particles and frequently pushing of neighboring grains. The pitting cavities shown in Figure 3.8 indicate the severe chemical dissolution experienced during diagenesis. Magnification of the fine-grained particles, acting as bridges and coating between grains, did not reveal any particular orientation and reflected the flocculation of the flakes. The fine grains could be a direct result of either the dissolution of the feldspar grains (silicates of aluminum) in the matrix, which also produces amorphous silicates, or the of diagenesis of the overlying shale as common sources of the fine- grained particles in sandstones (Pettijohn et al. 1972).

The X-rays of the targeted particles noted during the SEM examination could not clearly identify the particles' mineralogy due to the coating of the detrital grains with fine-grained materials, which was expected as a reflection of the sedimentary rocks' nature. Therefore, it was decided to conduct the analysis by using a mapping technique. The analysis was accomplished by collecting grains with different colors attached to a metallic scanning electron microscope stub (Figure 3.9). The particles were numbered and labeled as clear, white, red, purple, and black. The clear and white particles numbered (1-5) were basically silicates with an insignificant portion of aluminates (mainly quartz particles) (see Figures 4.10a and 4.10b). Several investigators

(e.g., Pettijohn et al. 1972; Friedman et al. 1978) have indicated that quartz minerals dominate the mineralogical composition of orthoquartzite family. Particle No. 6 was a white kaolinite particle, as evidenced by the nearly equal ratio between the silicates and the aluminates and the existence of potassium (Figure 3.10c), which is a typical clay mineral in sandstones of the lower Cretaceous age in Alberta (Carrigy and Mellon 1964). Particle No. 7 was a red feldspar particle, which was presumably the source of the kaolinite, and indicated that the source of the parent rock as either igneous or metamorphic rock (Figure 3.10d). Particle No. 8 was purple in color and identified as belonging to the amphibole family (Figure 3.10e). Particle No. 10 was a black chert identified by the existence of multiple quartz crystals in the same particle (Figure 3.10f) and particle No. 11 was a black titanium particle which is typically found in sedimentary rocks, as shown in Figure 3.10g (Pettijohn et al. 1972). The results of the X-ray analysis suggested that the black bands were fine rock fragments and chert particles, reflecting the depositional layering common to sedimentary environments. In order to confirm these observations, thin sections were also prepared to provide additional insights into this deposit's depositional history.

3.5.3 *Thin sections under petrographic microscope*

Thin sections were prepared to provide details of the grain shapes, the post-depositional processes, and the layering of the black bands. Sections were prepared with dimensions of 75x55mm after impregnation with a blue-dyed

epoxy resin under a high vacuum. These sections were oriented both parallel and perpendicular to the layering. Examination of the thin sections was carried out by using a petrographic microscope, under different magnifications, connected to a Nikon digital camera. In the well-sorted bands, the particles ranged from angular to very angular with frequent sub-angular particles (Figure 3.11). Approximately 90% of the particles were quartz particles with 5% polycrystalline quartz originating from a metamorphic rock, 4% were very fine -grained rock fragments (mostly chert), and less than 1% were clay minerals, bridging quartz particles as shown in Figure 3.12. The observed contact nature was generally short and long contacts with a few tangential and interpenetrative contacts.

In the less-sorted bands, most particles were angular in shape. Variable fine rock fragments were present in the band under examination. Their presence implies the variability in the parent rocks from which the sediments were transported. This less-sorted band was composed of approximately 60% monocrystalline quartz, 5% polycrystalline quartz, 10% siliceous siltstone, 20% chert, and 5% feldspar and clay minerals.

Figure 3.13 shows a partially clean fractured quartz grain. The deposit appears to have been subjected to overburden loads that were substantial enough to cause this fracturing and, hence, to influence the deposit's structure. The clean cracks within the quartz grains implies either that the clay

minerals were deposited following the compaction stage or that they were too weak to prevent loss of porosity due to compaction by early cementation. In the poorly sorted layers, evidence that compaction embedded the quartz particles between the fine-grained rock fragments was found, reflecting the influence of the burial depth after deposition (Figure 3.14). The overgrowth was distinguished by dust lines as one of the diagnostic aids for identifying the overgrowth (Pittman 1972) (see Figure 3.15). A fresh skeletal grain derived by dissolution is also apparent in Figure 3.16, suggesting that the process of dissolution occurred after the compaction stage and is still in progress; otherwise, the compaction process would have destroyed this skeleton.

Based on these results, the characteristics of this formation are consistent with locked sand characteristics as summarized by Barton et al. (1986) and introduced by Dusseault and Morgenstern (1979). The formation is basically a lower cretaceous middle McMurray Formation that is best described as over-consolidated un-cemented and well sorted friable sand with interlocked fabric, stems from deposition of authigenic silica, and has insignificant portions of clay minerals, acting as bridges and coating, that are responsible for the friable nature of its weak interlocked fabric. The fine particles presumably originated due to either the diagenesis of the overlying shale formation or the feldspar inherited from parent rocks. The clean cracks within quartz grains imply that clay minerals were deposited following the compaction stage or its

weak nature, preventing the loss of porosity due to compaction by early cementation.

3.5.4 Fabric orientation using Digital Image Analysis (DIA)

Baghini (1998) investigated the anisotropy of the same block samples used in the current investigation. The SEM images were analyzed quantitatively and qualitatively by using public-domain digital image analysis software (NIH 1997). The density slices technique, using a range of gray level thresholds, was used for segmentation instead of standard digital image techniques to overcome non-uniform lighting and noise in images. Both vertical and horizontal sections cut from the original block were investigated to identify the particle shapes as well as the preferred orientations of the long particle axis and their intensity. The vector mean direction (θ_m) and vector magnitude (V.M (%)), as recommended by Curray (1956), were used to analyze two-dimensional orientation data to establish the preferred orientations and their intensity (the measure of disperse), respectively. The axial ratio adopted by Oda (1972a and 1972b) was employed for particle-shape identification. The three parameters can be expressed as follows:

$$[3.1] \quad \theta_m = \frac{1}{2} \tan^{-1} \left[\frac{\sum \sin 2\theta_k}{\sum \cos 2\theta_k} \right],$$

$$[3.2] \quad \text{V.M}(\%) = \frac{100}{\sum a_k} \sqrt{(\sum a_k \sin 2\theta_k)^2 + (\sum a_k \cos 2\theta_k)^2},$$

$$[3.3] \quad \text{Axial ratio} = \frac{1}{\sum a_k} \sum_{k=1}^{a_k} \left(\frac{L_2}{L_1} \right)_k,$$

where θ_k is the direction of the observed quantity (the orientation of the particle long axis with respect to a horizontal reference) on an image from -90 to $+90$; L_2 and L_1 are the length of short and long axis, respectively; and a_k is the number of observations. The complete random orientation of particles gives a vector magnitude value of 0. On the other hand, a 100% value means a complete preferred orientation. Selected results for vertical as well as horizontal sections are reproduced in polar coordinates in Figure 3.17. For vertical sections, the particles show a unimodal distribution while horizontal sections show a polymodal distribution. The vertical sections show a mean vector of (1-10) and a magnitude of 18-25%. The horizontal sections show a mean vector of (15-39) and a magnitude of 7.5-15.5%. Axial ratios range from 0.68-0.71. The axial ratios indicate the high angularity of this formation as revealed from particle analysis. Baghini (1998) compared these values with those of other sand formations and concluded that diagenesis has no effect on particle shape and orientation.

The vertical sections show an almost horizontal preferred orientation. The low intensity of the vector magnitude indicates the absence of perfect horizontal orientation for the particles. In terms of porosity, the vertical sections do not exhibit any preferred flow path. For the horizontal sections, there is no

preferred orientation for all sections, and particle-orientation randomness is higher than those of vertical sections. As Baghini (1998) showed, the mechanical response of vertical and horizontal samples is different due to their inherent anisotropy. Vertical cores tend to give higher strength and higher dilation than horizontal cores, basically because the preferred orientation of contacts' normal parallel to the major principal stresses' direction (particularly in the triaxial compression state) (Oda and Konishi 1974).

3.6 Experimental procedures and measuring technique

The uniaxial compression test was favored for two reasons: (a) it directly reflects the effect of the interlocking on the strength as an intrinsic strength component independent of the confining pressures; and (b) by testing this formation in its natural dry state, one avoids the problems associated with pore fluid, which could possibly alter this formation's delicate micro-structure.

3.6.1 LVDTs and data acquisition

The axial and lateral strains were measured by using AC-operated miniature 100 MHR series LVDTs manufactured by Schaevitz. These transducers possess a nominal linear range of ± 2.54 mm and a linearity of 0.15% over the full range at 10 kHz and, hence, have the ability to capture small deformations. An additional advantage is the light weight of the transducer (6 g.) and relatively small dimensions: a length of 25.4 mm and a diameter of 9.4 mm. The primary coil is excited with an input voltage of 3 V RMS, which

requires a special signal conditioner connected to a DT800 datataker, produced by dataTaker Ply Ltd., acting as a data acquisition and data logger system. The DT800 has the ability to scan data from each sensor channel at 40Hz, and has an analog-to-digital converter with a resolution of 18 bits and a noise rejection of 50dB. Communication is done through a connection with a personal computer to store and upload the data in real time. The DT800 has a range of $\pm 10\text{V}$ and a resolution of 1mV for the input voltage and 1 μV for the displayed output voltage (1mV input can be displayed up to 0.001mV output voltage). Another signal conditioner (a LVM 110 produced by Schaevtiz sensors with a range of $\pm 5\text{V}$) was provided to augment the system with an additional three LVDTs from the same manufacturer. Figure 3.18 presents a schematic diagram of these components.

3.6.1.1 LVDT calibration

The LVDTs were calibrated by using a small micrometer, as shown in Figure 3.19. The transducers were calibrated at their middle range to achieve the best accuracy and to avoid the nonlinearity at the transducers' extremes. All transducers showed linearity for their full working range; an example is shown in Figure 3.20. Figure 3.21 shows the relationship between the residuals, defined as the difference between the actual and ideal output reading, normalized to the full range in percentage (accuracy) and the output voltage. The accuracy of the transducers is less than 0.1%, which translates into a percentage error in strain of less than 0.0078% for a 63.5 mm specimen and

a deformation magnitude of less than 0.005 mm based on the transducers' full scale. Although this kind of transducer (LVDT) has infinite resolution, the transducer's resolution is limited by the data acquisition system. For the selected data acquisition system, the transducer's resolution can be as low as 0.02 μm (0.072 mV for $\pm 10\text{V}$ LVDTs), which translates into 0.00003% strain for a 63.5 mm specimen.

3.6.1.2 *Stability characteristics and proof testing*

One of the central issues related to transducers is their stability for long-term performance, or the variation of the output readings with time. The stability is understood as an integrated effect of the noise in the system (e.g., the cables' type and length, the A/D converter's resolution, the extent of the amplification and the filtration in the signal conditioner) as well as the sensitivity of the transducers to the temperature. The stability was quantified for all LVDTs at room temperature in a stationary position, as shown in Figure 3.22. The temperature was monitored by using YSI 44007 thermistors, excited through the Datalogger (15 mW), with a working range of -80°C to 120°C and a resolution of 0.1°C . The thermistor was calibrated by immersing it in a container filled with ice and water while continuously stirring the thermistor in the container to identify the thermistor's zero drift, which was recorded to be in the range of 0.06°C (see Figure 3.23).

Table 3.4 summarizes the calibration factors and the nomenclatures for the whole set of LVDTs. Figure 3.24 shows the variation of the deformation as a function of time for one of the calibrated transducers. The variation of the output voltage for the whole set of transducers corresponds to a deformation of $0.15 \mu\text{m}$ - $0.8 \mu\text{m}$ (0.00023-0.00125% strain for a 63.5 mm specimen). The temperature's variation recorded in the laboratory was in the range of $0.8 \text{ }^\circ\text{C}$ for almost 24 hrs (Figure 3.25). Figure 3.26 shows a linear relationship between the variations in the deformation as a function of the temperature variation for one of the calibrated LVDTs. Table 3.4 shows that the temperature corrections factor ranges from $-0.14 \mu\text{m}/^\circ\text{C}$ (compression) to $1.15 \mu\text{m}/^\circ\text{C}$ (extension). However, as the table also shows, this range of variation was insignificant when translated into strain.

The reliability of the LVDTs in capturing deformation characteristics was assessed by comparing the calculated modulus with the known Young's modulus of a PVC cylindrical sample. The PVC sample has a Young's modulus of approximately 2.76 GPa and a compressive strength of 48 MPa. Figure 3.27 shows a typical result from the measurement system in the form of deviator stress versus axial strain. The calculated Young's modulus for the PVC sample was 2.8 GPa, with a relative standard error of 1.4% which confirms the transducers' reliability in measuring deformation.

3.6.2 Sample preparation and mechanical configuration

3.6.2.1 Sample preparation and dimensions

Block samples were cut in the laboratory by using a core barrel to reduce them to the required size of cylindrical samples. Extremes were flattened by using a blade to give the best smooth surface for testing purpose. The cutting of such samples required extreme patience due to the friable nature and brittleness of this kind of sand. Frequently, irregular end surfaces were unavoidable, and, hence, it was decided to abrade them by scarping meticulously on a horizontal metal plate while they were encased within a hollow aluminum cylinder. The preparation of samples for testing was done on a Perspex plate to reduce friction, and, hence, to prevent the loss of sample during the preparation stage. The extreme brittleness of the samples made it impossible to make efficient use of the blocks. The preparation of one sample took between 5-6 hrs. Cresswell (2001) recommended, for the preparation of Reigate Silver sand, spraying a hairspray on samples, forming a small shell around them. This shell could be removed through flushing when the samples were mounted in a triaxial cell. One of the concerns with this approach is how penetrative the hairspray is in the sample. Another concern is the effect of flushing, which can result in the loss of particles due to their friability. Wetting can also result in a reduction in the strength of this kind of sand (Richards and Barton 1999, Cresswell and Barton 2003).

Most of the tested samples in this investigation had a slenderness ratio ranging from 1:1 (63.5 mm diameter x 63.5 mm height) to 2:1. The size of tested specimens was intentionally selected to be large (compared to the regular sizes of previously tested locked sands; 36 mmx76mm) for two reasons: (a) the bigger the sample, the more representative the results; and (b) preparation disturbance effects have more adverse effects on small-sized samples than on larger ones. A total number of 12 samples were extracted from the blocks. The presented results in this chapter and the subsequent one are representative of the obtained results obtained from the experimental work.

3.6.2.2 *Mechanical configuration set-up*

A Honeywell sensotec loading cell model 41, with a load range of 19.8 kN, was used to measured the axial force. The load cell has a non-linearity of 0.2% of the full scale and a calibration factor of 0.0011 kN/mV. The load cell has a stiffness of 4.33×10^5 kN/m and was excited with an input voltage of 10 VDC. The output voltage was recorded by using the datataker and was transformed to a force unit using the calibration factor. A conventional uniaxial compression test loading fame, provided by ELE International, was used to compress the specimen between two platens. The loading frame has a capacity of 49 kN and is provided with a conventional strain-controlled hydraulic system to raise the lower platen at a constant rate of strain against a stationary upper platen that is connected to the load cell. The capacities of

the load cell and the loading frame used in this investigated were considered satisfactory for this kind of sand.

Figure 3.28 presents an illustration of the various components of the original measuring system's configuration. The adopted configuration (provided by SBEL (Structural Behaviour Engineering Laboratories Ltd.) included two loading platens with a circumferential groove into which the deformation rings were fixed by means of spring-loaded screws/pins, three LVDTs (two vertical LVDTs and a horizontal spring-loaded one), two aluminum deformation rings provided with LVDTs' mounting holes, and horizontal spring-loaded pins to fix the rings to the platens' grooves, two non-magnetic threaded connecting rods containing a screw nut in the upper extreme (attached to the upper ring) to lock the vertical LVDT into the core position with respect to the core body attached to the lower ring holes, a light aluminum chain connected at each extreme with two blocks attached to each other by using a guide bar (linear bearing), and an O-ring connecting two steel posts. The horizontal spring-loaded LVDT was attached to one block and fixed into position by using a vertical spring loaded pin and facing a screw attached to the other block for adjusting the initial reading.

The original mechanical configuration was later modified to facilitate the set-up procedure as well as to obtain replicate measurements for comparison. In particular, with the original system it was not possible to ensure either the

verticality of vertical LVDTs or the parallelism of upper and lower deformation rings. The inability to ensure the verticality of LVDTs or the parallelism of the deformation rings can result in a stick-slip problem, and, hence, the development of friction between the cores and the LVDT's core bodies.

For these reasons, the existing system was modified. The system was augmented with 3 additional LVDTs (two for lateral deformation and one for vertical deformation) besides the two existing vertical LVDTs and the chain's spring-loaded LVDT. The upper and lower deformation rings were re-fabricated to carry the three vertical LVDTs as well as two vertical aluminum alignment posts. For each of the horizontal LVDTs (H-LVDTs), the core body was fixed into position by using an extended mobile arm connected to a vertical post fixed on the loading frame. The LVDT core was attached to a short threaded rod that extended horizontally to a small puncture created on the sample's surface at one end. The other end carrying the LVDT core was adjusted to be at the center of the LVDT core body to record the relative movement between the core and the fixed core body. Figure 3.29 presents an illustration of the mounting procedure and the final configuration. The mounting procedure was carried out in the following steps:

1. Step I included placing the lower platen with the lower deformation ring attached to it as well as the vertical LVDTs' core bodies fixed into position by using the spring-loaded pins. Afterwards, the chain was attached to the

sample's circumference at the sample's mid-height and the blocks (carrying the horizontal spring-loaded LVDT) were connected by using an O-ring. Next, the sample was placed on the lower platen. Direct interference between the platen and the sample's end was alleviated by using a thick layer of silicone grease sandwiched between two latex membranes (Figure 3.29a).

2. Step II consisted of placing the upper platen with the upper deformation ring on the specimen and locking the system to the frame by applying a minimal sitting load. The system locking was necessary to prevent sample movement during the adjustment of the LVDTs' verticality. Once the system had been locked, the two vertical alignment bars were connected to the upper and lower deformation rings and fixed into position by using the spring-loaded pins (machined in the rings' sides) (Figure 3.29b).
3. Step III consisted of placing the connecting rods with the cores attached to the lower end, tightening the upper end to the upper deformation ring, and removing the adjustment bars. This stage ensured the verticality of the connecting rods to prevent the potential friction between the cores and the cores' bodies (Figure 3.29c).
4. Step IV included the mounting of the two horizontal LVDTs to complete the set-up. The horizontal LVDTs' bodies were fixed into position by using two fixed arms extended from the vertical posts on the loading frame. Once fixed, a short rod with the core attached was inserted into the core body by meticulously rotating the free end into the sample. Once a

puncture had been created in the sample, the LVDTs' core bodies were again adjusted to prevent friction (Figure 3.29d).

Samples were then sheared at a constant strain rate of 0.05% and data were logged every 10 seconds to capture both the pre- and post-peak response. The system was also augmented with an unguided free armature external LVDT to compare the local and external axial strain measurements.

3.7 Significance of local strain measurements

In this section, the significance of local strain measurements is described by comparing the local and external axial strains as well as the strain rates. The tested specimen had a slenderness ratio of 1:1 and sheared at a constant strain rate of 0.05%. The axial strain was averaged based on the reading of the three vertical LVDTs for comparison with the externally measured axial strain.

3.7.1 Local vs. external stress-axial strain curves

Figure 3.30 shows the deviator stress ($q = \sigma_1 - \sigma_3$; $\sigma_3 = 0$)-axial strain (ϵ_1) relationship based on local and external axial strain measurements. In general, the specimen attained peak strength of 2.0 MPa at axial strains of 0.32% and 0.57% according to the local and external measurements, respectively. Both strain systems showed an extremely brittle post-peak response. According to ISRM (1978b), the specimen tested is classified as a very low strength rock (weak rock), which lies close to the soil boundaries (the

soil boundary is characterized by a uniaxial compressive strength < 0.25 MPa) and constitutes the boundary between soils and hard rocks. The source of the high uniaxial strength is presumably due to the relatively extensive area of contacts as well as the increase in the coordination number, or the number of contacts per particle. That is, particles interact as surfaces rather than point contacts (as in dense sands), and, hence, the rotational resistance is enhanced. As a consequence, in the absence of bonding, individual contacts can sustain considerable contact forces, moments, and stress obliquity due to interlocking (Creswell and Powrie 2004).

The external measurements had a tendency to overestimate the axial strain at any stage during the shear evolution and particularly in the pre-peak regime, as has been found with other geomaterials (e.g., Ibrahim and Di Benedetto 2005; Yimsiri et al. 2005). Only at very large axial strains did the internal and external measurements converge together. In the pre-peak regime, both stress-strain curves were initially concave followed by a linear region. The size of the initial concave portion of the stress-strain curve was significantly overestimated by the external strain measurement as compared to the local one. Within this zone, the nonlinear behaviour suggested the closure of the existing fractures and the soft inclusions identified in the microstructural studies. In the linear portion of the stress-strain curves where the elastic zone is traditionally identified, the axial stiffness based on the local strain measurement was approximately 0.87 GPa and almost 2 times the

externally measured axial stiffness. In the post-peak regime, both measuring systems indicated strain weakening; however, the local one exhibited a slightly less abrupt reduction in strength.

3.7.2 Local and external axial strain rates

Figure 3.31 shows the variation of the instantaneous local and external strain rates ($d\varepsilon/dt$) with the axial strain as well as the local axial stress - axial strain curve. Despite assigning a controlled constant rate of strain (0.05%) at the start of the test, the external and local strain rates deviated from that assigned value. The external values were higher than the assigned value at the initial stages due to the beddings effect (the existence of irregular surfaces at sample ends) and coincided with the assigned values upon closure of the fractures throughout shearing evolution. On the other hand, the local strain rate value was recorded to have smaller values, up to 40% of the average value, before the closure of the fractures. Upon attaining an intimacy between the grains and full contact, a linear increase in the axial strain rate occurred as a reflection of the initiation of the unlocking mechanism, followed by a sudden increase around peak strength (around a 64% increase in the axial strain rate). This sudden increase probably signified the unlocking of most sample elements and the initiation of localization in the post-peak regime. In this regime, the strain rate oscillated above the specified value, reflecting, on average, an increase in the strain rate and non-uniformity of the strain rate as an indication of progressive rupture of the interlocked fabric. Consequently,

local strain rates are indicative of the failure process at all stages of shearing as compared to the constant values estimated from the external measurements. Extreme caution is necessary when assigning strain rates for the sake of comparing geomechanical responses or when the creep rates are of special interest. Hayano et al. (2001) made a similar recommendation in their study of soft rock, in which the creep axial strains measured locally were noticeably less than the externally measured ones.

In the present study, the interlocking between particles was found to be the main source of such high strength in the absence of virtual bonding between sand particles. For proper characterization of such geomaterials, experimental results' interpretations founded on external measurements are in error, particularly when coupling strength and deformation is targeted. Moreover, local axial strain measurements provide more insight than external measurements into failure mechanisms.

3.8 Summary and conclusions

This chapter provided a synopsis of the common characteristics of locked sands to create a framework to identify such deposits. Based on this framework, experimental work was presented to identify the fit of the High-Hill oil-free sand formation into the locked sands category. The experimental work was comprised of micro-structural observations, mineralogical composition analyses, and fabric analysis. A uniaxial compressive strength test was also carried out to reveal the

strength of this formation's unbonded coherent mass and to explore the significance of the local strain measurements. The experimental set-up as well as the calibrations, proof testing, and stability tests were presented and discussed in detail.

Based on the investigation, the High Hill river sand formation is a lower Cretaceous McMurray Formation. The dense nature of the samples, their mineralogical composition, and micro-structural observations are in keeping with characteristics of locked sands. The formation is best described as overconsolidated, well-sorted, unbonded, friable, locked sand with characteristics similar to those of oil-sands' middle members. This deposit was subjected to mild post-depositional diagenesis rendering transformation of its detrital sand-sized particles into a weak rock with interlocked fabric through deposition of authigenic silica. SEM images and petrographic analysis of thin sections disclosed that grains' contacts were dominated by tangential, long-to straight, occasionally interpenetrative and sutured contacts created via overgrowth and dissolution. The minor percentage of kaolinite in the form of coating or bridges is responsible for the friable nature of this deposit. The fine particles presumably originated from either the diagenesis of the overlying shale formation or the feldspar inherited from parent rocks. The deposit appears to have been subjected to overburden loads that were substantial enough to cause this fracturing, and, hence, to influence the deposit's structure. The clean cracks within quartz grains imply that clay minerals were deposited following the compaction stage or its weak nature

to prevent loss of porosity due to compaction by early cementation. The existence of a fresh skeletal grain by dissolution suggests the process of dissolution occurred after the compaction stage (otherwise, the compaction process would have destroyed this skeleton) and that the dissolution is still in process.

A local measurement system for both axial and lateral strain was used and modified for this friable material in conjunction with a routine external axial measurement system for proper characterization. The formation exhibits a uniaxial compressive strength of 2 MPa and hence may be classified as a weak rock. The interlocking between particles is the main source of such high uniaxial strength and axial stiffness in the absence of bonding between sand particles. The axial stiffness based on local strain measurement is approximately 0.87 GPa and almost 2 times the externally measured axial stiffness. In the post-peak regime, both measuring systems indicate strain weakening; however, the local one exhibits a slightly less abrupt reduction in strength. In terms of the failure mechanism and strain rate effects, the variation in local axial strain rate provides more insight than externally measured axial strain rate into failure evolution stages. This discrepancy is most critical when the strain rate effects are a concern or when the creep behaviour is targeted. Hence, for proper characterization of such geomaterials and their oil-sand counterparts, interpretations founded on commonplace external strain measurements are less reliable, particularly when the coupling strength and deformation is targeted.

Local strain measurements are necessary for proper behavioral characterization and failure mechanism interpretation of locked sands.

Table 3.1 Summary of physical and mechanical characteristics of locked sands.

Authors	Formation	Porosity (%)	RD (%)	Density (kN/m ³) and/or D50(mm)	Peak/secant friction angle, normal stress and strain at failure (direct shear test- dry else stated)	Strength parameters, confining pressure and strain at failure(Triaxial tests)	Uniaxial compressive strength and youngs modulus	Contact Nature/ or TI(%).
Dusseault and Morgenstern (1979)	Athabasca oil sand (Lower Cretaceous)	31.3	>100-135	18.5-23.3 D ₅₀ =0.2mm	44°-46° (100-800kPa) 1.5-2.5%	$\phi_{sec} = 50^\circ$		Mostly concavo/convex contacts with tangential contacts
Dusseault and Morgenstern (1978)					$\phi_{sec} = 39.9^\circ$ - 66.4°* (98-785kPa)			
Barton et al. (1986)	Barton sand (Eocene)	35.6	119	17	40.6°-47° (50-900kPa) 1.9-5.5%	C= 0 kPa $\phi = 40^\circ$ (100-700kPa) 2.9-7.9%	q _u 5-134 kPa	Tangential contacts- slight concavo/convex contacts (TI=54%)
Barton et al. (1986)	Bagshot sand(Eocene)	36.4	110	16.8	44.2°-50.3° (50-900kPa) 2.5-7%	C= 0-77.4 kPa $\phi = 41.5^\circ$ -50° (100-700kPa) 0.56-2.8%	q _u 20 kPa	Tangential and straight contacts- slight concavo/convex contacts (TI=41%)
Dusseault and Morgenstern (1979)	Swan River sandstone (Lower Cretaceous)	34.5	125		63°-70° (100-800kPa) 1.0-2.5%	-----	-----	Mostly concavo/convex contacts
Dusseault and Morgenstern (1979)	St. Peter friable sandstone (Ordovician)	27	135	19 kN/m ³	63°-69° (100-800kPa) 0.5-1.75%	$\phi = 57^\circ$ -69° $\phi_{field} = 56^\circ$	q _u =0.6-2 MPa E=1GPa v=0.2-0.33	Mostly concavo/convex contacts
Dittes and Labuz (2002)		27			$\phi = 57^\circ$ -63° $\psi = 45^\circ$ (19kPa)- 21° (152 kPa) $\phi_{sat} = 46^\circ$			
Dusseault (1980)	Itacolumites (flexible sandstone)				Much stronger than locked sands with no cementation due to intense diagenesis			High concavo/convex contacts due to quartz overgrowth
Richards and Barton (1999)	Folkstone bed sands							
	Uniform (including Reigate silver sand) Thin bedded(iron oxide)	28-35 25.7-33.6		D50 = 0.2-0.3mm	C= 7.7-179.4 kPa $\phi_{sec} = 34^\circ$ -65.5° C= 20.5-65.4 kPa $\phi_{sec} = 34^\circ$ -65.5°^A (0.15-0.8mm for zero normal stress) (0.6-2mm for normal stresses up to 674kPa)		q _u =101.9-725.2 kPa q _u =85.5-218 kPa	TI=29.7-65.4% TI= 27.3-66%
Cresswell and Powrie (2004)	Reigate silver sand	32	136 %	18.62		(50-800kPa) 0.4%-0.8% M=2.25 (p=0-1000 kPa)	q _u =108-101.9 kPa	TI=26%
Creswell and Barton (2003)						$\phi = 70^\circ$ (low stress and 40° and 10 MPa) (0.25mm at 1.5 kPa-2.5mm at 10Mpa)		

* function of bedding plane orientations

^ depends on the orientation of bedding. General trends that cohesion tends to decrease for parallel beddings and friction increases for normal beddings.

* TI (Tangential index defined as the proportion of tangential contacts amongst all grain/grain contacts)

Table 3.2 Summary of Depositional environment as well as lithology of McMurray Formation members.

McMurray Formation stratigraphy members	Depositional Environment	Lithology of Athabasca Oil sands " McMurray formation sand members"*
Upper	Low energy shoal/ Shore-face (coastal plain sediments)	-Fine-grained quartzose sands to clayey silts (Group III). - doesn't belong to the orthoquartzite family due to existence of significant percentage of clay minerals and non-quartz clastics.
Middle	Marine influenced estuarine channel complex	-Fine to medium grained quartzose sands. - Well sorted sand and largely fall in Group II.
Lower	Braided to meandering fluvial sand channels (paleovalley- fills)	- medium grained sand and pebble conglomerates fall within Group I

Table 3.3 Summary of main physical properties of High Hill river formation.

Main Physical properties
<p>Grain size analysis (ASTM D 422): -D₁₀= 0.075mm, D₃₀= 0.132mm, D₆₀=0.19mm - Coefficient of curvature = 1.22, Coefficient of uniformity= 2.53. - Percentage of fines < 10%. - Poorly graded fine to medium sand</p> <p>Average porosity and unit weight: Average porosity=34% Average unit weight= 17.6 kN/m³</p> <p>Relative density (ASTM D 4254, D 2847): Average maximum void ratio = 0.86, average minimum void ratio = 0.61, average in-situ void ratio= 0.54 Average relative density (RD%)= 125%</p>

Table 3.4 Calibration factors for transducers.

Transducer	Measurement	Calibration factor (mm/mV)	Temp. calibration factor ($\mu\text{m}/^\circ\text{C}$)	Change in strain/1°C reduction for a 2.5" sample (%)
LVDT V1*	Vertical strain	0.000245	0.59	0.0009
LVDT V2*	Vertical strain	0.000191	0.3	0.0005
LVDT 3AV	Vertical strain	0.000828	-0.14	-0.0002
LVDT 3HCHAIN*	Circumferential strain	0.00020	0.24	0.0004
LVDT 1AH	Horizontal strain	0.000828	1.15	0.0018
LVDT 2AH	Horizontal strain	0.000938	1.04	0.0016

* Original set of transducers.

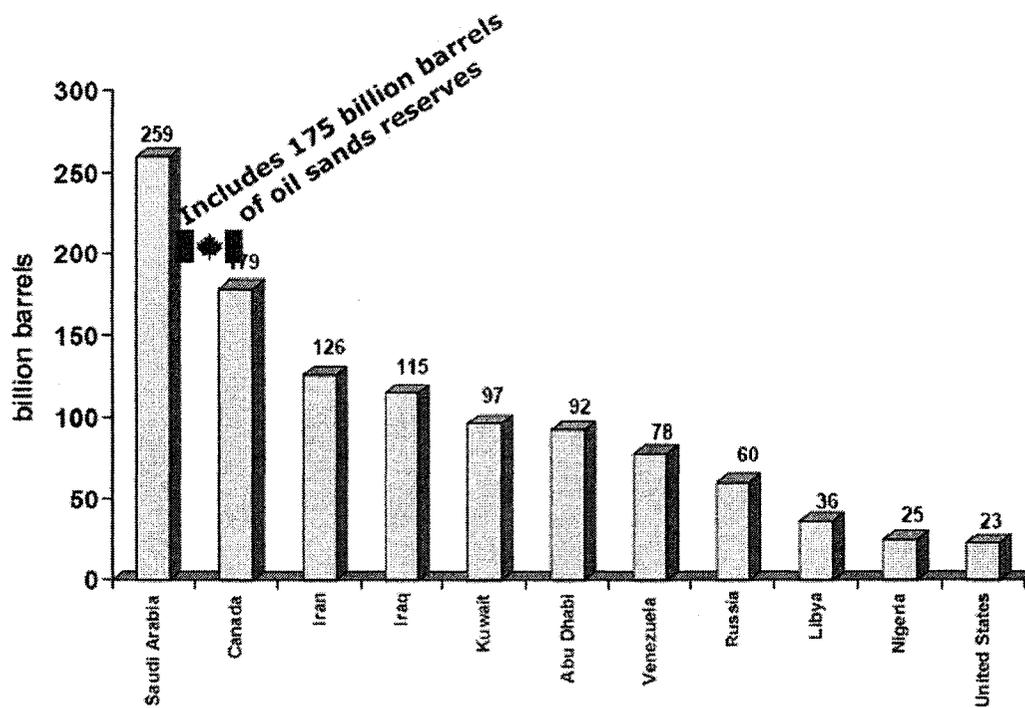


Figure 3.1 World 's Crude oil reserves
 (<http://www.syncrude.ca/users/folder.asp?FolderID=5987>)

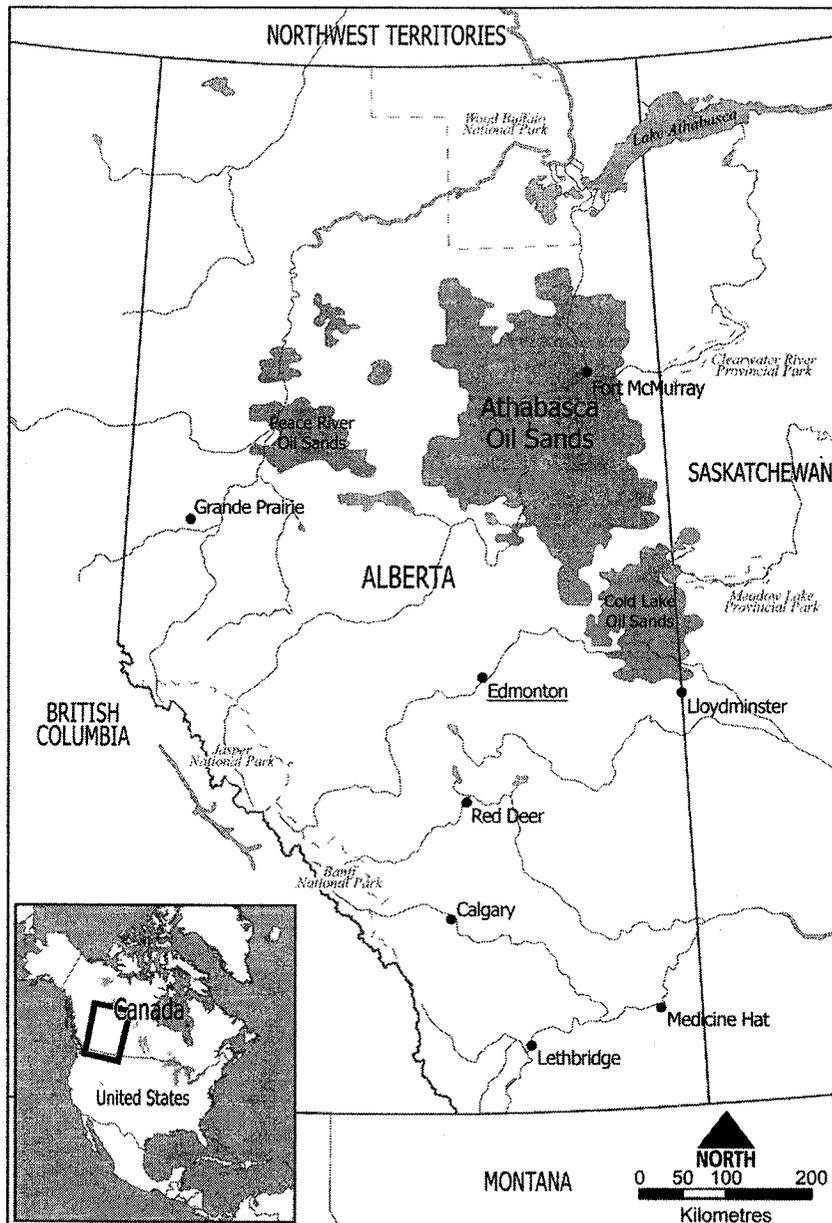


Figure 3.2 Distribution of Alberta Oil sands
 (http://upload.wikimedia.org/wikipedia/commons/7/7a/Athabasca_Oil_Sands_map.png).

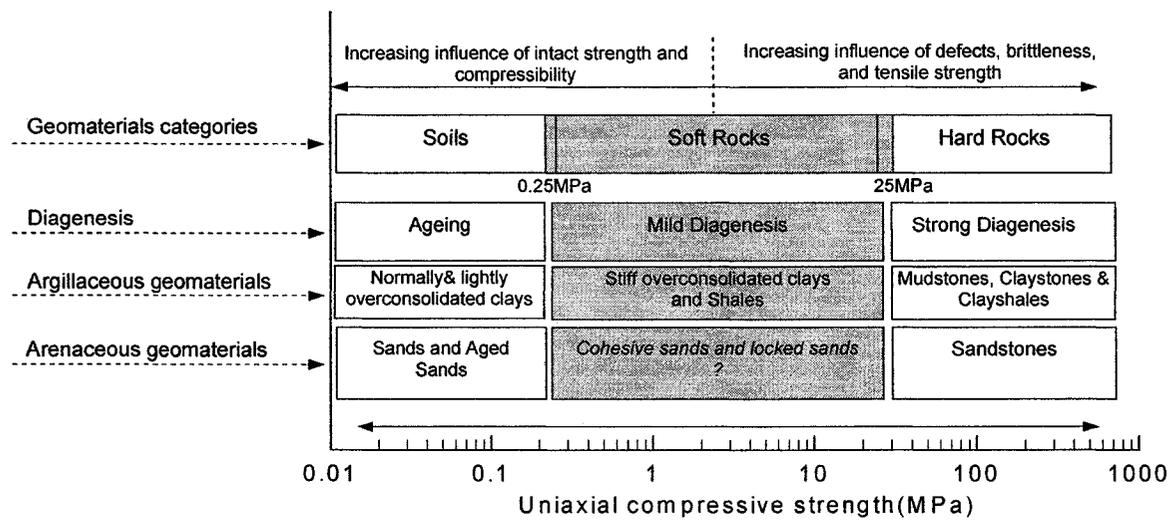


Figure 3.3 Geomaterials categories according to Diagenetic processes (modified from Barton 1993 and according to ISRM 1978).

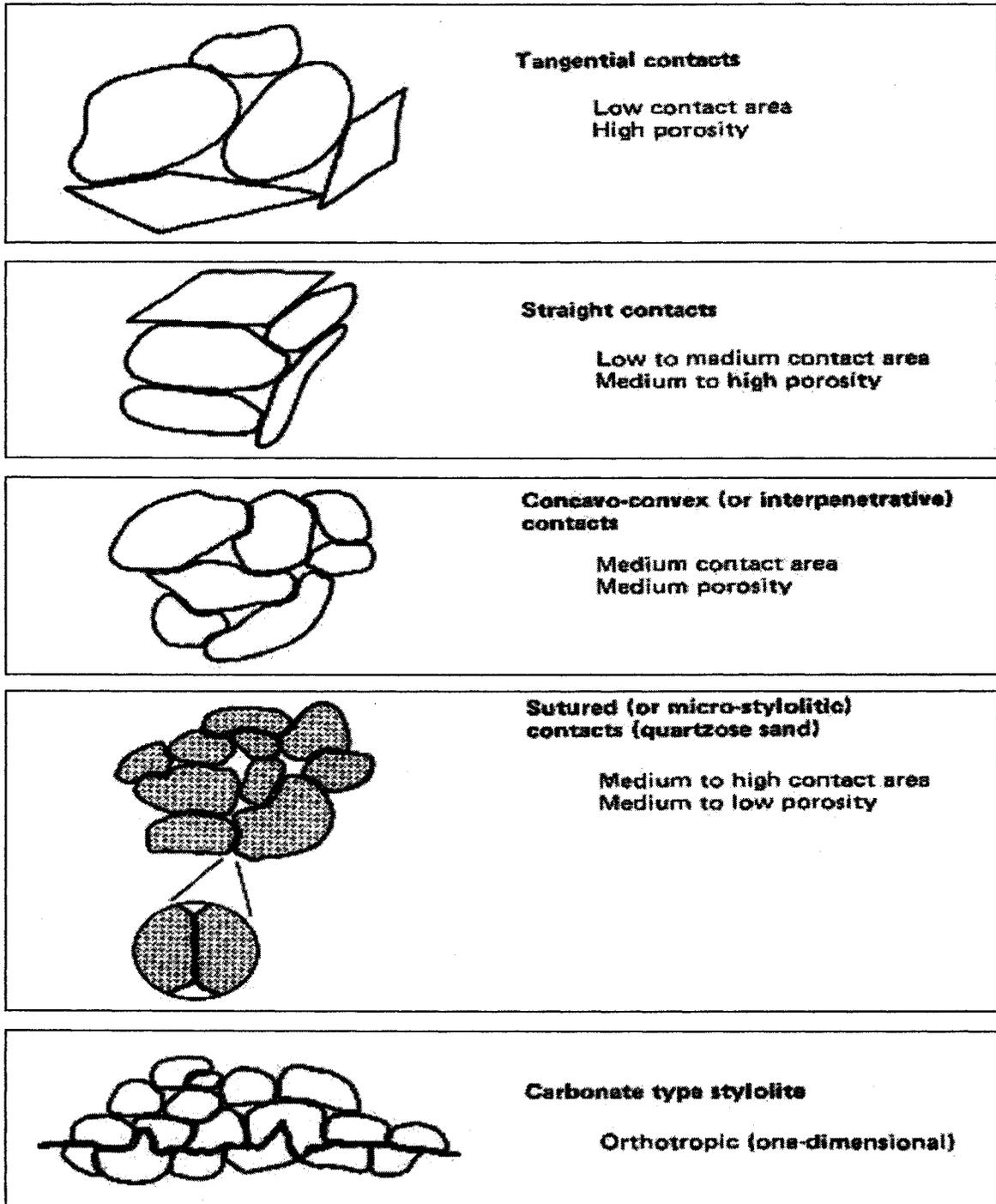


Figure 3.4 Classification of contacts in sandstones (modified from Siever 1959).

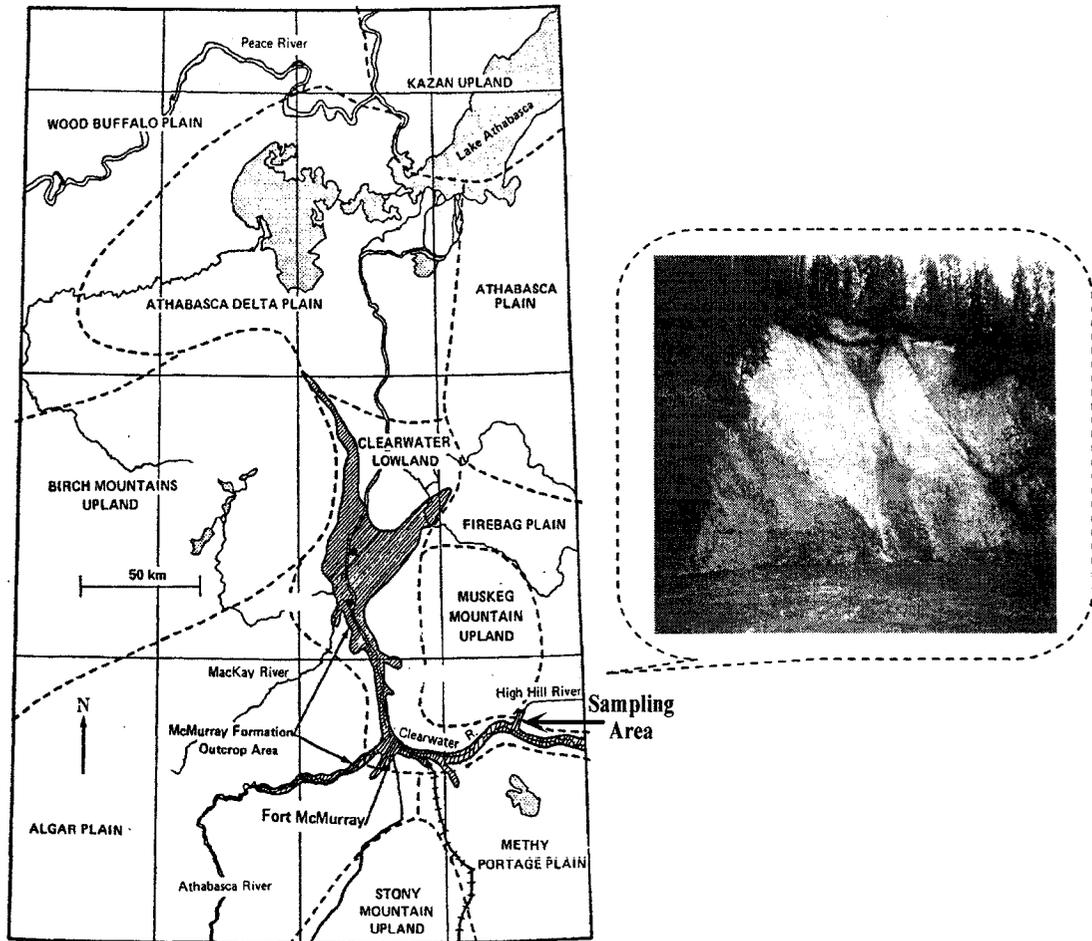
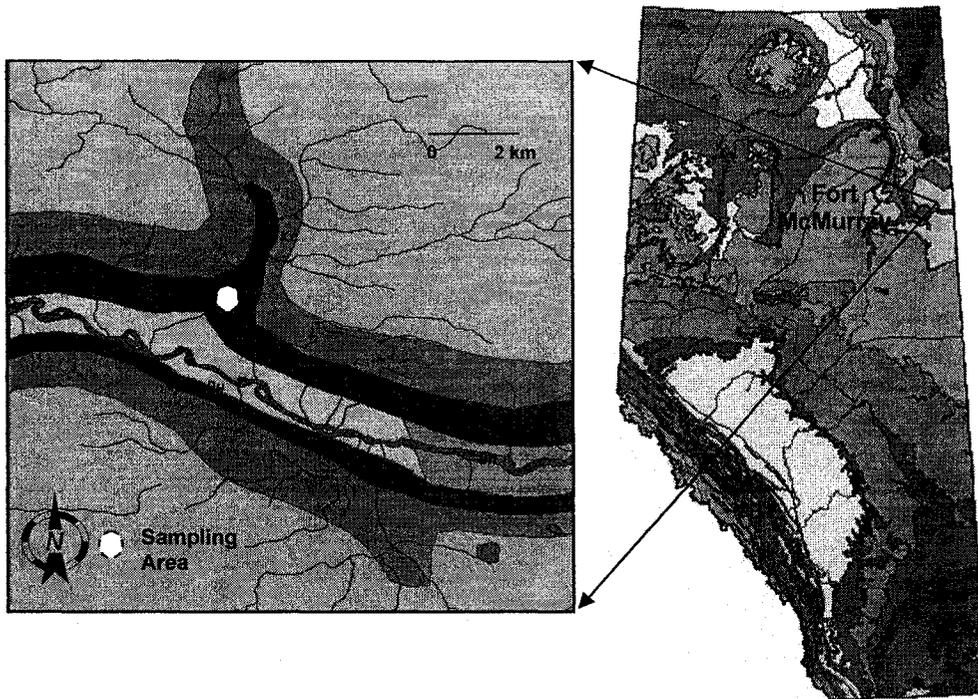


Figure 3.5 Physiography of the Athabasca oil sands and location of sampling area (modified from Dusseault 1977).



Color Code for investigated vicinity	Bedrock Geology
(Kpl)	LOWER CRETACEOUS PELICAN FORMATION: fine-grained quartzose sandstone, silty and glauconitic in lower part; marine
(Km)	LOWER CRETACEOUS MCMUARRY FORMATION (~130 Ma; Million years ago): Thick-bedded quartzose sandstone and siltstone; oil- impregnated, grey silty shale interbeds in upper part; non-marine to detalic
(Kc)	LOWER CRETACEOUS CLEARWATER FORMATION: dark grey, fossiliferous, silty shale, laminated siltstone and fine-grained cherty sandstone; glauconitic sandstone (Wabiskaw member) near base; marine.
(Dw)	UPPER DEVONION WATERWAYS FORMATION (pre-cretaceous erosion surface): grey and greenish grey shale and argillaceous limestone interbedded with grey and grayish brown, fine-grained to coarse clastic limestone; marine.
(DM)	MIDDLE DEVONION (undivided): arkosic sandstone (La Loche Formation); grey-green silty and sandy shale, reddish shale; dolomite; minor anhydrite (Mc Lean River Formation); brown and buff, massive, porous dolomite; brown to grey thin-bedded dolomite; dolomite limestone (Methy formation); gypsum; anhydrite; grey-green silty and dolomitic shale; minor dolomite (leached Prairie Evaporite Formation); marine and evaporitic

Figure 3.6 Bedrock geology in the investigated vicinity (AEUB and AGS 2005).

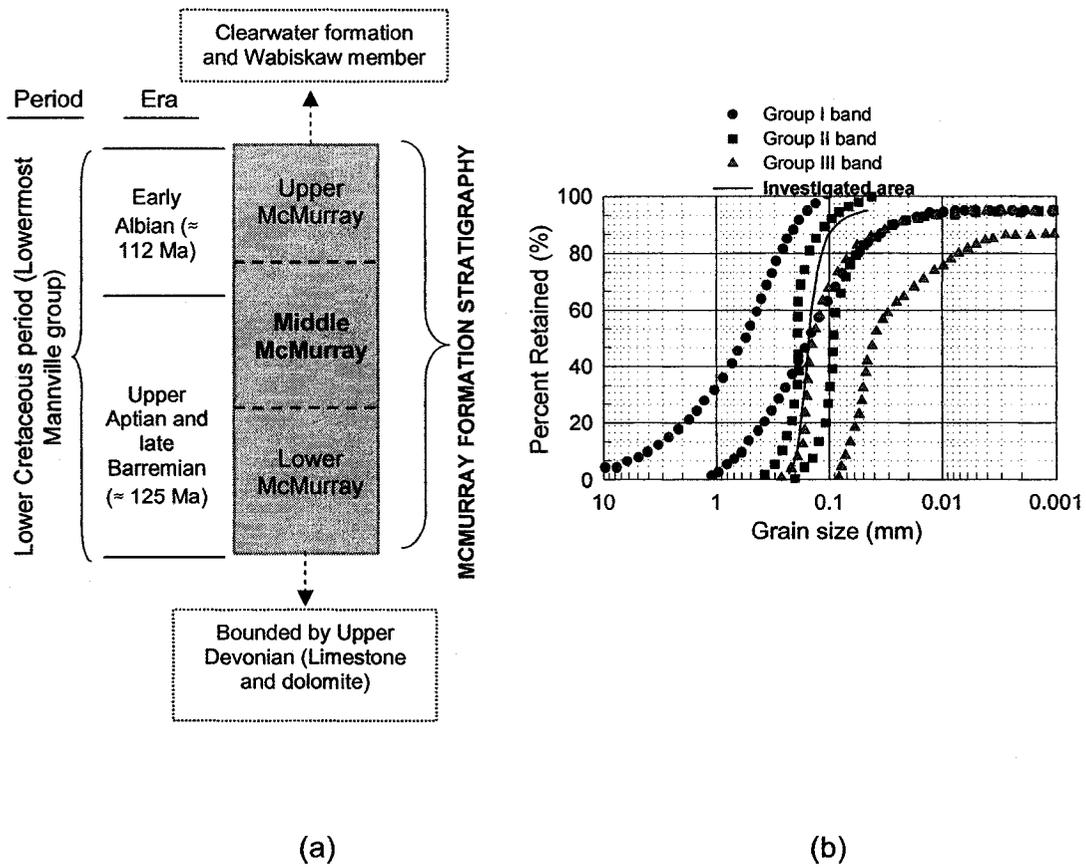


Figure 3.7 (a) An illustration of tripartite classification of McMurray formation members; and (b) Grain size distribution of Athabasca oil sands (modified from Carrigy 1967).

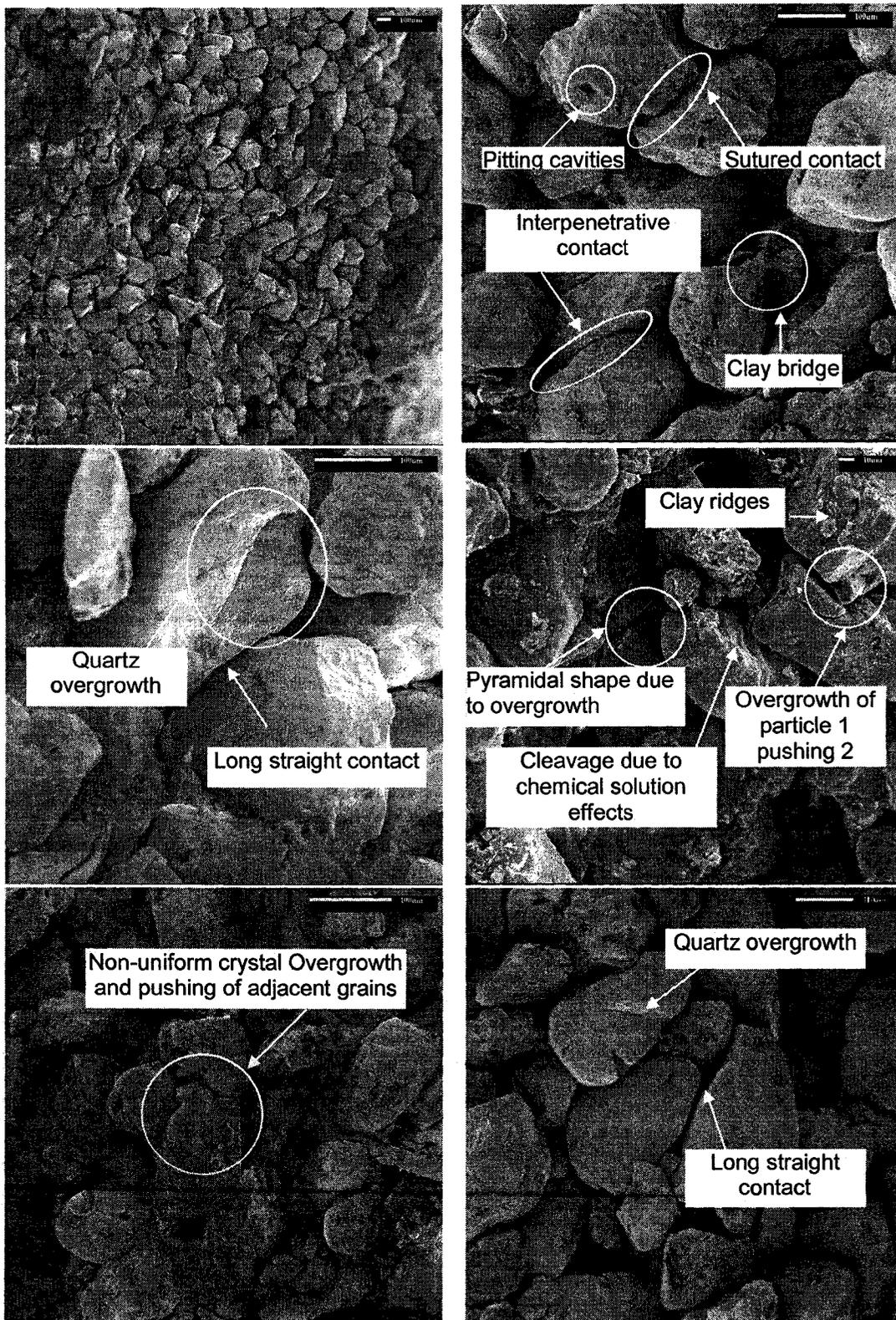


Figure 3.8 Selected SEM micrographs of High-Hill sand formation.

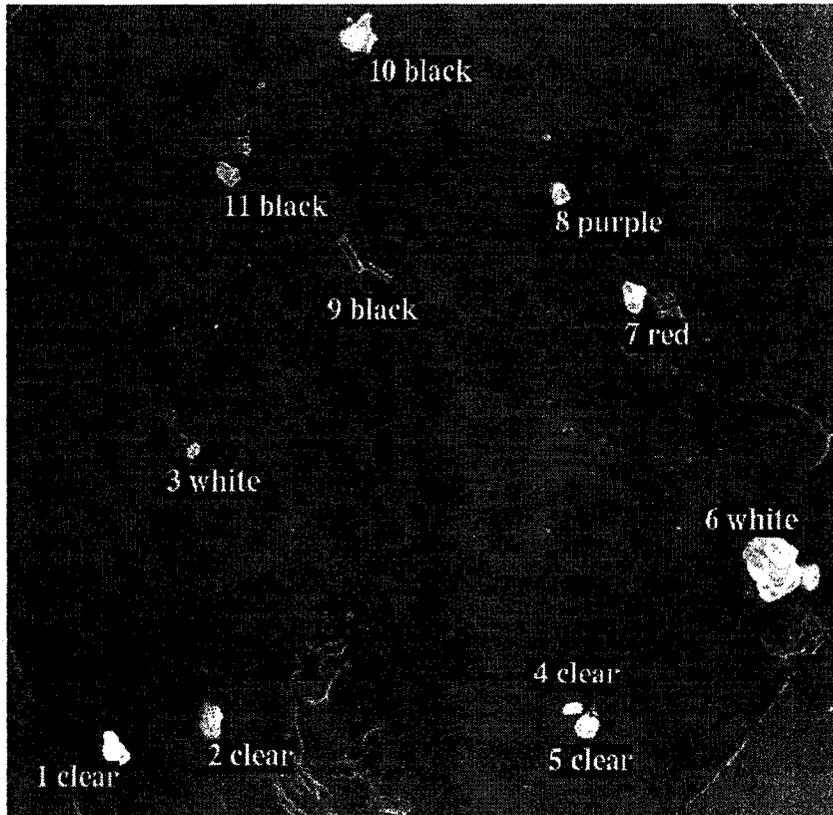
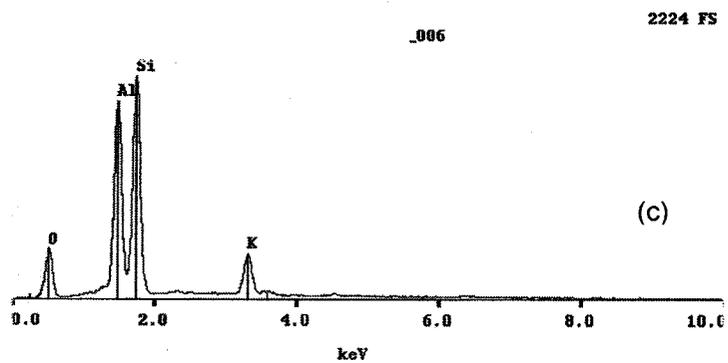
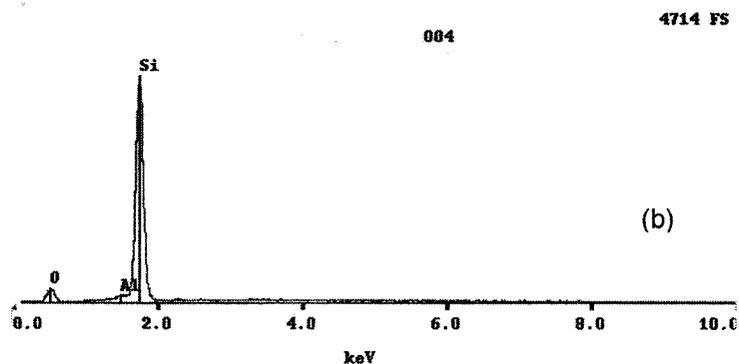
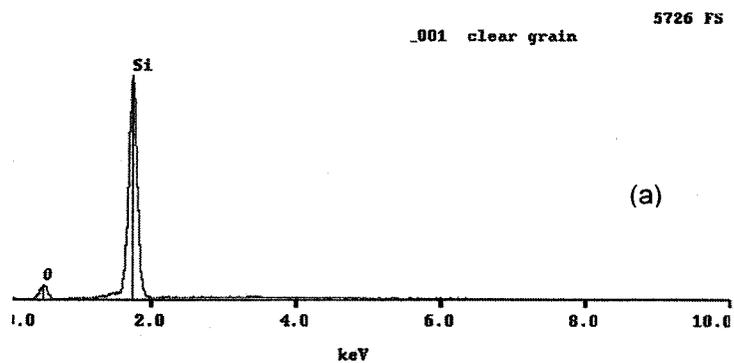


Figure 3.9 Map for collected particles under optical microscope.



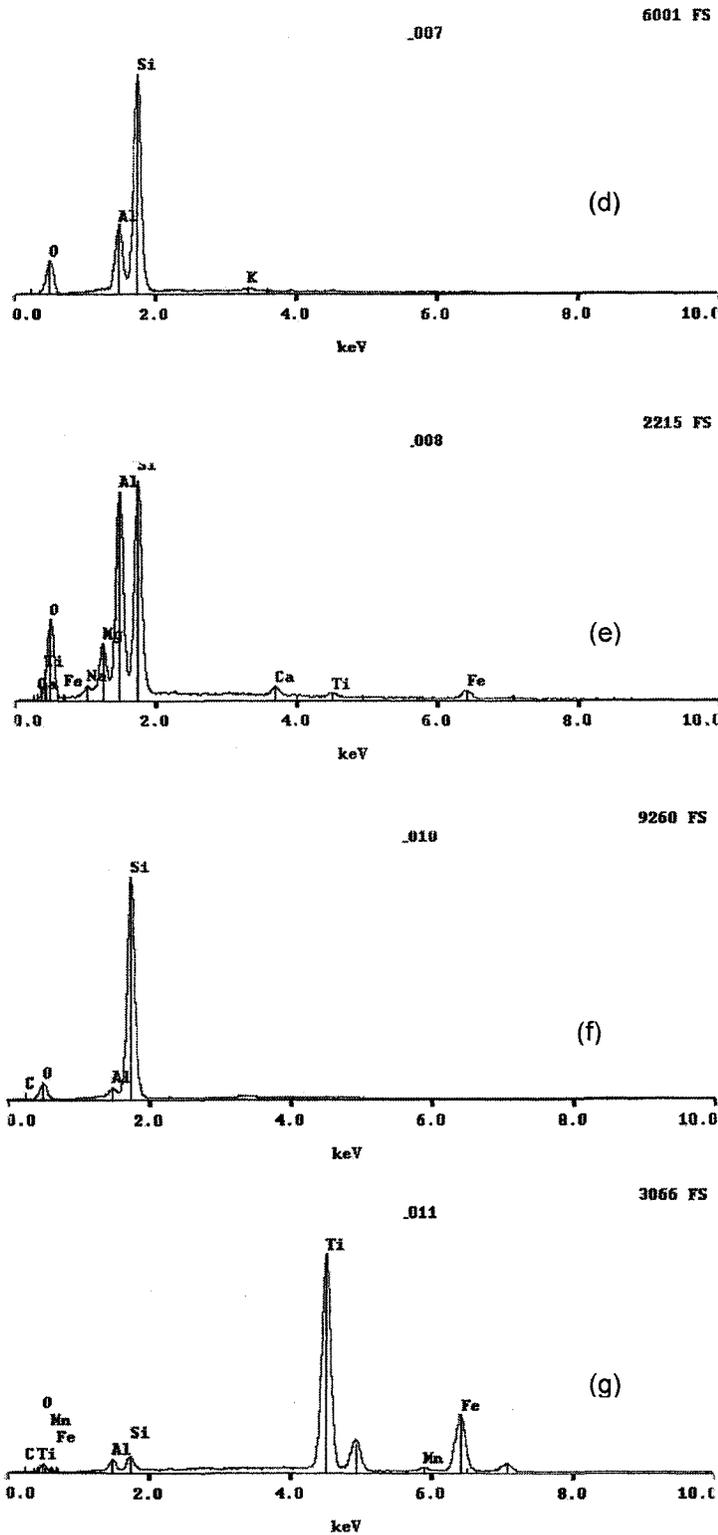


Figure 3.10 Energy dispersion X-ray results for mapped particles: (a-b) quartz; (c): kaolinite; (d) feldspar; (e) amphibole; (f) chert; and (g): titanium



Figure 3.11 Micrograph of well -sorted band under PPL: width of view 3.5 mm, illustrating quartz grains as white, very fine grained rock fragments, and clay as grey and pores in blue.

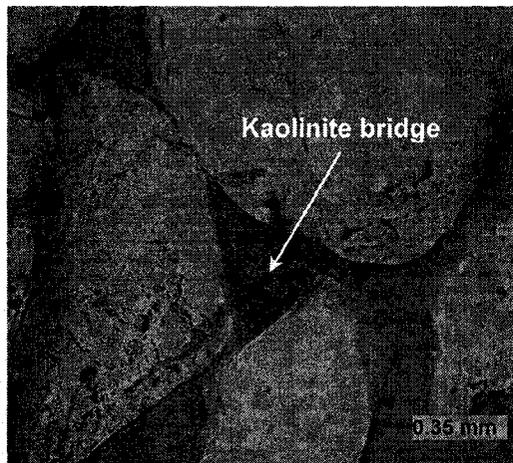


Figure 3.12 Micrograph under PPL showing kaolinite bridge between three quartz particles: width of view 0.35 mm.

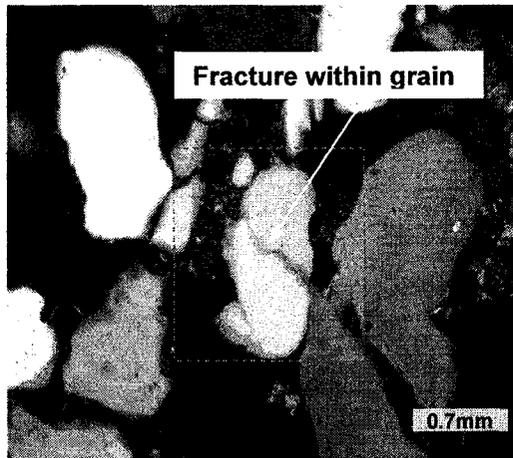


Figure 3.13 Micrograph under CPL showing particle fracture: width of view 0.7 mm.

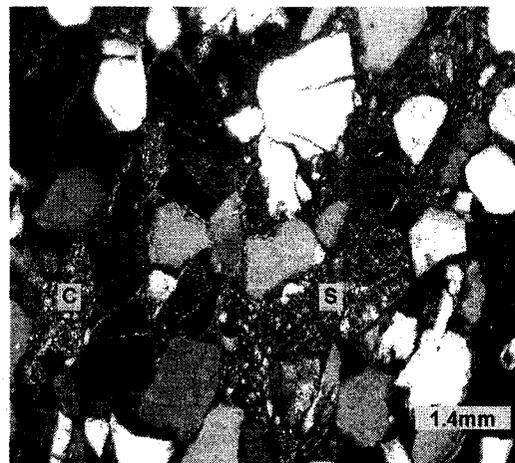


Figure 3.14 Micrograph of less-sorted band under CPL: width of view 1.4 mm, illustrating quartz grains as white and light to dark grey; very fine grained rock fragments as black with white spots, siliceous siltstone (S), chert (C); and pores as black.

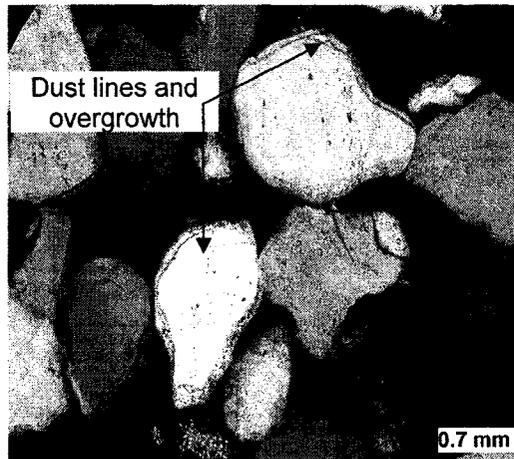


Figure 3.15 Micrograph under CPL: width of view 0.7 mm, illustrating quartz overgrowth, quartz grains as white and light to dark grey, and pores as black.

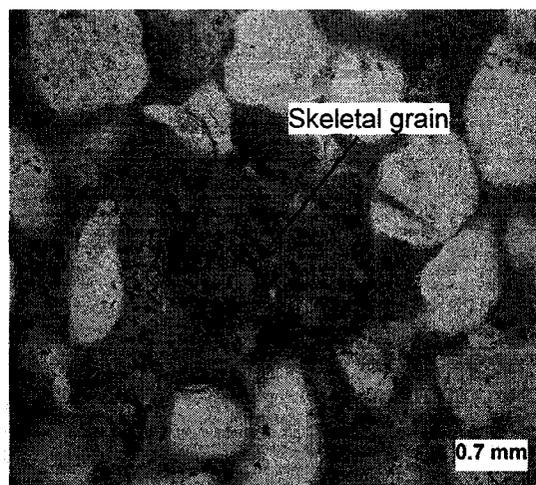
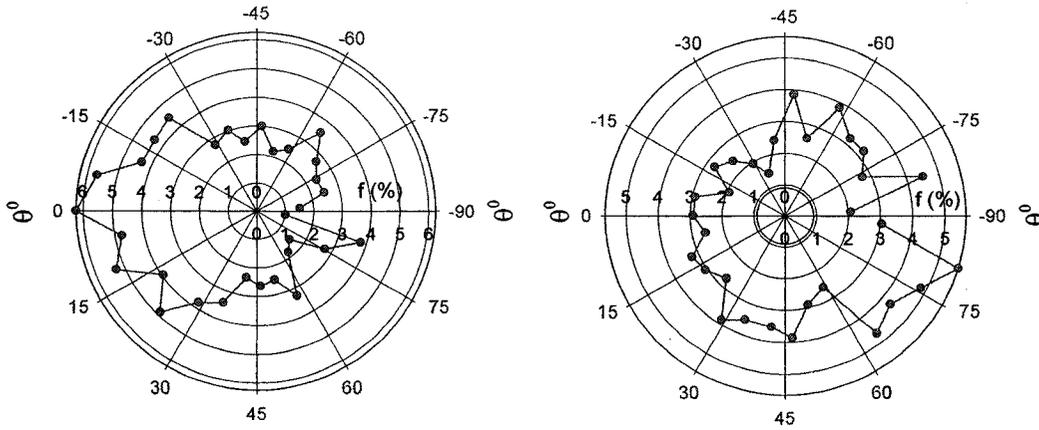


Figure 3.16 Micrograph under PPL: width of view 0.7 mm, illustrating existence of a skeletal grain due to dissolution during diagenesis.



(a) Vertical section

(b) Horizontal section

Figure 3.17 Distribution of particles orientation in horizontal and vertical direction: (a) Horizontal section; and (b) Vertical section.

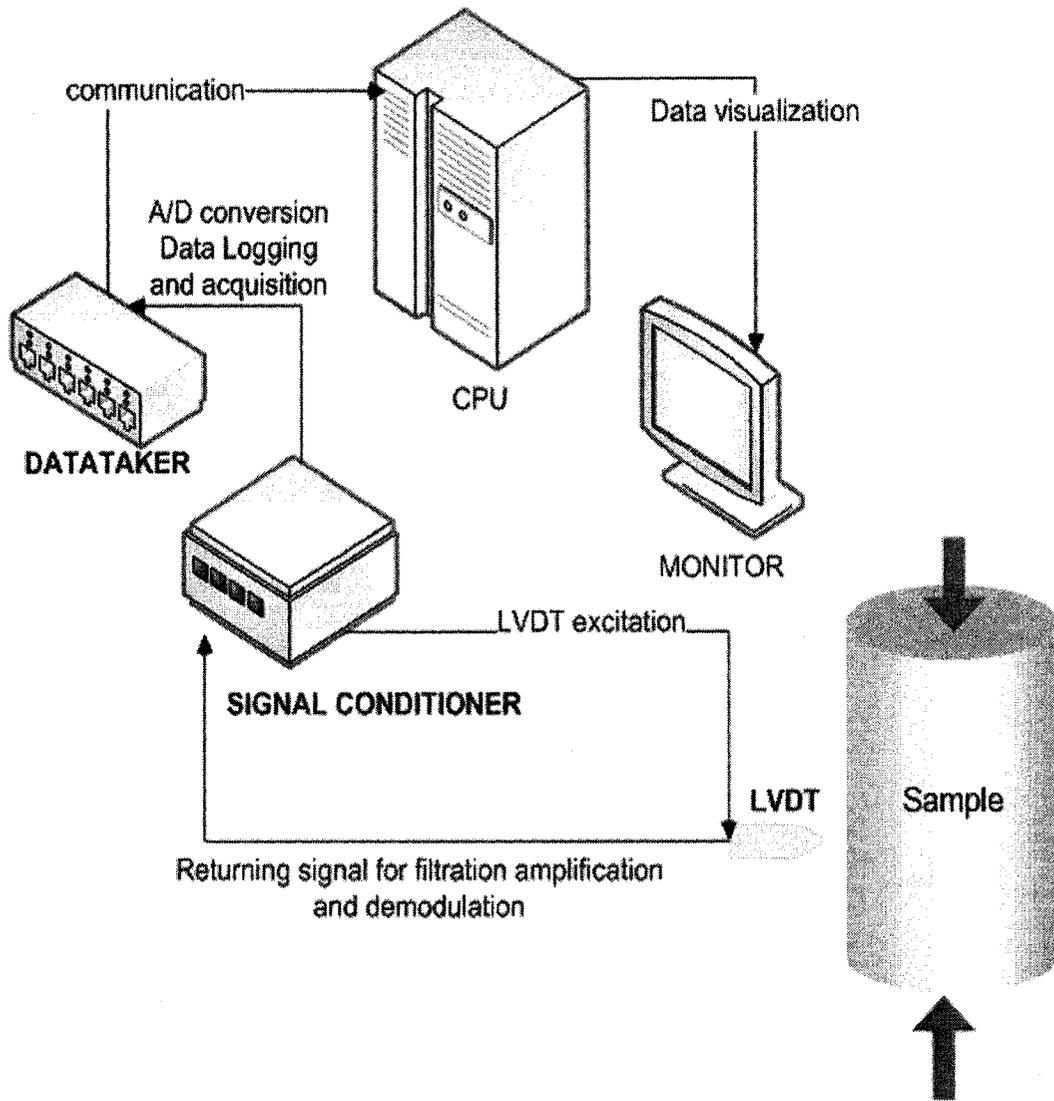


Figure 3.18 Signals cycle and laboratory components.

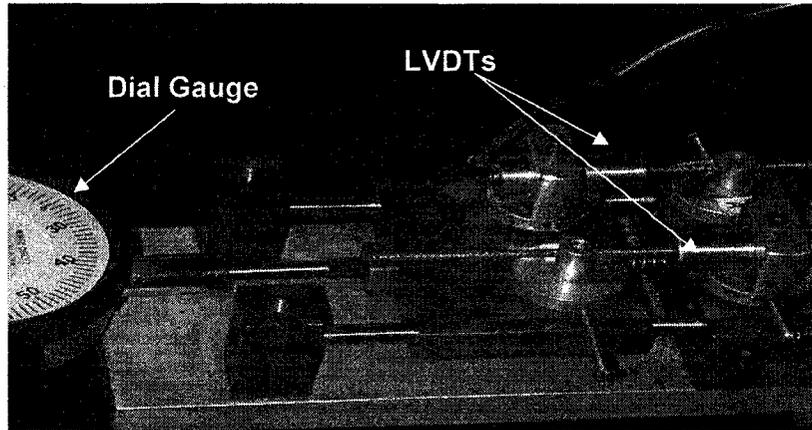


Figure 3.19 Micrometer for LVDTs calibration.

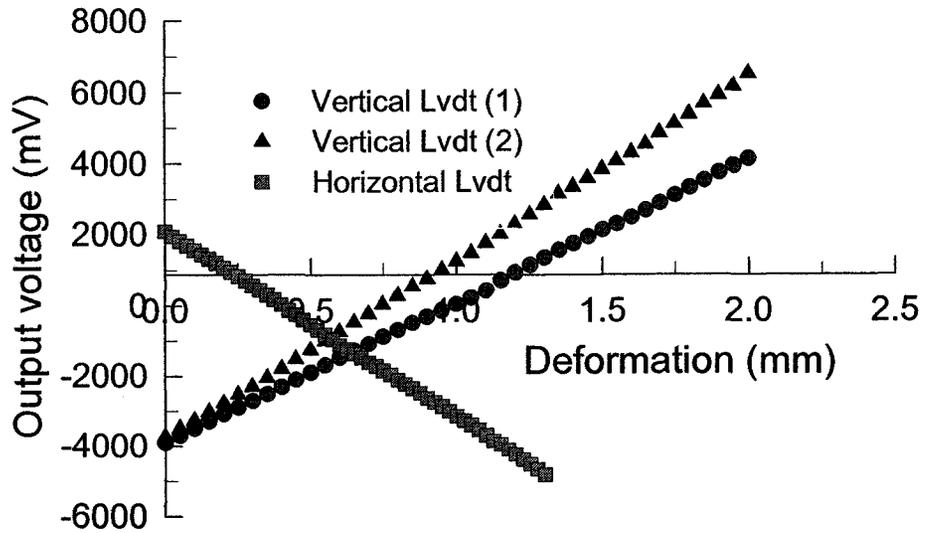


Figure 3.20 A sample of LVDTs calibration charts.

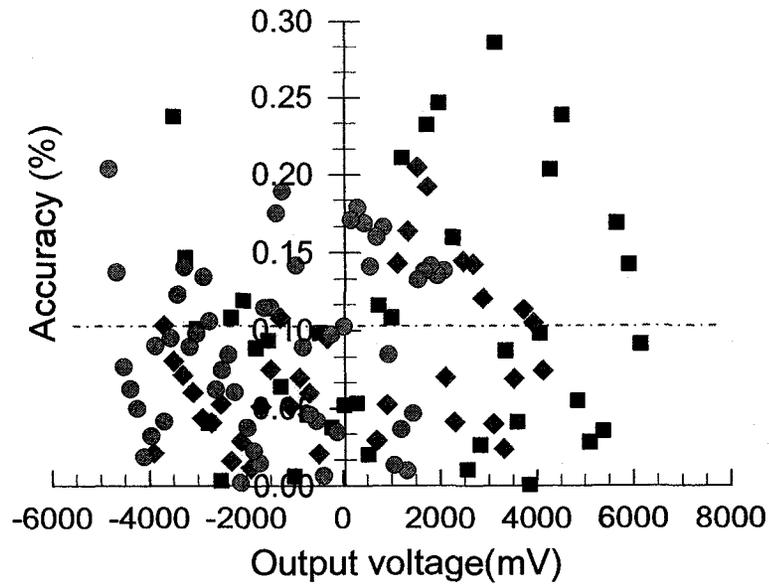


Figure 3.21 Accuracy versus Output voltage.

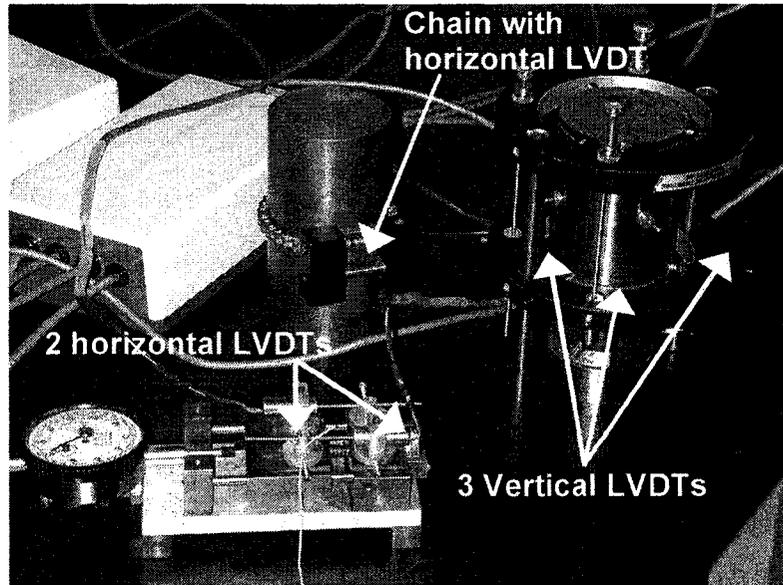


Figure 3.22 Transducers setup for stability test.

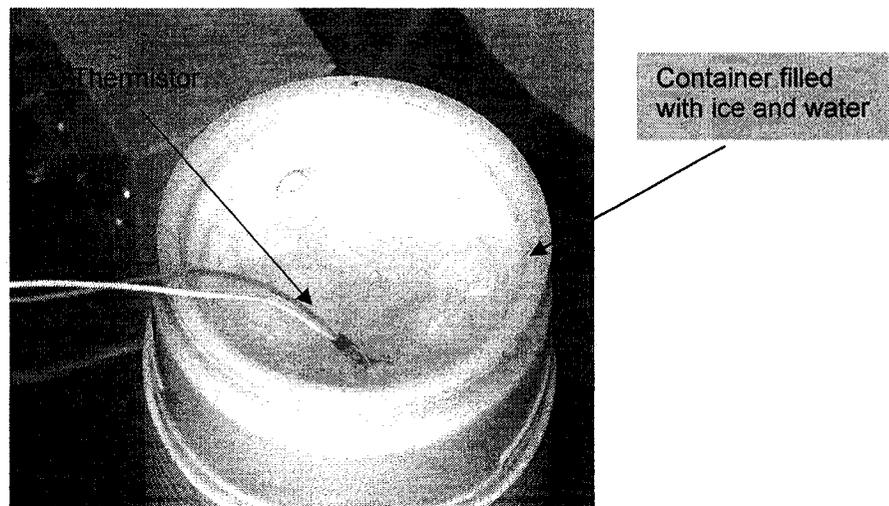


Figure 3.23 A picture of the thermistor embedded in ice water container for calibration purpose.

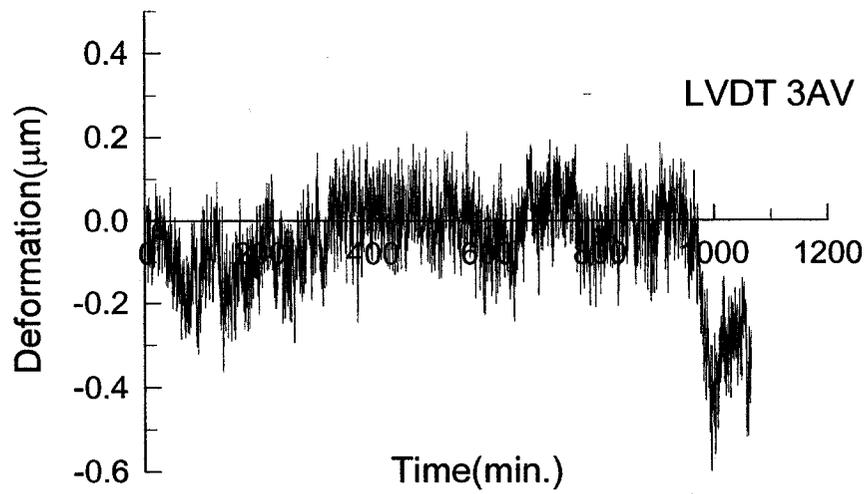


Figure 3.24 Stability of an LVDT for almost 24 hrs.

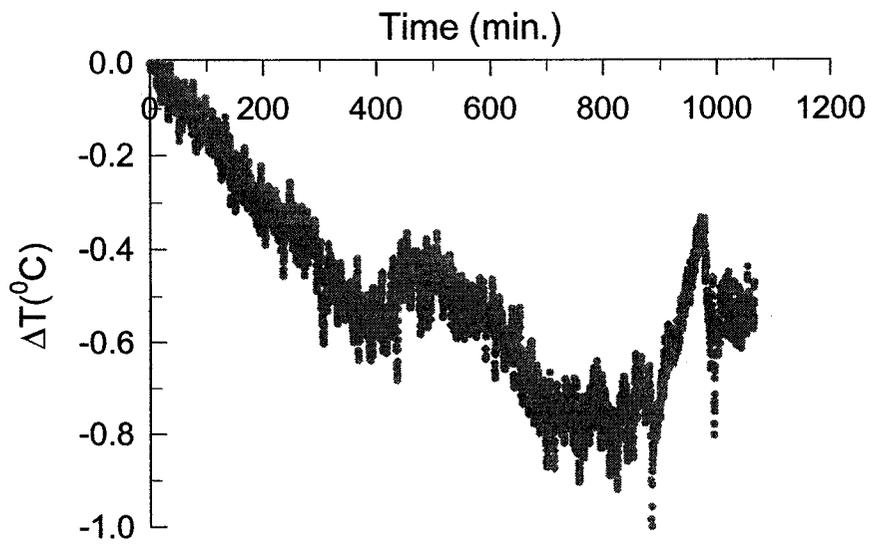


Figure 3.25 Variation of temperature with time.

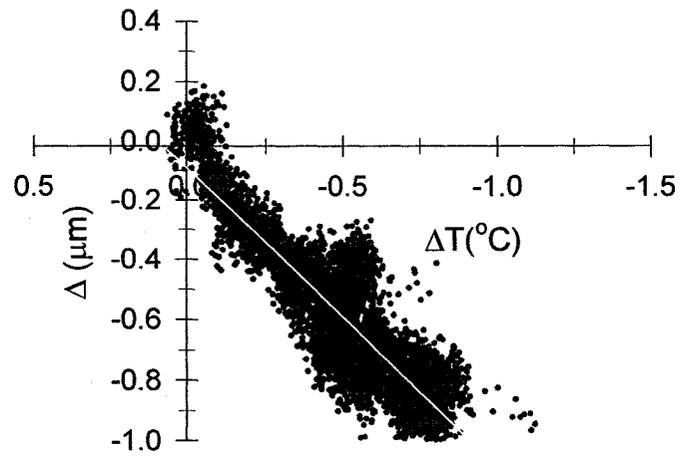


Figure 3.26 An illustration for variation of deformation with temperature for 24 hrs.

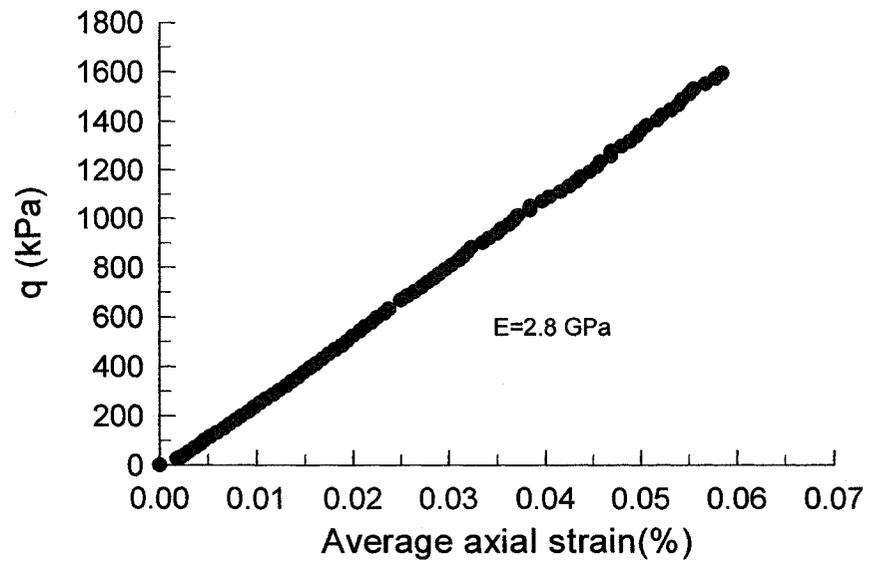


Figure 3.27 Proof testing results of a PVC sample.

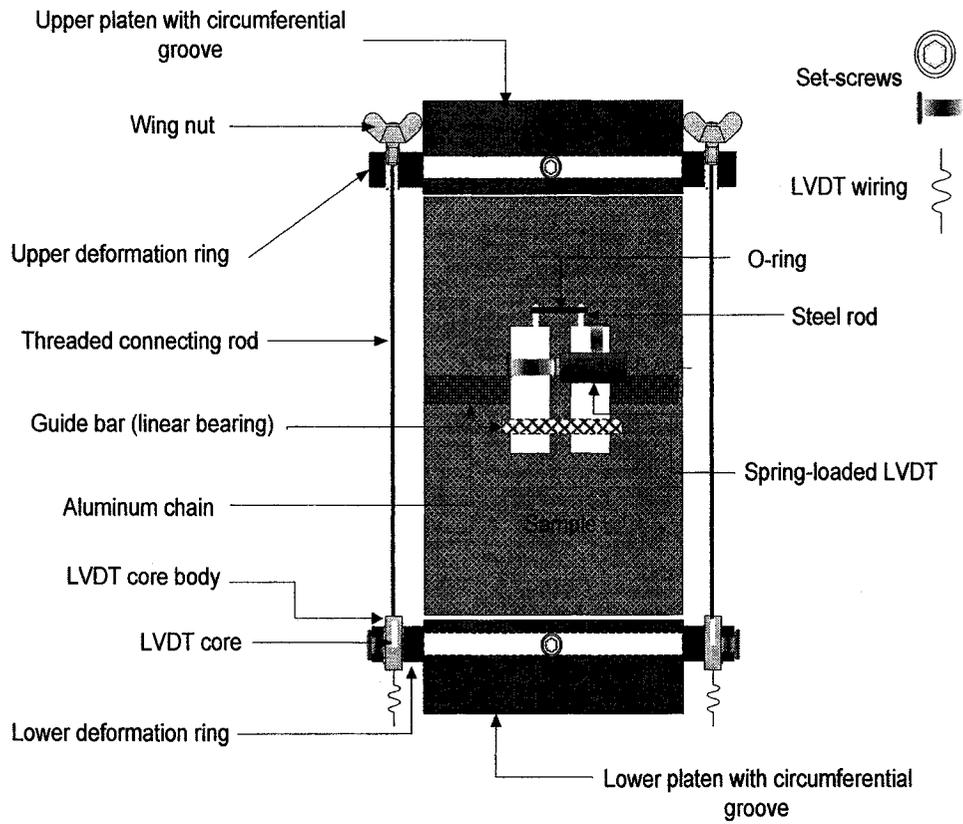


Figure 3.28 Original mechanical configuration and accessories.

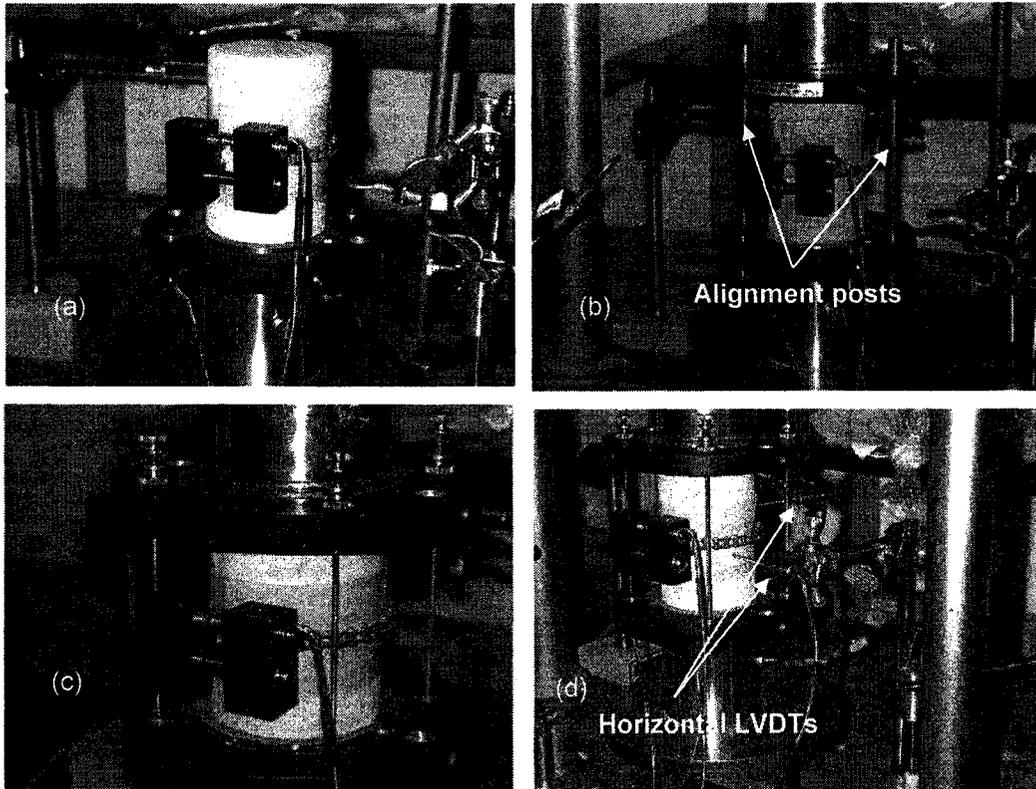


Figure 3.29 Mechanical configuration, accessories and mounting steps (modified set).

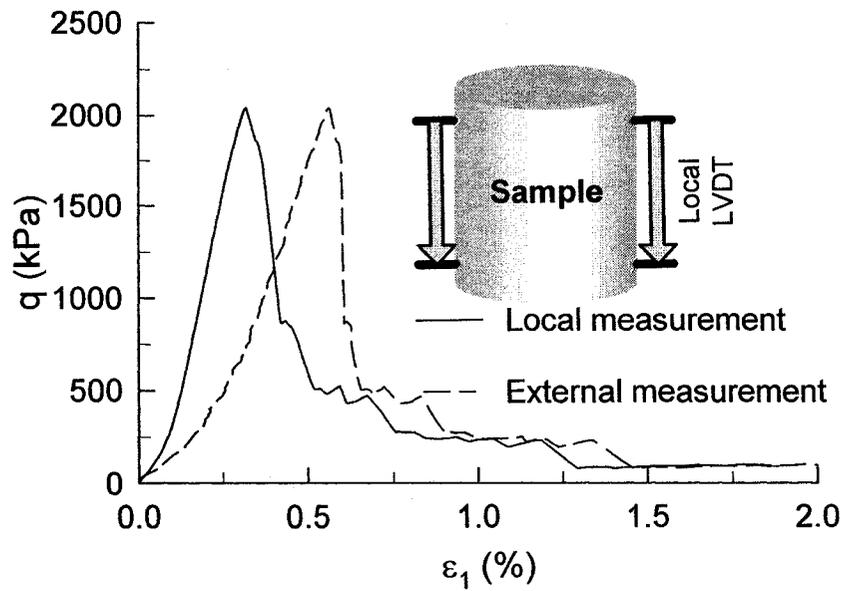


Figure 3.30 Local and external axial stress vs. axial strain curves.

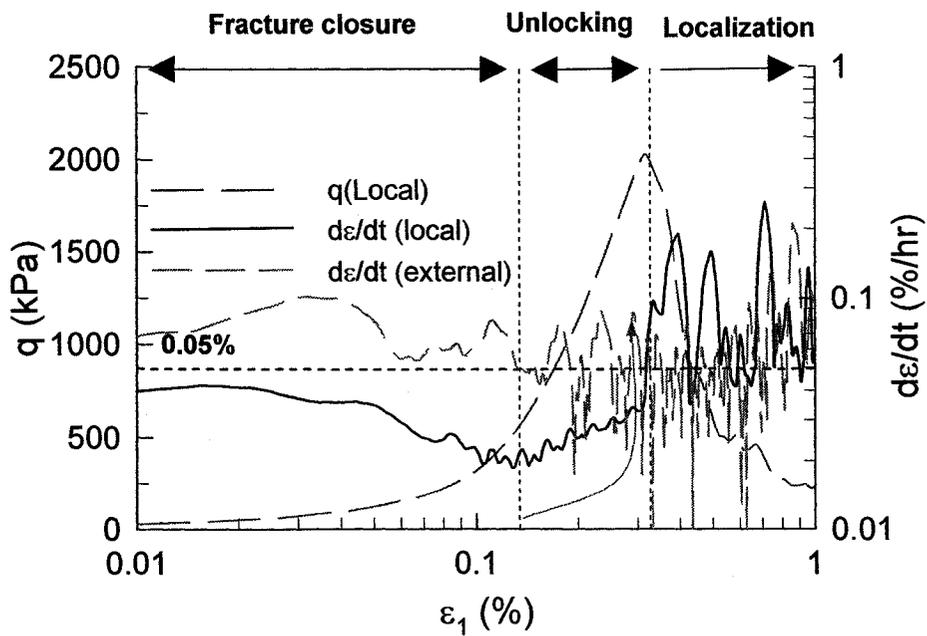


Figure 3.31 Comparison of external and local axial strain rate.

CHAPTER 4: MOBILIZED STRENGTH OF HIGH-HILL RIVER LOCKED SAND FORMATION AND FAILURE MODES

4.1 Introduction

As shown in Chapter 3, locked sands are classified as *arenaceous* (predominately quartz sand particles) weak sedimentary rocks that constitute the boundary between soils and rocks. The source of the high strength of locked sands compared to dense sands stems from the mild diagenesis that forms a coherent interlocked sand mass without the cementation commonly encountered in sandstones (Barton 1993). Previous investigations of the geomechanical response of locked sands typically focused on the ultimate strength and volume change characteristics by using external strain measurements (Dusseault and Morgenstern 1979; Barton et al. 1986; Richards and Barton 1999; Wong 1999; Dittes and Labuz 2002; Creswell and Barton 2003). It is now recognized that local axial and lateral strain measurements in soils and rocks are indispensable and can reveal important information in the pre-peak regime (e.g., Shibuya et al. 1994; Martin and Chandler 1994; Scholey et al. 1995; Fujii et al. 1998; Hayano et al. 1999; Bésuelle et al. 2000; Yimsiri et al. 2005; Sulem and Ouffroukh 2006). For locked sands, Abdelaziz et al. (2006) contrasted local and external axial strain results under uniaxial compression test conditions and showed that external measurements overestimated axial strains at various strain levels. Local strain results were shown to be more valuable for comprehending the failure mechanism, its evolution, and strain rates' effects. Recently, Creswell and Powrie (2004) carried out an experimental investigation of locked sand by using local

strain measurements. These researchers recommended the onset of dilation threshold as a strength threshold candidate to prevent the destruction of the interlocked fabric. However, no attention was paid to the variation in local lateral strain and its significance. For hard rocks, Martin and Chandler (1994) showed that the pre-peak damage thresholds of a hard rock (granite), determined from the local axial and lateral strains, were true material properties independent of the scale and loading rate. Furthermore, Martin and Chandler (1994) argued that the peak strength was related to the testing boundary conditions, i.e., the length-to-diameter ratio, the sample size, and the loading rate. These investigators also recommended the onset of dilation to be the long-term strength for granite instead of the peak strength. Roscoe (1970) stressed the value of understanding failure's precursory strains in soils, because the eventual failure surfaces and their orientation were determined well prior to the peak strength. Desrues and Chambon (2002) highlighted the importance of appropriately simulating the pre-peak response of soils up to the onset of localization instead of matching peak strength, which was merely a sign of the established failure.

Recent experimental and theoretical work on the failure kinematics of dense sands revealed that the shearing resistance was in principle governed by sliding, rolling, and mostly rotational resistance at grain contacts (e.g., Oda and Kazama 1998; Wan and Guo 2004; Alshibli and Alramahi 2006). Investigations of the shear strength of dense sands and locked sands indicate that sheared specimens tended to form columns/block of particle groups parallel or sub-

parallel to the major principal stress's direction in the pre-peak regime (Oda and Kazama 1998; Oda et al. 2004; Rechenmacher 2006; Cresswell and Barton 2003). These columns tend to dilate and buckle at peak stress to form a persistent inclined shear band creating large voids through the rotation of individual particles or blocks of particles. This dilatant mode of deformation through creation of voids in the direction of major principal stress is analogous to the concept of the dilation bands recognized in unconsolidated sands and weak sandstones (Bésuelle 2001; Du Bernard et al. 2002; Shultz and Siddharthan 2005) and the dilatant microcracks in brittle rocks (e.g., Schulson et al. 1999; Renshaw and Schulson 2001). For locked sands, the reported experimental results indicted a dominant shear band at failure (e.g., Dusseault and Morgenstern 1979; Cresswell and Powrie 2004). No experimental evidence has been reported for such column structures when cylindrical specimens are tested under compression.

The first part of this chapter deals with the evolution of failure by identifying damage thresholds at various stages of deformation, and the relationship between onset of dilation and localization for a traditional 2:1 specimen with lubricated ends. Based on the experimental results, mobilized strength parameters are partitioned in terms of the interlocking, friction angle, and dilation angle contributions with the plastic shear strain. The second part examines the formation of the vertical columns, under a uniaxial compression test, where samples are tested by using a slenderness ratio of 1:1 and a low strain rate. In

the third part, multistage uniaxial creep test results will be presented and discussed in terms of the relationship between the onset of dilation and the long-term strength. Finally, the results of direct shear tests at low normal stress levels will be presented and discussed to show the minimum mobilized strength.

4.2 Effect of testing conditions on laboratory results

Testing conditions in terms of specimen geometry, end effects (end confinement effects), and strain rate are vital in the assessment of experimental results, i.e., the strength, strain, onset of localization, and failure modes. Bishop and Green (1965) investigated the effects of the slenderness ratio and end effects on the strength of dense sands under triaxial compression test conditions. To prevent end effects on strength, Bishop and Green (1965) advocated the use of slenderness ratios greater than 2 (lubricated or non-lubricated) or, alternatively, 1:1 specimens with thick lubricated ends between two membranes. However, Frost and Yang (2003) used optical image analysis to show for dense sands, that despite using a slenderness ratio greater than 2, the strains and void ratio were not homogeneous within the tested specimens under triaxial conditions. Higher values of the void ratio and strain at the specimens' mid-height (the weakest elements) were noted as compared to the values at the specimens' extremes. Based on their findings, Frost and Yang (2003) recommended a slenderness ratio of 1:1 with thick lubrication at the specimen/platen interfaces to achieve homogeneity in the strains and void ratio within tested specimens during shearing. Desrues et al. (1996), using computer tomography for dense sands tested in triaxial tests, found that for tested specimens with a slenderness ratio of

1:1, the onset of localization was delayed compared to that of long specimens with a slenderness ratio of 1.9:1. Desrues et al. (1996) concluded that constraints favoring symmetry, such as a lubricated end-platen and a small slenderness ratio induced multiple localization modes instead of a persistent single band. Samieh and Wong (1997) compared the behavior of oil locked sand specimens with different slenderness ratios and end conditions in triaxial compression tests at low confining pressures. Samieh and Wong (1997) found that specimens with a slenderness ratio of greater than 1:1 failed with the formation of a single shear band. Shorter specimens with free ends resulted in a more stable and uniform deformation than that of longer specimens.

Another influential factor affecting the strain's homogeneity and the deformation's mode is the applied strain rate. Bhandari and Inoue (2005) recently explored the significance of low strain rates by using image analysis to capture strain rate effects on an artificial weak rock tested under biaxial conditions. Bhandari and Inoue (2005) found that low strain rates promoted uniform strain distribution among a sample's elements compared to high strain rates where localization became concentrated in narrow zones confined to the affected elements. In view of these studies, the formation of a single localized shear band (a preferred localization band) seems to preclude the propagation of reported pre-failure columnar structures in dense and locked sands due to early localization (loss of strain uniformity within specimen). Hence, in order to experimentally capture such reported column structures, specimens with a slenderness ratio of 1:1 and

thick lubricated ends tested under a low strain rate are preferable to achieve a uniform strain field, and, hence, render multiple localization zones at failure. However, these conditions are considered herein only to reveal such pre-peak structures, and are rarely encountered in the field as stresses and strains are never homogeneous due to either the inherent heterogeneity of geomaterials or the anti-symmetry of our civil engineering applications (complex boundary conditions).

4.3 Damage thresholds for brittle rocks

Damage thresholds are well recognized for low-porosity (<1%) brittle rocks and are measured by using electrical wired strain gauges mounted on the test specimen and, occasionally, with acoustic emission measurements (e.g., Brace et al. 1966; Bieniawski 1967*b*; Lajtai and Lajtai 1974; Martin and Chandler 1994; Eberhardt et al. 1998; Chang and Lee 2004). Each threshold marks a distinct stage of damage recorded by the local stress-strain responses. Despite the difference in origin and fabric between locked sands (interlocked fabric) and hard rocks (bonded fabric), these thresholds can be exploited as damage indicators as manifested by the variation in the axial and lateral strain responses. Such thresholds are defined for hard rocks as follows:

4.3.1 Crack closure stress

The crack closure stress is the axial stress level at which original open cracks, oriented at large angles to the direction of the axial loading close.

Crack closure stress is manifested by the shift of the stress-axial strain curve from an incremental rate increase to a constant rate increase (the axial stiffness levels off). Once the majority of the existing cracks are closed, linear elastic deformation commences. However, this initial portion of the stress-axial strain response may not always be observed and may be obscured by bedding effects due to the irregular surfaces at specimen ends during preparation.

4.3.2 *Damage initiation stress*

The damage initiation stress represents the initiation of the extension-stress-induced tensile micro-cracks. Damage initiation stress marks the onset of the stable micro-crack growth aligned with the direction of the major principal stress. The identification of this stress level can be captured only by detailed examination of the stress-lateral strain response, since the damage in the axial direction has no effect on the axial strains (Lajtai et al.1990).

4.3.3 *Damage coalescence threshold*

The damage coalescence threshold marks the reversal of the volumetric strains from compression to dilation. Beyond this stress level, crack propagation and interaction occur and result in the progression of strain localization and formation of the macro-scale shear bands. Martin and Chandler (1994) showed that this threshold is independent of sample size

and loading rate and should be regarded as a material property instead of the peak strength, which is dependent on the loading rate and scale effects.

4.4 Experimental results

The tested specimen was prepared with a slenderness ratio of 2:1, as traditionally used in our laboratories, and a diameter of 63.5 mm. The axial and lateral strains were monitored by using two vertical LVDTs and an aluminum chain at the mid-height of the specimen. End effects were alleviated by using two latex membranes placed at the specimen/platen interface to reduce friction effects. The specimen was sheared at a strain rate of 0.1%/hr. In the reported results, the compressive strains are positive while the dilative (extensile) strains are negative.

4.4.1 Axial stress vs. axial-strain curve

Figure 4.1a shows the axial stress ($q = \sigma_1 - \sigma_3$; $\sigma_3 = 0$)- average axial strain ($\epsilon_{1avg.}$) response for the tested specimen. The axial strain (the axial shortening with respect to the original height) is calculated based on the averaged vertical strain values of the two LVDTs. The specimen failed at an axial strain of 0.28% and attained peak strength at 1.8 MPa. According to the ISRM's (1978) classification of rocks, locked sand is classified as a very low-strength rock situated at the boundary between soils and weak rocks. The source of high uniaxial strength presumably stems from the relatively extensive area of grain contacts as well as the increase in the coordination number, i.e., the

number of contacts per particle. Thus, particles interact as surfaces rather than point contacts (as in dense sands), and, hence, the rotational resistance is enhanced. Therefore, in the absence of bonding, individual contacts can sustain considerable contact forces, moments, and stress obliquity due to this sand's interlocked fabric (Cresswell and Barton 2003).

In Figure 4.1a, the axial stress-axial strain curve is partitioned into three major stages. In the first stage (OA), the material undergoes an incremental increase in the axial stiffness up to point A. This point marks the closure of the inter-granular pores and the intra-granular cracks as well as the compression of soft inclusions in the specimens, as manifested by the microstructural studies. In stage AB, the axial stress-axial strain curve becomes linear with an average constant axial stiffness (E) of 0.85 GPa and an average Poisson's ratio (ν) of 0.2. At point B, the material undergoes a significant nonlinear change in axial stiffness up to the peak strength (C). Beyond point C, the strength drops abruptly. The inability to track the post-peak response results from the lack of servo-control in the post-peak region.

4.4.2 Axial stress vs. lateral strain curve

Figure 4.1b shows the relationship between the axial stress (q) and the lateral strain (ϵ_3) recorded by the chain LVDT located near the middle of the sample. The lateral strain exhibits lateral expansion all the way up to the peak strength. The axial stress-lateral strain curve can also be divided into three

stages based on changes in the nonlinearity. In stage OA, the axial-stress-lateral strain curve is characterized by negligible nonlinearity (almost constant). Beyond point A, an increase in nonlinearity occurs during stage AB. In the last stage (BC), a significant increase in the lateral strains and significant nonlinearity up to failure occur.

4.4.3 Volumetric strains and inception of localization

Figure 4.1c shows the volumetric strain (ε_v)-axial strain (ε_1) relationship. The volumetric strains integrate the variation of both the local axial and lateral strains ($\varepsilon_1+2\varepsilon_3$). The material exhibits an initial compression followed by a reversal in the volumetric strain (the onset of dilation). At the initial stage and up to point A, the volumetric strain is compressive and linear. Beyond this point, the volumetric strain remains compressive but non-linear up to point B. Point B marks the volumetric strain's reversal and the transition from compression to dilation up to the peak strength (C). Rowe (1969) showed that for extremely dense sands (with relative densities close to 100%), the maximum dilation factor ($1-d\varepsilon_v/d\varepsilon_1$) is equivalent to 2 and decreases with the reduction in the relative density. For the investigated sand, the maximum-recorded dilation factor is equivalent to 3.6 and is attributable to the particles' unusual contact characteristics, which are reflected in the exceptionally elevated relative density. According to the theory of internally constrained continua introduced by Goddard and Bashir (1990) and also according to Reynolds (1885), the dilatancy is a direct result of the internal and external

kinematic constraints. In granular materials, internal constraints are due to granular packing, so that the boundary grains prevent rearrangement during shearing. In bonded geomaterials, the bonds provide the source of internal kinematic constraint, which hinders rearrangement. In locked sands, the constraints are due primarily to the interlocked fabric. Therefore, higher energy is required to overcome such interlocking, and a higher dilation angle is expected as compared to that of dense sands with regular constraints (due to packing).

Localization of deformation in geological materials results in non-uniformity in strains. Consequently, constitutive models based on continuum approaches are less reliable once localization occurs. The non-uniformity of strains is considered by many researchers to be a good indicator of the onset of localization (e.g., Bésuelle et al. 2000; Sulem and Ouffroukh 2006; Abdelhamid et al. 2000; Labuz and Dai 2000). Fig. 4.1d depicts the relationship between axial strain differences ($\varepsilon_1 - \varepsilon_{avg}$), divergence of individual values from the mean value, and the average axial strain. During the early stage of loading, the readings are constant, signifying the relative uniformity of the axial strains. Immediately prior to the volumetric strain reversal, a significant divergence from the mean value occurs at point B', marking the onset of localization (the loss of strain uniformity). In this test, localization commences prior to the peak strength and in the hardening regime. This observation has been unequivocally reported in the literature as opposed to

the assumption of localization of deformation in the post-peak regime for most constitutive relations. For instance, localization in the hardening regime was reported by Desrues et al. (1985) for dense sand tested under biaxial condition, by Bésuelle et al. (2000) for Vosges sandstone under triaxial compression condition, by Sulem and Ouffroukh (2006) for Fontainebleau sandstone, and by Ord et al. (1991) for Gosford sandstone. Abdelaziz et al. (2005) demonstrated that the onset of dilation is an excellent indicator of localization for other geomaterials including dense sands, overconsolidated clays, locked sands, sandstones, and granites. This threshold is of paramount importance as strain measurements become less reliable once exceeded so that efforts to simulate geomaterials' behavior beyond this threshold within the framework of continuum approaches will be problematic.

4.4.4 Damage and failure evolution

4.4.4.1 Detection of damage thresholds for locked sand

Damage threshold will be established and discussed for the tested specimens by using the previously established framework. Figure 4.2a shows the relationship between the tangential axial stiffness ($E_{tan.}$), the incremental change in the axial stress with an incremental change in the axial strain, and the axial stress in a logarithmic plot. The tangential stiffness tends to increase nonlinearly up to point A, which marks the crack closure strength threshold (σ_{cc}) where the specimen's elements come into contact. After this point, the tangential stiffness in the axial direction is almost constant (0.85 MPa) up to

point B' where the initial reduction occurs in tangential stiffness, marking the initiation of damage coalescence. In addition, point B' is identical to the stress level where localization commences, as shown previously in Figure 4.1d.

Figure 4.2b shows the relationship between the tangential volumetric stiffness, the incremental change in axial stress with an incremental change in the volumetric strains, and the axial stress. The testing's initial stage up to point B' shows a constant slope reflecting the correspondence between the axial stress and the volumetric strain. At point B', volumetric stiffness accelerates nonlinearly and therefore can represent the initiation of damage coalescence (σ_{dl}) due to an increase in the density of unlocked sites throughout the specimen. At point B, volumetric stiffness discontinuity occurs, representing the change in the sample's behaviour from compression to dilation. This result implies the acceleration of damaged sites coalescence as well as the uncontrolled progression of the damage (the damage coalescence strength threshold (σ_{dc})).

Figure 4.2c shows the relationship between the tangential lateral stiffness (the incremental change of the axial stress with incremental change in the lateral strains) and the axial stress. Up to point A, a negligible variation in the lateral stiffness takes place. Beyond this point, the maximum lateral stiffness degrades nonlinearly and marks the damage initiation threshold (σ_{dl}) due to the progression of fabric unlocking. At point B, the tangential lateral stiffness

starts to decrease linearly, signifying the complete interaction of the unlocked sites as well as the significant progression of lateral unlocking.

From the previous analysis, several thresholds were identified. The damage initiation of the interlocked fabric and the crack closure stress thresholds coincided and occurred at 28% of the peak strength. Sporadic unlocking at several sites within the specimen resulted in an increase in the number of damaged sites. At the damage localization stress threshold, these damaged sites started to coalesce to form a persistent band where localization commenced. This threshold was shown to take place immediately before the volumetric strain reversal at 55% of the peak strength. After initiation of the damage coalescence, the damage started to accumulate in an accelerated manner at damage coalescence stress threshold, which marked a reversal in the volumetric strains. This threshold occurred at 67% of the peak strength, as identified from the current analysis. Table 4.1 provides a summary of the damage thresholds and the elastic parameters.

4.4.4.2 Evolution of failure in locked sand

Figure 4.3 shows the deviator stress vs. the plastic axial, lateral, and volumetric strains. The plastic axial and lateral strains can be approximated as the difference between the total strain and the elastic strain by using Hooke's elasticity law ($E=0.85$ GPa, $\nu=0.2$). However, this is only a simplification as the material can exhibit a non-linear elastic response, which

can be only captured through unloading and reloading during shear evolution.

The plastic volumetric strain can be expressed as

$$[4.1] \quad \varepsilon_v^p = \varepsilon_1^p + 2\varepsilon_3^p,$$

where ε_1^p and ε_3^p are the plastic axial and lateral strains, respectively.

The pre-peak observations of shearing can be divided into three stages. Stage I is marked by the closure of cracks, pores, and fractures within the quartz particles created due to the huge burial depth as well as the compression of soft inclusions: clay coating and bridging. This process causes plastic compressive strains, with negligible sliding between the adjusted particles. The second stage marks the initiation of fabric unlocking through lateral expansion creating inter-granular voids (damaged sites) parallel or sub-parallel to the axial direction, similar to the inter-granular microcracking encountered in sandstones (Sangha et al. 1974; Santarelli and Brown 1989; Menendez et al. 1996). In stage II, the plastic axial strains are quite constant, indicating the progression of sporadic fabric unlocking and damage in the axial direction through lateral expansion. The same observation has been reported for hard rocks during the extension of the cracks in the maximum principal stress direction (e.g., Lajtai 1998). In stage III, vertically unlocked damaged sites coalesce to form columns through which the load is transferred to the sample in conjunction with an increase in the

plastic axial strains (refer to Figure 4.3). Once coalescence occurs, the formation of macro-fractures (vertical and inclined bands) commences. It is argued that the formation of inclined bands is dictated by these columns' ability to buckle, which is a function of the rotational resistance of the particles constituting each column. In locked sand, the interlocking between particles is considerable, and, hence, the dominant admissible mechanism is the separation of these columns horizontally with slight buckling and the formation of inclined bands with low vertical angles (the higher the degree of constraint, the lower the failure surface's vertical angle due to the inability to buckle). The failure is presumably a combined effect of significant geometric dilation between columns, columns buckling, and the falling of the failed portions between columns, and, hence, the loss of the columns' load-carrying capacity. The inclination of the formed bands with the axial stress direction ranged from zero vertical bands to 15° (slightly inclined bands that are dominant). The examination of the failed specimen did not reveal the full formation of these columns, so that the early localization inhibited its complete formation in the axial direction and promoted the formation of inclined bands and the buckling of the specimen. Figure 4.4 presents a schematic illustration of the proposed mechanism of failure.

4.4.5 Mobilized strength components in locked sands

One of the earliest attempts to partition strength parameters was done by Taylor (1948) for dense sands. Taylor, using direct shear test results of dense

sands, attributed the dissipated energy at the peak to the energy dissipated by the basic friction between the particles and the energy needed to increase in volume (dilation). Thurairajah (1961) carried out a series of drained and undrained triaxial tests on sand and kaolinite. After integrating out the effect of dilatancy and change in soil grains' elastic deformation, Thurairajah (1961) found that the dissipated work in plastic deformation was related to the critical state friction angle (a constant deviator stress with no change in volumetric strain) at all stages of the test path. Rowe (1962) used the principle of the minimum energy ratio (the existence of a preferential sliding direction at particle contacts so that the internal energy absorbed is minimum) to separate the basic frictional component and the dilatancy component for a wide range of dense sands. According to his approach, the peak strength of dense sands is the end-product of a basic friction angle and a dilation component that is absent for loose sands. Ladanyi and Archambault (1970) and Barton (1973) proposed a similar concept to express the strength of jointed rocks.

Muhunthan et al. (2004) determined, by using drained triaxial compression tests for Ottawa sand at different relative densities, the shear resistance against particle-to-particle frictional deformation (the true friction resistance) after applying corrections to the measured shear strength for plastic volumetric and stored elastic deformations. Muhunthan et al. (2004) indicated that the shear resistance from particle-to-particle frictional deformation is the sand's critical state strength, which is mobilized and stabilized at a low strain

before the peak. Wong (1999) applied the stress dilatancy relation proposed by Rowe (1962) to reveal the source of the strength in oil-rich locked sands. The normalization with respect to dilation did not reveal a constant stress ratio, which Rowe (1962) had found for dense sands. Wong (1999) indicated that the stress- dilatancy relationship did not account for the interlocking component and rolling in locked sands. In addition, the strength in the pre-peak regime was dominated by the interlocking strength component. This component mobilized at a low strain well before the peak and started to degrade prior to the peak strength.

According to preceding discussion, the strength parameters for dense and locked sands can be expressed as the sum of three components: (a) the basic frictional component (ϕ_f) governed by the surface texture and mineralogy and bounded by the interparticle friction angle (ϕ_m) and critical state angle (ϕ_{cv}); (b) the intrinsic interlocking component, which is responsible for its coherent mass due to the interpenetrative, tangential, sutured contacts, and which resembles the role of bonding in weak and hard rocks; and (c) the dilation component controlled by the degree of angularity and dictated by the freedom of grains to override each other after the occurrence of significant unlocking.

4.4.5.1 Mobilized strength parameters at damage thresholds

The generalized form of the 2D stress dilatancy relationship for geomaterials can be expressed as the sum of three terms (Rowe 1962; Ueng and Chen 2000) as follows:

$$[4.2] \underbrace{\left[\frac{\sigma_1}{1 - \frac{d\varepsilon_v}{d\varepsilon_1}} \right]}_{\text{Basic resistance}} = \underbrace{\sigma_3 \tan^2 \left(\frac{\pi}{4} + \frac{\phi_f}{2} \right)}_{\text{Frictional resistance}} + \underbrace{\frac{dE_b(1 + \sin \phi_f)}{d\varepsilon_1 \left[1 - \frac{d\varepsilon_v}{d\varepsilon_1} \right]}}_{\text{Particle Breakage}} + \underbrace{2c_f \sqrt{\tan^2 \left(\frac{\pi}{4} + \frac{\phi_f}{2} \right)}}_{\text{bonding resistance}},$$

where σ_1 and σ_3 are the major and minor principal stresses, respectively; ϕ_f is the basic friction angle; $d\varepsilon_v/d\varepsilon_1$ is the dilatancy rate; $d\varepsilon_1$ is the axial strain increment; c_f is the intrinsic bonding resistance; and dE_b is the energy consumption due to particle breakage.

The basic resistance integrates out the effect of dilatancy by normalizing the major principal stress with the dilatancy factor $(1 - d\varepsilon_v/d\varepsilon_1)$. This resistance reflects the energy dissipated through the frictional resistance between particles, the particles' breakage, and overcoming the bonding between particles. The first right hand side term in equation [4.2] is considered to be controlled by mineral type and surface texture (small asperities). The second term expresses the energy consumption due to particle breakage (Ueng and Chen 2000). However, Ueng and Chen (2000) found that particle breakage is

insignificant for strong particles under low confinement stresses. Herein, this term will be ignored because the majority of the grains (quartz) of locked sands are strong in nature. The third term represents the bonding resistance, which is the resistance in excess of the basic frictional resistance, particle breakage, and dilation. The stress dilatancy relationship is applicable for sands not only at peak but also in the pre-peak regime (Rowe 1962; Shibuya et al. 1994).

Equation [4.2] can be reduced to equation [4.3] by using the current geometric and boundary conditions for the tested specimen ($\sigma_3=0$ and replacing the bonding component ($2c_f$) with an interlocking component (I)).

$$[4.3] \quad \frac{\sigma_1}{1 - \frac{d\varepsilon_v}{d\varepsilon_1}} = I \sqrt{\tan^2 \left(\frac{\pi}{4} + \frac{\phi_f}{2} \right)}$$

Equation [4.3] can be utilized to reveal the interlocking's contribution at any strain level. The basic friction angle is substituted by using the critical state friction angle for quartz sand (30°) (Rowe 1962; Rowe 1969; Bolton 1986).

The dilation angle (ψ) can be approximated by using the following equation (Vermeer and De Borst 1984):

$$[4.4] \quad \psi = \arcsin \left[\frac{\frac{2d\varepsilon_3^p}{d\varepsilon_1^p} + 1}{\frac{2d\varepsilon_3^p}{d\varepsilon_1^p} - 1} \right] = \arcsin \frac{d\varepsilon_v^p}{-2d\varepsilon_1^p + d\varepsilon_v^p}$$

The friction angle (ϕ) can be calculated by using Rowe's stress-dilatancy relationship for frictionless geomaterials, as derived by Vermeer and De Borst (1984) and expressed as follows:

$$[4.5] \quad \sin \phi = \frac{\sin \phi_f + \sin \psi}{1 + \sin \phi_f \sin \psi}$$

Figure 4.5 shows the approximate distribution of the interlocking component, dilation angle, and mobilized friction angle with the plastic shear strain (ε_s^p), which is calculated by using the following equation:

$$[4.6] \quad \varepsilon_s^p = \frac{2}{3}(\varepsilon_1^p - \varepsilon_3^p)$$

Figure 4.5 reveals that the shear strength parameters are not mobilized simultaneously. The interlocking contribution tends to be consumed in the pre-peak regime and levels off before attaining the peak strength. The maximum interlocking strength component is recorded at the damage initiation stress where the unlocking of the interlocked fabric starts to develop

sporadically throughout the specimen. Once unlocking takes place, friction starts to mobilize due to the interaction of the particles at damaged sites. At the damage coalescence strength level, 40% of the interlocking component is lost, and only the basic friction angle is mobilized. Beyond this damage threshold, a continuous reduction of the interlocking component occurs with the mobilization of the dilation angle due to the severe dilation of the interlocked fabric at this stress threshold. The dilation angle increases exponentially from zero at the damage coalescence strength to 36° at the peak where the particles or the blocks of particles become less constrained after unlocking. The maximum friction angle increases dramatically from zero at the damage initiation strength threshold, where sliding occurs between newly unlocked particles or in a block of particles, to 30° (the basic friction angle) at the damage coalescence strength and reaching 57° at the peak strength. The friction angle at the peak strength agrees with the reported maximum friction angles for the arenaceous McMurray Formation (Dusseault and Morgenstern 1978; Dusseault and Morgenstern 1979) .

With regard to design approaches, Cresswell and Powrie (2004) argued that due to the large difference between locked sand's peak and critical state strengths, a geotechnical design based solely on the critical state friction angle would be unduly conservative even with a factor of safety of 1. Alternatively, Cresswell and Powrie (2004) recommended a design based on the limiting stress envelope corresponding to the onset of dilation and no-

tension cut off, with a factor of safety on the peak stress ratios to avoid progressive failure. The results from this investigation support their recommendation, as the onset of dilation is also an excellent indicator of localization.

4.5 Development of columnar structures and dilation bands

A specimen, prepared with a 1:1 slenderness ratio, was sheared at a slow strain rate of 0.05%/hr to achieve a relatively homogeneous deformation throughout the specimen to allow for the propagation of the pre-peak's vertical structures to failure, and to delay the onset of localization. A thick layer of grease sandwiched between two latex membranes was placed at the specimen/ platen interfaces to alleviate the end effects.

Figure 4.6 shows the stress-strain and the volumetric strain curves for a specimen with a slenderness ratio of 1:1. During the shearing of the 1:1 specimen, dilation was initiated well before reaching the peak strength and took place at the same stress level as that for 2:1 specimen (the damage coalescence threshold). The specimen deformed uniformly during the shearing evolution, and the localization commenced slightly after peak strength was attained. This result was predictable due to the distributed dilation, which corresponded to a delay in the onset of localization as compared to that in 2:1 specimens. Furthermore, the short specimen attained a peak strength of 2 MPa, which is 10% greater than the recorded value for the long specimen, and failed with a less abrupt drop in

strength due to the low strain rate and the strain's relative uniformity as compared to that of the long specimen.

Figure 4.7 shows the rendered failure surfaces based on the previously mentioned testing conditions. Failure of the sample occurred through the formation of vertical columns extending in the direction of the axial stress, as numbered in the figure. These axially oriented columns tended to form the load-carrying members. The presence of such columns suggests that the lateral extension and unlocking at discrete locations along the specimen's vertical axis were the main cause of failure. Examination of the column-to-column interface revealed the existence of the columns' curvature (see Figure 4.7). This curvature signifies the possibility of horizontal rotation of such columns after their formation so that the rotational resistance is governing the shearing behavior, as several investigators have suggested for dense sands (Oda and Kazama 1998; Oda et al. 2004; Wan and Guo 2004; Alshibli and Alramahi 2006). As well, Cresswell and Barton (2003) speculated that this process is a key component of locked sand's failure.

4.5.1 Theoretical prediction of dilation band occurrence

The failure of geomaterials in laboratories or the field is customarily associated with the formation of localized deformation bands classified into two categories summarized by Borja and Aydin (2004). The first category includes mixed-mode deformation bands consisting of both shear strains

(parallel to the band boundary) and volumetric strains (the volume increases or decreases perpendicular to the band boundary) along the band boundaries. Examples of mixed-mode bands are dilative and contractive shear bands. The second category consists of pure shear bands (shear along the band boundary), compaction bands (compressive volumetric strains), and dilation (extension) bands (dilative volumetric strains). The prediction of the onset of localized bands as well as their orientation with the maximum principal stress plane for dilative porous rocks has been studied by several investigators (e.g., Rudnicki and Rice 1975; Perrin and Leblond 1993; Issen and Rudnicki 2000; Bésuelle 2001).

The developed rigorous mathematical formulation for band formation was subjected to two conditions (Rudnicki and Rice 1975): (a) kinematic compatibility so that the velocity field was continuous at the inception of localization and (b) stress equilibrium so that the traction rates were continuous across the boundary. By combining these conditions along with a constitutive relation that describes material behavior, and assuming an identical constitutive modulus for the material inside and outside the band at the onset of localization, the following equation is obtained:

$$[4.7] (n_j C_{ijkl} n_l) g_k = 0 ,$$

where the subscripts are index notations; n_i and n_l are the normal vector projections on the principal stress directions; C_{ijkl} is the modulus tensor; and g_k is a vector function of the distance across the band and related to the velocity gradient difference (inside and outside the band) within the band that vanishes at boundaries. One possible solution for the above equation was obtained for $g_k=0$ (a homogeneous strain field). An alternative solution can be achieved for $g_k \neq 0$ (the determinant of the coefficient is zero), resulting in the bifurcation criterion. By solving equation [4.7], it was realized that the hardening modulus is a function of the band orientation. The critical hardening modulus (the slope of the axial stress- plastic axial strain curve) and the associated orientation were obtained by assuming that the plane of the deformation band is parallel to the intermediate principal stress direction. The solution was shown to be valid for the following range of $(\beta+\mu)$ expressed as follows (Issen and Rudnicki 2000; Bésuelle 2001):

$$[4.8] \underbrace{(1-2\nu)N - \sqrt{4-3N^2}}_{\text{compaction bands}} \leq \frac{2}{3}(1+\nu)(\beta+\mu) \leq \underbrace{(1-2\nu)N + \sqrt{4-3N^2}}_{\text{dilation bands}},$$

where N is the normalized stress ratio (s_{22}/τ); τ (equivalent stress) = $\sqrt{\frac{1}{2}s_{ij}s_{ij}}$

and s_{ij} (deviator stress) = $\sigma_{ij} - \frac{1}{3}\delta_{ij}\sigma_{kk}$ (δ_{ij} is the Kronecker delta, which =1 if $i=j$

and =0 if $i \neq j$), and equivalent to $(1/\sqrt{3})$ for the axisymmetrical compressive

stress path; $\mu = \tan \phi$ (the slope of yield surface) ; and $\beta = -d\varepsilon_v^p/d\varepsilon_s^p$ (dilatancy factor).

If the lower inequality is violated, compaction bands are manifested with the band normal parallel to the major principal stress direction. On the other hand, if the upper inequality is violated, dilation bands are manifested with the band normal parallel to the minor principal stress direction. The dilation bands are described as the separation of the material's mass into walls being filled by parent material without shear offsets). Within the specified range, the mode of deformation can be described as a pure shear band, a dilative or contractive shear band. Figure 4.8 shows the deformation bands boundaries of occurrence and orientations in terms of μ and β for axisymmetrical compression conditions at the onset of localization based on the previous equation. The same figure indicates the data points estimated at the onset of localization, where divergence of local axial strains occurs, for both long and short specimens.

In the case of a high slenderness ratio, the data point lies at the upper boundary of the dilative shear bands with an orientation of 85° , which corresponds to the experimental observations. On the other hand, for a low slenderness ratio, dilation bands are observed. These results are in agreement with Chau's (1993) theoretical study of bifurcations in pressure-sensitive circular cylinders under axisymmetric compression. Chau (1993)

found that the first possible mode of localization was always the buckling of long specimens as compared to short specimens where a diffused mode of localization occurred.

The previous results showed that the samples with a high slenderness ratio hindered the propagation of columnar structures due to buckling effects and the early localization of the deformation. Hence, dilation became restricted to a distinct shear band. For short specimens, multiple localized sites were observed, and, hence, dilation was not localized in a distinct band, allowing for propagation of vertical columns and multiple dilatant bands. Multiple dilation bands separated these columns, resulting in higher dilation as compared to that in long specimens.

4.6 Multistage uniaxial creep test

The design of geotechnical structures not only necessitates the stability under short-term conditions to avoid tensile/shear failures but also long-term stability under constant loading conditions to avoid excessive deformations. Creep potential has been recorded for a wide range of geomaterials including sands (e.g., Lade and Chi-Tseng 1999; Kuwano and Jardine 2002; Bowman and Soga 2003), weak sedimentary rocks (Maranini and Brignoli 1999; Fuji et al. 1999) and hard rocks (Lin et al. 2004; Lau and Chandler 2004). In this section, the results of a multistage uniaxial creep test conducted on the locked sand are described and discussed. The chief objective of this test was to compare the creep potential

before and after reaching the onset of dilation stress threshold to verify its significance as a safety threshold for long-term strength.

4.6.1 Testing procedure

A cylindrical specimen was prepared with a height of 65mm and a diameter of 60 mm. The axial and lateral deformations were monitored by using three local LVDTs and a circumferential chain, respectively. End effects were alleviated via a thick layer of silicone grease sandwiched between two latex membranes at the specimen/platen interfaces. Once the specimen and measuring devices set-ups were completed, the specimen was sheared at a constant strain rate of 0.05%/hr. Data were logged every 20 seconds during both shearing and creep stages. The test program was comprised of three stages of creep followed by unloading, once the peak strength was attained. Upon reaching the creep stress level, the deviator stress was halted for at least 24 hrs monitoring the variation in axial as well as lateral deformation were monitored. At the end of each creep stage, the specimen was re-sheared at the same strain rate to attain the next level of the creep stress level. The creep stress levels were intentionally selected so that the first creep stage (at 750 kPa) was conducted below the onset of dilation threshold (1200-1250 kPa) as determined from the monotonic uniaxial compression tests. The second and third creep stress levels (1500 kPa and 2000 kPa) were carried out above the onset of dilation threshold.

4.6.2 Stress- strain and volumetric strain curves

Figure 4.9a shows the deviator stress vs. the average axial strain relationship during the shearing and creep stages. The conducted stages of creep as well as the onset of dilation threshold are marked on the figure. The tested specimen failed at a uniaxial strength of 2.45 MPa, which is approximately 1.23 of the monotonic uniaxial test without creep stages for similar testing conditions. This result was expected due to the ageing effect under a constant deviator stress that could result in an increase in the specimen's ability to carry more load, as other investigators have shown this phenomenon for granular materials (e.g., Murayama et al. 1984; Mitchell 1986; Mesri et al. 1990; Schmertmann 1991). Creep effects were not only evidenced through the increase in uniaxial strength value, but also were reflected as an increase in initial stiffness upon reloading for the next creep stage (see Figure 4.9a). During creep stages, a slight reduction in the deviator stress was recorded. The reduction in deviator stress was more pronounced in the second and the third stages of creep. On the other hand, no reduction of deviator stress was noted in the first stage below the onset of dilation threshold. It is postulated that during the course of shearing, the zones of the material became stressed. With the progression of the creep, some of the stressed zones had a propensity to relax, and, thereby, distribute stresses to their unstressed neighboring elements. This redistribution could reduce the overall deviator stress. Upon reloading, the unstressed elements carried some of the imposed stresses, and, hence, an increase in initial

stiffness was the result upon attaining the previous stress-strain path. However, the reduction in stresses could be otherwise due to the stiffness of the loading frame.

Figure 4.9b shows the deviator stress vs. lateral strain relationship. During the course of the experiment, the lateral strains exhibited extension. The same observation regarding the initial increase in the lateral stiffness was recorded upon reloading to the next stage of creep, as the figure reveals. Figure 4.9c depicts the volumetric strain- axial strain relationship. The volumetric strains remained contractive up to almost 1200 kPa, followed by dilation up to the peak strength. However, the maximum rate of dilation was not coincident with the peak resistance. The volumetric strains during the creep seemed to be a continuation of the original volumetric strain relation.

4.6.3 Creep strains and strain rates

In soil mechanics, the variations of creep strains are traditionally investigated in terms of axial strains. Three stages of creep are usually described: primarily, transient, and tertiary creep. During the primary creep, strains are accelerating with time, and strain rates are progressively reduced to reach the transient creep (the constant rate of strain). The third zone is the tertiary creep, where the strain rate accelerates in an unstable manner resulting in failure.

For efficient and reliable quantification of creep strains, temperature effects should be considered to avoid error with temperature variations. The temperature effects seemed to have minor effects on the LVDTs' readings during creep stages. The variation in temperature was in the range of 1- 2 °C. This is manifested as a deviation in average vertical and lateral strains of 0.0008% based on the previously shown temperature calibration factors. Table 4.2 presents the test results at each creep stage in terms of axial, lateral, and volumetric creep strains. Following the same convention, positive and negative strains indicate compression and expansion (dilation), respectively.

4.6.3.1 Axial creep strains and strain rate

Figure 4.10 shows the variation in axial strains with time during creep stages. Axial strains were calculated as the average of the three vertical LVDTs' readings. The axial strains at any stage of creep are characterized by a sudden increase in strain, followed by a reduction in the axial strain rate to a constant rate of strain, and finally reach a plateau. Figure 4.11 indicates the relationship between the axial strain rates and the elapsed time in a logarithmic scale. A monotonic reduction in axial strain rates occurs with elapsed time. For comparison, the axial strain rate at each stage is calculated based on the linear portion of the axial strain vs. the elapsed time relationships shown in Figure 4.11 and summarized in Table 4.2. The axial strain rate in the first creep stage is almost 58% lower than the creep rate in

the second and third stages. On the other hand, the difference in axial strain rates in the second and third stages is insignificant and in the order of magnitude of 10^{-6} %/min.

4.6.3.2 *Lateral creep strains and strain rate*

Figure 4.12 shows the variation in lateral strains with time during creep stages. The lateral strain curves at any stage of creep are characterized by a sudden increase followed by a reduction in the lateral strain rate to a constant rate of strain and eventually reaching a plateau. Nonetheless, the increase in lateral strain values during creep stages is more noticeable as compared to axial strain values. Figure 4.13 indicates the relationship between the corresponding lateral strain rates and time in a logarithmic scale. Like the measured axial strain rates, the lateral creep strains rate is reduced at each stage of creep. The same methodology for determining the axial strain rates and adapted for axial strains is repeated for lateral strain. Below the onset of dilation threshold, the lateral creep strain rate is insignificant and is comparable to the variation in axial strain rates ($2.6E-06$ %/min). Beyond the onset of dilation threshold, during the second and third stages of creep, an order of magnitude increase in the creep strain occurs by an order of magnitude. On the other hand, the difference in lateral strain rates in the second and third stages is insignificant and in the order of magnitude of 10^{-5} %/min. A comparison of the axial and the lateral strain responses suggests that the variation in lateral strain values and its corresponding rates are more

significant than those of axial strain values and corresponding rates for locked sands.

4.6.3.3 Volumetric creep strain

Figure 4.14 shows the relationship between the elapsed time and the volumetric strains. The volumetric strains indicate dilation during all creep stages. The volumetric strains initially increase dramatically followed by a reduction in the volumetric strain rates and eventually attain a plateau. The volumetric strains are negligible only in the first stage of creep. In the second and third stage of creep, the volumetric strains are more significant than those in the first stage and are increased by an order of magnitude as shown in the figure.

4.6.4 Implications of creep test

The results clearly indicate that creep resulted in an increase in the uniaxial compressive strength due to ageing effects. In the first stage of creep, the geomaterials behaved in an elastic-plastic manner (with no reduction in the creep stress level). On the other hand, the second and third stages showed a reduction in the creep stress level (strain weakening). Although the reduction in strength was insignificant (corresponding to 10-12% of the original creep stress level), it signified a variation in response above the onset of dilation threshold.

The tested geomaterials exhibited all stages of creep in the axial and the lateral directions but no creep failure (tertiary creep) was observed. The variation in the axial and lateral strain rates was insignificant below the onset of dilation (the first creep stage). The axial and lateral strain rates were shown to be comparable in the first stage of creep. Beyond this stage, the lateral strain rates increased by an order of magnitude; however, the variation in the axial strain rates became double. In general, no major variation occurred in the axial or lateral strain rates, once the onset of dilation threshold was attained. Before and after the onset of dilation threshold was reached, the volumetric strains were dilative. However, they were negligible below this threshold and increased by an order of magnitude in the second and third stages.

In fact, the dilative nature of locked sand under constant loading conditions implies the dominance of lateral damage of its interlocked fabric and its propagation in the axial direction. In the first stage of creep, the volumetric strains are negligible and indicate the existence of minor or sparse fabric unlocking (damage) as compared to the propagating and newly formed unlocked elements in the second and the third stages as also manifested by the reduction in load-carrying capacity and the increase in volumetric strain rates. Beyond the onset of dilation, the dilative volumetric strains significantly increased. This implicitly implies the propagation of unstable damage, whereas below the onset of dilation, negligible volumetric strains were

recorded. Based on the preceding discussion, the onset of dilation can be considered to be an indicator of the long-term strength of locked sands when tested under uniaxial compression test conditions. However, additional work should be carried out to generalize such conclusion under higher confining pressure.

4.7 Direct shear box (DSB) results at low normal stress

The short- and long-term experimental results from uniaxial compression tests showed that the onset of dilation can be regarded as an excellent indicator of both short- and long-term strengths. In this section, the direct shear test results for locked sand specimens tested under low normal stresses will be presented and discussed. The DSB test was carried out to reveal the strength that can be mobilized at low stress levels. To achieve this objective, four specimens in a dry state were prepared to be tested under normal stress ranges of 10, 100, 300, and 500 kPa, respectively.

The direct shear tests were conducted in an automated direct shear apparatus. Figure 4.15 presents an illustration of the various components of the apparatus is shown in. In this type of equipment, cylindrical samples can be used and fitted within the two halves of a box with an internal cylindrical hole of a diameter of 63.2 mm and a height of 43.12 mm. The lower part of the box can be displaced with respect to the upper one in either direction. The shear load, axial displacement, and horizontal displacements were recorded at each 10s intervals

and logged into a computer connected to the apparatus. Intact samples were sheared at a strain rate of 0.0025 mm/min.

4.7.1 Direct shear results

4.7.1.1 Shear stress vs. horizontal deformation

Figure 4.16 shows the nominal shear stress (τ_{xy}) vs. the horizontal deformation (δ_x) relationships at normal stress ranges of 10-500 kPa. The samples failed at an average horizontal deformation of 0.8 mm. After the maximum shear stress was reached, a sudden reduction in strength occurred, reflecting the material's extreme brittleness. However, this brittleness decreased with the increase in the normal stresses. Up to a normal stress of 300 kPa, the maximum mobilized strength was only slightly dependent of the normal stress and was presumably governed by the interlocking strength component. For normal stresses greater than 300 kPa, the material exhibited a significant increase in the maximum shear stresses, and, hence, marked the destruction of the interlocking and the initiation of the full mobilization of the fictional resistance.

4.7.1.2 Vertical deformation vs. horizontal deformation

Figure 4.17 shows the vertical deformation (δ_y) - horizontal deformation (δ_x) relationships at a normal stress range of 10-500 kPa, revealing that samples showed dilation throughout the shearing process. An initial increase in the dilation rate occurred in pre-peak regime, followed by a reduction in the

dilation rate in the post-peak regime. The dilation rates continued to increase, and, hence, the samples did not attain the critical state in the post-peak regime. This result was due to the rotation of the loading cap as well the persistence of a residual interlocking between the particles. At 10 kPa, the sample exhibited extreme dilation before attaining the peak, followed by a reduction in the dilation rate at the peak, whereas, for the range of normal stresses of 100-500 kPa, the maximum rate of dilation was coincident with the peak strength.

4.7.1.2 *Failure envelope at low normal stress*

Figure 4.18 presents a plot of the peak strength for all ranges of normal stresses. The failure envelope is bilinear and can be divided into two portions. The first portion is a linear envelope with an intercept of 287 kPa and a friction angle of 7° up to a normal stress of 300 kPa. The second portion can be approximated by a friction angle of 46° at higher normal stresses.

4.7.1.3 *Corrected shear stress vs. horizontal deformation*

The shear strength of granular materials is usually attributed to the combined effect of friction and dilation. In locked sands, the interlocking between grains results in a third component. In order to quantify the interlocking strength component, the integration of the dilatancy component is necessary. Once this component is integrated out, one can quantify the contribution of the interlocking component and the frictional resistance. The corrected shear

stresses to estimate the interlocking contribution can be calculated by using the following equation by integrating out the contribution of the dilation and the basic friction (Taylor 1948):

$$[4.9] \tau_c = \tau - \sigma_n \tan \psi - \sigma_n \tan \phi,$$

where τ is the total shear stress; τ_c is the corrected shear stress; σ_n is the normal stress; and $\psi = \text{atan}(-\delta y / \delta x)$ is the dilation angle.

Figure 4.19 shows the original shear stress-horizontal deformation curve and the corrected one for an intact sheared sample at a normal stress at 10 kPa. Provided the shearing resistance is dominated by the dilation and basic friction angle as in dense sands (e.g., Rowe 1962; Taylor 1948), the corrected shear stresses should be nil. However, as shown in Figure 4.19, the corrected shear stresses are identical to the original ones. Hence, the maximum shear stresses for locked sand cannot be expressed only as a function of the dilation and basic friction angle as they can be expressed in dense sands. This finding indicates that the interlocking component dominates the shearing resistance of intact locked sand samples at low normal stresses.

4.7.1.4 Mobilized strength under DSB

In a direct shear test, the principal stresses continuously rotate with the progression of shearing. At low normal stress levels, the minor principal

stress can be in tension, and failure can be due to tension (Lajtai 1969). This tension results in a progressive reduction in the interlocking component during the course of shearing.

In the uniaxial compression tests, it was found that the damage initiation stress level was around 510 kPa. It was argued that fabric unlocking commenced at this stress level and was manifested through initiation of the lateral strain nonlinearity. In direct shear tests, the initial portion of the failure envelope was approximated by a friction angle (ϕ) of 7° and an intercept of 287 kPa. The deviator stress ($\sigma_1 - \sigma_3$) can be expressed as

$$[4.10] \quad \sigma_1 - \sigma_3 = \frac{2\tau_{xy}}{\cos(\phi)}$$

Substituting the values in equation [4.10] renders a deviator stress level of 556 kPa, which is almost equivalent to the damage initiation stress level (510 kPa) estimated from the uniaxial compression test. Hence, the minimum mobilized strength in direct shear box tests can be taken as the damage initiation stress level, as estimated from uniaxial compression test results. This stress level is dominated by the interlocking component with a negligible mobilization of friction, as previously shown in uniaxial stress conditions.

4.8 Summary and conclusions

In this chapter, an experimental investigation was carried out on locked sand from the McMurray area. Local axial and lateral deformations, for the tested specimens under uniaxial compression test conditions, were monitored by using miniature LVDTs. The focus was on identifying the damage thresholds and the partitioning strength parameters with plastic strain. Finally, the reported columnar structures for locked sands were captured by altering the tested specimen's configuration and testing conditions. The existence of such columns due to the formation of dilation bands was substantiated by the theoretical consideration of the localization theory for porous rocks. A multistage creep test was also carried out to reveal the relationship between the onset of dilation and long-term strength of locked sand. Finally, the experimental results of DSB tests were presented and discussed to reveal the mobilized strength in DSB at low normal stress and its relation to the damage thresholds estimated from uniaxial compression test.

The results revealed that the exceptionally high uniaxial strength and axial stiffness were the by-products of the particles' angularity and the interlocking between particle contacts, which enhanced the ability to resist moments at the contacts. Stress-strain curves were partitioned in the pre-peak regime based on the local axial and lateral strains' response to describe the evolution of failure: the closure of cracks and pores, the initiation of unlocking, the onset of localization, and the coalescence and unstable propagation of damaged sites. The onset of unlocking (the damage initiation threshold) occurred at 28 % of the

peak strength and, according to our observations, was coincident with the crack closure strength threshold. The onset of localization and dilation occurred at 55% and 67% of the peak, respectively. The partitioning of the strength components with the plastic strain showed that the interlocking and frictional strength components were not simultaneously mobilized. The interlocking component degraded with the plastic strain, resulting in the progressive mobilization of frictional resistance. The interlocking component was consumed in the pre-peak regime and leveled off before attaining peak strength. At the peak strength, the resistance resulted from the frictional contribution (the basic friction and dilation) with less contribution from the interlocked component (17% of the peak strength) due to its consumption during the unlocking process. At the onset of dilation, 50% of the mobilized resistance was attributable to the interlocking strength component. The remaining resistance reflected the basic frictional resistance's contribution.

Columnar structures were captured by creating a relatively homogeneous strain field to hinder the onset of localization and to allow for complete formation of these columns. The pattern of failure is best described as the formation of vertical slabs/walls in the direction of the axial direction as a result of uniform unlocking through the lateral expansion and formation of dilation bands. The experimental results, compared to those in theoretical studies of failure modes for porous brittle rocks, revealed that the dilation bands were the dominant mode of deformation (a predominantly open-mode of deformation propagating in the

direction of the major principal stress), and, hence, were responsible for such columnar failure patterns.

Under constant stress loading conditions, the material exhibited a creep potential that became more significant beyond the onset of dilation. Creep lateral strains and their rates were more significant than axial strains. Creep lateral strain rates were one order of magnitude greater than axial strain rates beyond the onset of dilation. Creep volumetric strains were dilative at all stages of the creep test. This result supports the premise that significant lateral unlocking occurs and its propagation in the axial direction during creep stages. Hence, the onset of dilation was recommended to be an indicator of the long-term strength of such a formation, based on the conducted creep tests.

The results of a DSB testing indicated that the mobilized strength under low normal strength was independent of the normal stress levels. The mobilized strength was dominated by the interlocking strength component and was equivalent to the damage initiation strength threshold estimated from the uniaxial compression test results.

Based on our experimental results, in boundary value problems constructed in or on locked sands, the damage coalescence and initiation thresholds are recommended to be the upper and lower strength boundaries. However, more work is needed to confirm such conclusion under higher confining pressures.

Table 4.1 Elastic deformation parameters and damage thresholds.

Elastic parameters	Tangential axial stiffness (E= 0.85 GPa)
	Poisson's ratio ($\nu=0.2$)
Damage thresholds	Crack closure threshold ($\sigma_{cc}= 510$ kPa)
	Damage initiation threshold ($\sigma_{di} = 510$ kPa)
	Damage localization threshold ($\sigma_{dl} = 1000$ kPa)
	Damage coalescence threshold ($\sigma_{dc} = 1200$ kPa)

Table 4.2 Summary of testing steps and results.

Stage	q (kPa)	Axial strain rate (%/min)	Lateral strain rate (%/min)	Creep volumetric strain (%)
1	750	2.46 E-6	2.6E-06	-0.003
2	1500	4.7 E-6	1.50E-05	-0.028
3	2000	4.0E-6	1.27-05	-0.08

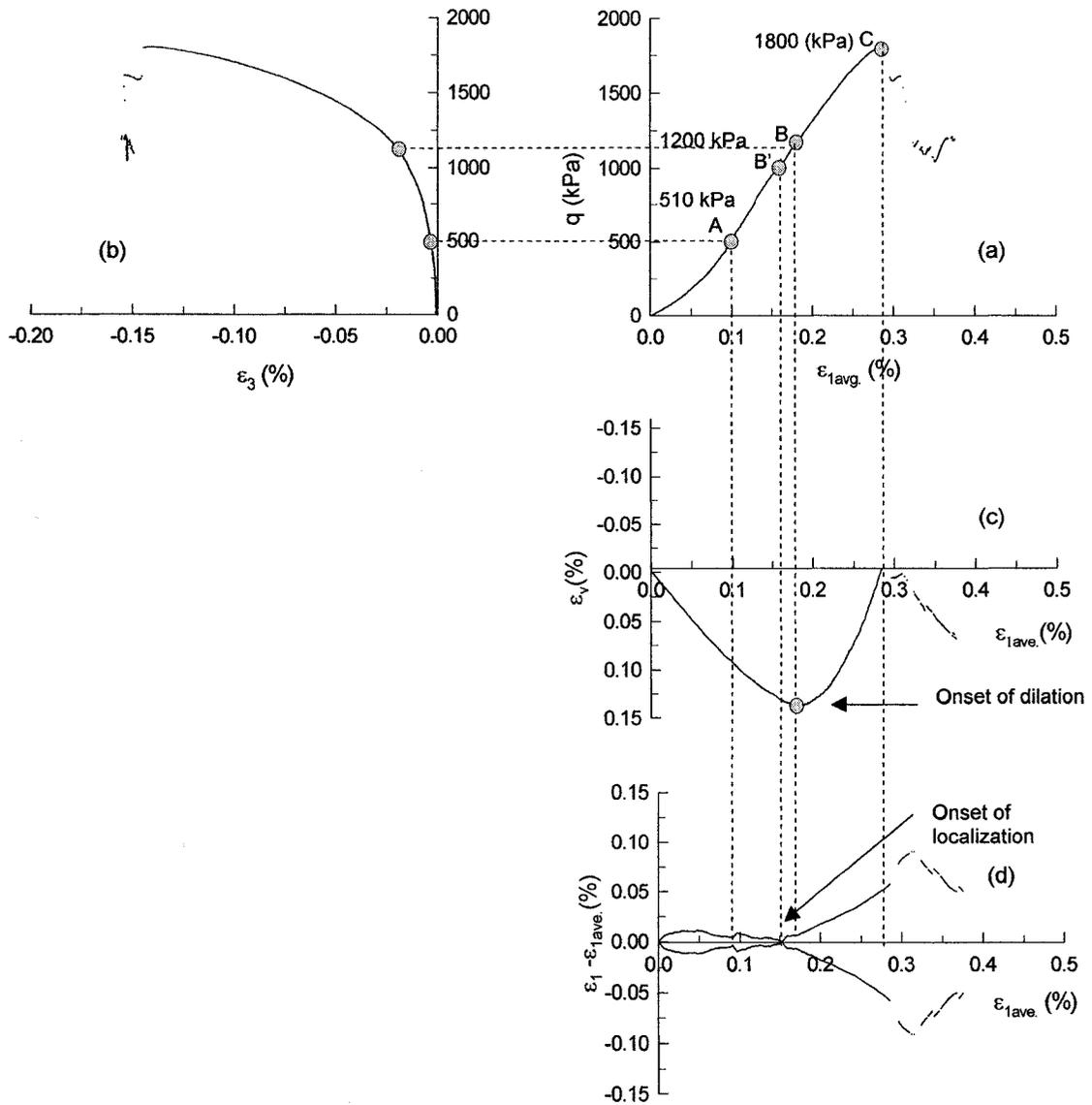


Figure 4.1 Experimental results illustrating variation of deviator stress, axial, lateral and volumetric strains during shear evolution.

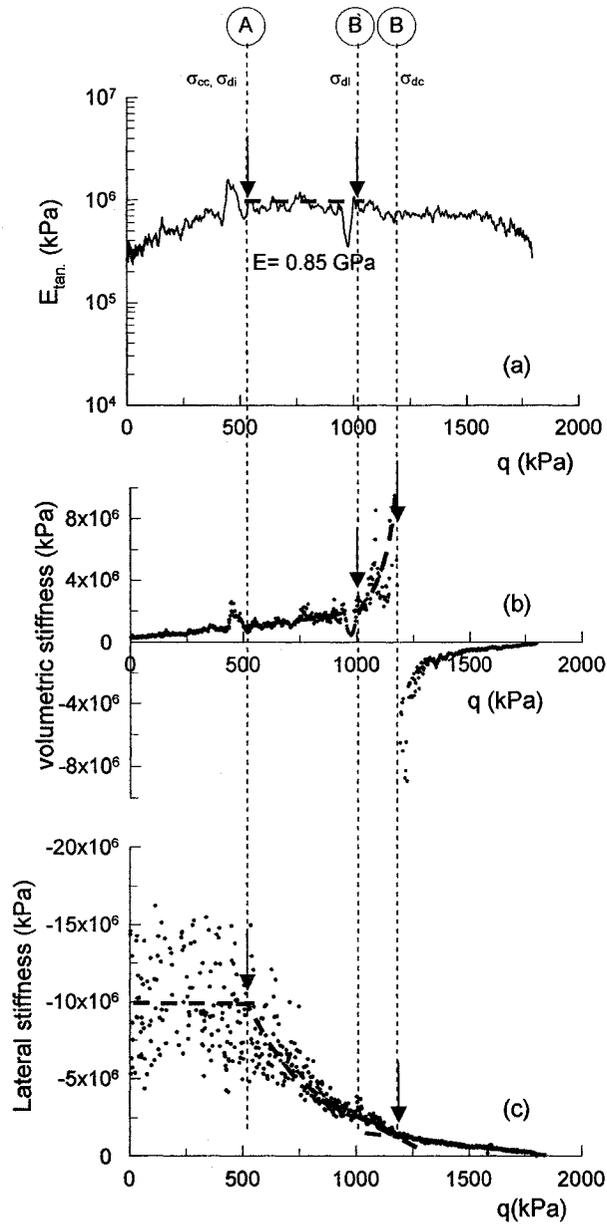


Figure 4.2 Variation in tangential axial stiffness, volumetric stiffness, and lateral stiffness during shear evolution.

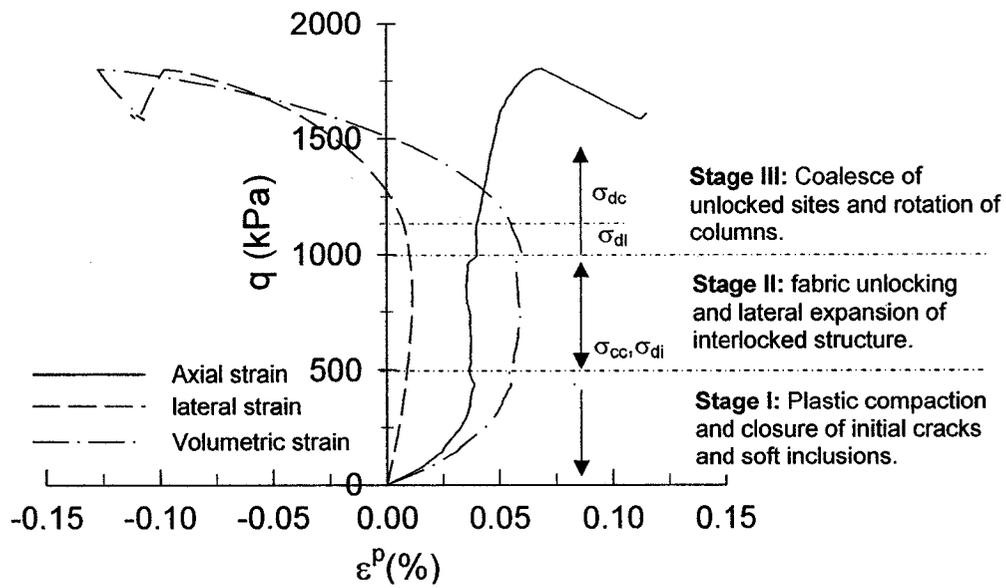
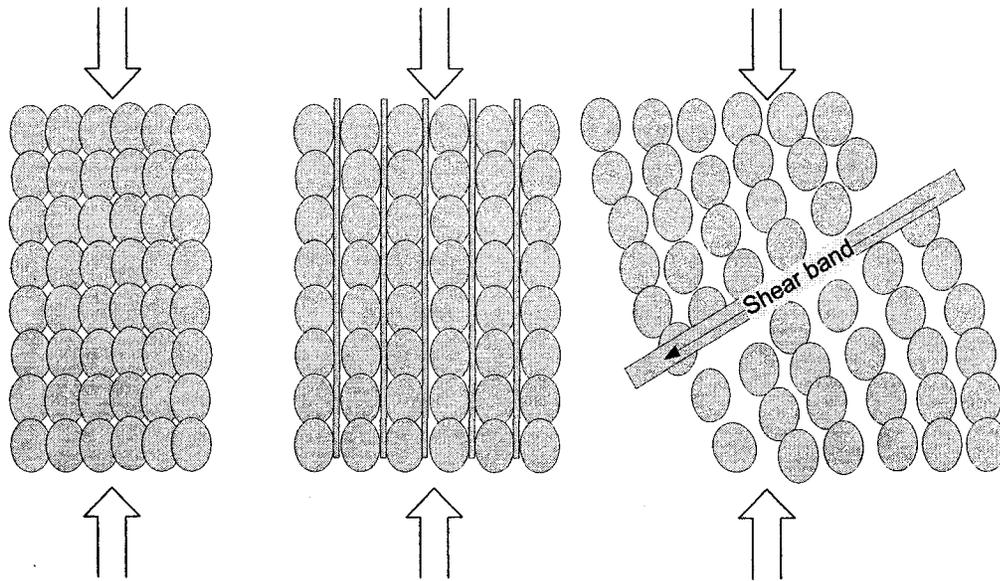
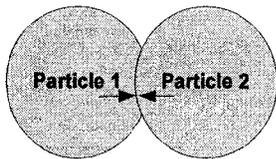


Figure 4.3 Relationship between the deviator stress, plastic axial, lateral, and volumetric strains.

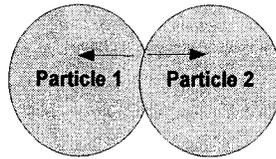


Interlocking at contacts



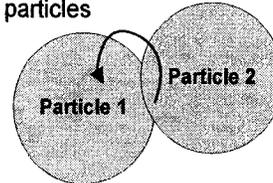
(I) Initial configuration and void flattening

Fabric unlocking



(II) Lateral expansion and formation of vertical columns to allow for particle arrangement

Dilation and rotation of particles



(III) Buckling of columns and formation of predominant shear band

Figure 4.4 Mechanism of failure in cylindrical locked sand specimen (not to scale).

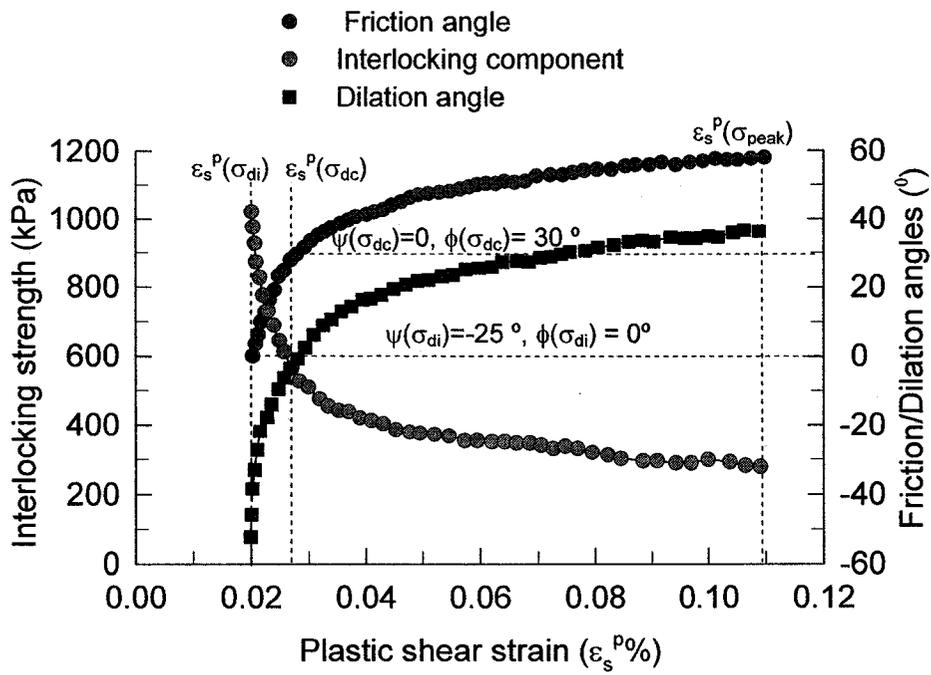


Figure 4.5 Mobilized strength components with plastic shear strain in terms of interlocking component, friction angle, and dilation angle.

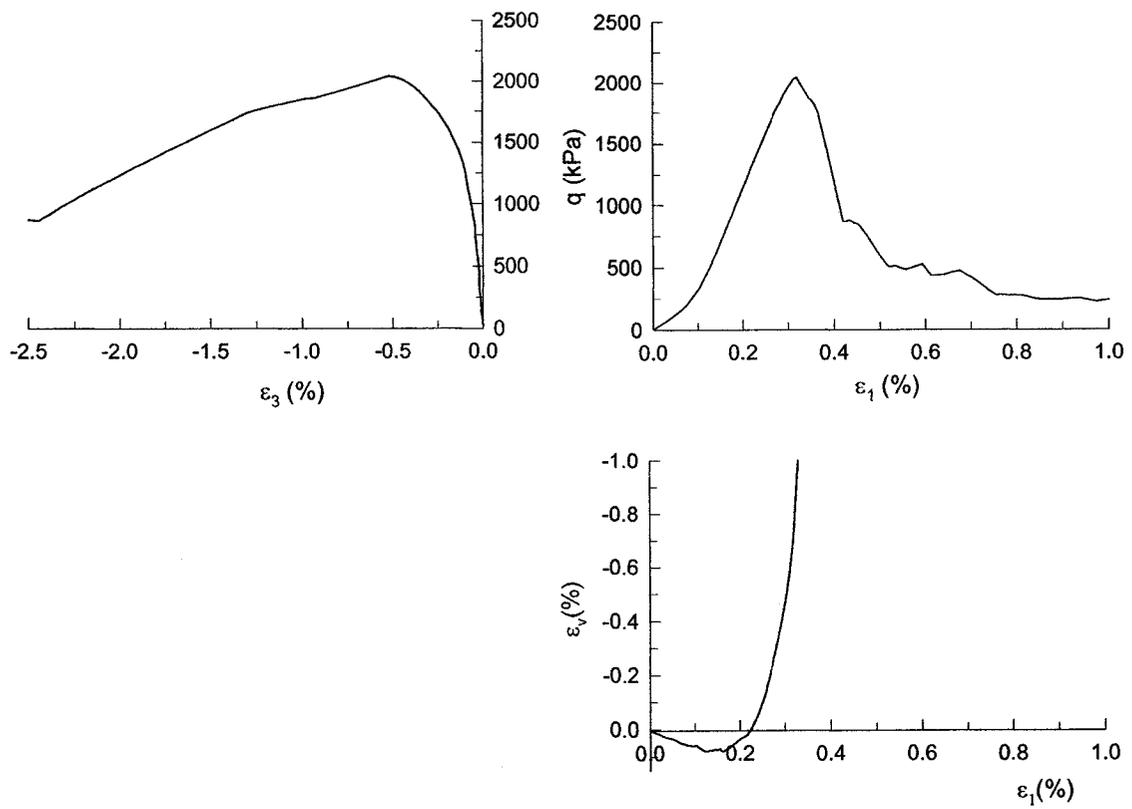
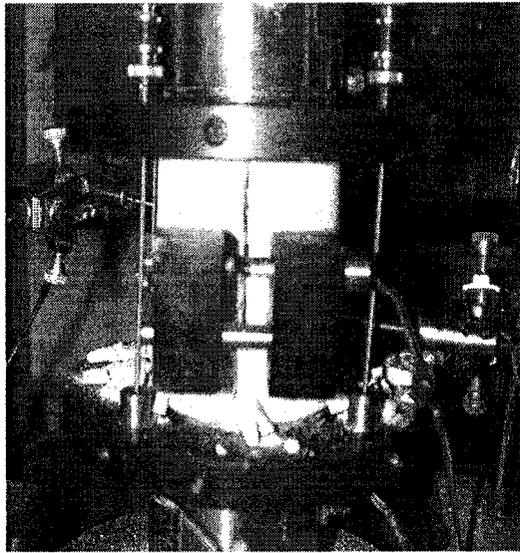
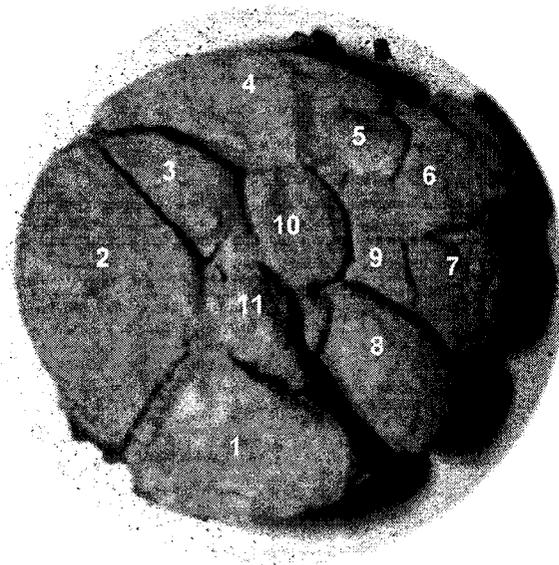


Figure 4.6 Stress-strain and volumetric strain curves for a specimen with a slenderness ratio of 1:1.



(a)



(b)

Figure 4.7 Formation of axially extended columnar structures under homogeneous strain field (observe curvature at columns interface): (a) longitudinal view of failure pattern; and (b) cross sectional view.

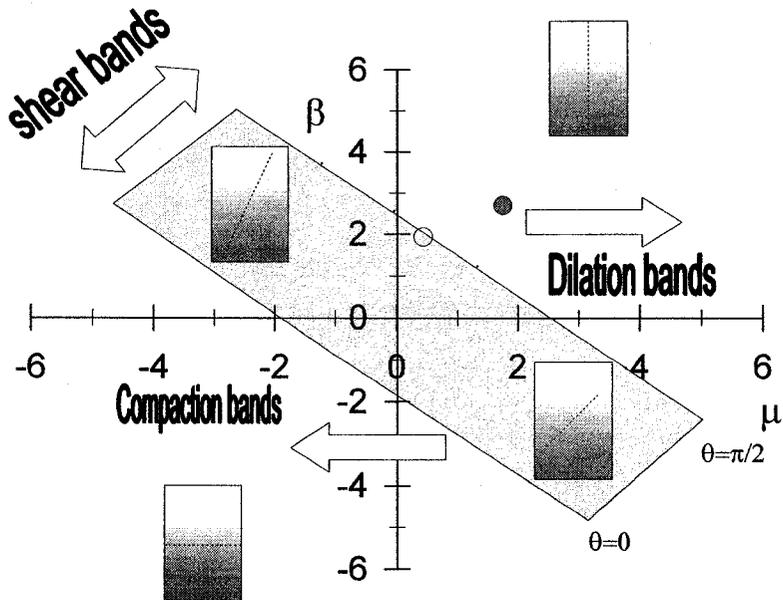
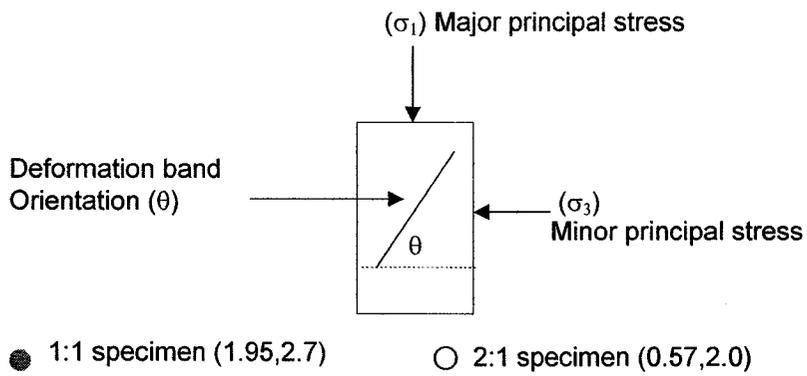


Figure 4.8 Deformation bands boundaries and orientation for axisymmetrical compression stress path derived from bifurcation criterion for porous rocks (Poisson's ratio $\nu=0.2$).

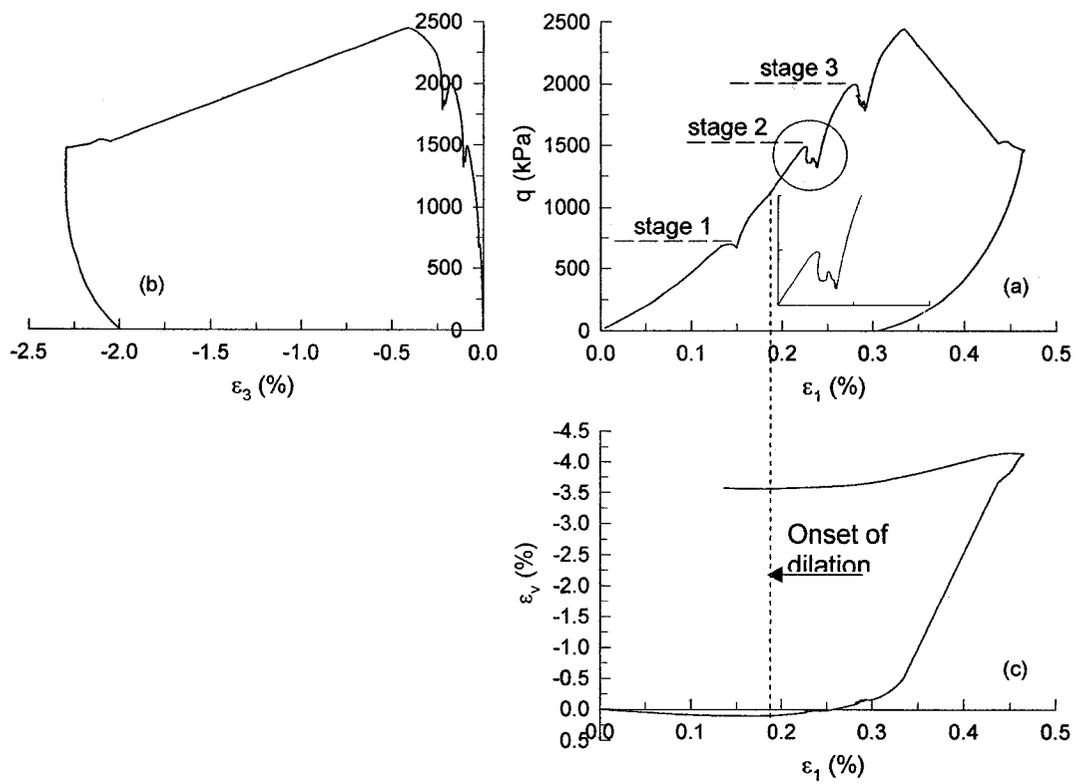


Figure 4.9 Stress-strain and volumetric strain curves during creep test.

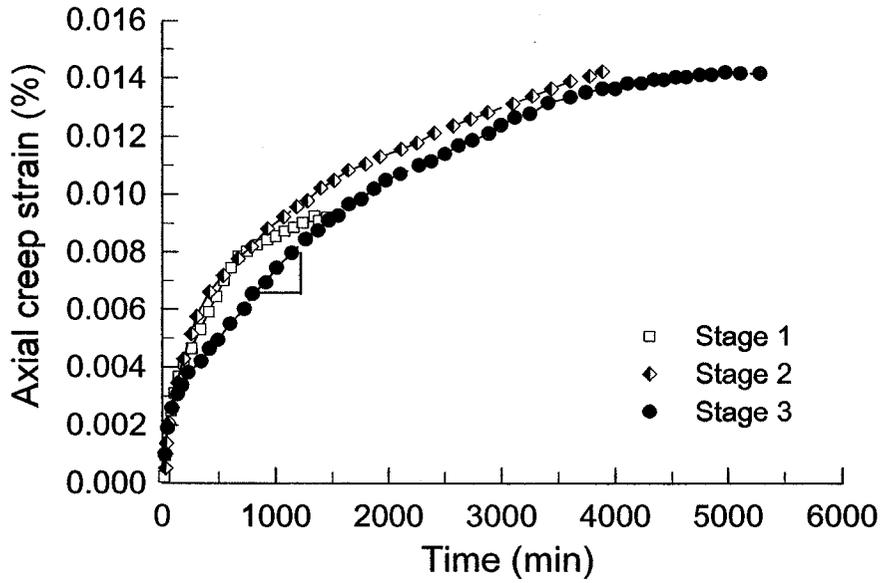


Figure 4.10 Variation of axial strain with time during creep stages.

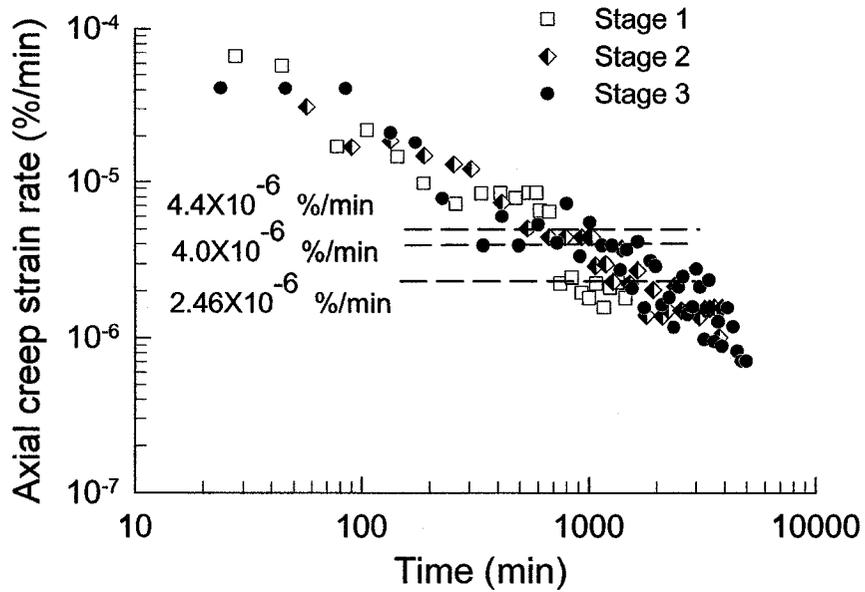


Figure 4.11 Variation of axial strain rates during creep stages.

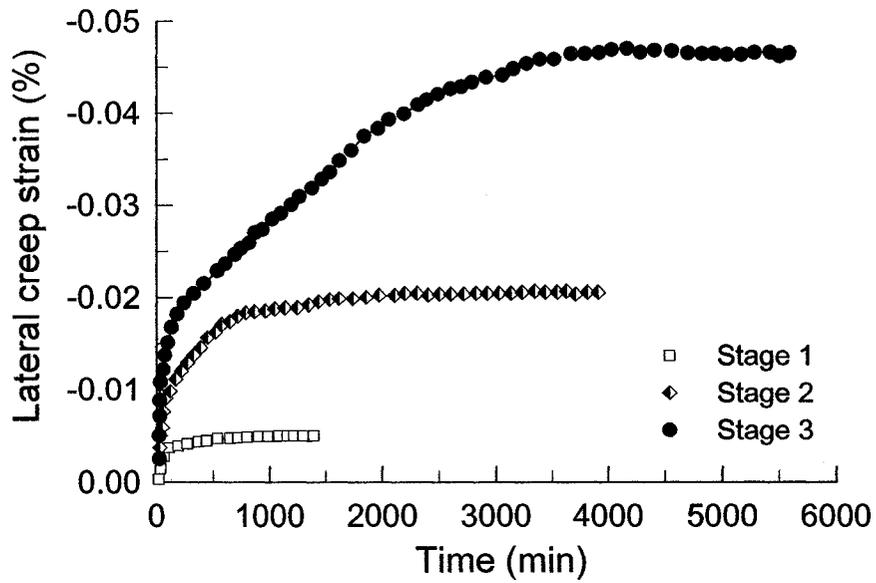


Figure 4.12 Variation of lateral strain with time during creep stages.

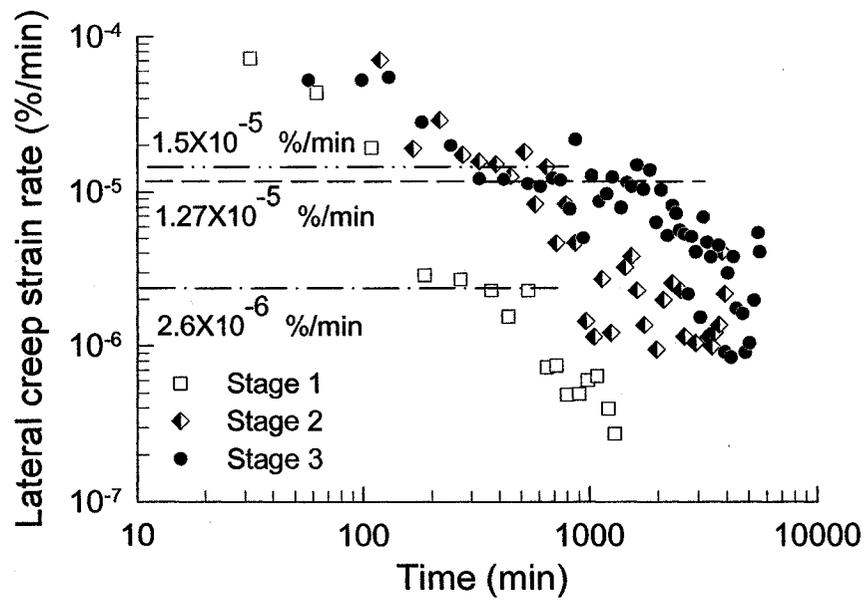


Figure 4.13 Variation of lateral strain rates during creep stages.

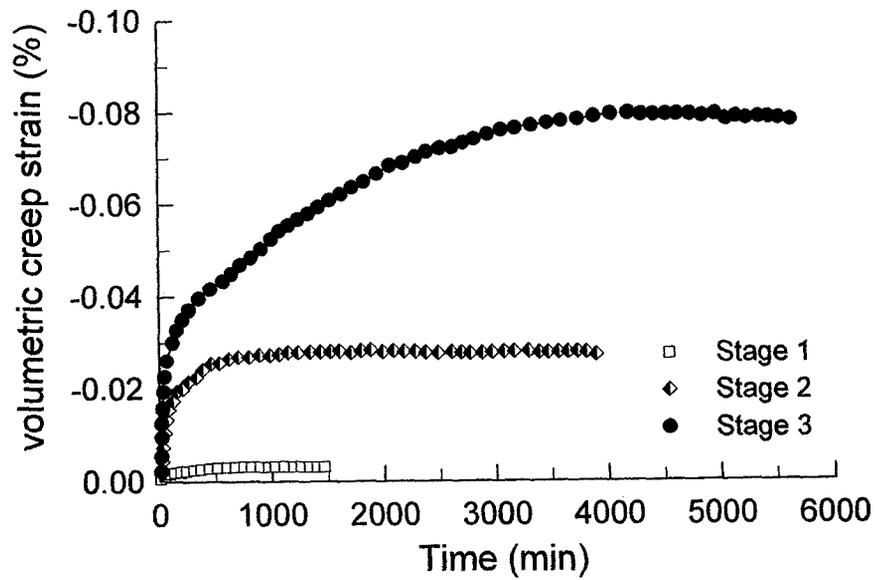


Figure 4.14 Variation of volumetric strains during creep stages.

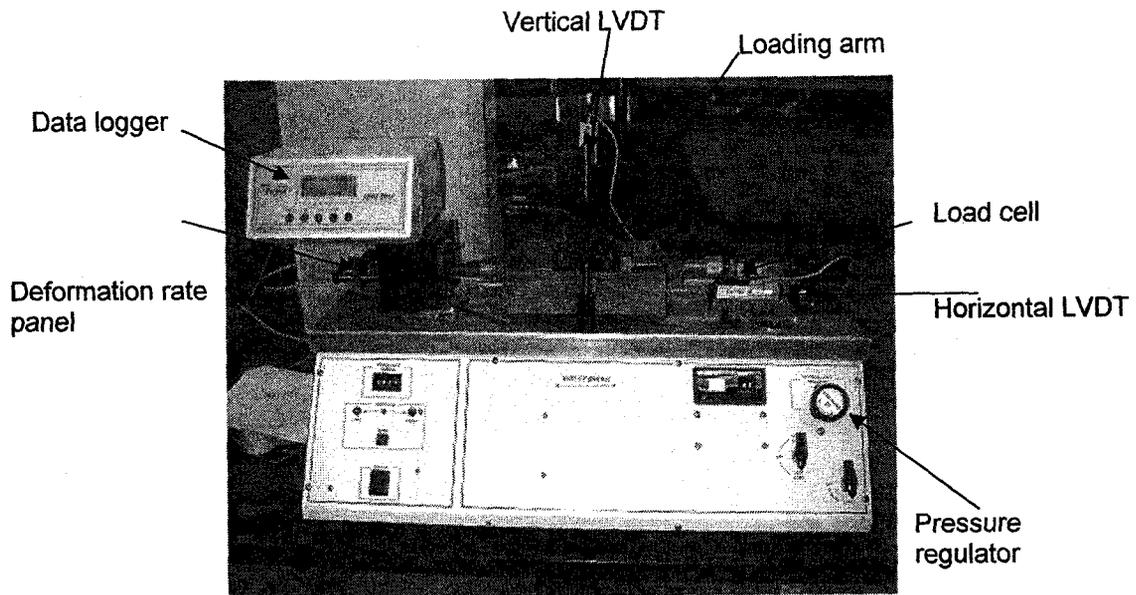


Figure 4.15 Direct shear test apparatus and its components.

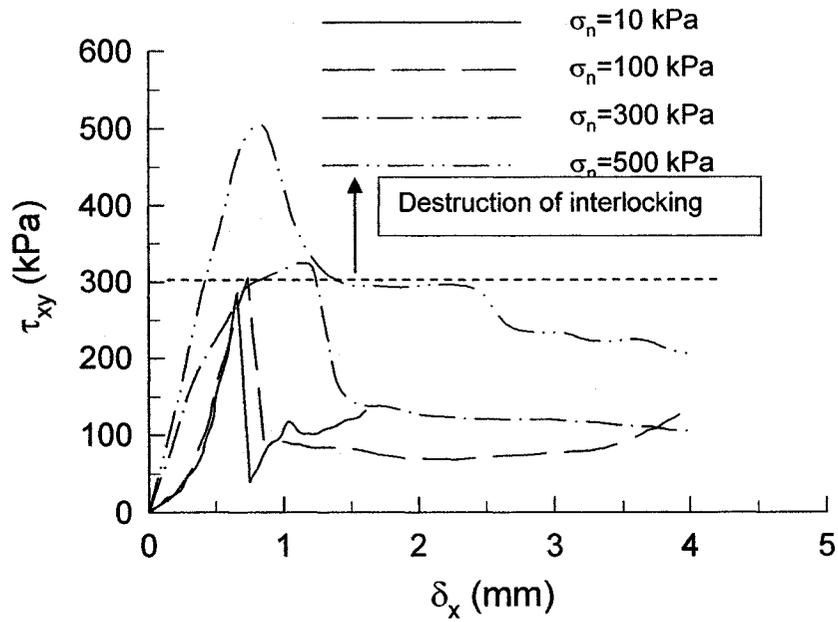


Figure 4.16 Shear stress vs. horizontal deformation in DSB.

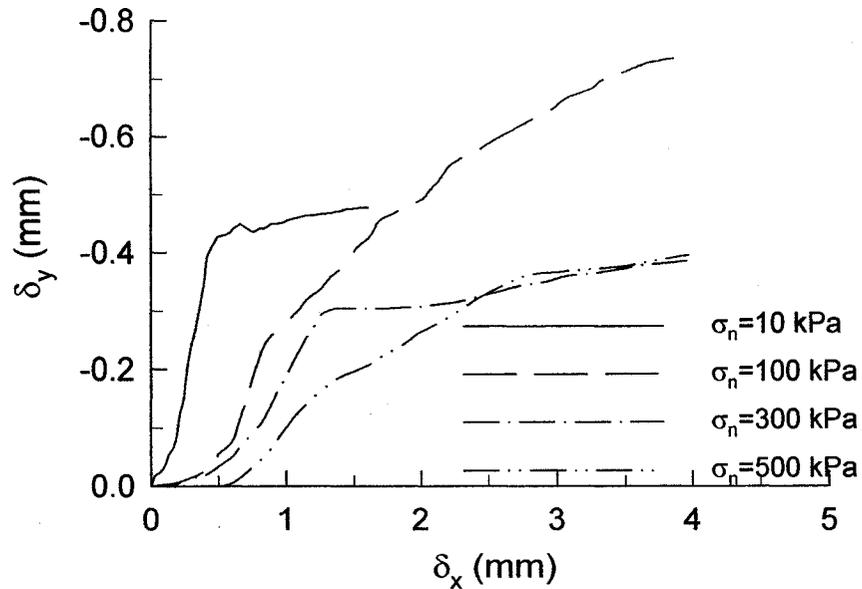


Figure 4.17 Vertical deformation vs. horizontal deformation in DSB.

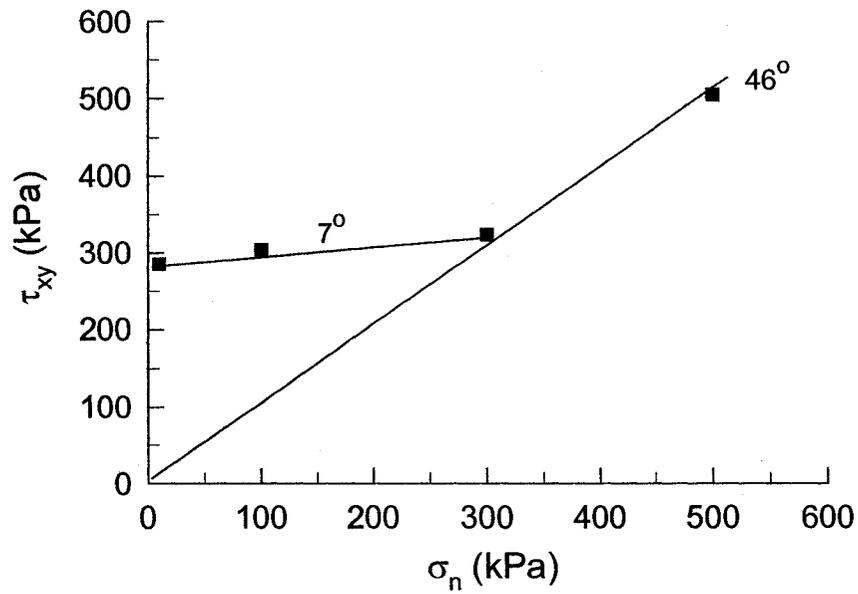


Figure 4.18 Failure envelope for locked sand from direct shear tests.

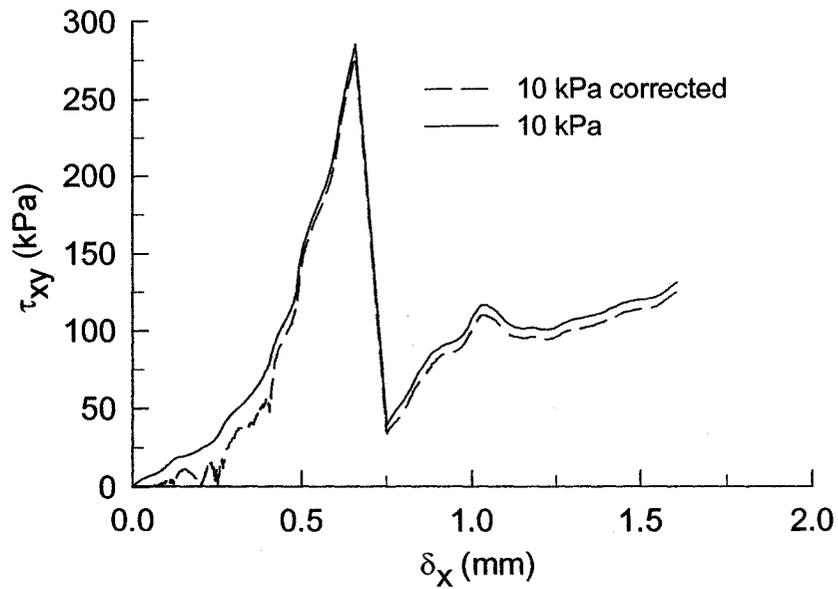


Figure 4.19 Corrected shear stresses for a specimen tested under a normal stress of 10 kPa.

CHAPTER 5: MOBILIZED STRAIN-DEPENDENT STRENGTH PARAMETERS FOR TWO NATURAL STIFF CLAYS

5.1 Introduction

In the previous chapter, the experimental analysis of a quasi-frictional material with interlocked fabric (locked sand) test results under uniaxial compression conditions revealed the non-simultaneous mobilization of strength parameters (friction and interlocking) and the degradation of the interlocked fabric with strain. In this chapter, this finding is used to interpret the mechanical behaviour of various bonded geomaterials at different confining pressures and to capture the transition between soils (frictional materials) and hard rocks (bonded materials) based on the mutual contribution of the strength parameters to the overall shearing resistance.

To achieve this objective, stress strain curves for two natural stiff clays (bonded quasi-frictional materials) were analyzed to reveal the contribution of the strength parameters at different strain levels by using a simplified partitioning approach. Shear strength parameters are partitioned in terms of the cohesion/bonding and the friction angle in Mohr Coulomb (M-C) space for the two clays, and in terms of normalized friction and bonding parameters in q-p space for Nanticoke clay as it produces the extreme case of bonding brittleness distinguished in most bonded geomaterials. A parametric study is carried out to reveal the similarity in the behaviour of different bonded geomaterials and to capture the transition between soils and rocks according to the mutual contribution of the bonding and frictional

resistances components to the overall resistance. In the rest of the chapter, the terms “bond” and “cohesion” are interchangeably employed to express resistance in excess of the frictional resistance, i.e., the combined effect of dilation and basic frictional resistance. In addition, all results are generally expressed in terms of effective stresses so that there is no differentiation between primed and unprimed parameters unless otherwise stated.

5.2 Historical development of strength parameters in clays

Deformations rather than a soil's limiting strength govern the design of many structures. The limiting soil strength is often taken into consideration by using allowable strengths, which correspond to small but arbitrary selected strains in the stress strain curve obtained from laboratory tests. On this basis, two classes of strength parameters can be distinguished (Bishop 1971). The first arises from the Mohr-Coulomb failure hypothesis (M-C), which is widely applied in soil mechanics and assumes the simultaneous mobilization of strength parameters. The second class emerges from the basic shear strength properties of soils and which is not readily applicable or observable in laboratory tests but helps to enhance our understanding and provide a physical basis for soil behavior.

Hvorslev (1960) carried out the first investigation into the physical meaning of strength parameters. He showed that the shear strength of remolded saturated clays could be partitioned into two components regardless of their history (their consolidation state): (a) effective cohesion, which is a function of void ratio and (b) effective friction, which is a function of effective stresses. Gibson (1953)

pointed out that the effect of volume change should be accounted for when estimating the effective strength parameters. Nonetheless, no significant variations in parameters were observed after correction for remolded clays. Lambe (1960) proposed a mechanistic picture for the mobilization of strength parameters, arguing that the strength is comprised of three components which were not mobilized simultaneously: the cohesion, the basic friction, and the dilatancy component. According to his argument, cohesion is mobilized at low strain and does not contribute to the peak resistance. Schmertmann and Osterberg (1960) carried out the first experimental program to partition shear strength parameters. Schmertmann (1963) did not agree with Hvorslev (1960) that the variation in void ratio represents the variation in structure (fabric and bonding) in natural soils. In contrast, Schmertmann (1963) proposed that the shear strength parameters should be expressed as a function of the strain rather than the void ratio. He found that the shear strength parameters were not mobilized simultaneously (cohesion was mobilized at low strain while friction required sufficient strain to be mobilized).

De Mello (1988) showed the difference between the two classes of shear strength criteria, as indicated by Bishop (1971), by plotting the stress- strain curves on the same figure and correlating the shear strength at different strain levels (the second criterion) as compared to the maximum deviator stresses (first criterion). He showed that the available strength at working strain levels was much lower than it was according to the maximum deviator stress concept. Even

if a factor of safety was applied to the maximum resistance, the available limiting strength would not be representative of the actual mobilized strength. De Mello's (1988) study indicates that the maximum shear strength criterion should not be used and that strain should be incorporated as an independent parameter. Conlon (1966) also pointed out that the classical Mohr Coulomb criterion could not be used to predict the failure mechanism for bonded stiff clays in slope stability problems. Conlon (1966) showed that using the shear strength parameters, as a function of strain, is more beneficial for understanding and interpreting the failure mechanism of natural stiff clays. According to his view, the stability of slopes in stiff clays depends on the degree of friction mobilization relative to the degradation of the cohesive component, which cannot be captured by assigning constant parameters. Therefore, it can be concluded that the strength parameters need to be partitioned according to the strain to capture the failure mechanism in natural stiff clays that exhibit strain weakening due to bond degradation.

5.3 Partitioning of strain-dependent strength parameters in Mohr-Coulomb space

Several investigators have referred to plastic strain as the most significant parameter for capturing the changes in the microstructure from a micromechanical viewpoint (e.g., Vardoulakis and Sulem 1995; Mroz and Rodzik 1995; Sulem et al. 1999). To partition the shear strength parameters as a

function of the plastic shear strain, several assumptions must be made. The main assumptions and limitations to be considered are:

1. The Mohr-Coulomb criterion is valid at any strain level.
2. The cohesion/bonding component captures the integrity of soil bonding at various strain levels.
3. The frictional resistance represents the change in the fabric and comprises both the basic friction angle (derived from mineral texture and roughness) and the dilatancy component (the volume change component).
4. The mobilized strain-dependent parameters are considered to be fundamental parameters for a particular field stress path, and their distributions are influenced by laboratory boundary conditions and measurement techniques (local vs. external measurements).
5. This approach is applied to stiff clays and assumes that negligible particle crushing effects occurs under low stress levels and that the material deforms plastically (the elastic deformation is negligible compared to the plastic deformation).

Figure 5.1 shows a schematic representation of a typical stress- strain curve for drained triaxial compression results. The results are represented in terms of the plastic shear strain (ϵ^p) (the material was assumed to deform plastically, i.e., plastic deformation is dominant). Consequently, one can assume that the total strain increment is equivalent to the plastic strain increment as elastic

strains effects have been shown to be insignificant (Leroueil 1997). In Figure 5.1, the dashed lines represent different levels of plastic strain. At each strain level (marked as $\varepsilon^p_1, \varepsilon^p_2, \varepsilon^p_3$), the intersection of the dashed line with the stress strain curves at different confining pressures ($\sigma_{31}, \sigma_{32}, \sigma_{33}$) renders three points (as shown in the figure) with different deviator stresses (for instance, points 1,2, and 3). At each point, the principal stresses are known, and, hence, a single envelope can be drawn by obtaining the best fit at this strain level (see Figure 5.2). For each envelope, the inclination of the best fit line (α) and the intersection with the ordinate (a) can be obtained.

According to the Mohr-Coulomb criterion, the fitted data points can be expressed as a linear relationship by using the following equation:

$$[5.1] \quad \frac{\sigma_1 - \sigma_3}{2} = c \cos \phi + \frac{\sigma_1 + \sigma_3}{2} \sin \phi,$$

where σ_1 and σ_3 are the major and minor principal stresses, respectively; c is the cohesion intercept; and ϕ is the friction angle.

The cohesion and friction angles can be then obtained by using the intercept (a) as well as the inclination ($\tan \alpha$) of the fitted linear relationship by using the following equations:

$$[5.2] \quad c = \frac{a}{\cos \alpha}, \text{ and}$$

$$[5.3] \quad \phi = \sin^{-1}(\tan \alpha).$$

The values of cohesion and friction can be then plotted with the plastic shear strain, as will be shown later. Hence, multiple yield functions can be constructed at any strain and can be generally written as follows (Vermeer and De Borst 1984)

$$[5.4] \quad f(\sigma, \varepsilon^p) = f(\sigma, \varepsilon^p) + f(\sigma_n \tan \phi, \varepsilon^p),$$

where $\varepsilon^p = 2/3 (\varepsilon_1^p - \varepsilon_3^p)$ for triaxial compression conditions; ε^p is the plastic shear strain; and σ_n is the normal stress.

5.4 Partitioning of strain-dependent strength parameters for two natural stiff clays

5.4.1 Pietrafitta clay

Callisto and Rampello (2002) carried out a series of drained triaxial compression tests on natural stiff Pietrafitta clay. Figure 5.3 presents the stress-strain curves at different confining pressures. As indicated in the previous sections, the shear strength parameters are to be calculated at different strain levels. Fitting lines, with a determination coefficient of at least 0.95, which is statistically acceptable, are indicated in Figure 5.4. Figure 5.5 shows the distribution of the cohesion and friction with the plastic shear strain,

revealing that both strength components are not mobilized simultaneously. The cohesive component is fully mobilized at zero plastic strain followed by degradation to the large strain value or residual value. The friction component is mobilized with the plastic strain to a maximum value and then degrades to a slightly lower value, which cannot be taken for granted as the residual value due to strain localization. However, as Adachi and Oka (1993) and Oka et al. (1998) stated, a reduction in load-carrying capacity cannot be attributed solely to the strain localization. A reduction in the post-peak regime is also a reflection of bond degradation. At any certain strain level, the available strength is a combination of both components, which cannot be described properly by using only two constants. The cohesive component is mobilized to its maximum value (54.1 kPa) up to a strain level of 2.1%, followed by a slight reduction of the component with strain. The reduction percentage of the cohesive component is almost 23%. On the other hand, the friction component attains its maximum value (15.3°) at a strain level of 4% and degrades with a percentage reduction of 41%. Figure 5.5 also shows that the percentage of reduction in the frictional resistance is higher than for cohesion. The reduction of the load-carrying capacity can be attributed to the reduction of both components and is a function of the confining pressure. At low confining pressures, the cohesive component is prevalent, and, hence, the reduction rate of the cohesive component governs the reduction of the strength. At high confining pressures, the reduction of the frictional resistance dictates the reduction of the strength. The stress-strain curves reveal that at

low confining pressures, the reduction of the deviator stresses is not significant due to the slight reduction of the cohesive component. Conversely, at moderate confining pressures, the reduction of deviator stresses is appreciable due to the significant reduction in frictional resistance. At higher confining pressures, the reduction of the frictional resistance is not as influential as at moderate confining pressures.

5.4.2 Nanticoke clay

Nanticoke clay is a stiff, moderately, overconsolidated clay from Southern Ontario. Lo (1972) studied the strength of Nanticoke clay at a strain rate of 0.05%/hr under several confining pressures (Figure 5.6). Figure 5.6 reveals that strain weakening (reduction from peak to residual values) decreases with the increase in confining pressure. Based on the results of the triaxial compression tests, the data were analyzed to obtain the multiple yield envelopes (see Figure 5.7). Figure 5.8 presents the distribution of the parameters with the plastic strain. Both components of strength are shown to be neither mobilized nor degraded simultaneously, as mentioned previously. The cohesive component is fully mobilized at zero plastic strain followed by full degradation (indicating the brittleness nature of this clay). The friction component is mobilized progressively with the plastic strain to a maximum value and then degrades to a slightly lower value. The cohesive component is mobilized to its maximum value (71 kPa) up to a strain level of 0.75%, followed by an abrupt reduction in the component to zero at a strain level of

2.0 %. The reduction percentage of the cohesive component is almost 100%. On the other hand, the friction component attains its maximum value (21°) at a strain level of 3% and degrades with a percentage reduction of 19%. Figure 5.8 also shows that the reduction in the cohesive component is more significant than that in the frictional component.

Bishop (1967) introduced a brittleness index to characterize the brittleness in the post-peak regime in terms of the peak and residual strength. Hajiabdolmajid (2001) introduced a strain-dependent brittleness index ($I_{B\varepsilon}$) to express the ability of brittle geomaterials to withstand deformation during the entire failure process (pre-peak to post peak) as well as their susceptibility to brittle fracture and progressive failure. The strain-dependent brittleness index is intended to capture the relative mobilization and degradation of the strength components and is expressed as follows:

$$[5.5] \quad I_{B\varepsilon} = \frac{\varepsilon_f^p - \varepsilon_c^p}{\varepsilon_c^p} \quad I_{B\varepsilon} \leq 0; \quad 0 \leq \varepsilon_f^p \leq \varepsilon_c^p$$

$$I_{B\varepsilon} > 0; \quad \varepsilon_f^p > \varepsilon_c^p,$$

where ε_f^p is the plastic shear strain at which frictional resistance is fully mobilized, and ε_c^p is the plastic shear strain required for bonding resistance to attain the residual value. This brittle clay has a strain-dependent brittleness index ($I_{B\varepsilon}$) of 1.4. This value expresses the lag in frictional strength

mobilization compared to rate of cohesion degradation. Hence, at low confining stresses and small plastic strains, the loss of strength will be significant. On the other hand, only a slight reduction in strength will occur at high confining pressures.

5.5 Bonding and frictional resistance in q-p space

Traditionally, the shear strength of geomaterials is expressed in terms of either M-C space or q-p space. In q (deviator stress) -p (mean effective stress) space, the ordinates are expressed as follows:

$$[5.6] \quad q = \frac{(\sigma_1 - \sigma_3)}{2}, p = \frac{(\sigma_1 + 2\sigma_3)}{3}.$$

The strength parameters are partitioned in q-p space for Nanticoke clay in order to interpret the parameters in terms of the critical state soil mechanics (Schofield and Wroth 1968). For purely frictional materials, the slope of the yield surface is described in terms of a frictional parameter at large strain, i.e., M_{crit} (the stress ratio at the critical state, where shear stress levels off with no variation in the volumetric strain rate), which is expressed as follows (Schofield and Wroth 1968):

$$[5.7] \quad M_{crit} = \frac{q}{p}.$$

For bonded geomaterials, the shear strength expression deviates from the previous equation. In general, the shear strength equation at any strain level can be expressed as follows:

$$[5.8] \quad q(M, p_t) = M(\varepsilon^P)[p + p_t(\varepsilon^P)],$$

where p_t is the tensile mean strength of bonds at any strain level and M is the slope of the yield surface and describes the frictional resistance. At low strain levels, the tensile strength is maintained due to the integrity of bonds' initial strength and is represented by p_{t0} . At large strain levels, the tensile strength degrades and disappears, reflecting the full consumption of the bonds during shear evolution. Consequently, the slope of the yield surface (M) progressively becomes equivalent to M_{crit} , and the shear strength equation becomes equivalent to equation [5.8].

Multiple yield envelopes in q - p space are estimated as in M - C space at different plastic strain levels in order to obtain the distribution of the parameters (M , p_t). These two parameters (M , p_t) are normalized with respect to the stress ratio at the critical state (which is assumed to be reached at large strain) and the initial tensile strength (the initial intersection of the envelope with the horizontal axis), respectively. The mathematical relation describing the evolution of the frictional

resistance in a normalized space for Nanticoke clay can be written as follows and is illustrated in Figure 5.9.

$$[5.9] \quad \frac{M}{M_{\text{crit}}} = 2.51 \exp(-0.177 \varepsilon^p) - 2.507 \exp(-0.8597 \varepsilon^p),$$

where M expresses the variation of frictional resistance with plastic strain levels. At large strain levels, this parameter reduces to the critical state parameter (M_{crit}) where bonding effects vanish.

The second expression is related to the tensile strength at various strain levels, which can be expressed as follows and is illustrated in Figure 5.10.

$$[5.10] \quad \frac{p_t}{p_{to}} = 1.05 \exp(-3.912 \varepsilon^p).$$

5.5.1 Incremental form of the strength equation in q - p space

The differentiation of equation [5.7] with respect to plastic shear strain renders the following equation:

$$[5.11] \quad \frac{\delta q}{\delta \varepsilon^p} = \left[p \frac{\delta M}{\delta \varepsilon^p} + \frac{\delta p}{\delta \varepsilon^p} + M \frac{\delta p_t}{\delta \varepsilon^p} + p_t \frac{\delta M}{\delta \varepsilon^p} \right].$$

Based on [5.9] and [5.10], the distribution of the parameters with plastic shear strain can be in general expressed as follows:

$$[5.12] \quad \frac{M}{M_{\text{crit}}} = c(\exp(-d\varepsilon^P)) - e(\exp(-f\varepsilon^P)),$$

$$[5.13] \quad \frac{p_t}{p_{to}} = a \exp(-b\varepsilon^P),$$

where a,b,c,d,e, and f are fitting parameters.

The differentiation of equations [5.12] and [5.13] with respect to the plastic strain increment leads to the following two expressions:

$$[5.14] \quad \frac{\delta M}{\delta \varepsilon^P} = M_{\text{crit}} \left[-cd(\exp(-d\varepsilon^P)) + ef(\exp(-f\varepsilon^P)) \right],$$

$$[5.15] \quad \frac{\delta p_t}{\delta \varepsilon^P} = -ab(\exp(-b\varepsilon^P)).$$

Substituting equations [5.12], [5.13], [5.14] and [5.15] into equation [5.11] and simplifying the equation renders the frictional and bonding contributions in incremental form as follows:

$$[5.16] \quad \delta q = \delta q_1 + \delta q_2,$$

where

$$\delta q_1 = \underbrace{(p' M_{crit.} (-cd(\exp(-d\varepsilon^P)) + ef(\exp(-f\varepsilon^P))) + M_{crit.} (c(\exp(-d\varepsilon^P)) - e(\exp(-f\varepsilon^P))))}_{\text{frictional component}} \delta p' \delta \varepsilon^P,$$

and

$$\delta q_2 = \underbrace{a M_{crit.} p_{to} (\exp(-b\varepsilon^P)) ((\exp(-d\varepsilon^P))(-cd - bc) + (\exp(-f\varepsilon^P))(eb + ef))}_{\text{bonding component}} \delta \varepsilon^P.$$

The above equation can be used to render an approximate estimate of the frictional and bonding resistances at various plastic strain levels as will be shown in the following section.

5.6 Parametric study of the bearing of confining stresses on stress strain curves

5.6.1 Contribution of frictional and bonding resistance at different confining stresses

Figure 5.11 shows the stress-strain curves based on the incremental forms of equation [5.16] to reveal the effect of confining stresses (10-150 kPa) on the behaviour of a bonded geomaterial when an initial tensile strength (p_{to}) of 500 kPa and a critical stress ratio ($M_{crit.}$) of 1.1 are assumed. The constant mean effective stress path, (constant p), is simulated to capture the effect of shearing on the stress-strain curves, i.e., to rule out volumetric effects due to the increase in mean normal stress. Figure 5.11 illustrates the frictional resistance, bonding resistance, and total resistance at various confining

stresses. The figure shows that the rate of frictional resistance mobilization with strain increases with the increase in confining stresses. At low confining stresses, the frictional resistance has minor effect on the peak resistance, and most of the available resistance results from the bonding resistance. The reduction in load-carrying capacity is due to the degradation of bonding, and the mobilized residual friction dominates the residual strength. At intermediate confining stresses the peak is due to both components; however, they are not the maximum components. A plateau is revealed, and two peaks characterize the stress-strain curves. The first is at low strain where bonding resistance starts to degrade followed by either strain hardening or strain weakening depending on the rate of frictional resistance mobilization relative to cohesive resistance degradation with strain. If full frictional resistance can be mobilized, strain hardening occurs and is accompanied with a second peak. The initial weakening is due to the degradation of the bonds while the second one is due to degradation of frictional resistance. Where strain hardening occurs, frictional resistance dominates peak strength with a negligible contribution from the bonding resistance. At high confining stresses, no observed plateau is observed and the frictional resistance mobilization rate is greater than bonding resistance degradation rate. These results are in agreement with Conlon's (1966) conceptual framework to explain slope failures in stiff clays and their mechanical behaviour at different confining pressures (see Figure 5.12).

5.6.2 Failure, bonding, and plateau envelopes

Figure 5.13 shows the peak resistance, bonding resistance (bonding contribution at peak), and plateau strength boundaries. The failure envelope is characterized by nonlinearity, and the bonding envelope can be approximated by an elliptical shape. Malandraki and Toll (2000) captured experimentally the elliptical shape of bonding resistance for artificially cemented soils under different stress paths. At low stress levels, the bonding resistance dictates the shearing resistance. This contribution progressively reduces so that the frictional resistance becomes dominant at high confining stresses. The extension of the failure envelope at high stress levels leads to zero strength at zero confining stress. This is a typical behaviour of sheared unbonded geomaterials, i.e., frictional materials. The point at which bonding resistance vanishes marks the transition from bonded to unbonded geomaterials, in which the transition from over-consolidation state to normally consolidation state in conventional Soil Mechanics terms. Based on the two boundaries, three zones can be identified. The first describes the behaviour of geomaterials, in which the bonding is the main contributor to strength (rocks). The third zone is typical of geomaterials where friction dominates the behaviour (soils). The middle zone is a transition zone where a quasi-frictional behaviour is manifested due to the mutual contribution of both components.

5.6.3 *Resemblance of stress-strain curves' shapes and failure envelopes of bonded geomaterials with different origin*

Conlon (1966) experimentally captured the same behaviour in terms of pre-peak wavy stress strain curves and nonlinear failure envelope for a very stiff sensitive clay deposit of fine-grained estuarine sediments in Canada, tested in a triaxial compression (see Figure 5.14). Lajtai (1974) also carried out direct shear tests on solid plaster with two voids to simulate the behaviour of jointed weak rocks and also captured the wavy stress strain curves and nonlinearity (see Figure 5.15). Figure 5.16 shows the stress strain curves at low stress range for two different materials. The first material is a weakly artificially cemented soil investigated by Lo et al. (2003) (see Figure 5.16a). The second one is based on the experimental results obtained by Lagioia and Nova (1995) in a study of a calcarenite, which is a coarse-grained weak rock formed in a marine deposition in the lower Pleistocene epoch and characterized by the existence of calcium carbonates (see Figure 5.16b). Sangrey (1972) also observed the pre-peak initial yield, due to bond destruction in Labrador clay (sensitive cemented clay) based on drained triaxial tests. Feda (1995) also reported a series of triaxial compression tests on undisturbed cemented Neogene lacustrine clay. Feda (1995) pointed out that during the pre-peak stage of testing, destruction of bonds was also reflected in the pre-peak wavy form of stress strain curves, and, particularly at a low stress range.

Mitchell (1970) investigated the behaviour of Leda clay, which is a highly cemented sensitive soil, under triaxial compression test conditions and identified three zones (see Figure 5.17a) based on the nonlinearity of the failure envelope. In the first zone, fissures separating clay blocks govern the behaviour and result in high dilation and, hence, a high friction angle. In the second zone, the bonding resistance governs the behaviour, and in the third zone, the intrinsic frictional behaviour dominates. In the present study, only the second and third zones identified by Mitchell (1970) were captured. In this investigation, the focus is on the simulation of shear strength of intact materials with no visible fissures. Lefebvre and Rachelle (1974) also captured the nonlinearity due to bonding degradation effects in a sensitive Champlain clay deposit in Saint-Laurent lowlands east of Quebec City on the south shore of the Saint-Laurent River in Canada (see Figure 5.17b).

5.7 Remarks on mobilized strength of bonded geomaterials

At low strain levels and at low-to-medium confining stresses, the mutual contribution of bonding and frictional resistance governs the mobilized strength. For quasi-frictional materials, at working strain levels, in the pre-peak regime, the strength is only a fraction of the full bonding and frictional resistances. This can be approximated by the plateau strength envelope, due to the balance between frictional strength mobilization and bonding strength degradation. At this strength level, only a fraction of the bonding and full frictional strength can be relied on, whereas at the peak strength frictional resistance dominates due to strain

hardening (Figures 5.11c and d). Various investigators, based on their experimental investigations of various bonded geomaterials, indirectly stated this idea and opposed the concept of the full mobilization of the frictional resistance.

Conlon (1966) stated that the strength available at the instant of failure of a slope in bonded stiff clay is due to bonding resistance, as full frictional resistance cannot be totally mobilized prior to failure. Moreover, Mitchell (1975) carried out drained constant mean effective stress tests on Leda clay to reveal the operational strength for slope failure. Mitchell (1975) revealed that the onset of dilation is a good approximation of the operational strength for sensitive clays. Cresswell and Powrie (2004) carried out experimental work on locked sand, where bonding between particles results from the interlocking of grains rather than the cementation bonds between particles. Cresswell and Powrie (2004) showed that the onset of dilation marks a major degradation of interlocking and should be considered as the limiting strength based on a factor of safety for peak values, given the proximity of the two envelopes at different confining stresses. Cecconi and Viggiani (2001) carried out triaxial compression tests on a bonded coarse-grained weak rock. These investigators pointed out that the onset of dilation marks the onset of bond destruction and that the peak resistance corresponds to a lower bonding contribution compared to that of the onset of dilation.

It is highly recommended to reconsider either the plateau strength (whenever it can be identified from the experimental results) or the onset of dilation as an upper strength threshold for all bonded geomaterials, as shown for locked sand in the previous chapter. However, further research is needed to identify the relationship between the onset of dilation as a macroscopic threshold for destruction and the field operational strength of bonded geomaterials. Comparing the operational strength based on the back analysis of well-instrumented failure case histories can reveal the relationship between the onset of dilation and the operational strength.

5.8 Summary and conclusions

This chapter investigated the contribution of the frictional resistance and the bonding resistance at various strain levels for two natural stiff clays using a simplified approach. The partitioning of the strength parameters was presented in two different stress spaces: M-C space and q-p space. In M-C space, partitioning was presented in terms of the cohesion and the friction angle for the two investigated clays. The results revealed that the cohesion and friction angle are not simultaneously mobilized (as the conventional M-C hypothesis implies for the peak strength). The cohesion distribution's trend indicated progressive degradation with plastic strain level. On the contrary, friction angle distribution builds up progressively with shear strain. In q-p space, partitioning was also carried out in terms of two normalized parameters corresponding to the bonding and the friction strength parameters for Nanticoke clay. An incremental equation

based on normalized strength parameters was proposed to reveal the contribution of the frictional and the bonding resistance at various strain levels. An example of the results of this equation was also presented to show the bearing of the non-simultaneous mobilization of strength parameters on stress-strain curves and failure envelope. It was shown that strain hardening and weakening are the by-products of the balance between cohesive strength degradation and frictional strength mobilization at different confining pressures. Three yield surfaces were captured and sufficiently described the failure evolution and failure envelope's nonlinearity, the bonding contribution and its degradation, and balance between the frictional strength mobilization and bonding strength degradation (plateau strength). The significant influence of bond degradation was manifested by the initial reduction in resistance in the pre-peak regime at low confining pressures. The experimental results for various bonded geomaterials, with different origin, revealed that this initial plateau in the stress-strain curves was a characteristic behavior. In addition, the transition between soils and rocks was also captured by using such framework with supporting evidence from the literature on different bonded geomaterials.

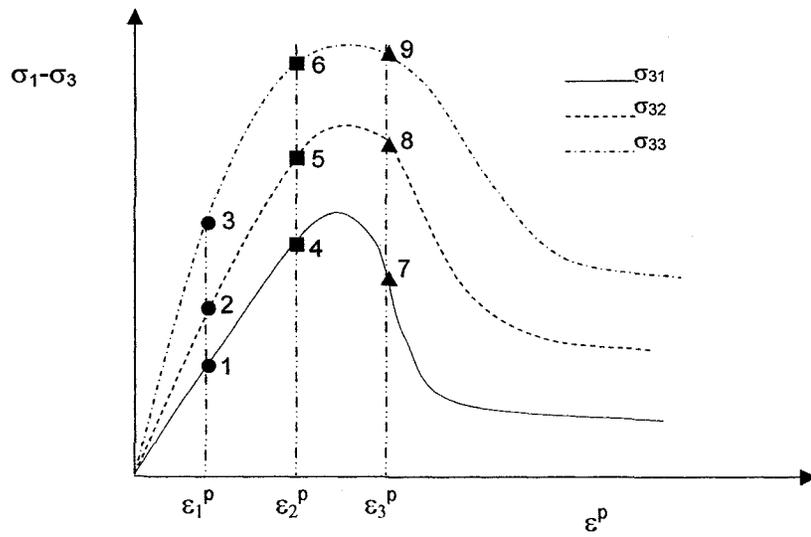


Figure 5.1 A schematic diagram of three stress-strain curves.

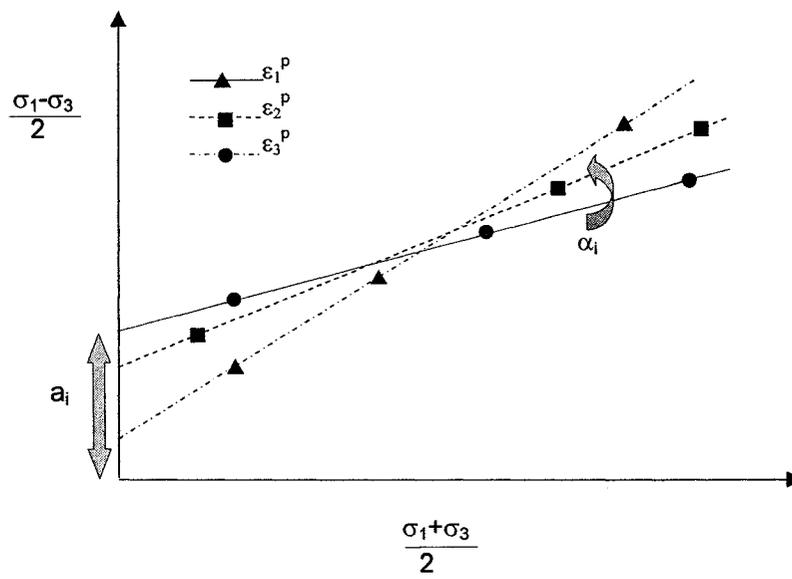


Figure 5.2 Multiple envelopes at different strain contours.

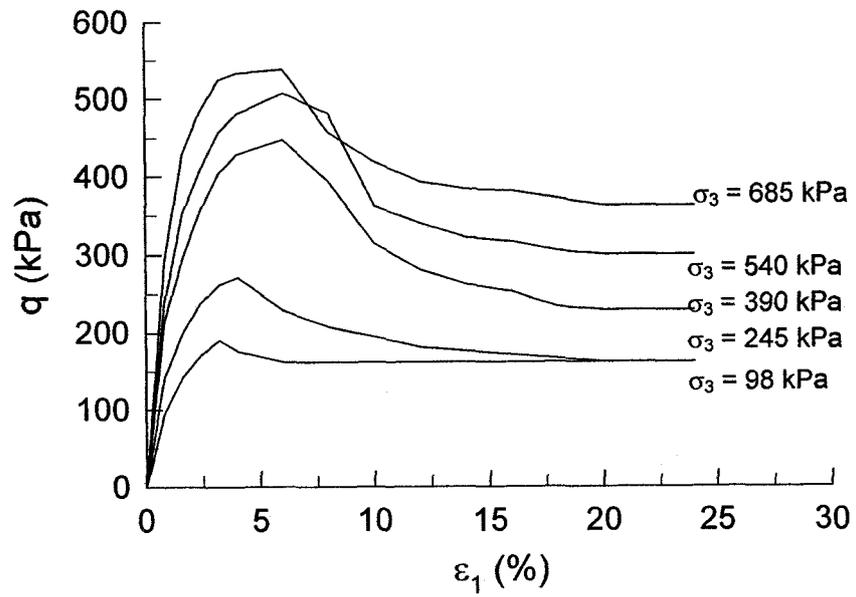


Figure 5.3 Pietrafitta clay stress-strain curves (modified from Callisto and Rampello 2002).

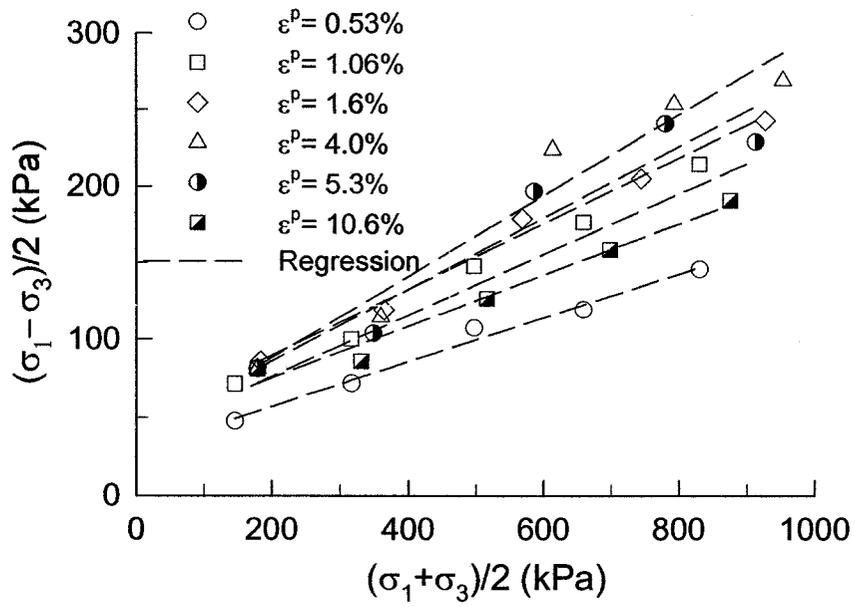


Figure 5.4 Multiple-Yield envelopes in terms of plastic shear strain contours (Pietrafitta clay).

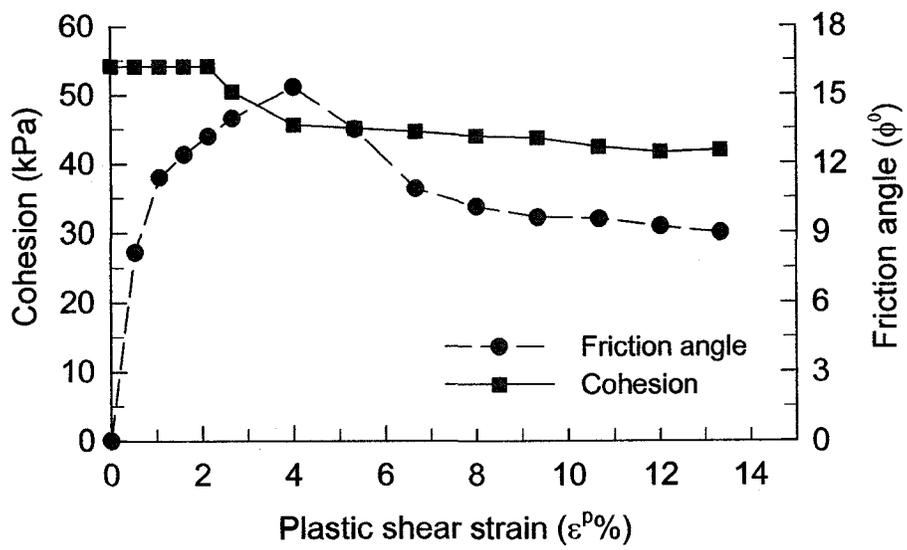


Figure 5.5 Distribution of strength parameters with plastic shear strain level (Pietrafitta clay).

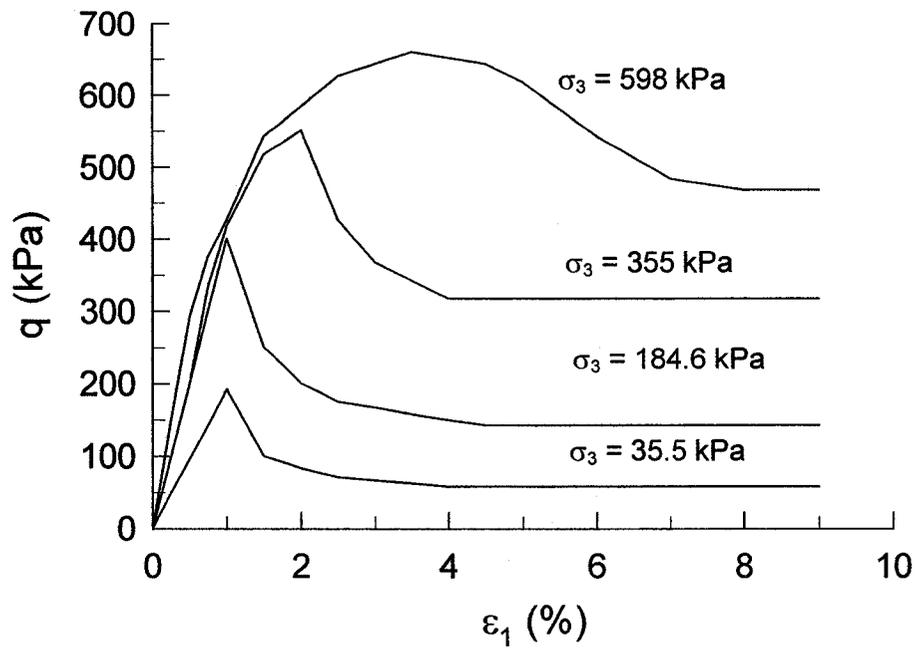


Figure 5.6 Nanticoke clay stress-strain curves (modified from Lo 1972).

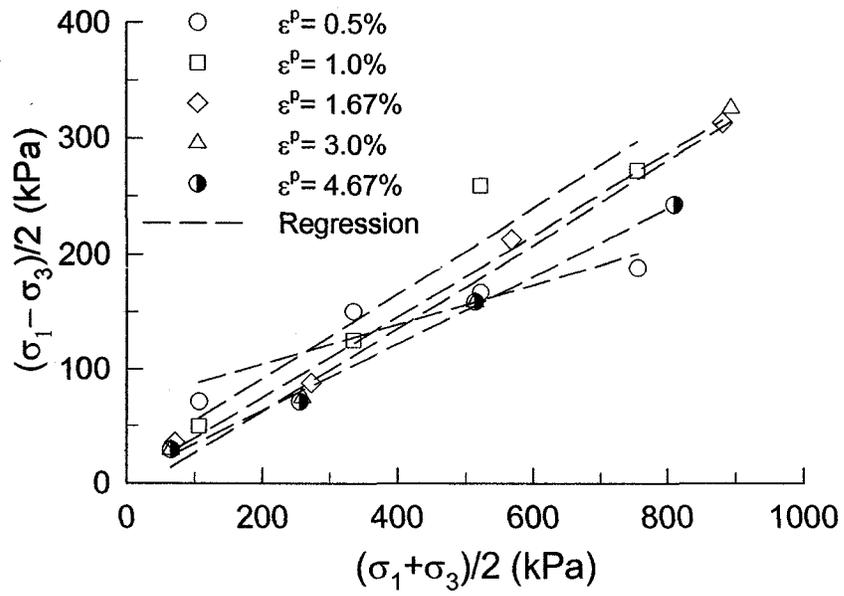


Figure 5.7 Multiple-Yield envelopes in terms of plastic shear strain contours (Nanticoke clay).

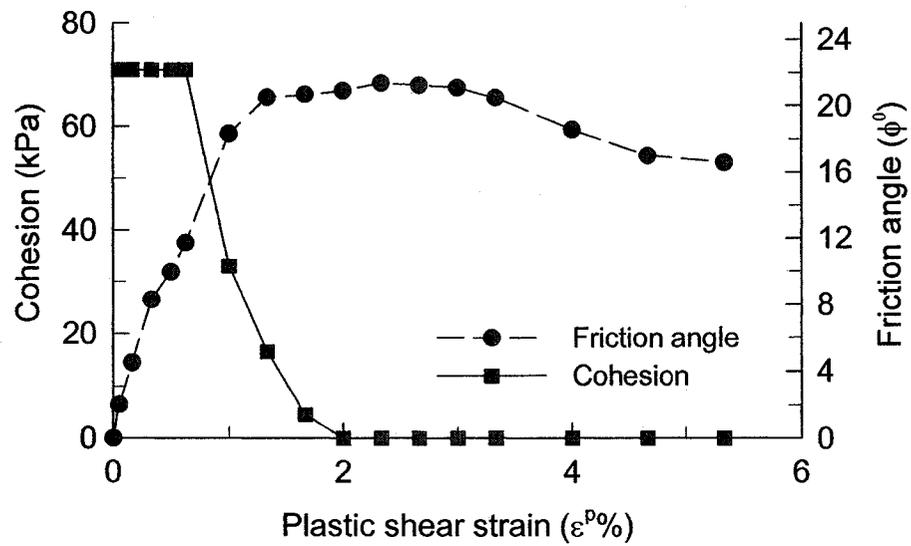


Figure 5.8 Distribution of strength parameters with plastic shear strain level (Nanticoke clay).

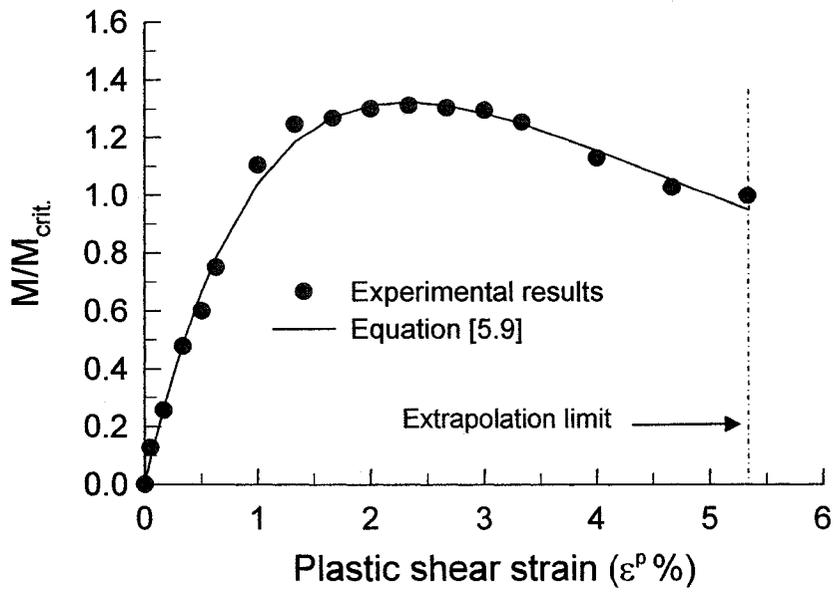


Figure 5.9 Distribution of normalized frictional strength parameter with plastic shear strain.

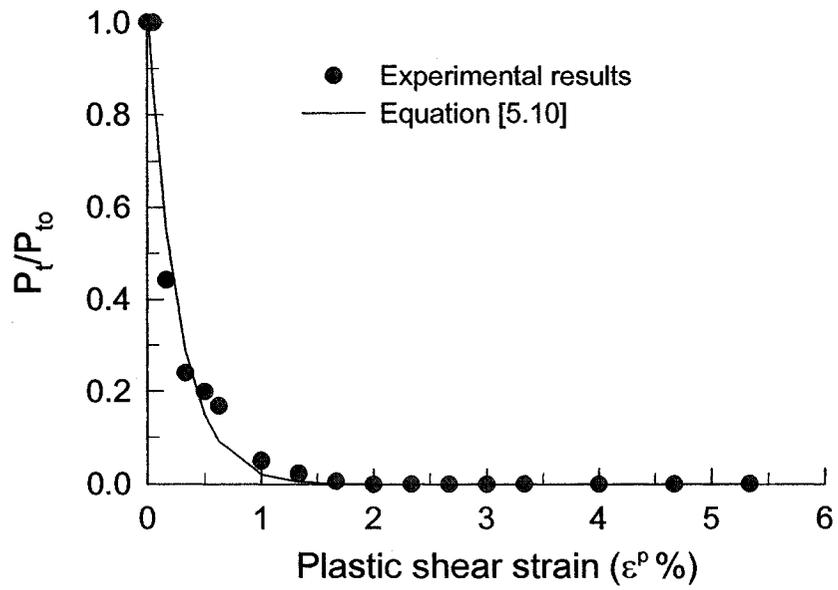


Figure 5.10 Distribution of normalized tensile strength with plastic shear strain.

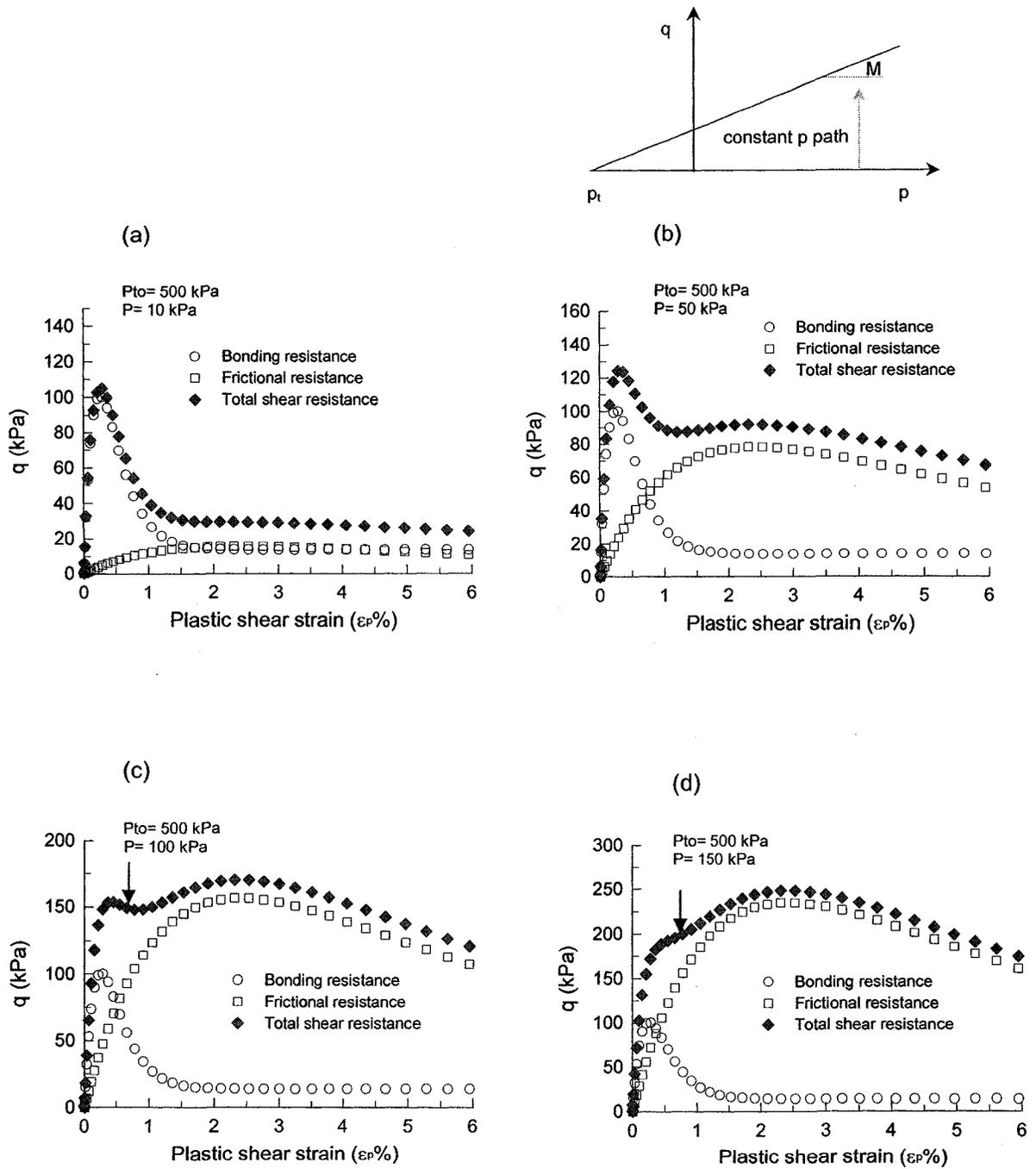


Figure 5.11 Stress-strain curves at various confining stress based on the incremental of equation [5.16].

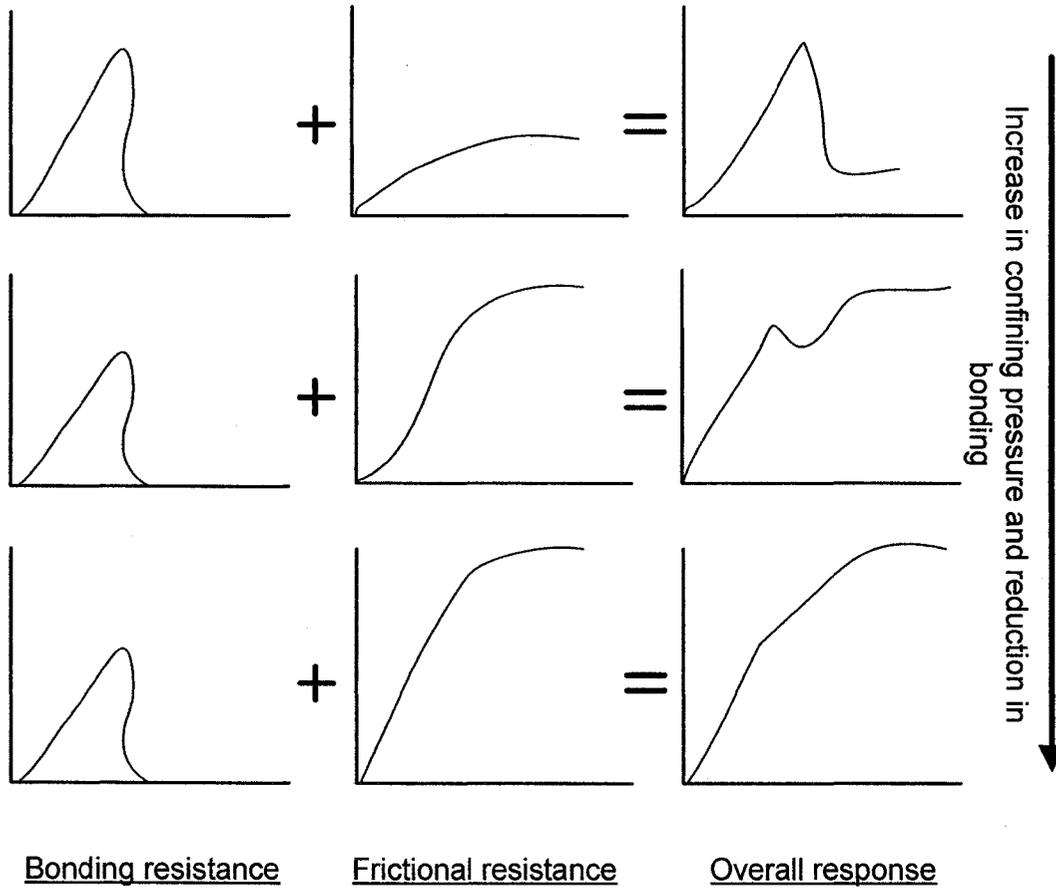


Figure 5.12 Conceptual framework to visualize contribution of frictional and bonding resistance to stiff clay shearing response (modified from Conlon 1966).

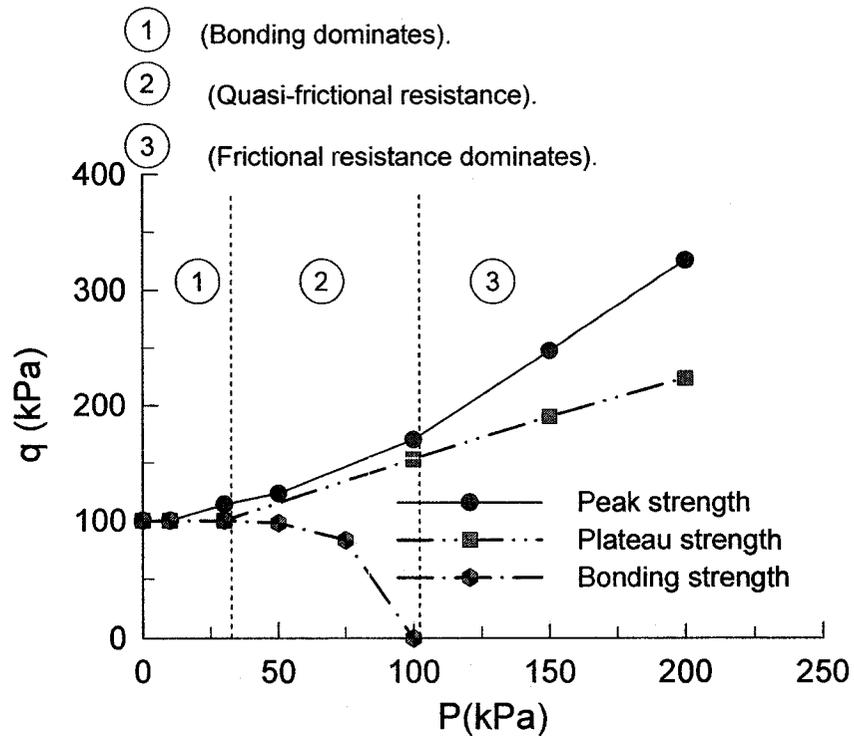
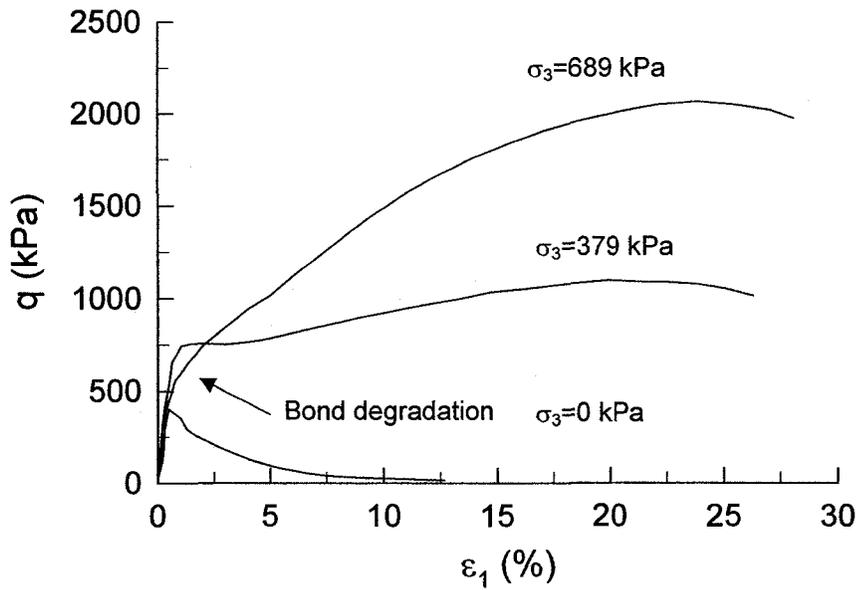


Figure 5.13 Failure, Bonding strength and plateau strength envelopes.

(a) Stress strain curves



(b) Failure envelope

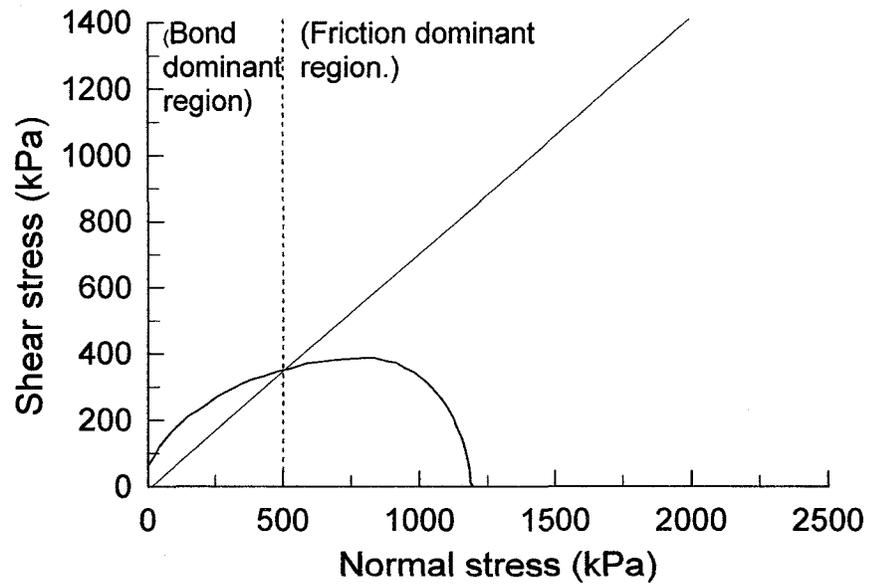


Figure 5.14 Stress-strain curves and failure envelope for Toulnustouc Stiff clay (modified from Conlon 1966).

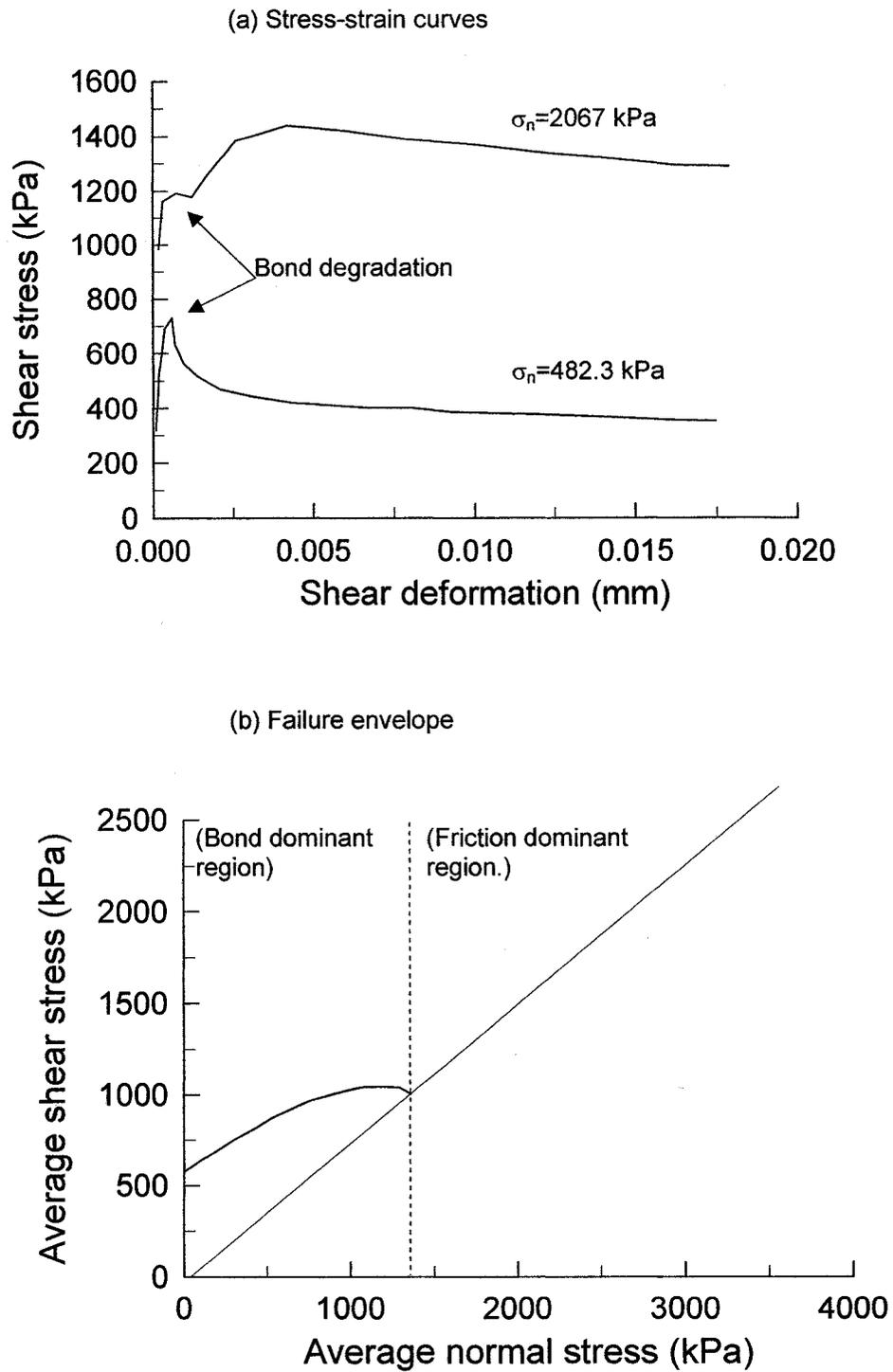


Figure 5.15 Stress-strain curves and failure envelope for Jointed weak rock (modified from Lajtai 1974).

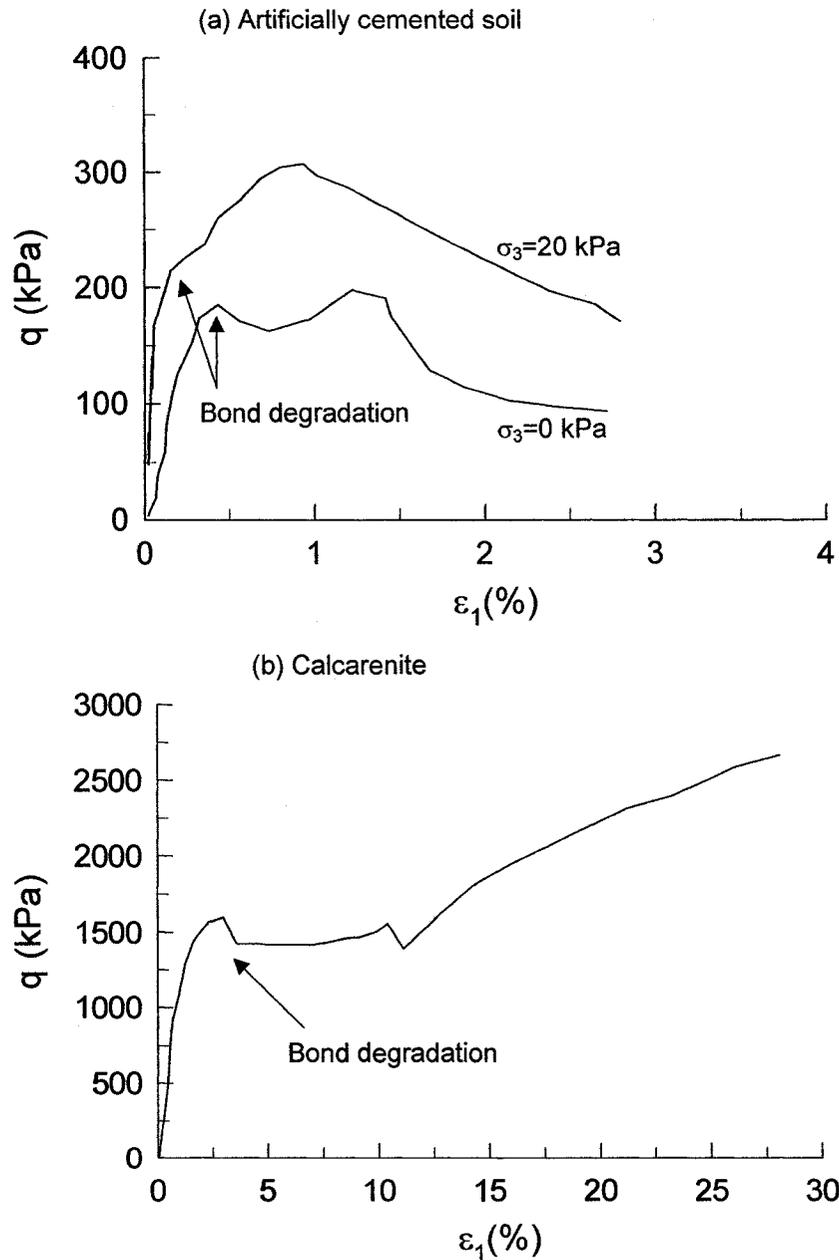


Figure 5.16 Stress-strain curves of Artificially cemented soil and Calcarenite (modified from (a) Lo et al. 2003; and (b) Lagioia and Nova 1995).

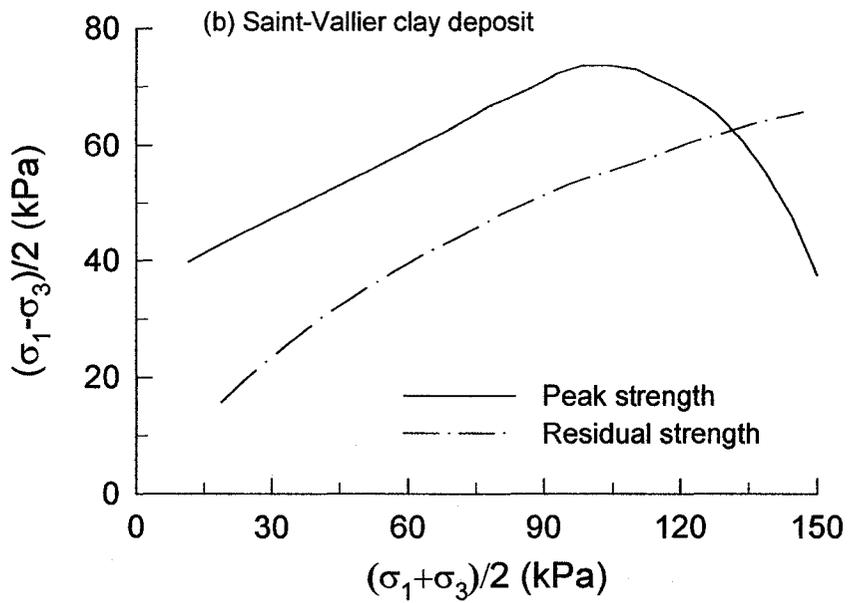
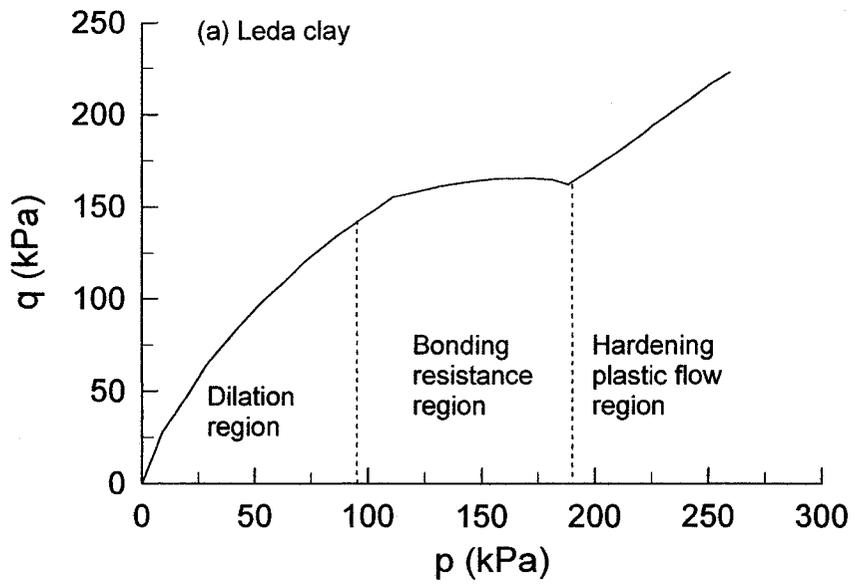


Figure 5.17 Failure envelopes for cemented sensitive clay (modified from (a) Mitchell 1970; and (b) Lefebvre and Rochelle 1974).

CHAPTER 6: SIGNIFICANCE OF STRAIN-DEPENDENT STRENGTH PARAMETERS IN THE ASSESSMENT OF SLOPE STABILITY PROBLEMS

6.1 Introduction

Back analysis of failed slopes in bonded clays (stiff clays and eastern Canadian clays) revealed that the average mobilized strength at failure (the operational strength) was less than the laboratory peak strength and greater than the residual strength (e.g., Skempton 1964, 1970; Bjerrum 1967; Burland et al., 1977; Lefebvre and La Rachelle 1974; Chandler 1984; Lefebvre 1981; Cooper et al. 1998). Realization of such phenomenon introduced the concept of progressive failure as the probable cause of this, where mobilized strength along slip surface varied according to the experienced level of strain from peak to residual strength. Despite our theoretical understanding of such phenomenon, no comprehensive work has been presented numerically to explain clearly the deviation of the average operational strength from the laboratory peak strength due to progressive failure. In recognition of the importance of strain-dependent parameters as shown in the previous chapters, a parametric study is presented in this chapter to reveal the significance of strain-dependent strength parameters in the analysis of slope stability problems as compared to conventional peak strength parameters, and, in particular to capture the progressive failure of slope as a function of strain.

The distribution of strength parameters (refer to Chapter 5) and the peak's constant strength parameters of Nanticoke clay were used to simulate the failure

of a homogeneous slope. The peak strength parameters were determined from the experimental results presented in Chapter 5 as conventionally interpreted in M-C space. Two computational tools have been used for the investigation in this chapter. The first computational tool was a commercial 2D finite difference program, FLAC version 4.0 (*Fast Lagrangian Analysis of continua*) developed by Itasca (2000). The essential element of this program is the explicit numerical scheme in which dynamic equations of motion in conjunction with incremental constitutive laws are solved over a very small time step. In this way, variation of field variables propagates as a real physical disturbance, and no iteration procedure is required to arrive at the solution. The program is most suitable for analyzing nonlinear behavior of materials and related instability and failure phenomena. FLAC has been used in these analyses to reveal the slip surface associated with peak strength parameters and the strain-dependent strength parameters (SDP). The second computational tool was a commercial limit equilibrium slope stability program (SLOPE/W, Geo-Slope International Ltd 2004). SLOPE/W is user friendly software that has been used to simulate complex slope geometries. This program enables the users to visualize the minimum factor of safety, the critical slip surface, and the distribution of normal stresses and shear stresses along the slip surface using different limit equilibrium methods. SLOPE/W has been used in this investigation to reveal the factor of safety and the critical slip surface obtained by using the peak strength parameters.

6.2 Limit equilibrium and deformation analysis

Limit equilibrium analysis is widely used in the field of geotechnical engineering to assess the stability of slopes. An implicit assumption in limit equilibrium analysis is that the soil exhibits constant strength regardless of experienced deformation. As LEM has no consideration for deformation, it is assumed that the selected strength is mobilized everywhere along the slip surface regardless of the deformation (Duncan, 1996). In stability problems, it is the custom to use limit equilibrium to render the factor of safety. Morgenstern and Sangrey (1978) state that one of the uses "... of the factor of safety is to provide a measure of the average shear stress mobilized in the slope." They go on to state that, "This should not be confused with the actual stresses." In addition, due to the additional assumptions for determinacy in LEM, global equilibrium is satisfied rather than equilibrium everywhere in the soil mass. Duncan (1996) and Krahn (2003) provided a comprehensive summary of the available methods and limitations of limit equilibrium according to the inherent assumptions regarding the inter-slice forces and equations of equilibrium. Fredlund and Krahn (1977) pointed out the main difference between the limit equilibrium methods and showed the similarity between the different methods and the calculated factors of safety regardless of LEM assumptions. Krahn (2003) stated that LEM relies only on the principal of statics. The method says nothing about strains and displacements, and as a result it does not fulfill displacement compatibility. In addition, the stresses calculated in LEM do not represent the actual case, and localized shear stress concentrations are not captured. Predicted stresses in

LEM are only based on the assumption of a unique factor of safety for all slices, where the slice normal and shear stresses are derived primarily from the slice weight. Tavenas et al. (1980) and Wright et al. (1973) pointed out the importance of incorporating the actual stresses in slope stability analysis. These researchers also realized that the definition of the factor of safety assumes a certain stress path, which is not necessarily developed in the field. In addition, for a given slope, the critical slip surface associated with the minimum factor of safety may be different from the critical slip surface associated with critical acceleration (Sarma and Bhava 1974). Yu et al. (1998) also pointed out that the collapse mechanism due to the rigid plastic assumption in LEM is usually kinematically inadmissible.

The second technique used in the assessment of the stability of slopes is the application of the deformation analysis by using either the finite element methods or finite difference methods. The numerical analysis is a very powerful tool in enhancing our understanding of civil engineering activities. Griffith and Lane (1999) applied the finite element method to calculate the factor of safety using elasto-plastic Mohr-coulomb criterion. Dawson et al. (1999) also applied finite difference method by using the strength reduction technique to estimate the factor of safety. In the deformation analysis methods, no prior assumptions are required regarding the shape, the location of the failure surface, or the inter-slice forces as in LEM. Failure occurs naturally through the zones within the soil mass in which the soil shear strength is unable to sustain the applied shear stresses.

In deformation analysis, the stresses are also closer to reality and stress concentrations are considered in the analysis. The potential failure surface for a simulated slope is approximated by using the non-convergence criterion as introduced by Zienkiewicz (1971). In the context of this approach, the shear strength parameters in a nonlinear analysis are reduced until numerical instability occurs. The factor of safety is calculated as the ratio of the available shear strength and the factorized strength that results in a converged solution prior to instability. Slope stability analysis using the strength reduction has been carried out successfully by Dawson et al. (1999) and Griffith and Lane (1999) for homogeneous slopes.

6.3 Slope geometry and material properties

A 1:1 slope with a height of 20m is used for these analyses. In FLAC version 4.0 (2000), the whole domain was divided into 100x60 square elements before the excavation stage. Displacement-type boundary conditions were assigned at the model boundaries. The domain's left and right boundaries were restrained in horizontal direction. The base of the domain was restrained in both vertical and horizontal directions (see Figure 6.1). Before the excavation stage, the material behaviour was assumed to be elastic by prescribing both the bulk and shear modulus to estimate the initial stress distribution upon reaching an equilibrium state under gravitational stresses. The excavation was carried out by the removal of the excess grid zones from the domain. After excavation, the model was run to reach an equilibrium state while prescribing an elastic model. The distributions of

the vertical stress under gravitational stresses before and after the excavation are shown in Figure 6.2 and Figure 6.3, respectively. According to FLAC conventions, negative normal and shear stresses indicate compression and anti-clockwise distortions, respectively. The excavation stage resulted in a local shear stress concentration in the vicinity of the toe as shown in Figure 6.4. Once the initial stress conditions were established, the constitutive model was altered from the elastic model to either the elasto-plastic M-C model or the non-associated strain hardening/softening M-C model. In FLAC, it is possible to assign strength parameters as a function of plastic strain using the table option. In the subsequent simulations, the peak strength parameters as well as the strain-dependent strength parameters for Nanticoke clay (refer to Chapter 5) are used for illustration. Table 6.1 lists the elastic properties as well as the peak and residual strength parameters for Nanticoke clay. It is assumed that the difference between the peak and residual strength, in a conventional elasto-plastic analysis, is due to the cohesive component. As a matter of fact, the experimental work for Nanticoke clay was not carried out to large strains to reveal the possibility of a reduction in the friction angle, and, probably due to the limitation of the experimental setup.

Although the elastic strains were assumed to be negligible, when the strain-dependent strength parameters were estimated, in Chapter 5, high elastic properties were assumed for the model to avoid any inconsistency. Hence, the

elastic strains estimated based on these elastic parameters would be negligible. It should also be noted that all the parameters are in terms of effective stresses.

6.4 Peak strength parameters simulations

6.4.1 *Finite difference simulations*

Once the initial stress conditions were established, the material was assigned M-C peak strength parameters. The equilibrium of the model was examined by recording the unbalanced forces (the maximum of the unbalanced nodal forces). The reduction of unbalanced forces with the number of steps is an indication of reaching a steady state. Non-convergence occurs when the recorded unbalanced forces are either high or when large fluctuation in the unbalanced forces occurs. In deformation analysis approach, the slip surface evolves naturally without any assumptions for inter-slice forces as in conventional LEM. The rendered slip surface from deformation analysis give a factor of safety of 1 based on the assigned strength parameters. When the peak strength parameters were assigned to the model, no slip surface was detected. In order to get the potential slip surface, the strength reduction factor (SRF) technique was used, based on the peak strength parameters. The reduced strength parameters were calculated using the following equations:

$$[6.1] \quad \text{SRF} = \frac{c_p}{c_m} = \frac{\tan(\phi_p)}{\tan(\phi_m)},$$

where c_p is the peak cohesion, c_m is the assumed mobilized cohesion, ϕ_p is the peak friction angle, and ϕ_m is the assumed mobilized friction angle.

Using the SRF technique, the strength parameters have been incrementally decreased by applying a reduction factor to the peak strength parameters. In each case, a simulation was run to obtain the potential slip surface. The last converged model defined the potential slip surface and the factor of safety. This approach is in agreement with the definition of the factor of safety (fos) as the factor by which the strength is reduced to reach equilibrium. Table 2 summarizes the reduction factors, unbalanced forces, and the convergent solutions. Figure 6.5 and 6.6 show the unbalanced forces history for reduction factors of 1.43 and 1.44, respectively. It can be seen from the figures that the model's convergence with a SRF of 1.44 was impossible, when compared to a SRF of 1.43. There was an increase in unbalanced forces that fluctuated between 1.5-4 kN/m for a SRF of 1.44 as compared to stabilized unbalanced forces (0.8 kN/m) for a SRF of 1.43. The potential slip surface based on a SRF of 1.43 is shown in Figure 6.7.

6.4.2 *Limit equilibrium simulation*

SLOPE/W (Geo-Slope International Ltd 2004) is commercial limit equilibrium software to assess the stability of slopes. SLOPE/W was used to obtain the minimum factor of safety and the critical slip surface associated with the peak strength parameters by using M-C plasticity model. The critical slip surface

obtained from LEM is not necessarily the potential slip surface. It is merely defined as the surface which renders the minimum factor of safety. This analysis was carried out to compare the factor of safety and the critical slip surface from LEM with that obtained from the numerical simulation. Table 3 summarizes the factors of safety based on different analysis methods. Figure 6.8 shows the slope geometry, the minimum factor of safety, and the critical slip surface. The factor of safety ranged from 1.42 to 1.48. The factor of safety associated with the critical slip surface (1.45), based on the methods that satisfy both force and moment equilibrium, is relatively in agreement with the factor of safety obtained from the numerical simulation (1.43). Although the factor of safety is almost identical from the LEM and numerical simulation, the slip surfaces are relatively different. The limit equilibrium approaches renders state of stresses that may not represent the actual stresses in order to assure a unique factor of safety for all slices. Figure 6.9 shows the distribution of the normal stresses, mobilized stresses, and available strength at the base of the slices. The first observation is that the ratio between both of the mobilized and available strength is the same for all slices. The second observation is that in order to render a unique factor of safety, tensile normal stresses are manifested at the crest of the slope, which indicates an upward movement of the slope and a negative frictional resistance. For reduced strength parameters which corresponded to the factor of safety, the same critical slip surface was obtained from LEM. In both cases, the normal and mobilized stress distributions were identical at the base of slides.

6.5 Strain-dependent strength parameters simulation

The rendered parameters from the analysis of Nanticoke clay in the previous chapter were implemented in FLAC's strain hardening/softening model. Although it is known that the softening curve is inversely related to the element size (smaller elements give a thinner shear band and results in more softening), no effort was done to investigate such phenomenon in this analysis. The aim of this analysis is neither to estimate the thickness of the band nor to quantify the shear strains within the band. The potential slip surface based on the simulation is shown in Figure 6.10. The plasticity indicator indicates a combination of tension and shear failure along the slip surface. The tensile failure is more evidenced near the toe and the crest of the slope (represented as circles).

Figure 6.11 shows the evolution of the potential slip surface with the plastic strain. The potential slip surface evolves from the toe, where high shear stresses are recorded, progressively to the slope's crest with the increase in plastic strain level. Figure 6.12 shows the distribution of the unbalanced forces with the number of iterations. The model showed a reduction in the unbalanced forces to a threshold followed by an increase in unbalanced forces. It was assumed that initiation of global failure commenced at the point where the unbalanced forces accelerated. In order to provide a better understanding of the problem, the model was run to the assumed instability point. Marker points were assumed along the potential slip surface as shown schematically in Figure 6.13. These points served

to monitor the evolution of strength parameters and stresses along the potential slip surface.

6.5.1 *Deviation of operational strength from peak strength*

Figure 6.14 and Figure 6.15 show the distribution of the mobilized cohesion and friction angle along the potential slip surface as a function of plastic strain. As shown on the figures, the plastic strain varies along the potential slip surface. The plastic strain tends to be higher at the toe and reduces progressively towards the crest of the slope. The experienced plastic strain at the toe is two orders of magnitude greater than the plastic strain at the crest. The results of the model comply with the assumptions of non-simultaneous mobilization of strength parameters. The cohesive component tends to degrade with progressive mobilization of friction angle. At the toe, the cohesive component is completely degraded, while the friction angle is fully mobilized followed by a decrease in its value. On the other hand, the cohesive component is fully mobilized at the crest of the slope with a negligible contribution of the friction angle. At any intermediate point along the slip surface, the mobilized strength is due to the mutual contribution of the friction angle and cohesive resistance that correspond to the plastic strain level. The contribution of the cohesive component increases from the toe to the crest, while the friction angle tends to decrease.

Figure 6.16 shows the distribution of the normal, shear stresses, and available strength along the slip surface from the toe to the crest. The vertical, horizontal, and shear stress estimated from the model were used to calculate the normal and shear stress at selected points along the slip surface. The normal and shear stresses on the slip surface were calculated using the following equations:

$$\text{[6.2]} \quad \sigma_n = \sigma_{xx} \sin^2(i) + \sigma_{yy} \cos^2(i) - \sigma_{xy} \sin(2i), \text{ and}$$

$$\text{[6.3]} \quad \tau_n = \sigma_{xy} (\sin 2(i) - \cos 2(i)) - \left(\frac{\sigma_{yy} - \sigma_{xx}}{2} \right) \sin 2i,$$

where σ_n , τ_n are the normal and shear stresses on the slip surface, respectively; i is the inclined angle of a segment along the slip surface with the horizontal direction; and σ_{xx} , σ_{yy} , and σ_{xy} are the normal and shear stresses acting in x - and y - coordinate directions.

The available shear strength was calculated based on the distribution of the normal and shear stresses using the following equation:

$$\text{[6.4]} \quad \tau_a = c + \sigma_n \tan(\phi),$$

where τ_a is the available strength; and c and ϕ are the cohesion and friction angle, respectively.

It can be seen from figure 6.16 that the normal and shear stresses at the toe are higher than the corresponding values at the crest. Hence, the model accounts for stress concentration due to excavation as compared to LEM. The available strength is almost identical to the mobilized shear stress along the slip surface except at the crest. This indicates that at the onset of instability, the upper part of the slope is not experiencing failure.

Figure 6.17 shows the stress path experienced at the monitoring points (refer to Figure 6.13). Points a and b mark the initial and final stress conditions, respectively for each stress path. It can be seen from the figure that the stress path differs along the slip surface and is quite complex. A part of the complexity is due to the fact that there is a continuous rotation of principle stresses with the evolution of the deformation (the maximum stress rotation occurs at the toe). At the toe (point 6), the material initially experiences an increase in shear stress at constant mean stress followed by a reduction in shear stress where there is a reduction in the major principle stress with a slight reduction in the minor principle stress (passive mode of deformation). On the contrary, at the crest (point 1), the material initially experiences an active mode of deformation (an increase in major principle stress accompanied by constant or a slight reduction in minor principle stress) followed by an increase in shear stresses due to the reduction of both major and minor principle stresses. At the intermediate points, the stress path can

be approximated as an increase in shear stress at constant normal mean stress. In addition, it can be observed that none of the stress paths overshoots the peak strength envelope. The mobilized shear stresses are much less than the laboratory peak strength. This is partly due to the complexity of the stress paths experienced along the slip surface as compared to the simplified laboratory stress paths. Despite the difference in stress paths, one can simply plot the relationship between the mobilized shear stress and the plastic strain to approximate the field stress –strain curve as shown in Figure 6.18.

Based on the distribution of strength parameters (refer to Figures 6.14 and 6.15, the average friction angle and cohesion along the slip surface are 15° and 29 kPa, respectively. The averaged normal and shear stresses are 117 kPa and 57.4 kPa, respectively (refer to Figure 6.16). The global factor of safety based on the averaged values renders a factor of safety of 1.0. Using the peak strength parameters ($c=71$ kPa and $\phi=17^\circ$), the mobilized strength on the slip surface is 54% of the peak strength. The residual strength parameters render a factor of safety of 0.62. According to the LEM and the numerical simulation using peak strength parameters, the mobilized shear stresses on the slip surface are 70% of the peak strength. In addition, using the averaged strength parameters in LEM renders the same slip surface that is predicted from the SDP model with a factor of safety of 0.82. The deviation of the factor of safety from unity is simply due to the underestimation of the

normal stresses estimated from the LEM (there is no consideration for stress concentration at the toe).

In summary, assuming that the SDP's slip surface is the actual slip surface, the peak strength parameters tend to overestimate the factor of safety, while the residual strength underestimates the factor of safety. Hence, the mobilized strength is greater than the peak strength and less than the residual strength. Based on the previous analysis, the mobilized strength at the assumed instance of failure is 54% of the peak strength. This average value of mobilized strength accounts for the progressive failure of slope, and in agreement with the operational strength ranges estimated based on back analysis of failed slopes in highly plastic stiff clays (Mesri and Abdel-Ghaffar 1993). Hence, the incorporation of non-simultaneous mobilization of strength parameters in slope stability analysis renders a plausible explanation for the deviation of the mobilized strength from the peak strength and efficiently captures the evolution of failure. However, more work is needed to confirm such conclusion based on well-documented case histories.

6.6 Comparison between slip surfaces' positions

Figure 6.19 shows the potential slip surfaces obtained from the previous analyses. The slip surface, associated with peak strength parameters, estimated from Limit equilibrium analysis is relatively shallower than the one estimated from the finite difference code (FLAC). This is an agreement with Potts et al. (1997)

using the finite element method and the limit equilibrium method. The SDP's slip surface is shallower than both slip surfaces obtained from LEM and FLAC using constant parameters. If it is assumed that the cohesion degrades with the progressive mobilization of friction as a function of strain, the slip surface is anticipated to be shallower than the one obtained using constant parameters. It was shown in the previous chapter that the concept of non-simultaneous mobilization of strength parameters reflects the failure envelope non-linearity. According to Jiang et al.'s (2003) investigation on the effect of failure envelope non-linearity on the stability of slopes, the incorporation of nonlinearity of failure envelope results in shallower slip surface as compared to linear failure envelope. This is in agreement with the results presented in this investigation.

6.7 Summary and conclusions

In this chapter a parametric study was carried out to reveal the significance of non-simultaneous mobilization of strength parameters in the assessment of the stability of slopes. The focus was to explain the deviation of the operational strength from the peak strength. For this purpose, two commercial software were used for the simulation of a 45° homogeneous slope with a height of 20 m. The finite difference software (FLAC) was used to reveal the potential slip surfaces associated with peak strength parameters and the strain dependent strength parameters by considering the hardening/softening model. The factor of safety, based on peak strength parameters, was determined by using the strength reduction technique. The slope stability analysis was also carried out by using

limit equilibrium software (SLOPE/W) to reveal the factor of safety and the slip surface associated with peak strength parameters. The distribution of the strength parameters, normal and shear stresses, using the hardening/softening model, along the potential slip surface were also estimated.

The results of the finite difference model and LEM showed that the factor of safety associated with peak strength parameters was 1.43 and 1.45, respectively. The LEM's slip surface was relatively shallower than the one estimated based on the strength reduction technique. The strain-dependent strength parameters' slip surface was shallower than both of those estimated from LEM and FLAC. The numerical simulation using the hardening/softening model captured the progressive failure of the simulated slope. The mobilized shear strength parameters values varied along the slip surface according to the experienced plastic strain level. The plastic strain at the toe was at least one order of magnitude greater than the plastic strain at the crest. For an average value of normal and shear stresses, the peak strength parameters overestimated the factor of safety. On the other hand, the residual strength parameters underestimated the factor of safety. The mobilized strength to render a factor of safety of 1 was 54% of the peak strength and accounted for the progressive failure of the slope. This was only captured through incorporation of the strain-dependent strength parameters. It was concluded that this conceptual framework is a powerful framework to explain the discrepancy of the mobilized strength in

the field from the peak strength. However, more work is needed to confirm such conclusion based on well-documented case histories.

Table 6.1 Summary of elastic and M-C constant strength parameters for Nanticoke clay.

Parameter	Value
Bulk Modulus (kPa)	1.19x10 ⁶
Shear Modulus (kPa)	0.10x10 ⁶
Peak cohesion (kPa)	71
Peak friction angle	17
Residual cohesion(kPa)	0
Residual friction angle	17

Table 6.2 SRFs and average unbalanced forces.

SRF	Average Unbalanced force (kN/m)	Convergence
1.35	0.2	Yes
1.43	0.8	Yes
1.44	1.5-4.0	No

Table 6.3 Factor of safety based on LEM.

Method	Factor of safety (fos)
Janbu method	1.42
Bishop method	1.45
Ordinary method	1.48
Morgenstern-price method	1.45
Spencer	1.45

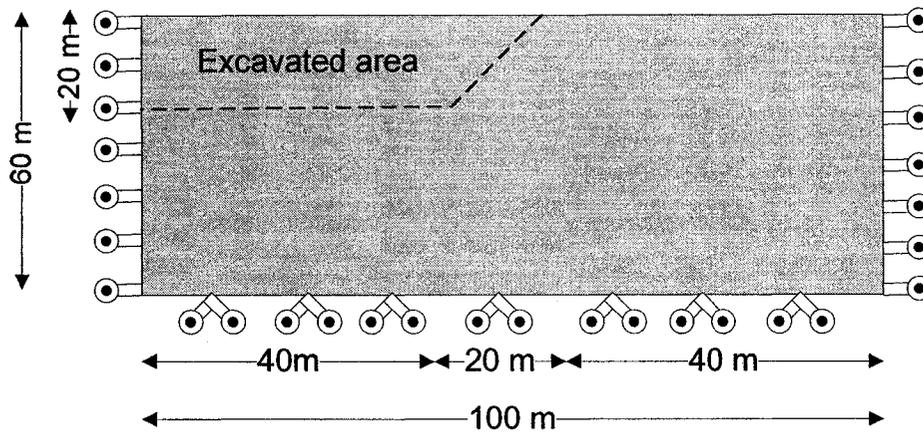


Figure 6.1 Schematic diagram of excavated slope.

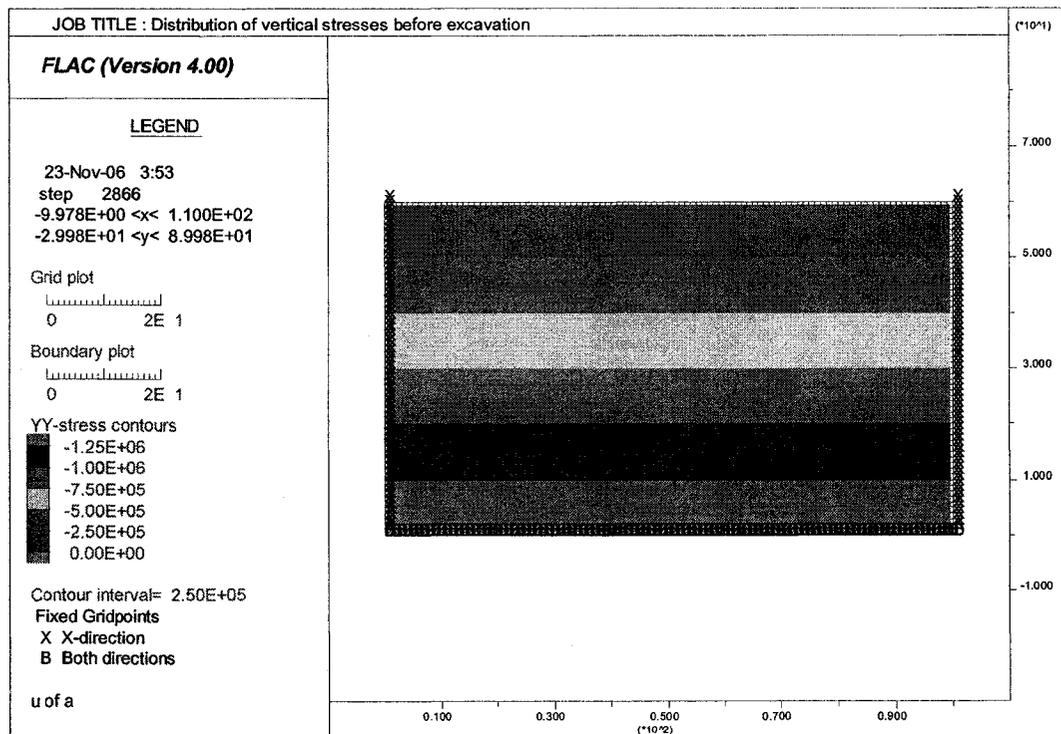


Figure 6.2 Vertical stress contours before excavation.

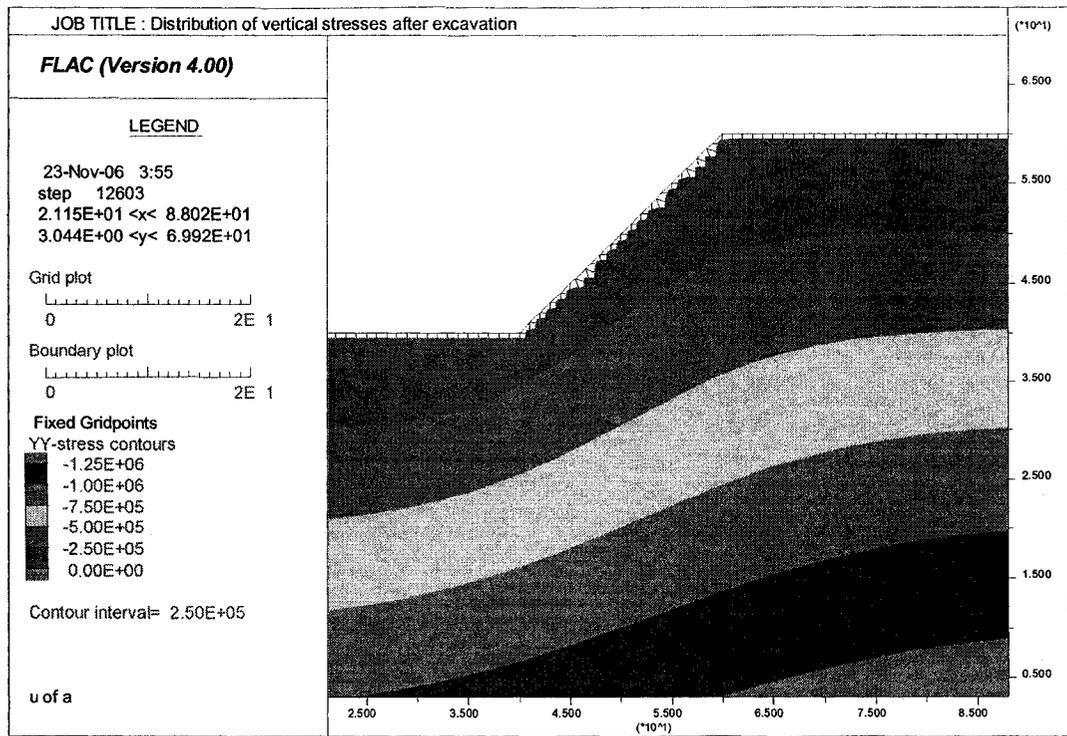


Figure 6.3 Vertical stress contours after excavation.

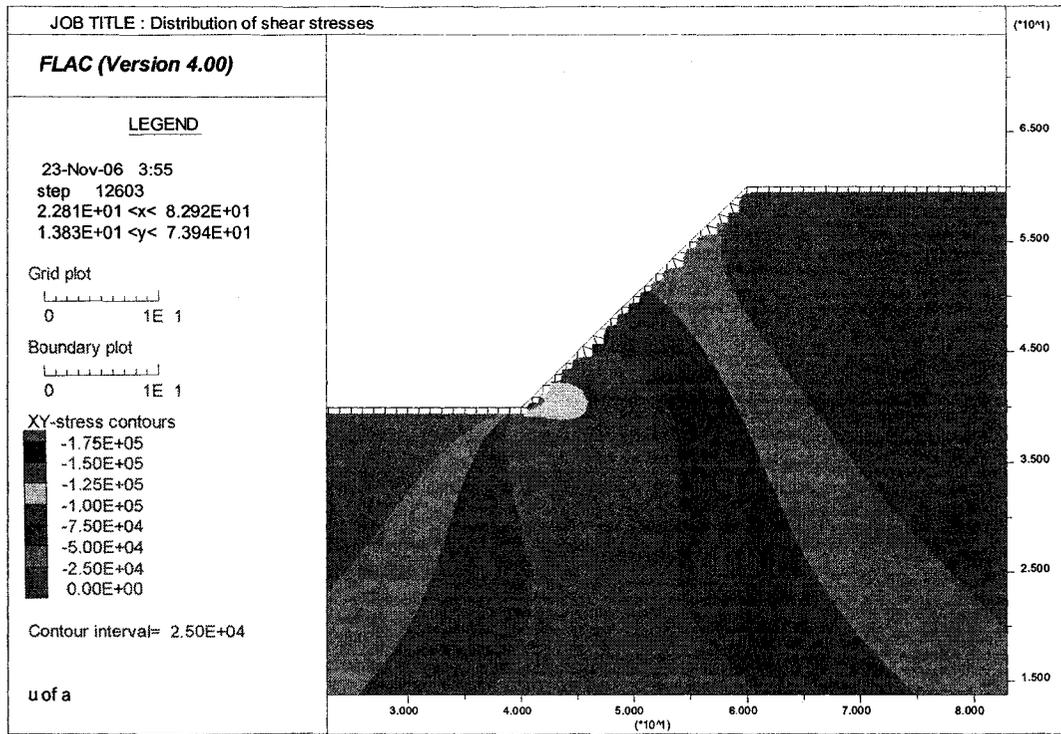


Figure 6.4 Shear stress contours after excavation.

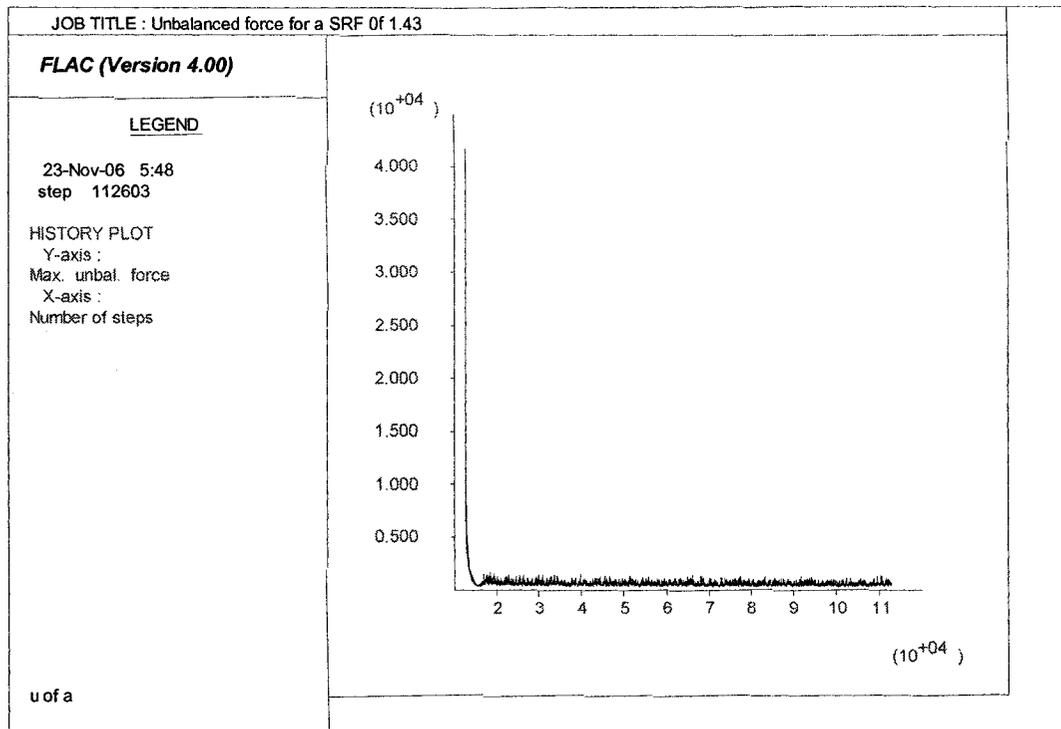


Figure 6.5 Unbalanced forces versus no of steps based on a SRF of 1.43.

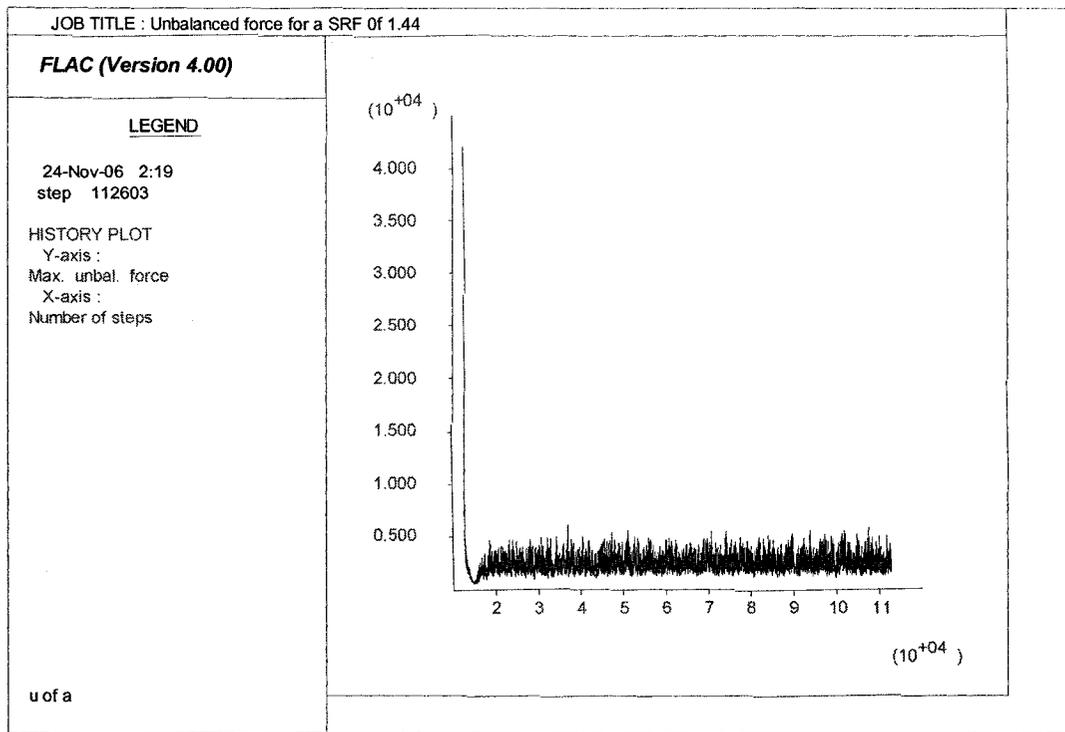


Figure 6.6 Unbalanced forces versus no of steps based on a SRF of 1.44.

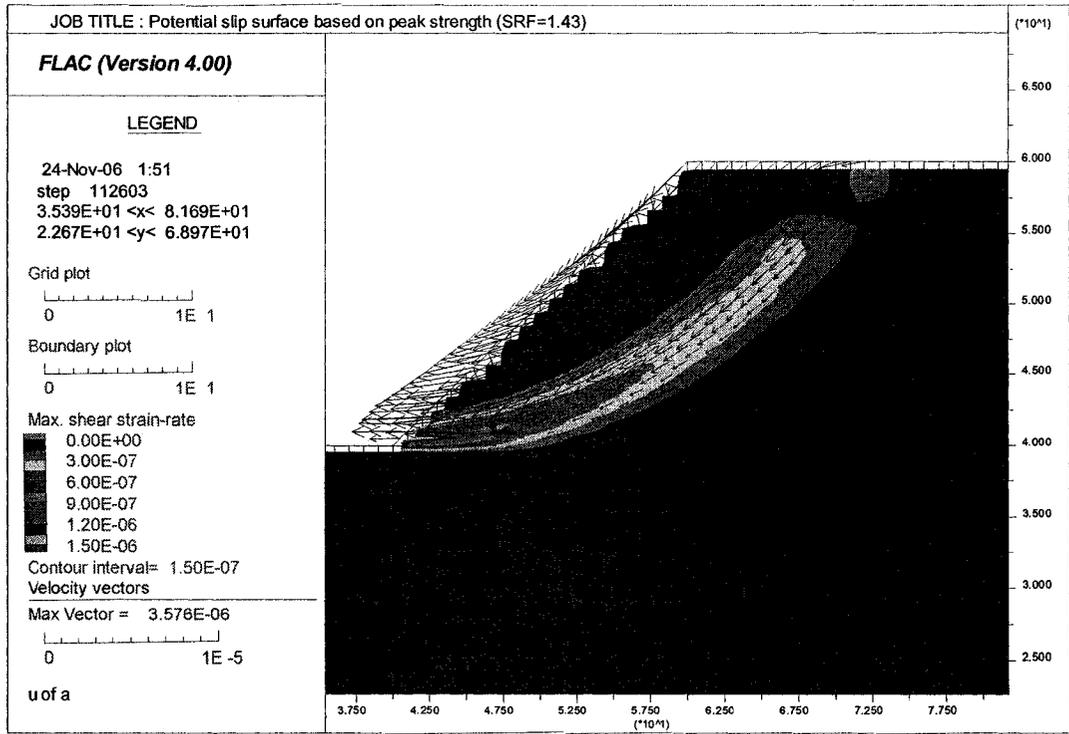


Figure 6.7 The potential slip surface obtained from the SRF technique implemented in FLAC.

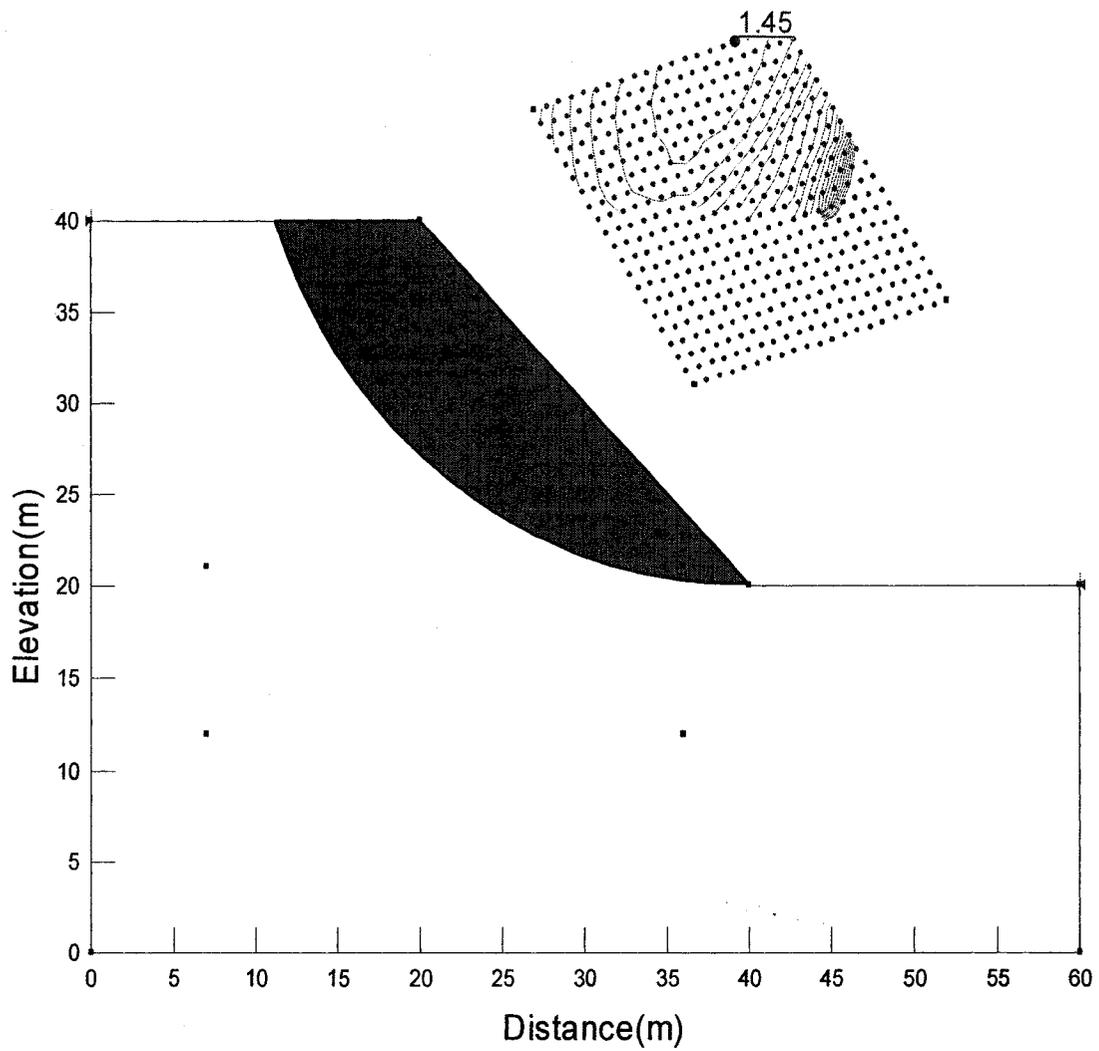
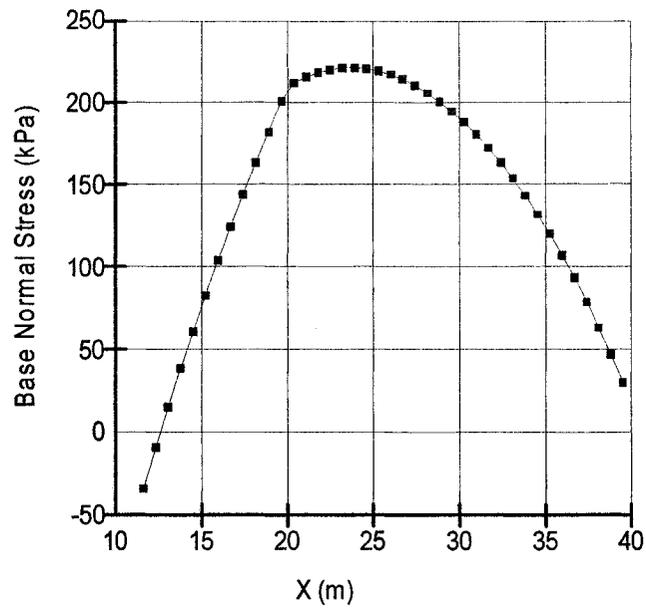
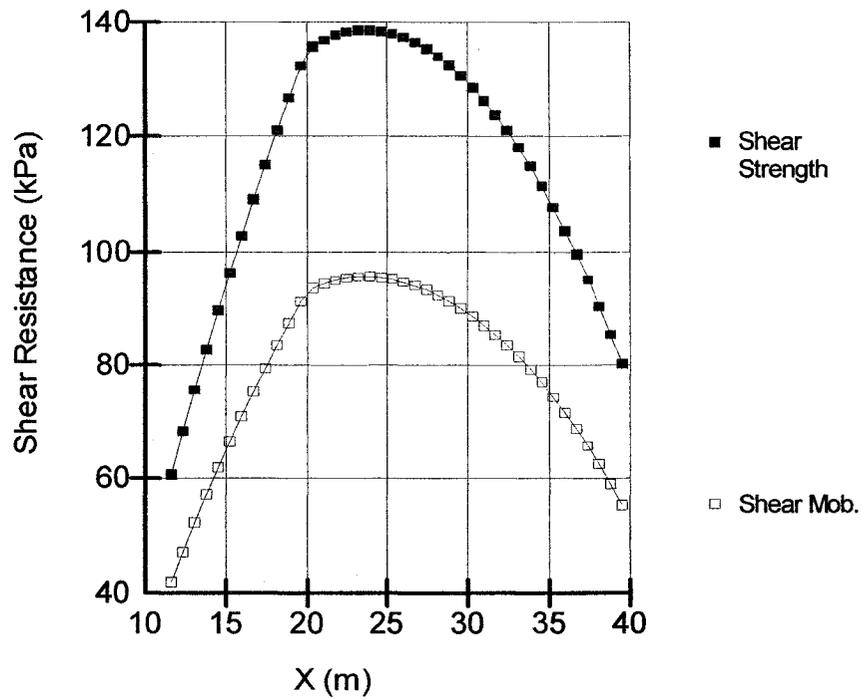


Figure 6.8 The minimum factor of safety and the critical slip surface obtained from LEM.



(a)



(b)

Figure 6.9 The distribution of normal and shear stresses from LEM: (a) the normal stress distribution; and (b) the available and mobilized strength.

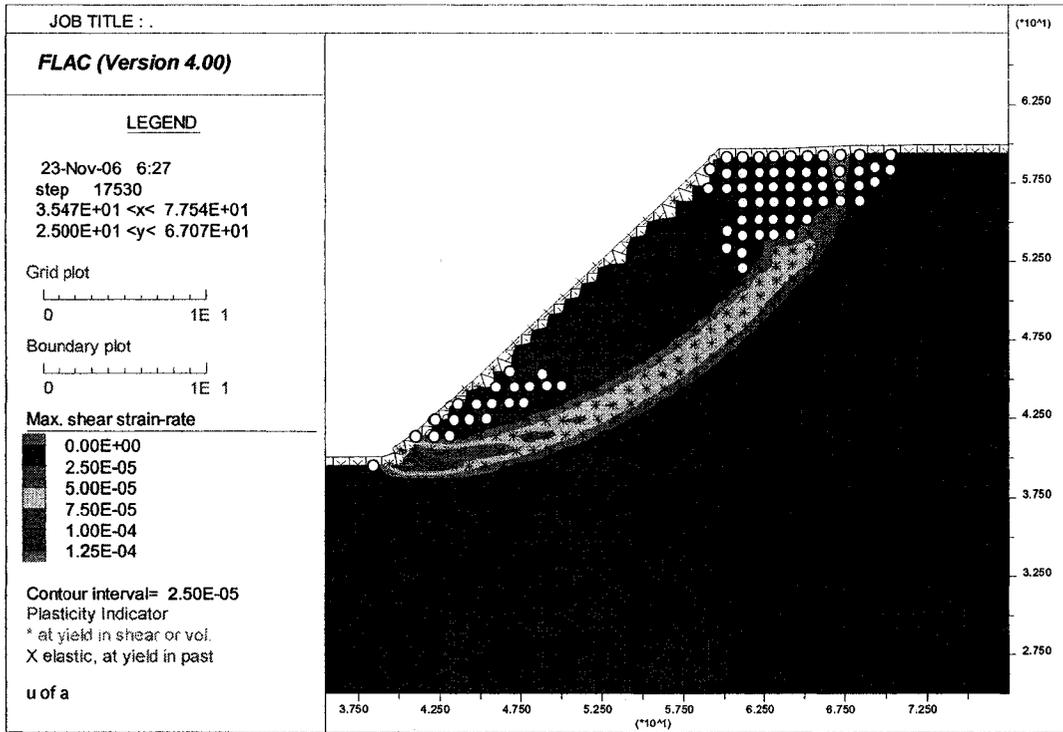
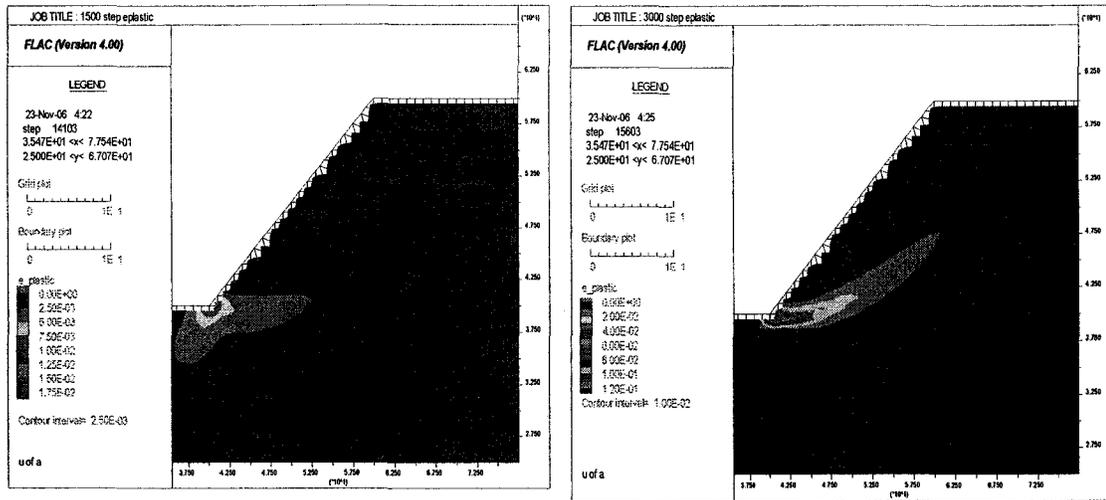
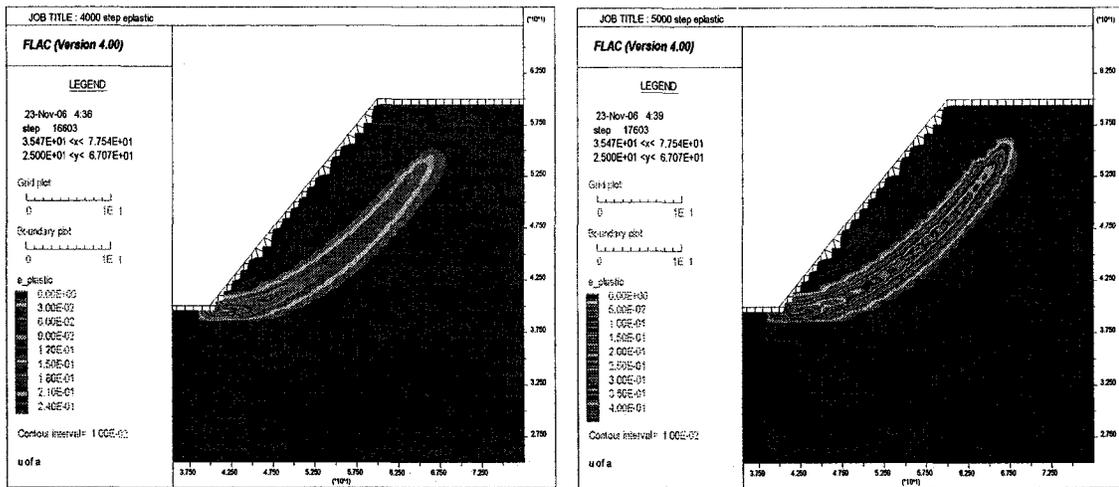


Figure 6.10 The potential slip surface obtained from hardening/softening model implemented in FLAC.



(a)

(b)



(c)

(d)

Figure 6.11 The evolution of potential slip surface at different plastic strain levels: (a) (0-0.017); (b) (0-0.12); (c) (0-0.24); and (d) (0-0.4).

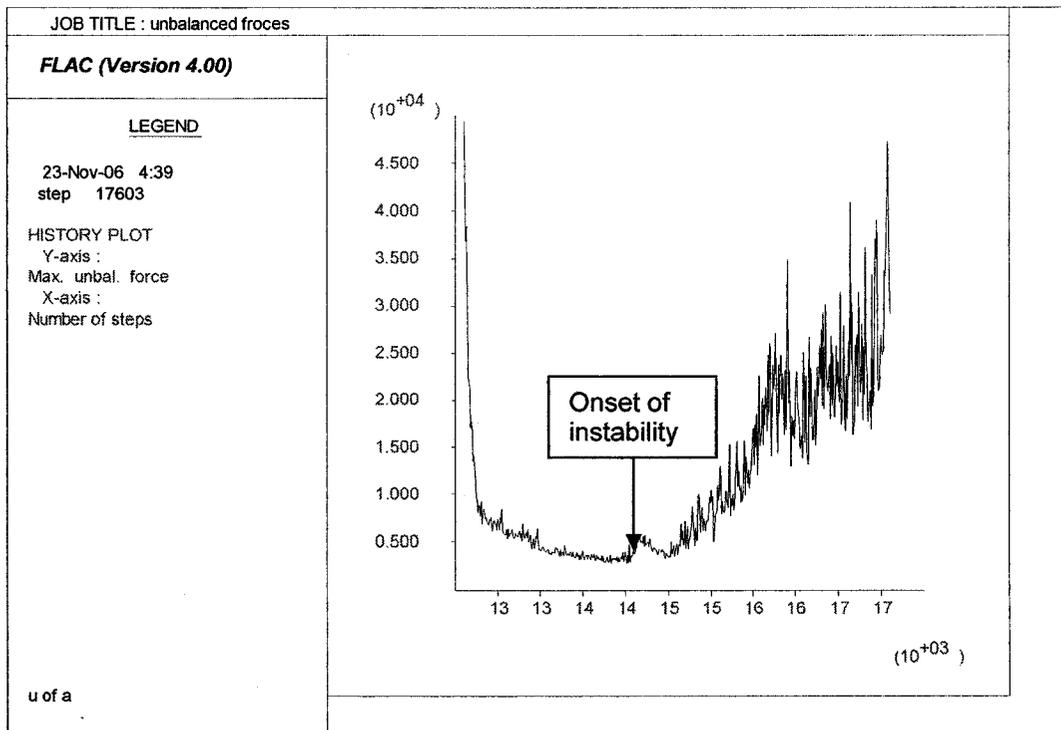


Figure 6.12 The unbalanced forces from hardening/softening model implemented in FLAC.

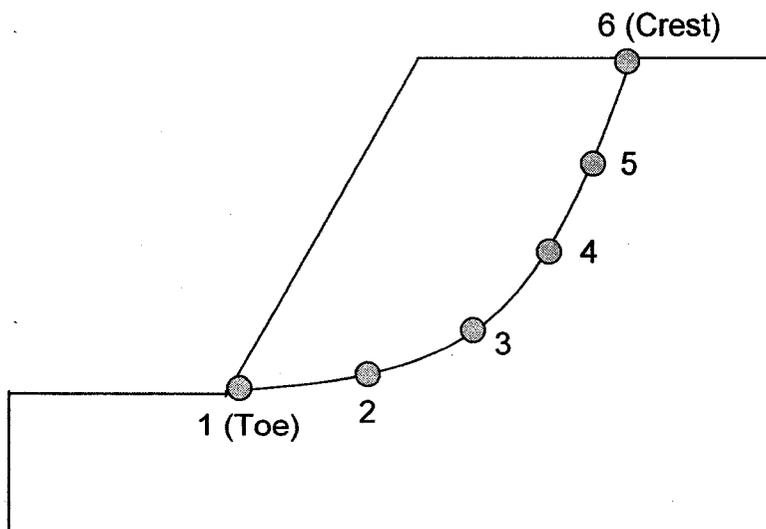


Figure 6.13 Monitoring point along the potential slip surface.

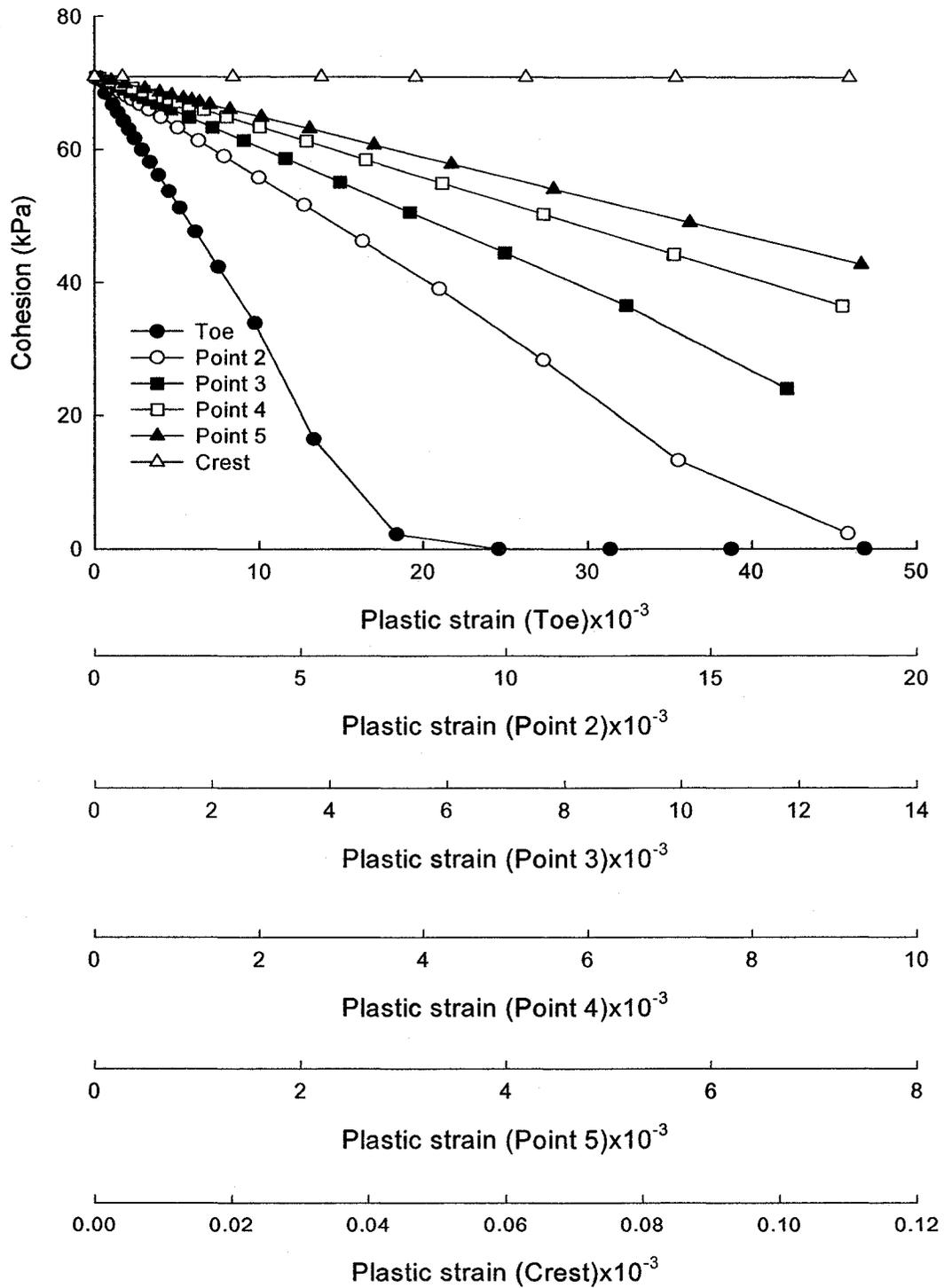


Figure 6.14 Distribution of cohesion along the potential slip surface.

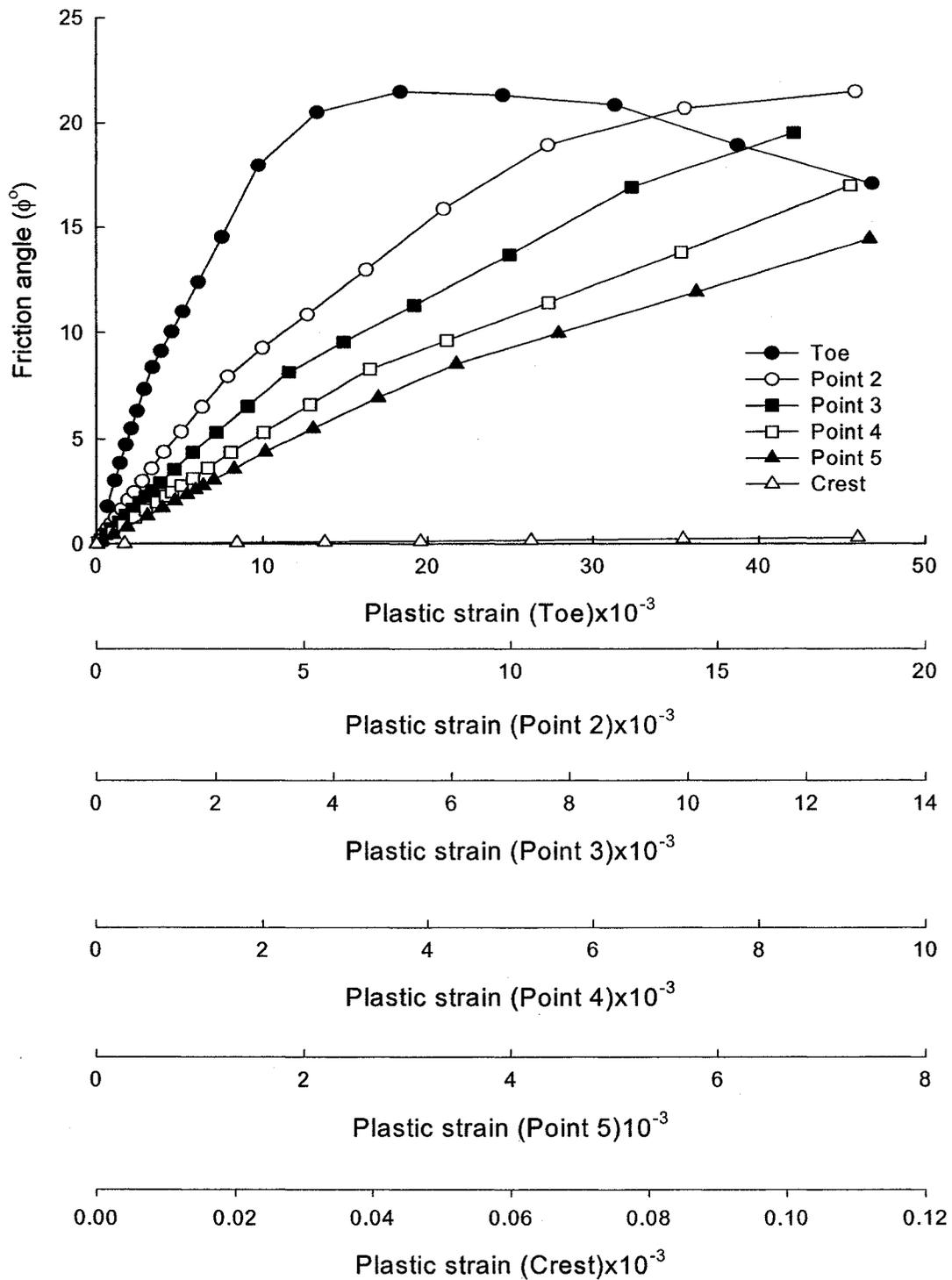


Figure 6.15 Distribution of friction angle along the potential slip surface.

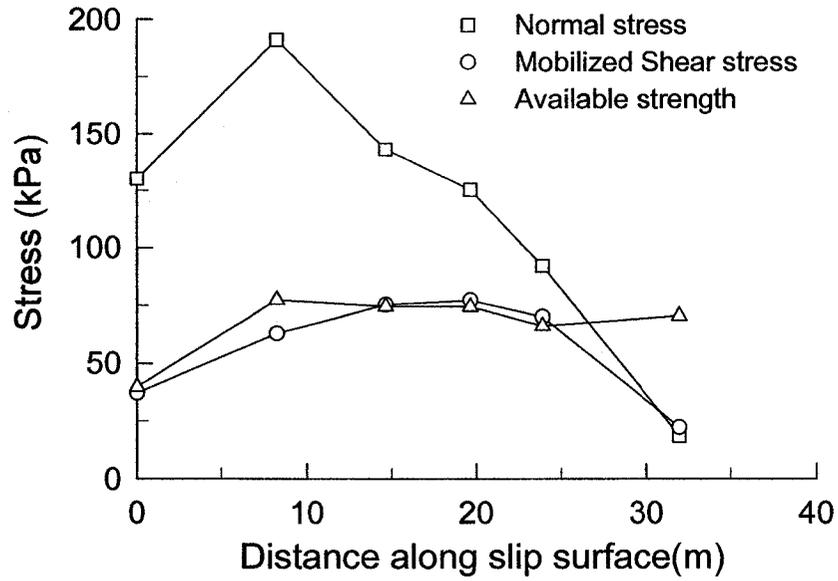


Figure 6.16 Distribution of normal, shear stresses, and available strength along the slip surface associated with SDP.

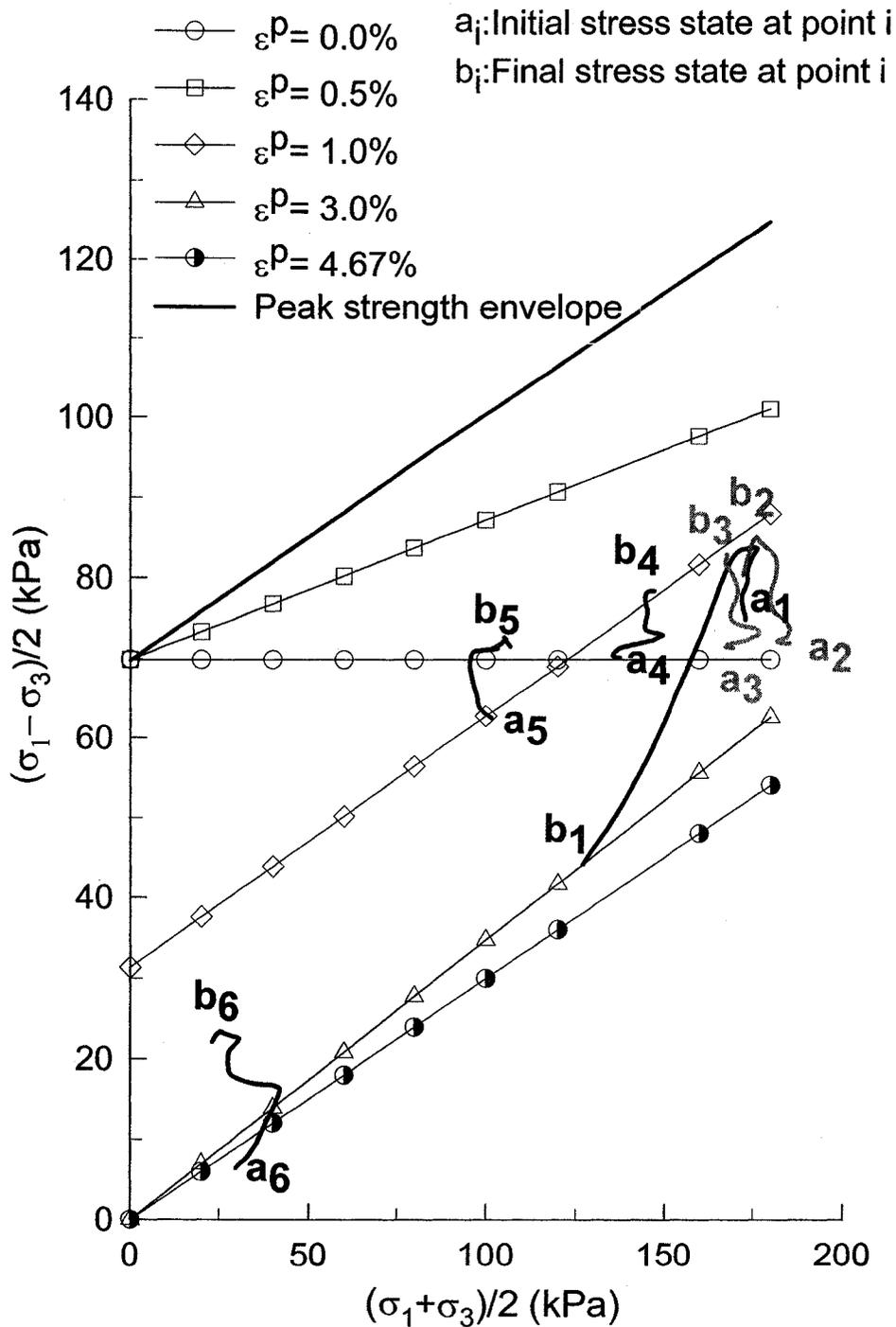


Figure 6.17 Stress paths of the monitoring points along the slip surface.

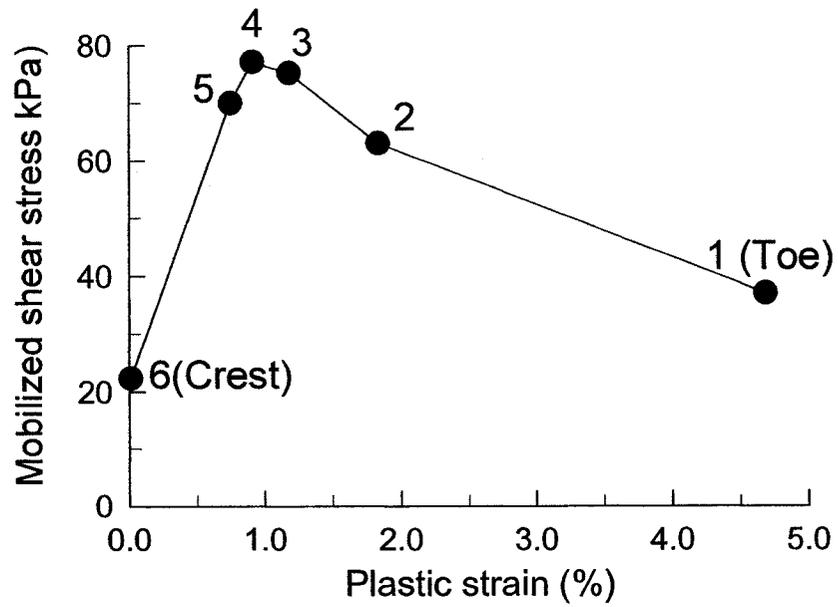


Figure 6.18 Idealized field stress-strain curve.

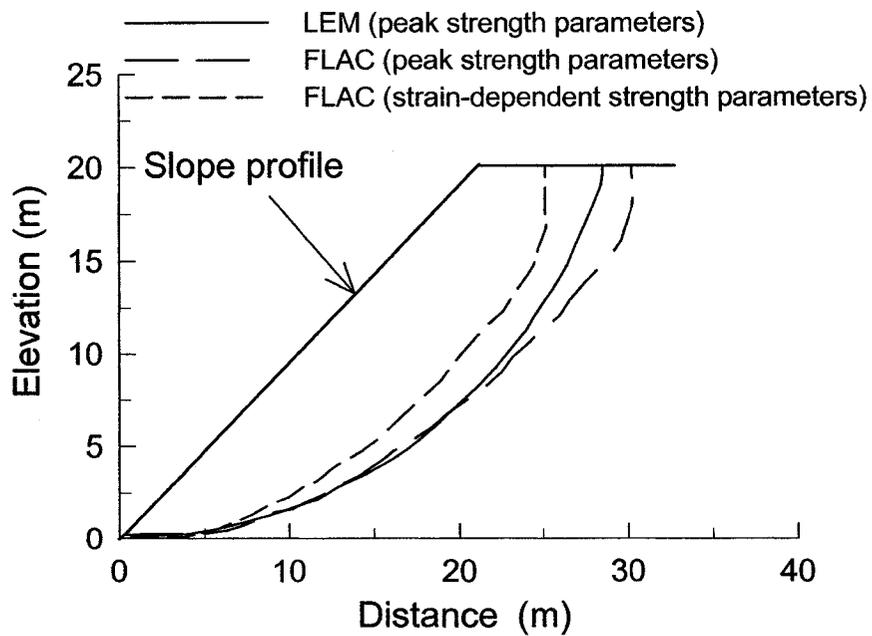


Figure 6.19 A comparison between the positions of the slip surfaces based on peak strength and strain-dependent strength parameters.

CHAPTER 7: SUMMARY, CONCLUSIONS, AND RECOMMENDATIONS

7.1 General

This chapter presents a summary of the research objectives, and the conclusions reached during this investigation. Recommendations for future research are also provided.

7.2 Objectives of the thesis

The experience gained by back analysis of civil engineering applications in stiff clays and weak rocks has shown that the operational strength can not be expressed in terms of peak or residual strength. Despite the wealth of information on the strength of geomaterials, no generalized framework exists for identifying the operational strength, particularly in the intermediate zone between soils and rocks (quasi-frictional materials). The main challenge in determining the operational strength for such materials lies partly in capturing the relative contribution of the interlocking/bonding and frictional resistance at working stress and strain levels, or in the pre-peak regime.

The main objectives of this investigation can be summarized in the following points:

1. To examine high-quality laboratory results for various geomaterials to understand failure precursors in the pre-peak regime, and to capture the onset of failure strength thresholds for soils and rocks.
2. To conduct a comprehensive experimental study on a McMurray fabric-dominant weak rock (locked sand) to reveal the following:
 - a. The source of the high uniaxial strength of such a sand formation in the absence of bonding between particles and to confirm this formation's fitting into locked sands category;
 - b. The significance of local strain, compared to our routine external strain measurements for McMurray sand formations, in terms of failure mechanism, deformation modulus, and post-peak response;
 - c. The pre-peak damage thresholds for locked sand to capture the damage evolution based on the variation of local axial and lateral strains;
 - d. The contribution of the frictional and the interlocking strength parameters with the plastic shear strain;
 - e. The relationship between the onset of localization and the onset of dilation as a candidate for the strength threshold;
 - f. The mechanism and kinematics of failure for locked sand;
 - g. The relationship between the onset of dilation and the long-term strength (onset of creep) of locked sand as a possible candidate for long-term strength; and
 - h. The upper and lower strength thresholds for locked sand.

3. To develop a simplified approach for partitioning the shear strength parameters (cohesion and friction angle) with plastic strain for stiff clays, and to show the significance of non-simultaneous mobilization of strength parameters in revealing:
 - a. The resemblance in bonded geomaterials' shearing responses at low confining pressure irrespective of these geomaterials' origin; and
 - b. The transition between soils (frictional materials) and rocks (bonded materials) based on the mutual contribution of frictional and bonding resistance to overall resistance.
4. To highlight the significance of the concept of non-simultaneous mobilization of strength parameters in capturing the progressive failure of slopes and to interpret the deviation of the operational strength of stiff clays from the peak strength with such conceptual framework.

7.3 Summary and conclusions

This section summarizes this thesis' research and conclusions. The first part of the thesis included a critical review of high-quality test results for various geomaterials to reveal the significance of salient pre-peak strength thresholds as compared to the peak and residual strengths. The second part provided an investigation of the behaviour of a locked sand (a quasi-frictional material with an interlocked fabric) to show the significance of local strain measurements, to identify pre-peak damage thresholds based on variations of the local strains, to

provide an understanding of failure's evolution and mechanism, to show the relative contribution of strength parameters with strain, and to give recommendations about the operational strength of such geomaterials as an example of quasi-frictional materials with interlocked fabric. The third part suggested a simplified approach to visualize the relative contribution of bonding and frictional resistance to the overall resistance for bonded geomaterials at any strain level. This concept was then employed to show the resemblance of bonded geomaterials' behaviour at low confining stresses regardless of their origin and to capture the transition between soils and rocks based on the mutual contribution of the frictional and bonding resistance to the overall resistance. The fourth part of the thesis showed the significance of the strain-dependent strength parameters in capturing the progressive failure of slopes and in revealing the discrepancy between the operational strength and the peak strength using numerical simulation.

7.3.1 Pre-peak strength thresholds for geomaterials

High-quality shear strength and volume change data for various soils and rocks were compiled and reviewed to reveal the salient pre-peak strength thresholds by using microstructural observations and localized strain measurements. The first part of the discussion revealed the relationship between the onsets of dilation (the volumetric strain reversal stress level) and localization (the loss of strain-uniformity) for various geomaterials. In the second part, three examples were provided to investigate the relationship

between the onset of dilation and the long-term strength. In the third part, another threshold was introduced for bonded geomaterials: extensile induced dilation (the initiation of bond degradation). Based on data analysis, the following conclusions were drawn.

1. Geomaterials display a marked change in their behaviour prior to reaching the laboratory peak strength.
2. The onset of shear-induced dilation (volumetric strain reversal) is an excellent indicator of the onset of localization for the investigated geomaterials regardless and can be directly estimated from global measurements.
3. The peak and residual strengths measured in our laboratory are a direct result of progression of localization and can be merely considered as merely signs of failure.
4. The ratio of the stress level required for initiating shear-induced dilation normalized to the peak strength ranges from 0.76 to 0.86 and is independent of fabric and/or bonding nature.
5. For the investigated geomaterials, the onset of shear-induced dilation is an excellent indicator of the onset of creep. Beyond this threshold, creep potential becomes significant.
6. The onset of shear-induced dilation may be considered as an upper strength threshold for the investigated geomaterials.

7. The onset of extension-induced dilation marks the initiation of bond degradation for bonded geomaterials and can be captured only by local lateral strains measurements.
8. The use of pre-peak strength thresholds (the onset of shear-induced dilation and the onset of extension-induced dilation) seem promising to determine the upper and lower strength thresholds for geomaterials.

7.3.2 High-Hill River characteristics and significance of local strain measurements

An up-to-date synopsis of the common characteristics of locked sands was presented. Based on the established framework, experimental work was presented to establish the fit of the High-Hill oil-free sand formation in the locked sands category. The experimental work was comprised of micro-structural observation, mineralogical composition analysis, and fabric analysis. Uniaxial compressive strength tests were also carried out to reveal the strength of the locked sand's unbonded coherent mass and to explore the significance of local strain measurements. The experimental set-up and the calibrations, proof testing, and stability tests were presented and discussed in detail.

It was shown that the High-Hill River sand formation is a lower Cretaceous McMurray Formation with coherent mass. The dense nature of the samples, the mineralogical composition, and the micro-structural observations were in

keeping with the characteristics of locked sands. This formation was best described as overconsolidated well-sorted unbonded friable locked sand with characteristics similar to those of oil-sands' middle members. The formation was subjected to mild post-depositional diagenesis that transformed its detrital sand-sized particles into a weak rock with interlocked fabric through deposition of authigenic silica. SEM images and petrographic analysis of thin sections disclosed that the grain contacts were dominated by tangential, long-to straight, occasional interpenetrative and sutured contacts created via overgrowth and dissolution. The small percentage of kaolinite in the form of coating or bridges was responsible for this deposit's friability. The fine particles presumably originated from either the diagenesis of the overlying shale formation or the feldspar inherited from parent rocks. The deposit appeared to have been subjected to overburden loads that were adequate to cause this fracturing, and, hence, influenced the structure of the deposit. The clean cracks within quartz grains implied that either the clay minerals were deposited following the compaction stage, or the clay minerals were so weak nature to prevent loss of porosity due to compaction by early cementation. The existence of a fresh skeletal grain by dissolution suggested that the process of dissolution occurred after the compaction stage and is still in process; otherwise, the compaction process would have destroyed this skeleton.

A local measurement system for both axial and lateral strains was used and modified for this friable material in conjunction with a routine external axial measurement system for characterization. The formation exhibited a uniaxial compressive strength of 2 MPa and, hence, was classified as a “weak rock”. The interlocking between particles was the main source of this high uniaxial strength and axial stiffness in the absence of bonding between sand particles. The axial stiffness based on local strain measurement was approximately 0.87 GPa and almost 2 times the externally measured axial stiffness. In the post-peak regime, both measuring systems indicated strain weakening; however, the local one exhibited a slightly less abrupt reduction in strength. In terms of the failure mechanism and strain-rate effects, the variation in the local axial strain rate provided more insight into failure evolution stages than externally measured axial strain rate. Hence, for proper characterization of such geomaterials and their oil-sand counterparts, the interpretation of experimental founded on commonplace external strain measurements are less reliable, particularly when the coupling strength and deformation are targeted. Local strain measurements are necessary to obtain proper behavioral characterization and to interpret the failure mechanism interpretation of locked sands.

7.3.3 High-Hill river geomechanical characteristics: damage thresholds, mobilized strength parameters, failure kinematics, and long-term strength

The geomechanical characteristics of the High-Hill River sand formation were explored in detail under uniaxial compression test conditions. The first part of the discussion was to identify damage thresholds that marked a significant variation in the stress-strain response identified by the variation in the local axial and lateral strains. Four thresholds were identified and discussed in terms of the failure mechanism during shearing. The relationship between the onset of dilation (the damage coalescence threshold) and the damage localization threshold was established by using local strain laboratory measurements results. The mechanism and kinematics of failure in locked sand was also described. For practical purposes, the contribution of frictional and interlocking strength parameters at different strain levels was obtained, and their mutual contribution was discussed. In the second part, the discussion explained how the columnar structures and kinematics of deformation for locked sand were captured by the creation of a relatively homogeneous strain field throughout the shearing process. The experimental observations were compared to investigate the dominant mode of deformation at the onset of localization. The formation of dilation bands as the dominant mode of deformation in the pre-peak regime was discussed and its significance during the progression of failure in locked sands was substantiated by using the theoretical framework previously established for

porous rocks. The third part presented a multistage uniaxial creep test to confirm the significance of the onset of shear-induced dilation as an excellent indicator of long-term strength. Finally, the direct shear test results were presented to reveal the maximum mobilized strength under pure shear conditions. The following conclusions were made in the chapter.

1. The exceptional high uniaxial strength and axial stiffness compared to those in dense sands were the by-products of the particles' angularity and the interlocking between particle contacts and was manifested as enhanced ability to resist moments at contacts.
2. The stress-strain curves were partitioned in the pre-peak regime based on the local axial and lateral strain responses to describe the evolution of failure, i.e., the closure of cracks and pores, initiation of unlocking, onset of localization, and coalescence and unstable propagation of damage.
3. No coupling occurred between the axial and lateral stiffness. The lateral strain exhibited extensile strains throughout the shearing process.
4. The maximum dilatancy factor, 3.6, for the investigated locked sand exceeded the maximum dilatancy factor for dense sands.
5. The onset of unlocking (the damage initiation threshold) occurred at 28 % of peak strength and was coincident with our observations of the crack closure strength threshold.
6. The onset of localization and dilation took place at 55% and 67% of the peak, respectively.

7. The interlocking and frictional strength parameters were not simultaneously mobilized. The interlocking component degraded with the plastic strain, resulting in the progressive mobilization of frictional resistance. The interlocking component tended to be consumed in the pre-peak regime and leveled off before attaining peak strength.
8. At peak strength, the resistance resulted from the frictional contribution (basic friction and dilation) with less contribution from the interlocked component (accounting for 17% of peak strength) due to its consumption during the unlocking stage.
9. At the onset of dilation threshold, 50% of the mobilized resistance was attributable to the interlocked fabric. The remaining resistance reflected the mobilization of the basic frictional resistance.
10. Columnar structures were captured by the creation of a homogeneous strain field to hinder the onset of localization and to allow for the complete formation of such columns/slabs. The pattern of failure was dominated by the formation of vertical slabs/walls in the direction of major principal stress as a result of uniform unlocking through lateral expansion and the formation of dilation bands.
11. The comparison of this study's experimental results to those of the theoretical studies of failure modes of porous brittle rocks revealed that dilation bands were the dominant mode of deformation (a predominantly open-mode of deformation propagating in the direction of the major principal stress) and whereby such columnar failure patterns were created

12. Under constant loading conditions, the material exhibited a creep potential that became more significant beyond the onset of dilation.
13. Creep lateral strains and their rates were more significant than axial strains. Creep lateral strain rates were one order of magnitude greater than axial strain rates beyond the onset of dilation.
14. Creep volumetric strains were dilative in nature at all stages of creep. This finding signified the progressive damage and damage propagation in axial direction during creep stages.
15. The results of DSB tests indicated that the mobilized strength under low normal strength was independent of the normal stress levels. The mobilized strength was dominated by the interlocking strength component and was equivalent to the damage initiation strength threshold estimated from the uniaxial compression test results.
16. In boundary value problems constructed in or on locked sands, damage coalescence and initiation thresholds are recommended to be the upper and lower strength boundaries. However, further work is required to investigate such argument at higher confining pressures.

7.3.4 Simultaneous mobilization of strength parameters for bonded geomaterials

For locked sand, it was shown that the concept of the non-simultaneous mobilization of strength is very important to interpret the failure mechanism and to capture failure evolution. This concept was extended to explain some of the peculiar pre-peak characteristics of bonded geomaterials, to reveal the

contribution of strength parameters at different confining pressures, and to capture the transition between soils and rocks.

The contribution of frictional resistance and bonding resistance was investigated at various strain levels for two natural stiff clays by using a simplified partitioning approach. The partitioning of the strength parameters was presented in two different stress spaces: M-C space and q-p space. In M-C space, the partitioning was presented in terms of the cohesion and friction angle for the two investigated clays. The results revealed that the cohesion/bonding and friction angle were not simultaneously mobilized (as the conventional M-C hypothesis implies based on the peak strength criterion). The cohesion/bonding distribution's trend indicated progressive degradation with the plastic strain level at low strain levels. In contrast, the friction angle built up progressively with the plastic shear strain and necessitated a higher strain to be mobilized. In terms of q-p space, partitioning was carried out in terms of two normalized parameters corresponding to the cohesion/bonding and friction strength parameters for Nanticoke clay. An incremental equation based on the normalized strength parameters was proposed to reveal the contribution of frictional and bonding resistance at various strain levels as follows:

$$[7.1] \delta q = \delta q_1 + \delta q_2,$$

$$\delta q_1 = \underbrace{\left(p' M_{\text{crit.}} (-cd(\exp(-d\varepsilon^p)) + ef(\exp(-f\varepsilon^p))) + M_{\text{crit.}} (c(\exp(-d\varepsilon^p)) - e(\exp(-f\varepsilon^p))) \right) \delta p'}_{\text{frictional component}} \delta \varepsilon^p, \text{ and}$$

$$\delta q_2 = \underbrace{a M_{\text{crit.}} p_{\text{to}} (\exp(-b\varepsilon^p)) \left((\exp(-d\varepsilon^p)) (-cd - bc) + (\exp(-f\varepsilon^p)) (eb + ef) \right)}_{\text{bonding component}} \delta \varepsilon^p$$

A parametric study was carried out to show the bearing of the non-simultaneous mobilization of strength parameters on the stress-strain curves and failure envelope. It was shown that strain hardening and weakening are the by-products of the balance between cohesive strength degradation and frictional strength mobilization at different confining pressures. Three yield surfaces were captured that sufficiently described the failure surface and its nonlinearity, the bonding contribution and its degradation effects, and the balance between bonding strength degradation and frictional strength mobilization as reflected in an initial “plateau” in pre-peak regime. The influence of bond degradation was manifested through the initial reduction in resistance in the pre-peak regime, followed by either strain hardening or weakening. Strain hardening was shown to be controlled by the ability to mobilize full frictional resistance after initiation of bonding degradation. The compiled experimental results for various bonded geomaterials of different origins revealed that this initial plateau in stress-strain curves is a characteristic behavior where bonding is a major source of strength. In addition, the transition between hard rocks and soils was also captured by using such framework, and was substantiated by experimental results of

different bonded geomaterials. The transition was described in terms of the contribution of bonding and frictional resistance to the peak strength.

7.3.5 The significance of strain dependent strength parameters in slope stability problems

A parametric study was carried out to reveal the significance of non-simultaneous mobilization of strength parameters in the assessment of the stability of slopes. The focus was to explain the deviation of the operational strength from the peak strength. For this purpose, two commercial software were used for the simulation of a 45° homogeneous slope with a height of 20 m. The finite difference software (FLAC) was used to reveal the potential slip surfaces associated with peak strength parameters and the strain dependent strength parameters by using the hardening/softening model. The factor of safety, based on peak strength parameters, was determined by using the strength reduction technique. The slope stability analysis was also carried out by using limit equilibrium software (SLOPE/W) to reveal the factor of safety and the slip surface associated with peak strength parameters. The distribution of the strength parameters, normal and shear stresses, using the hardening/softening model, along the potential slip surface were also estimated.

The results of the finite difference model and LEM showed that the factor of safety associated with peak strength parameters was 1.43 and 1.45,

respectively. The LEM's slip surface was relatively shallower than the one estimated based on the strength reduction technique. The strain-dependent strength parameters' slip surface was shallower than both of those estimated from LEM and FLAC. The numerical simulation using the hardening/softening model captured the progressive failure of the simulated slope. The mobilized shear strength parameters values varied along the slip surface according to the experienced plastic strain level. The plastic strain at the toe was at least one order of magnitude greater than the plastic strain at the crest. For an average value of normal and shear stresses, the peak strength parameters overestimated the factor of safety. On the other hand, the residual strength parameters underestimated the factor of safety. The mobilized strength to render a factor of safety of 1 was 54% of the peak strength and accounted for the progressive failure of the slope. This was only captured through incorporation of the strain-dependent strength parameters. It was concluded that this conceptual framework is a powerful framework to explain the discrepancy of the mobilized strength in the field from the peak strength. However, more work is needed to confirm such conclusion based on well-documented case histories.

7.4 General framework for simulation of boundary value problems in strain weakening materials

Based on the presented work in this thesis, it is recommended to consider either of the following approaches to carry out a numerical simulation:

1. Incorporating the onset of dilation as an upper strength threshold within an elasto-plastic model along with appropriate deformation parameters that are based on local strain measurements.
2. Incorporating the strain-dependent strength parameters in a strain hardening/softening model to capture the overall failure mechanism. However, the results should be treated carefully due to the dependence of the shear strains on the size of the grid. Initially, the model should be calibrated with high quality experimental data so that the model can be as close as possible to the real conditions.

7.5 Recommendations for future investigation

The following topics are suggested as a potential research subjects:

1. The effect of incorporating the concept of non-simultaneous mobilization of strength parameters in evaluating the stability of real boundary value problems: reanalyze well-documented case histories with high quality experimental data using such framework and compare the results with traditional design approaches.
2. The effect of strain rate, sample configuration (the slenderness ratio), and end effects (lubricated vs. non-lubricated ends) on the normalized dilation ratio: characterize the temporary strain hardening beyond the onset of shear-induced dilation for bonded weak rocks.

3. The significance of pre-peak thresholds as operational strength thresholds: reanalyze well-documented case histories augmented with high-quality experimental data obtained from local strain measurements.
4. The relationship between the ratios of onset of shear-induced dilation normalized to extensile-induced dilation and the ratio between the peak and residual strength for bonded weak rock under different confining pressure: conduct a comprehensive experimental investigation using local strain measurement and determine these thresholds using the framework in this thesis.
5. Exploiting discrete element method with bonded model, e.g. PFC to better understand the concept of the non-simultaneous mobilization of strength parameters and its relation to the overall mobilized strength: use the built-in bond model in PFC .
6. It was recommended in Canadian sensitive clays, that the mobilized strength in slope stability problems can be roughly estimated at 8% strain. However, this is only an approximation, and an interesting research subject is to reveal the relationship between the 8% strain strength threshold and damage thresholds: procure block specimens and conduct an experimental program at low stress levels to reveal the damage thresholds, creep limit, and its relation to the 8% strain strength level.

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Appendix A

Published conference papers

This appendix contains the conference papers published during the preparation of this thesis in a chronological order. The full citation for these papers is as follow:

1. Abdelaziz, T., Martin, C.D., Chalaturnyk, R. 2005. Mechanistic view of strength mobilization in stiff clays and weak rocks. In the Proceedings of the Fifth international Geotechnical Engineering conference, January 2005, Cairo, Egypt, pp. 171-182.
2. Abdelaziz, T., Martin, C.D., Chalaturnyk, R. 2005. The onset of dilation in soils and rocks. *In the 58th Canadian Geotechnical Conference & 6th Joint CGS & IAH-CNC Groundwater Specialty Conference, 11-14 September 2005, Saskatoon, Canada, Paper No. 583.*
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MECHANISTIC VIEW OF STRENGTH MOBILIZATION IN STIFF CLAYS AND WEAK ROCKS⁴

Abdelaziz, T. ¹, Martin, C.D.², Chalaturnyk, R.³

ABSTRACT

Research is underway to formulate a constitutive model that captures the effect of bonding for both shear strength and volume change for bonded natural geomaterials within the framework of critical state soil mechanics. In order to provide a preliminary understanding of mobilization of strength with strain Mohr- Coulomb strength parameters are partitioned with plastic strain levels for two natural bonded geomaterials. Results have shown that bonding strength and frictional strength are not mobilized simultaneously as implied in conventional Mohr-Coulomb hypothesis. Bonding strength is mobilized at small strains and degrades dramatically with plastic shear strain. At the same time mobilized frictional resistance entails larger plastic strain level to be mobilized. These results also reveal that the peak strength is a function of the interplay between the frictional and bonding resistance. At low confining pressures the bonding resistance is dominant while at high confining pressures frictional resistance is the controlling parameter. Strain weakening is a function of the rate of friction mobilization and bonding degradation. At low stress levels the strain weakening is due to the degradation of the bonding resistance while at high stress levels it is basically due to the reduction of the frictional resistance. At intermediate confining pressures it is the interplay between the two components of strength. Extreme caution should be considered when adapting strength parameters in solving boundary value problems using conventional approaches due to bond degradation and slow rate of frictional strength mobilization.

INTRODUCTION

Leroueil and Vaughan (1990) have unambiguously indicated that all natural geomaterials share the same characteristics irrespective their origin. In the field of geomechanics there is a class of materials that constitutes the boundary between soils and rocks; stiff clays and weak rocks. The complexity of such materials stems from the quasi-frictional nature of behaviour due to the existence of bonds as a main source of strength.

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Degradation of bonds with strain and strain weakening is a characteristic behaviour as indicated by several investigators through probing mechanical behaviour of artificially cemented soils, natural stiff clays and weak rocks (Cuccovillo and Coops, 1993; Maccarini, 1987; Lo and Wardani, 2002 and Lo et al., 2003; Burland, 1990; Cotecchia and Chandler, 1997; Cecconi and Viggiani, 2001; Anagnostopoulos et al., 1993). In these natural geomaterials bonding has been acquired during geological history and should be considered as an intrinsic parameter rather than a fitting parameter as for unbonded geomaterials (Burland et al., 1996; Leroueil and Vaughan, 1990, Adachi and Oka, 1993 and Oka et al., 1998). Debate regarding the source of shearing resistance has been an issue for decades. The first investigation into the physical meaning of strength parameters was carried out by Hvorslev (1960). Hvorslev showed that the shear strength of remolded saturated clays could be partitioned into two components. Effective cohesion which is a function of void ratio and effective friction which is a function of the effective stresses. Lambe (1960) also argued that the strength comprises three components; cohesion, friction and dilatancy. Cohesive component according to his perspective is mobilized at low strain and may not contribute to the peak resistance. Schmertmann and Osterberg (1960) opposed the idea of Hvorslev (1960) that void ratio is representative of natural soils. They showed that the shear strength parameters are not mobilized simultaneously; cohesion is mobilized at low strain while friction requires sufficient strain to be mobilized. De Mello (1988) has shown that the available strength at working strain levels is much lower than the maximum deviator stress concept using a certain factor of safety. Conlon (1966) pointed out that the stability of slopes in stiff clays depends on the degree of mobilization of friction relative to the degradation of the cohesive component, which cannot be captured by assigning constant parameters. The use of shear strength parameters, as a function of strain is beneficial to interpret the failure mechanism in natural geomaterials. Therefore, it can be concluded that there is a need to partition the strength parameters according to strain to capture the failure mechanism in natural stiff clays and in general bonded geomaterials.

Based on the forgoing it can be concluded that shear strength should be expressed as a function of strain. Integration of frictional resistance, bonding resistance, and volume

change is the main constituent of advanced constitutive models (Gens and Nova, 1993; Lagioia and Nova, 1993; Tamagnini and D'Elia, 1999; Lui and Carter, 2002; Lee et al., 2004). However it is argued herein that a simplified approach to partition strength components based on Mohr-Coulomb can shed the light into the mechanism of failure as a preliminary step for formulating more comprehensive models. Therefore, strength parameters will be estimated as a function of plastic shear strain using a simplified approach. Interpretation of shear strength data will be provided in light of distributions of cohesion and friction as intrinsic parameters for the investigated stress range.

PREDICTION OF STRAIN DEPENDENT STRENGTH PARAMETERS

Shear strength parameters are to be predicted as a function of plastic shear strain. Plastic shear strain can be considered as the most proper parameter to capture changes in microstructure from a micro-mechanical viewpoint (Vardoulakis and Sulem, 1995; Sulem et al., 1999). Values are to be estimated by fitting the strength data at different values of plastic strain. The end product is a multiple yield envelopes at different strain levels. Thereafter the values of cohesion and friction are plotted with plastic shear strain as will be shown later. In this way the yield function at any strain can be written as:

$$f(\sigma, \varepsilon^p) = c(\varepsilon^p) + \sigma \tan \phi(\varepsilon^p)$$

Where f is the yield surface, c is the mobilized bonding strength (cohesion), ϕ is the mobilized friction angle, and ε^p is the plastic shear strain. In the context of plasticity the cohesion and friction are the hardening and softening parameters that control the evolution of the yield surface. All discussion herein will be in terms of effective stresses unless clearly stated otherwise.

NANTICOKE CLAY

Nanticoke clay is a stiff fissured moderately overconsolidated from Southern Ontario. Lo (1972) studied the strength of Nanticoke clay obtained from 1.4 inch specimen performed at a strain rate of 0.05%/hr under several confining pressures (Figure 1). It can be observed that the peak stress difference occurs at increasing strains, and the amount of

strength drop decreases with the increase in confining pressure. Based on the results of the triaxial compression tests, the data were analyzed to obtain the multiple yield envelopes as shown in Figure 2. From this Figure the distribution of the parameters is shown in Figure 3. It can be observed that both of the components are not mobilized or degrade simultaneously. The cohesive component is fully mobilized at zero plastic strain followed by full degradation.

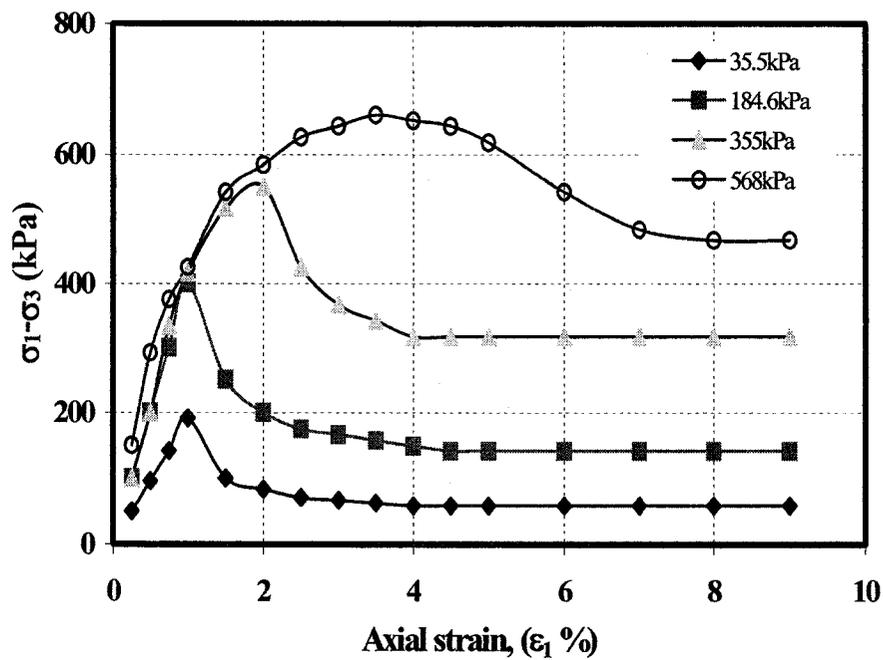


Figure 1. Nanticoke clay stress-strain curves.

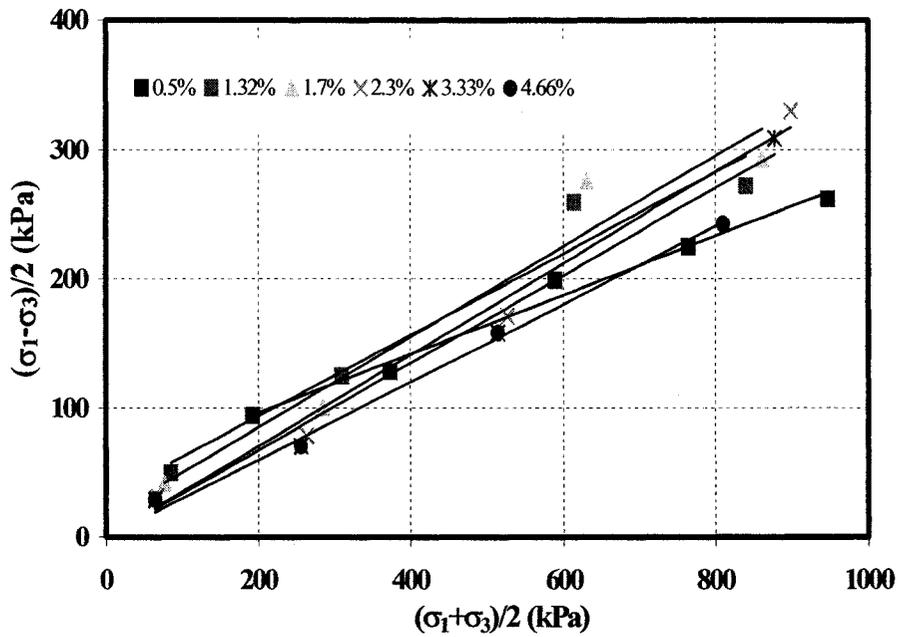


Figure 2. Multiple-Yield envelopes in terms of Plastic shear strain (Nanticoke clay).

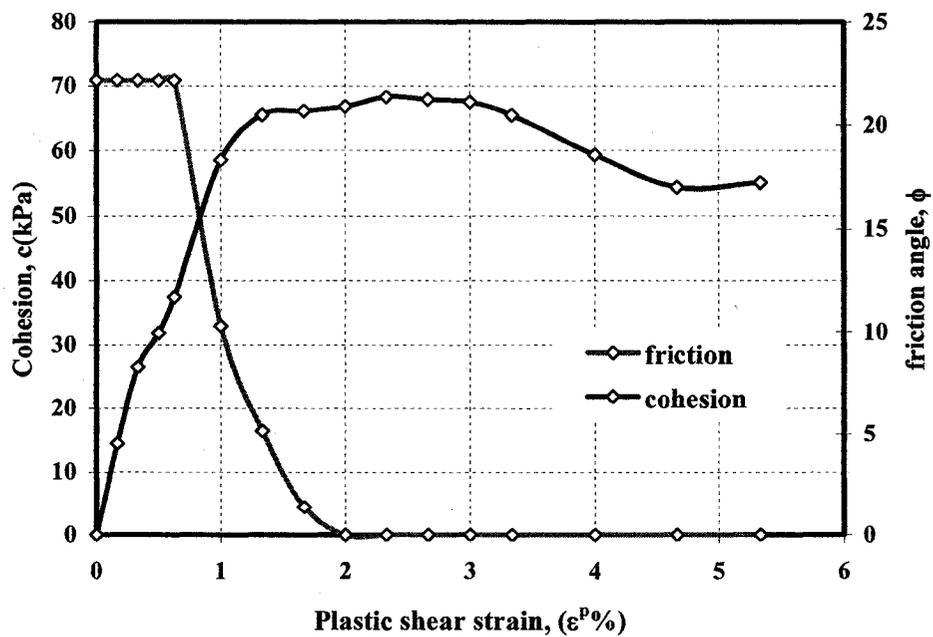


Figure 3. Distribution of shear strength parameters with plastic strain level (Nanticoke clay).

The friction component is mobilized with the plastic strain to a maximum value and then degrades to a slightly lower value. The cohesive component is mobilized to its maximum value (70.89 kPa) up to a strain level of 0.75%, afterwards reduction of the component is observed to zero at a strain level of 2.0%. The reduction percentage of the cohesive component is almost 100%. On the other hand, the friction component attains its maximum value (21.35°) at a strain level of 2.5% and degrades with a percentage reduction of 19%. It can be also seen the percentage of reduction in the frictional resistance is much lower than that of the cohesive resistance. It can be inferred from the distribution that at low confining stresses, the reduction of deviator stresses will be significant while a slight reduction of the load carrying capacity is expected at high confining pressures. The strain weakening is dictated by the balance between the rate of cohesion degradation and friction mobilization. At low confining pressures cohesive component is dominant owing to slow mobilization of frictional resistance the reduction in load carrying resistance can be attributed to reduction of cohesive resistance. At moderate confining stresses, the reduction is a function of the rate of cohesion degradation and friction mobilization. At higher confining pressures, the reduction is dictated by the rate of friction degradation and that is why there is negligible decrease in the load carrying capacity. Based on assigning constant parameters the peak strength parameters are $c=71.12$ kPa and $\phi =17.75$. Assuming a suitable factor of safety for our value boundary problem reveals an important issue. At low strain levels where our world lies and where it is required to limit deformations the mobilized friction is lower than the available one using the factor of safety. On the other hand the cohesive resistance is satisfactory as long as it is not lower than the actual mobilized cohesive resistance.

ANDALUSIAN BLUE MARL

Olalla and Cuellar (2001) carried out a series of undrained triaxial compression tests with pore pressure measurements for samples of carbonated overconsolidated clay known as “Andalusian blue marls” at various values of mean effective stresses. The aim of their investigation was to reveal the cause of Aznalcollar tailing dam failure where it was constructed on this formation. It has been shown that the failure of this tailing dam was primarily due to the extreme brittleness of the foundation. The stress-strain curves are

replicated as shown in Fig. 4. Fig. 5 shows the distribution of the cohesion and friction with plastic shear strain. It can be observed that both of the components are not mobilized or degraded simultaneously. Cohesive resistance is mobilized at zero plastic strain and dramatically vanishes with the increase in strain level. Frictional resistance increases gradually to a peak value at a strain level higher than that for complete vanishing of cohesive resistance followed by a reduction of the frictional resistance at large strain. It can be also seen that the available strength at any strain level is a combination of both components, which cannot be accurately described using only two constants. The cohesive component is mobilized to its maximum value (66.7 kPa), followed by dramatic reduction of with strain level. The reduction percentage of the cohesive component is almost 100% at

a strain level of 1.8%. On the other hand, the friction component attains its maximum value (39°) at a strain level of 2.3 % and degrades to a lower value (19.8°) at large strain. Closer look at the distributions and the stress strain curves reveals that the cohesive component has no contribution to the peak resistance as it degrades at a strain level less than peak strength strain level. However the stress levels that are shown here are high enough that no contribution for the cohesive resistance can be stated. The reduction in the load carrying capacity is related to the confining stresses and the reduction of the friction angle. Based on assigning constant parameters the peak strength parameters are $c=0$ and $\phi = 37.4^\circ$ based on peak values. Consequently, a mobilized values based on factor of safety of peak values can be highly conservative at low strain levels based on peak values at low strain levels. At low strain levels cohesion contributes to the overall resistance and which is ignored using a factor of safety based on peak values.

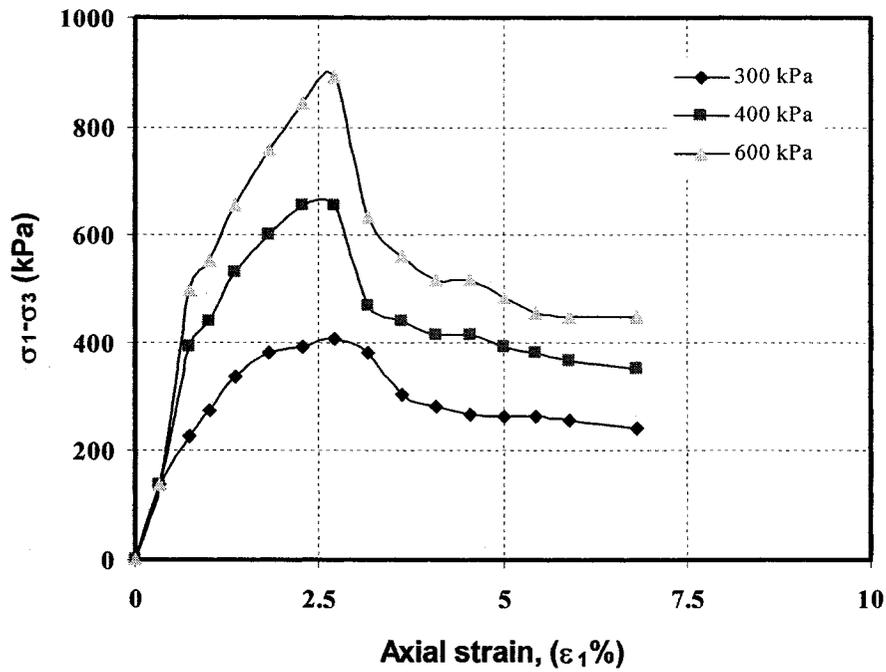


Figure 4. Blue Marl stress-strain curves.

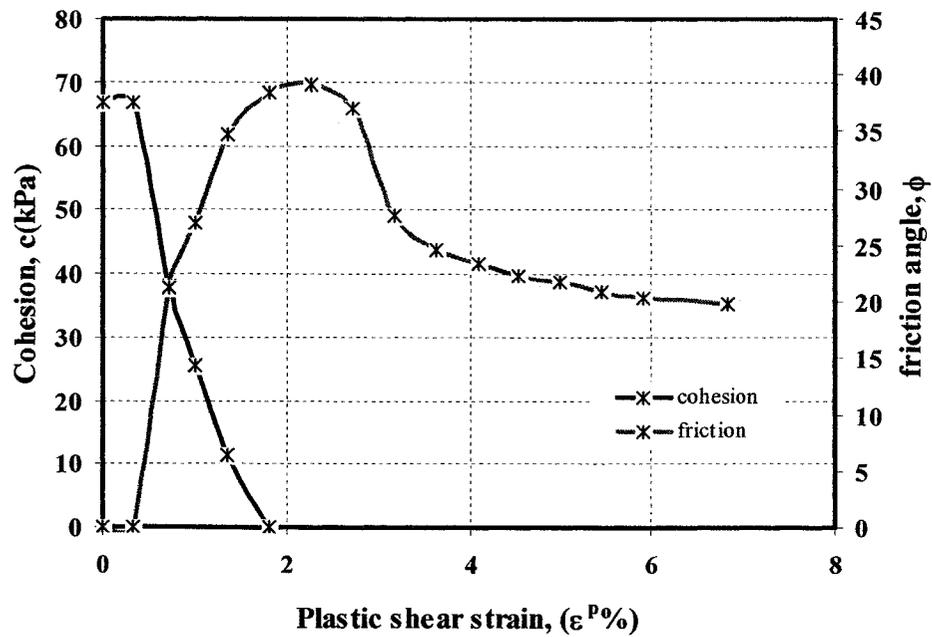


Figure 5. Distribution of shear strength parameters with plastic shear strain (Blue Marl).

MOBILIZATION OF STRENGTH PARAMETERS PHYSICAL PERSPECTIVE

During the shearing process the energy applied by the external agency is dissipated by overcoming the bonds between the particles (cohesion) and by sliding and rolling (friction)

and through the work done for the rearrangement of particles (volume change component). The two components are not simultaneously mobilized during the shearing process. One can speculate that the applied shear stresses should be enough to overcome the cohesive component between the particles prior to allow particles interaction. When the cohesive strength is attained the particles come into contact as particles or blocks. At this stage there should be some redistribution of the shear stresses (due to relative degradation of the cohesive component) that they can be transferred to the contact between particles. This redistribution leads to an increase of the shear stresses at contacts. The resistance to shear stresses is supplied at the contacts due frictional resistance due to interlocking of blocks and sliding of individual particles. Therefore, intuitively it is anticipated that the cohesive component is mobilized at low strains while frictional component demands higher strain level to be fully developed. Further the dilation component is seen to contribute to the frictional resistance at low strain due to the degradation of bonds and formation of blocks. Degradation of bonds and enhancement of dilation is recorded by several investigators for bonded geomaterials and artificially cemented soils (Cresswell and Baron, 2003; Cotecchia and Chandler, 1997; Cuccovillo and Coops, 1999; Lo and Wardani, 2002). Based on the distributions for the two investigated bonded geomaterials, it can be inferred that excessive shearing at contacts results in a reduction of the frictional resistance. One can speculate that the strength degradation in the post peak regime is dictated by the degradation of both cohesive and dilation component along with range of applied stresses (Trollope, 1967). Whenever decision needs to be taken regarding source of strength strain should be considered to determine strength parameters. However generalization of both components contributions cannot be done, as they are a function of stress path as well. At low strain levels and at low confining stresses where our world lies, bonding strength provides a contribution to peak strength while, less contribution from friction can be revealed. Therefore, using a factor of safety based on peak values can render erratic values as these values are nothing

but average of such distributions. In other words, erratic cohesive resistance and higher friction angles can be the end product of such approximation.

CONCLUSIONS

It is observed that the cohesive and frictional shear strength parameters are not mobilized simultaneously. Cohesion is mobilized at small strains reflecting the integrity of the material, as there is no significant degradation in the structure. Once slip starts, significant reduction of cohesion occurs with a gradual mobilization of frictional resistance. Frictional resistance requires larger strains to be mobilized as compared to cohesion. The strain hardening/weakening at various confining stresses is a function of balance between degree of cohesion degradation and frictional resistance mobilization. Strain weakening at low confining stresses is the response of degradation of the cohesive components at low strain with negligible contribution of frictional resistance due to the inability of confining stresses to close the micro-cracks. At higher confining pressures, the weakening is due to the reduction of the frictional resistance due to the excess interaction of the particles during slippage. From the observations of progressive mobilization of strength parameters, bilinear peak strength envelope is expected. The initial portion of the envelope is dictated by the cohesive component while the second portion is governed by the frictional resistance. Therefore, caution should be used when applying limit equilibrium analysis that utilizes a conventional linear Mohr-Coulomb failure envelope. Selected strength parameters can be misleading at low strain levels where our world occurs to limit deformation. In relation to more advanced constitutive models it can be seen that two hardening parameters should be employed or even two yield surfaces to effectively imitate the mechanical behaviour. One describes the degradation of the bonding strength while the other one takes into account the strength due to the basic frictional resistance as well as the volume change contribution to the overall frictional resistance.

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THE ONSET OF DILATION IN SOILS AND ROCKS

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ABSTRACT

Determining the operational strength of bonded geomaterials is a major challenge in geotechnical engineering. Peak and residual strength are customarily obtained from traditional laboratory tests for such purpose. However, bonded geomaterials also display marked changes in their behaviour before reaching the laboratory peak strength. Careful inspection of laboratory stress-strain curves, based on local strain measurements, acoustic emissions, and imaging techniques show that for a wide spectrum of geomaterials, the onset of shear-induced dilation marks the initiation of shear-band localization. This shear-induced dilation typically occurs at a threshold between 0.76 and 0.86 when normalized to the peak strength. Examination of the stress-strain curves for a bonded geomaterial (granite) using a localized measurement technique shows that extensile-induced dilation marks the initiation of bond degradation. The thresholds for shear-induced dilation and extension-induced seem promising to assess the operational strength of bonded geomaterials.

RESUME

La détermination de la force opérationnelle des géomatériaux cimentés est un défi majeur en géotechnique. Le maximum et la force résiduelle sont traditionnellement obtenus à partir de tests en laboratoire. Les géomatériaux cimentés montrent toutefois des changements marqués dans leur comportement, avant d'atteindre leur force maximale mesurée au laboratoire. L'analyse minutieuse des courbes de tensions-contraintes mesurées en laboratoire et basées sur, la mesure de tension locale, les émissions acoustiques et les techniques d'imagerie, montre que pour une large gamme de géomatériaux, le début de la dilatation induite par le cisaillement marque l'initiation de la bande de cisaillement. Cette dilatation induite par le cisaillement se produit typiquement à un seuil compris entre 0.76 et 0.86 de la force maximal. L'inspection des courbes de tensions-contraintes d'un géomatériau cimenté (granite), mesurées par une technique locale, montre que la dégradation du liant est initiée par la dilatation induite par rallongement. Les seuils de dilatation induite par le cisaillement semblent prometteurs quant à l'évaluation de la force de géomatériaux cimentés.

1. INTRODUCTION

The phenomenon of shear-induced dilation was recognized by Reynolds (1885) to describe the increase in volume when dense granular materials are sheared. Reynolds (1885) attributed this distinctive feature to geometrical constraints imposed by particle packing (fabric) for dense granular materials. Today it is recognized that shear-induced dilation is not only a characteristic behaviour of dense granular geomaterials but is also observed in over consolidated stiff clays, weak rocks, and hard rocks. As a consequent, shear-induced dilation can be considered as a direct result of overcoming different constraints (fabric and/or bonding) and reflects a behaviour limit of geomaterials.

Volume change (compression and dilation) in compression laboratory tests and in particular triaxial tests is measured using two techniques. The first technique is through monitoring the change of volume of fluids within specimen pores

and hence it represents a global volume change subjected to system compliance. The second technique depends on measuring, independently, axial and lateral strains from strain gauges mounted on the surface of the sample. The volume change can then be calculated using:

$$[1] \Delta V / V = \varepsilon_{axial} + 2\varepsilon_{lateral}$$

Cook (1970) proved that the volumetric strain of a sample measured by surface strain gauge was a pervasive volumetric property of rocks and not a superficial phenomenon. Hence, the volumetric strains obtained using this technique can be used to evaluate the contractant as well as dilatant behaviour in the testing of both soil and rock.

One of the consequences of shear-induced dilation is the formation of relatively thin shear zones, i.e., only the grains involved in the shear zone formation move significantly. The strength of the dense granular soil is directly linked to the

formation of these shear bands. However, it is not clear how strength should be interpreted from the traditional laboratory triaxial test because the peak strength of the sample is not correlated to the shear zone formation. The onset of the shear zone formation can be recorded by both the global as well as the local volume change measuring techniques.

In addition to the shear-induced dilation discussed above, brittle rocks also exhibit extension-induced dilation during compression testing due to bond degradation. Martin & Chandler (1994) showed, using triaxial compression tests at various confining stresses, that this extension-induced dilation was only observed in the lateral strain gauges and hence would not be observed in testing methods that relied on a global volume change. This may be considered as another strength threshold in this kind of geomaterials.

In this paper, the interpretation of shear-induced and extension-induced dilation in laboratory testing of soils and rock is reviewed. A framework is proposed for assessing the onset of dilation and its relevance in determining the strength of these soils and rock from laboratory test results.

2. SHEAR-INDUCED DILATION

Advances in laboratory testing techniques enabled several researchers to investigate the deformation mechanisms in compression tests for dense sands, stiff clays, weak rocks, and hard rocks. These researchers used local and global measurements, acoustic emissions, and computerized tomography to monitor the deformation process. In this section the results from their work are reviewed and discussed.

2.1 Granular soils and rocks

2.1.1 Dense sand

Finno and Rechenmacher (2003) investigated the deformation of medium to dense concrete sand in drained plane-strain compression test. Samples were prepared using vibratory method to achieve an initial void ratio of 0.543. Figure 1a shows deviator-stress versus axial-strain curve at a confining pressure of 98 kPa. It can be seen from Figure 1a that there is a strength reduction in the post-peak regime.

The onset of localization was identified via local lateral strain measurements using 4 horizontally mounted LVDTs as shown in the insert in Figure 3c. Divergence of local lateral strains with shearing progression at point A shows the onset of localization (Figure 1c). Notice that the onset of localization (Figure 1c) is coincident with the onset of dilation (volumetric strain reversal) in Figure 1b.

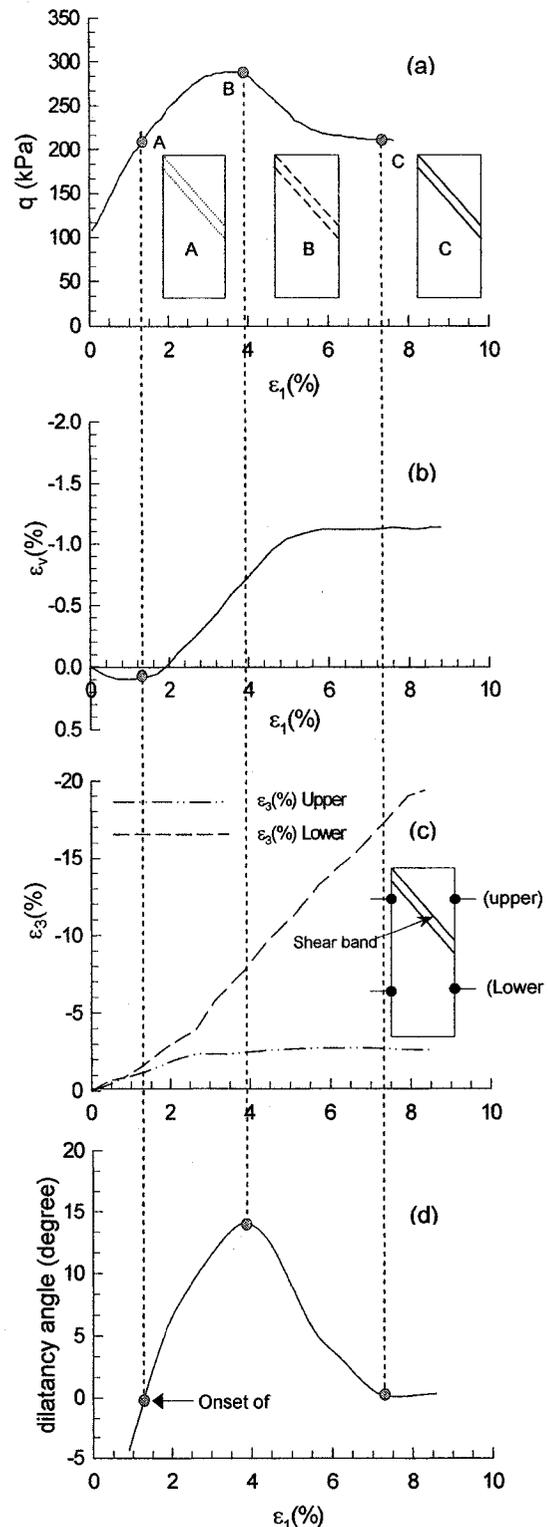


Figure 1. Stress-strain and volumetric strain curves for dense sand (Finno and Rechenmacher (2003)); (a) Deviator stress-strain curve, (b) Volumetric strain curve, (c) lateral strain curves, and (d) dilatancy angle.

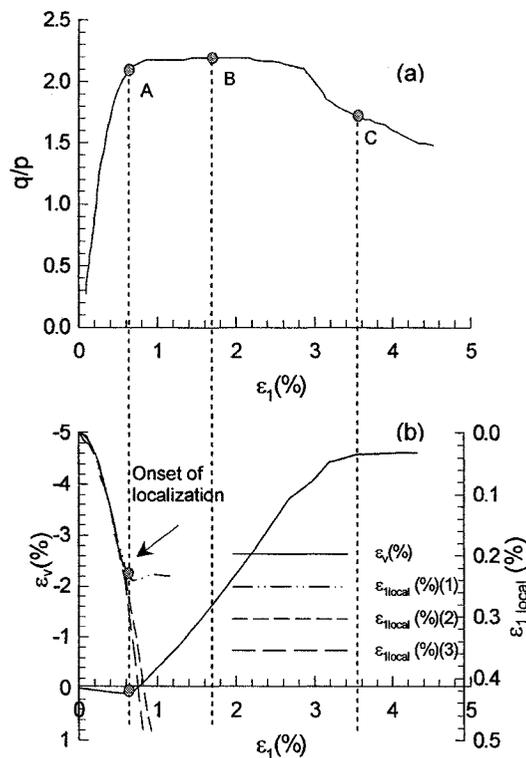


Figure 2. Stress ratio-strain and volumetric strain curves for interlocked sand (Cresswell and Powrie (2004)); (a) Stress ratio-strain curve, and (b) Volumetric and local strain curves.

This is also supported by Figure 1d. At peak strength (point B) the strain localization results in a partially formed shear band and subsequent shearing results in the formation of a continuous band at large strains (point C). It would appear that the shear band is fully developed at Point C where the volumetric strain remains constant, i.e., the critical state.

Figure 1d shows the evolution of the dilation angle with axial strain. It can be seen that at point A the dilatancy angle is equal to zero where localization commences. At point B the dilation angle reaches its maximum value followed by reduction in the post-peak regime. Finno and Rechenmacher (2003) clearly stated that the peak strength is a direct result of the localization and loss of homogeneity in sheared specimens of dense sand.

From these results it appears that for dense sands in plane strain testing the onset localization can be determined from the global volumetric strains and that the volumetric strain reversal is representative of the onset of shear-induced dilation, which also corresponds to the onset of shear localization, i.e. shear banding. In addition,

strength reduction in post-peak regime is a direct consequence of localization process.

2.1.2 Locked sand

Locked-sand was first investigated by Dusseault and Morgenstern (1978) who concluded that the fabric of the sand was responsible for its high strength and dilation angles at low confinement, and not bonding due to cementation. Cresswell and Powrie (2004) carried out triaxial compression tests on locked sand to investigate its stress-strain characteristics. The sand used in their tests had a tangential index (the ratio of tangential contacts to long straight contacts) of 26% and a relative density index of 136% and a void ratio of 0.47. Axial deformation was monitored using 4 axial LVDTs and the radial deformation was monitored with a radial strain caliper. Figure 2a shows the stress-ratio (q/p) versus axial-strain curve at a confining pressure of 100 kPa and Figure 2b depicts the volumetric strain as well as local axial strains measured using internal LVDTs. According to Cresswell and Powrie (2004) divergence of the local axial LVDTs in Figure 2b; point B; indicates the onset of localization where the strain field becomes discontinuous. This point also coincides with the point of shear-induced dilation (volumetric strain reversal) as shown in Figure 2b. Progression of deformation from point A (onset of localization) to B (partial formation of shear band) to C results in the complete formation of shear band as indicated formerly for dense sands. Cresswell and Powrie (2004) indicated that onset of shear induced dilation represents the stress level at which sand starts to unlock and loses its integrity. They also recommended this threshold to be the limiting strength envelope for locked sands.

2.1.3 Vosges Sandstone.

Besuelle et al. (2000) carried out triaxial compression tests on Vosges sandstone. This sandstone is a poorly cemented sedimentary soft rock and most of its characteristics are derived from particle contact details. It is a pink quartz sandstone (quartz = 93%) with few percent of feldspar and white mica, and a porosity of 22%. Axial and lateral strains were measured using three axial and four lateral transducers based on the concept of flexible strip with attached strain gauges. The localization was detected through the comparison of three axial strain measurements. If the difference between the mean axial strain (average of three measurements) and each of the three axial measurements remains constant, the strain field can be assumed to be homogeneous with no localization.

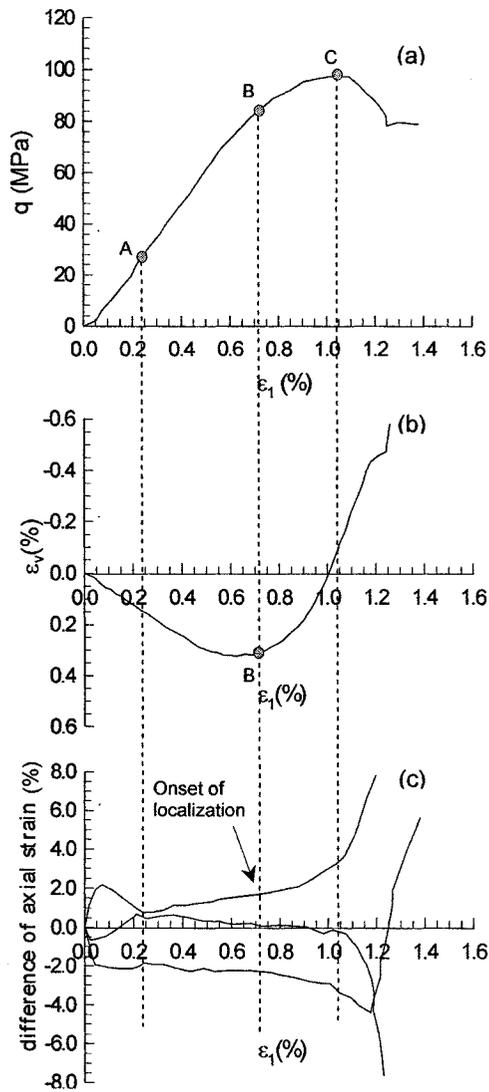


Figure 3. Stress strain volumetric strains curves for sandstone (Besuelle et al. (2000)); (a) Deviator stress-strain curve, (b) Volumetric strain curve, and (c) axial strain difference curves.

Figure 3a shows the stress-strain curve at a confining pressure of 20MPa. It can be observed that there is a strength reduction in the post-peak regime due to localization. Figure 3c shows the relationship between the differences of axial-strain and mean axial-strain. Besuelle et al. (2000) suggested that the first divergence of the three measurements was due to bedding errors in the experiment represented at point A. Once seating had occurred the difference in the three axial strains remained nearly constant up to point C where the difference of the three transducers increased dramatically. The original authors

selected this point (point C) as the onset of localization based on axial strains differences. Using the axial strains differences as well as lateral strain differences the onset of localization can be shown to commence at point B. Volumetric strain was also determined using axial and lateral strain measurements as shown in Figure 3b. Note that the onset of shear-induced dilation (volumetric strain reversal) was again coincident with the onset of localization (point B; Figure 3b). Besuelle et al. (2000) also showed that the onset of localization generally occurs at a progressively lower normalized deviator stress as the confining pressure is increased.

2.2. Fine grained soils and rocks

2.2.1 Overconsolidated Silty clay

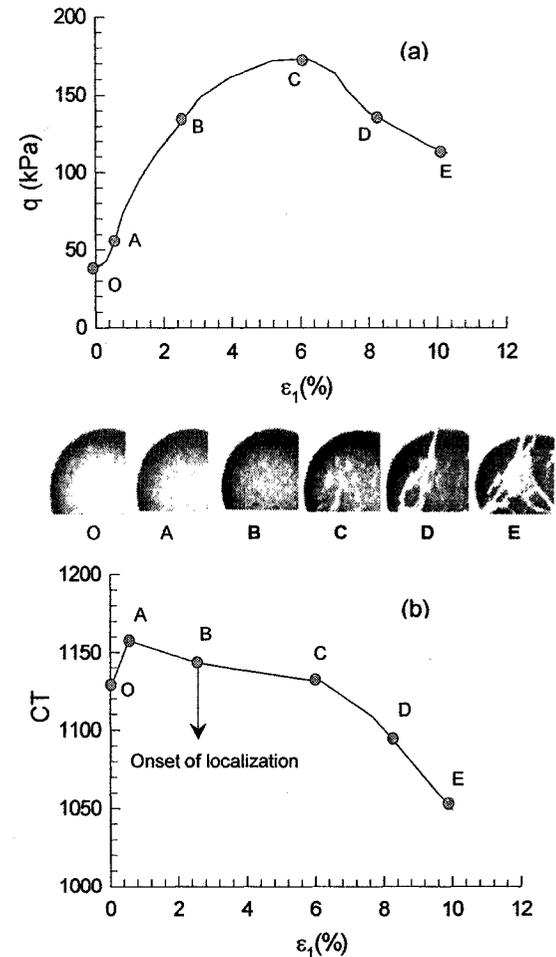


Figure 4. Stress-strain and variation of CT number with strain (Su et al. (2004)); (a) Deviator-strain curve, and (b) variation of CT number with strain.

Su et al. (2004) studied the deformation mechanism of undisturbed silty clay samples using

triaxial compression test under undrained conditions by means of computerized tomography. Scanning was recorded at different strain levels and for different slices during shearing. The CT images are quantified using CT number, which is a representative number for the density of the sample and has a linear relationship with density. The change of density can be observed from the changes of the brightness between two cross-sectional images at the same location. The white areas in the images are considered to represent damage and localization. Figure 4a shows deviator-stress versus axial-strain relationship at a confining pressure of 300 kPa. The sample has a void ratio of 1.043, plasticity index of 14.31, and water content of 37.72 %. The letters on the curve represent strain levels at which CT images are taken. Figure 4b gives the relation between the CT number and the axial strain obtained from the middle slice as it represents the most homogeneous location in specimen (away from boundary effects). Based on variation in CT numbers, the stress strain curve are divided into several portions; (i) micro-defects close up (O-A); (ii) new microcracks open and develop where damage starts (bond-degradation stress level; A-B); (iii) local deformation and damage starts (point B); (iv) softening and reduction in load carrying capacity. Based on Su et al. (2004) experiments localization starts to occur at point B where local defects and damage commences. X-ray portion shown at this axial strain level reveals the existence of separate locations of damage sites (indicated by the light grey points) that coalesce after this stress level to form the major shear band. At peak (point C) the shear band is partially developed and the clay lost its homogeneity. Beyond point C there is a significant reduction in CT number as a consequence of reduction in the density due to localization. Hence the peak strength (point C) is a direct result of localization and one can argue that point B represents the onset of shear-induced dilation and localization in these tests.

2.2.2. Bonded clay

Vallericca Clay is a bonded overconsolidated clay and can be considered as a weak sedimentary rock. It has a plasticity index of 31.6%, water content of 28.6% and a carbonate content of 30%. In order to investigate the formation of shear bands in this weak rock, Viggiani et al. (1994) carried out a plane-strain compression test using the same technique as Finno and Rechenmacher (2003). They also showed that the onset of localization took place pre-peak and was coincident with the onset of shear-induced dilation for this kind of geomaterials.

2.3 Hard crystalline rock

Lockner et al. (1992) studied the deformation mechanism of western granite by means of acoustic emissions. They conducted triaxial compression tests on granite samples with a

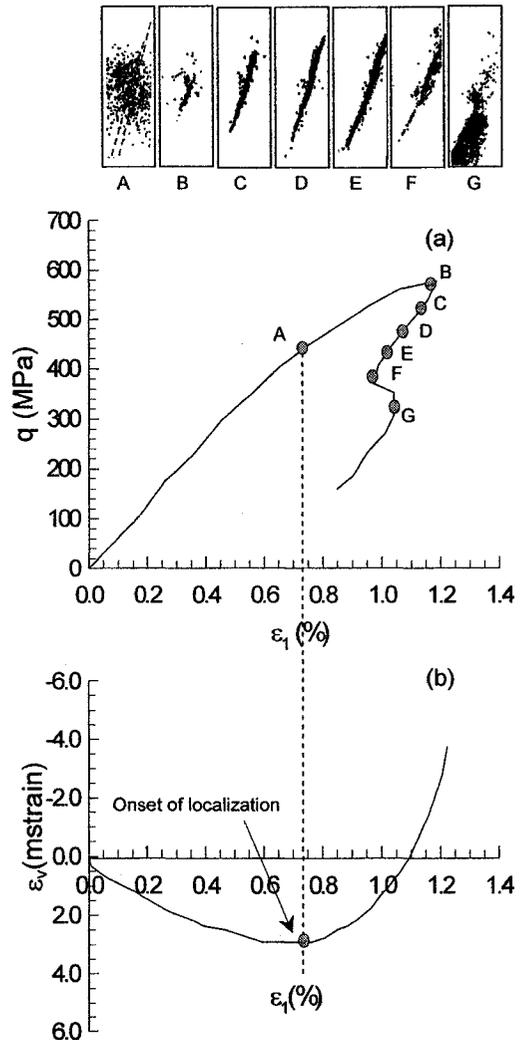


Figure 5. Stress-strain and volumetric strain for granite (Lockner et al. (1992)); (a) Deviator stress -strain curve, and (b) Volumetric strain curve.

porosity of 0.01 at a confining pressure of 50 MPa. The cell was provided with six piezoelectric transducers to monitor deformation mechanism using acoustic source locating techniques. Figure 5a shows the deviator-stress versus axial-strain curve for a granite sample as well as acoustic emission events recorded during deformation. The letters (A to G) on the deviator stress versus axial strain curve represent the strain levels at which

acoustic emissions are recorded. Figure 5b gives the volumetric strains-axial strain relationship.

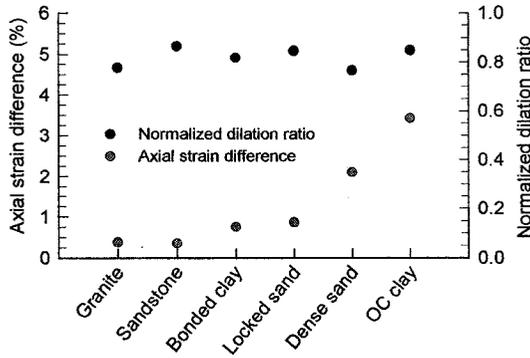


Figure 6. Normalized dilation ratio and axial stress difference for various geomaterials.

According to their acoustic emission observations, localization starts at point A where several distributed damage sites start to coalesce. This point is also coincident with the volumetric strain reversal based on their volume change measurements. At peak strength (point B) the shear band is partially formed and progression of deformation beyond this point results in the complete formation of shear band as shown through the transition from point B to G accompanied with a reduction in strength as a direct result of localization.

2.4 Summary

Based on the laboratory tests reviewed, it is clear that the failure process in a wide spectrum of geomaterials is dominated by the formation of shear bands. The onset of shear-induced dilation, i.e., volumetric strain reversal, appears to be an excellent indicator of the initiation of shear band formation (onset of localization). The peak strength, however, appears to be related to the development of the shear band throughout the sample.

The stress magnitudes associated with the onset of shear-induced dilation in the laboratory tests presented have been normalized to the peak strength. Figure 6 shows that for all geomaterials examined this ratio ranges from 0.76 to 0.86, regardless of whether it is dense sand or hard rock. Hence it would appear that the once a shear band initiates, failure is imminent regardless of the bonding or fabric. The normalized dilation ratio can be expressed as:

$$[2] \text{NDR} = \frac{q_{dil.}}{q_p}$$

While the ratio of the onset of shear-induced dilation to the peak strength may be independent

of bonding and fabric, the amount of plastic strain required to reach the peak strength once a shear band has initiated is not independent of material type. The axial strain difference (ϵ_{p-d}); between the peak strength (ϵ_p) and onset of shear-induced dilation ($\epsilon_{dil.}$) can be expressed as:

$$[3] \epsilon_{p-d} = \epsilon_p - \epsilon_{dil.}$$

Figure 6 also reveals an increase in axial stress difference between the peak and dilation stress levels where constraints between particles become less. For hard rocks, sandstone, locked sand and bonded clays the degree of constraints at particle contacts are high due to the existence of bonds and interlocking. As can be seen from Figure 6 the ability of this kind of geomaterials to sustain any deformation after onset of localization up to failure is insignificant (0.5%-1%). One can speculate that overcoming these constraints (bonds and interlocking) does not allow particles to adjust properly to resist shearing. On the other hand for dense sand and OC clay the degree of constraints is less. Therefore, these geomaterials can sustain higher degree of deformation before failure as compared as geomaterials with high degree of constraints and hence less degree of freedom.

3. EXTENSION-INDUCED DILATION

In the testing of hard rocks, surface mounted strain gauges are normally used to monitor the stress-strain response. Lajtai et al. (1990) showed, based on his observation during the middle stages of compression test, that only lateral dilation of the cylindrical sample was recorded with no axial shortening, and suggested that damage initiation in Lac du Bonnet granite was caused by extension-induced cracking. Martin and Chandler (1994) used damage-controlled triaxial tests on Lac du Bonnet granite to identify three characteristic stress thresholds as shown in Figure 7. Crack initiation stress caused by extension-induced dilation (30-40% of peak resistance). In zone (III), propagation of microcracks in the vertical direction form columnar structures separated by discrete microfractures. At 70-80% of the peak resistance, the onset of shear-induced dilation, these columns start to buckle and microcracks coalesce to form major macrofractures, which results in reversal of volumetric strains. In zone (IV) strain becomes localized in a small band and global failure occurs at peak where most of columns become distorted. Martin & Chandler proposed that the strain hardening observed from the onset of shear-induced dilation to the peak strength resulted from the testing methodology and was not a material property.

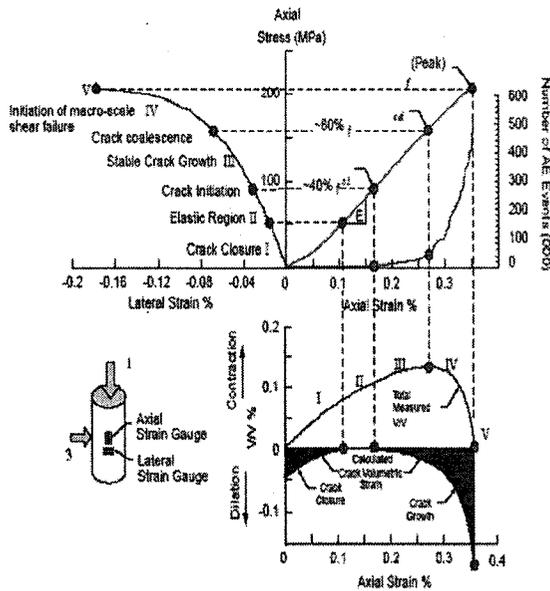


Figure 7. Stages of crack development for granite using damage tests (After Martin and Chandler (1994)).

Martin (1997) has shown that the operational strength around the Mine by test tunnel was coincident with the deviator-stress required to cause extension-induced dilation (crack initiation stress) in laboratory tests.

Vallejo (1987) carried out direct shear tests on brittle kaolinite clay with preexisting cracks. He observed that failure started in samples in form of secondary cracks developed from the tips of original cracks. These secondary cracks are developed in the region of tensile stresses in samples due to shear stresses. Secondary cracks tend to extend in a direction normal to maximum tensile stresses regardless of the direction of original cracks. Vallejo (1989) also studied the effect of fissures on the strength of brittle kaolinite. He showed that under compressive stresses there are local tensile stresses developed at crack tips. The fracture of specimens is dominated by the extension of cracks in the direction of major principle stress. The original fissures are acting as stress concentrators from which cracks propagate in the tensile stress regimes so that the final strength is dictated by bridges between artificial fissures.

In summary in bonded geomaterials (regardless degree of bonding; clays or hard rocks) there is another characteristic strength threshold after which bonds starts to degrade (extension-induced dilation). Degradation of these bonds under compression is due to heterogeneity of geomaterials where local tensile stresses develop. The cracks open laterally and extend in the direction of the major principle stress in triaxial

compression stress path in a stable fashion with formation of vertical columns separated by discrete vertical microcracks. At a certain stress level (onset of shear-induced dilation, i.e., volumetric reversal stress level;) these separated columns start to buckle (coalescence of extensile cracks) to form shear band and strain becomes localized in a band composed of these microcracks (Figure 8). It should be borne in mind that direct capturing of extension-induced dilation could be difficult to disclose from global volume change measurements. This requires local (axial and lateral) strain measurements so that the nonlinearity of lateral strains can be accurately captured.

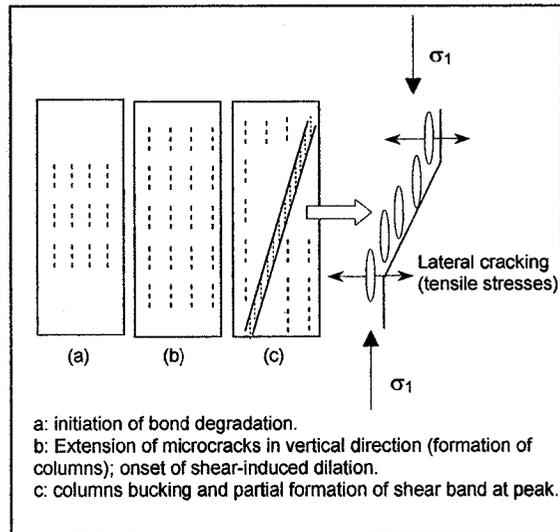


Figure 8. Analogue for deformation mechanism in bonded geomaterials.

4. CONCLUSIONS

A review of laboratory test results of dense and locked sands, sandstone, stiff clay, weak rocks, and hard rock was carried out. Based on this review the following conclusions can be drawn:

- 1- Onset of shear-induced dilation (volumetric strain reversal) marks the onset of shear-band localization for a wide spectrum of geomaterials. It can be directly estimated from global measurements.
- 2- The ratio of the stress level required for initiating shear-induced dilation normalized to the peak strength ranges from 0.76 to 0.86 and is independent of fabric and/or bonding.
- 3- The amount of plastic strain required to reach the peak strength from the onset of shear band initiation depends on fabric and/or bonding, i.e., type of geomaterial.
- 4- Onset of extension-induced dilation marks the initiation of bond degradation for

bonded brittle geomaterials and can only be measured using local lateral strains.

- 5- More attention should be paid to these pre-peak dilation thresholds for bonded brittle soils where local measurements can be carried out.

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GEOMECHANICAL BEHAVIOR OF A MCMURRAY SAND FORMATION USING LOCAL STRAIN MEASUREMENTS

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ABSTRACT

Block samples of locked sands were procured from High-Hill river bank East of Fort McMurray. Micro-structural observations revealed the existence of interlocked fabric due to overgrowth and dissolution. X-ray spectrum analysis indicated that Quartz is the major mineralogical constituent of the formation with no virtual bonding. In absence of bonding, existence of such interlocked fabric provided sufficient evidence to interpret its high uniaxial strength and axial stiffness. Local strain results showed that resorting to commonplace external strain measurements can be misleading where strength and deformation coupling is targeted in locked sands.

RÉSUMÉ

Des échantillons de bloc de sables verrouillés ont été obtenus de la banque de fleuve de Haut-Colline à l'est du fort McMurray. Les observations microstructurales ont indiqué l'existence du tissu enclenché dû à la surcroissance et à la dissolution. Radiographie et l'analyse du spectre a indiqué que le quartz est le constituant minéralogique principal de la formation sans la liaison virtuelle. En l'absence de la liaison, l'existence d'un tel tissu enclenché a fourni l'évidence suffisante pour interpréter sa haute résistance uniaxiale et rigidité axiale. Les résultats locaux de contrainte ont prouvé que le recours aux mesures externes banales de contrainte peut être fallacieux où l'accouplement de force et de déformation est visé en sables verrouillés.

1. INTRODUCTION

Nature provides a wide spectrum of geomaterials ranging from soils to hard rocks. At the intermediate zone there is a class of geomaterials that inherited both soils and rocks characteristics as a result of mild diagenesis. This zone can be classified into bond-dominant geomaterials (bonding is responsible for its coherent mass) and fabric-dominant geomaterials (interlocking is the binding agent between particles). Despite the wealth of information on bond-dominant geomaterials (e.g., comprehensive data compilation and analysis presented by Vaughn and Leroueil 1990), exploration of fabric-dominant geomaterials' behavior is comparatively neglected in geotechnical studies. An example of fabric-dominant geomaterials is *locked sands* where mild diagenesis is reflected in its unusual particles contact characteristics resulting in interlocked coherent mass as introduced by Dusseault and Morgenstern (1979).

In Northeastern Alberta, and particularly within 115 Km of Fort McMurray, Dusseault (1977) recorded the existence of oil-sand and oil-free sand natural steep slopes. Dusseault (1977) found no interrelationship between slopes height and its inclination with the oil content. Dusseault and

Morgenstern (1978) showed experimentally that the strength of oil sands was merely a reflection of its interlocked fabric and hence independent of oil content. Protulipac (1991) recommended oil-free locked sands as a reasonable substitute for oil sands to avoid sample disturbance due to gas expulsion where strength and deformation characteristics are concerned. As a matter of fact, entire experimental investigations on locked sands from the McMurray area have been carried out using global strain measurements to establish its strength and deformation characteristics. It is not generally recognized that local measurements on laboratory samples can reveal additional information that cannot be obtained using commonplace global measurements. The purpose of this work is to establish the characteristics of locked sand from McMurray area at small strains using local strain measurements and to compare the results to those obtained using global measurements.

2. GENERAL LOCKED SANDS CHARACTERISTICS

Locked sands are unique formations that constitute the sands/sandstones border and retaining properties neither entirely analogous to soil properties and not being as strong as rocks. This is schematically illustrated in Figure 1 in terms of diagenetic effects and uniaxial strength. As depicted, locked and cohesive sands are on the borderline between soils and hard rocks. The overlap between the boundaries is deliberate as it is intricate to demarcate such boundaries. With regard to arenaceous geomaterials, sand and sandstone behaviors have been explored extensively in soil mechanics and rock mechanics in terms of deformability and strength. In dense sands the strength is in primarily a function of density, which yields a weak interlocking between particles with relative densities less than 100%. In sandstone strength is primarily a function of the amount of interstitial cement between sand particles. Despite the lack of cementation in locked sand pores, it possesses strength properties much higher than dense sands and much less than sandstone. In addition, locked sands exhibit relative densities greater than 100% in contrast to dense sands (Dusseault and Morgenstern 1979). The deviation from dense sand strength properties stems from mild diagenetic processes that reduce sand mass porosity with negligible or no cementation (Barton, 1993). A comprehensive summary of the main characteristics of locked sands is provided in Table 1 and summarized in the following points:

(a) *Quartzose sands*: these formations are older than Quarternary in age, poorly cemented, and can be reduced to its grains component through rubbing (friable). They belong to the orthoquartzite family, which is characterized by the existence of at least 95% quartz rich detrital grains cemented with authigenic quartz (Pottjohn et al. 1972). They show a tendency to disintegrate when immersed in water (Barton et al. 1986). However this is dependent on amount of fines in the matrix. Frequent cross beddings with traces of fine grains (such as Kaolinite, illite, and iron oxide) are reported based on various investigations (Dusseault 1977, Richards and Barton 1999, Barton et al. 1986).

(b) *Microfabric characteristics*: grain-to-grain contacts are large with interpenetrative fabric and rugose texture. This is attributed to pressure solution at particle contacts and/or post-depositional quartz overgrowth. Interlocking is measured through the tangential index (TI, the percentile of tangential contact to the total contact). It has been shown that there is no relation between the TI and the porosity of such sands (Richards and Barton 1999).

(c) *Packing*: porosities are lower than dense sands and cannot be reproduced in laboratory and relative densities can be in excess of 100%. As a consequence it exhibits low compressibility and can be used to support foundations with small deformation.

(d) *Shear strength characteristics*: It is characterized by brittle behavior, small strain at failure, high friction angles of 57°-70°, extreme dilation, a higher residual friction angle of 30°-35° compared to usual Quartzose sands due to rough surface textures, and a reduction in strength and dilation rates due to wetting. High TI (less interlocking) results in a negligible cohesion intercept as indicted by Richards and Barton (1999). Fabric cohesion (cohesion intercept due to interlocking) is only observed in fine to medium grained sands due to interlocking. In contrast it is negligible in coarse-grained locked sands. Furthermore, there is no obvious relationship between cement content, if exists, and fabric cohesion for such materials.

3. CHARACTERISTICS OF HIGH-HILL RIVER SAND FORMATION

Along High Hill River and Cottonwood Creek, East of Fort McMurray, several exposed oil-free sand cliffs with heights up to 70 m exist. In the current investigation, block samples were procured from outcrops along High Hill River banks (Figure 2). The oil-free sand block sample was completely dry in situ presumably due to the existence of a shale cap at top of the sand formation. The block had dimensions of 22 cm by 22 cm by 15 cm and cut from exposed faces using electrical wire saw. These blocks were then contained in wooden boxes encased with expanding foam for transportation and storage.

3.1 Geology

Oil-free sand blocks retrieved from High-Hill River banks belong to the Lower Cretaceous McMurray Formation that constitutes the majority of Athabasca oil sands. This formation is overlain by the lower Cretaceous Clearwater Formation and underlain by the upper Devonian Waterways Formation (dolomite and limestone) According to the tripartite classification introduced by Carrigy (1959); as upper, middle, and lower members, the formation under investigation possesses akin characteristics to the middle member of McMurray Formation characterized by existence of fine to Medium grained quartzose sand (Carrigy 1967) that is formed in a marine influenced estuarine channel complex (e.g., Hein et al. 2000; Wightman and Pemberton 1997). A summary of McMurray

Table 1 Summary on physical and mechanical characteristics of locked sands.

Authors	Formation/ geological age	Porosity (%)	RD (%)	Density (kN/m ³) and/or D ₅₀ (mm)	Peak secant friction angle, normal stress range and strain at failure (direct test- condition unless stated)	Strength parameters, confining pressure and strain at failure (Triaxial tests)	Uniaxial compressive strength and Young's modulus	Contact Nature/ or TI (%)
Dusseault and Morgenstem (1979)	Athabasca oil sand (Lower Cretaceous)	31.3	>100-135		44°-46° (100-800kPa) 1.5-2.5%	φ = 50°	-----	Mostly long to straight contacts, concavo/convex contacts and tangential contacts
Dusseault and Morgenstem (1978)				18.5-23.3 D ₅₀ =0.2mm	φ _{sec.} = 39.9°-66.4°* (98-785kPa)			
Barton et al. (1986)	Barton sand (Eocene)	35.6	119	17	40.6°-47° (50-900kPa) 1.9-5.5%	C= 0 kPa φ = 40° (100-700kPa) 2.9-7.9%	q _u 5-134 kPa	Tangential contacts-slight concavo/convex contacts (TI=54%)
Barton et al. (1986)	Bagshot sand (Eocene)	36.4	110	16.8	44.2°-50.3° (50-900kPa) 2.5-7%	C= 0-77.4 kPa φ= 41.5°-50° (100-700kPa) 0.56-2.8%	q _u 20 kPa	Tangential and straight contacts-slight concavo/convex contacts (TI=41%)
Dusseault and Morgenstem (1979)	Swan River sandstone (Lower Cretaceous)	34.5	125		63°-70° (100-800kPa) 1.0-2.5%	-----	-----	Mostly concavo/convex contacts
Dusseault and Morgenstem (1979)	St. Peter friable sandstone (Ordovician)	27	135		63°-69° (100-800kPa) 0.5-1.75%			Mostly concavo/convex contacts
Dittes and Labuz (2002)		27		19 kN/m ³	φ = 57°-63° ψ=45°(19kPa)-21° (152 kPa) φ _{sat.} =46°	φ = 57°-69° φ _{field} = 56°	q _u =0.6-2 Mpa E=1GPa ν=0.2-0.33	
Richards and Barton (1999)	Folkstone bed sands Uniform (including Reigate silver sand)	28-35		D ₅₀ = 0.2-0.3mm	C= 7.7-179.4 kPa φ _{sec.} =34°-65.5°		q _u =101.9-725.2 kPa	TI=29.7-65.4%
	Thin bedded(iron oxide)	25.7-33.6			C= 20.5-65.4 kPa φ _{sec.} =34°-65.5° (0.15-0.8mm for zero normal stress) (0.6-2mm for normal stresses up to 674kPa)		q _u =85.5-218 kPa	TI= 27.3-66%
Cresswell and Powrie (2004)	Reigate silver sand	32	136%	18.62		(50-800kPa) 0.4%-0.8% M=2.25 (p=0-1000 kPa)	q _u =108-101.9 kPa	TI=26%
Creswell and Barton (2003)				D ₅₀ = 0.25mm	φ = 70° (at low stress and 40° for 10 MPa) (0.25mm at 1.5 kPa-2.5mm at 10MPa)			

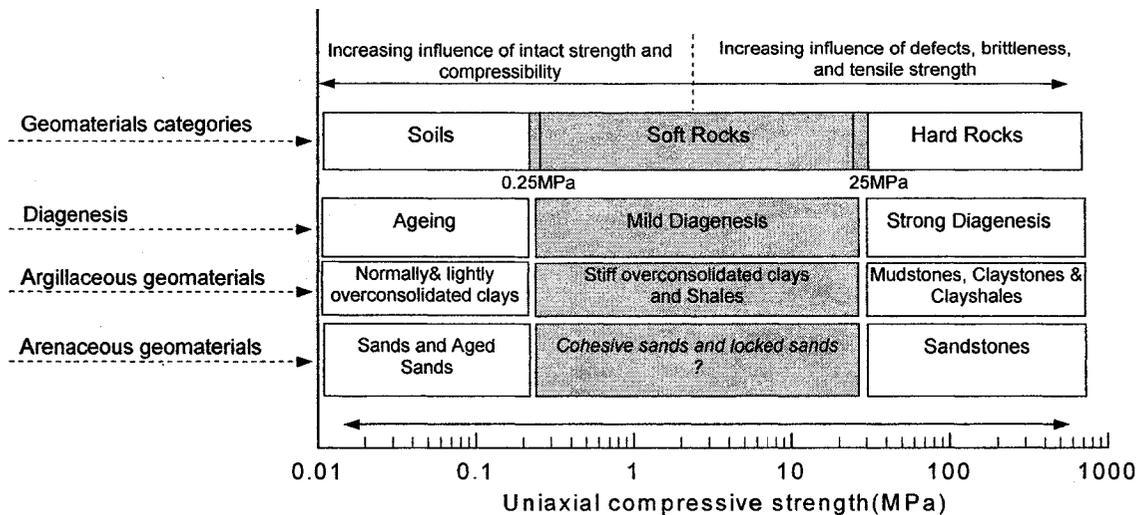


Figure 1 Classification of geomaterials based on degree of diagenesis and UCS.

Formation members' characteristics and depositional environment is shown in Table 2.

3.2 Index Properties

Sampled blocks appear as white to light grey in color, characterized by thin cross-black beddings, and rough texture. Under optical microscope, the majority of detrital grains are clear, white or translucent. Rubbing the sample surface with minimal finger tip pressure results in a degradation of the interlocked structure and hence reflecting extremely weak cementation if it exists. The grain size distribution is shown in Figure 3 as imposed on classification of McMurray's oil-sand formation introduced by Carrigy (1967). According to these groups' grain size distributions this formation is considered as a middle McMurray Formation member. According to the USCS it is classified as poorly graded sand (SP); uniform sand. The matrix contact (the proportion of fine grains) is estimated to be less than 10%. Hence, the intact structure cannot be attributed to either gradation effects as in dense sands or cementation by fine portions. It exhibits a relative density of 125% ($e_{max} = 0.86$, $e_{min} = 0.61$, $e_{in-situ} = 0.54$) indicating existence of exceptional grain packing that cannot be reproduced in laboratory and confirms its fittings into locked sands. Based on a locked sands classification scheme introduced by Barnes and Dusseault (1982), describing the degree of diagenesis and its reflection on contact nature based on of the

difference between the minimum and the actual porosities, this sand belongs to the second category (a porosity reduction of 2.7 %) characterized with small fabric cohesion, mild diagenesis reflected in tangential, long and occasional interpenetrative contacts. In summary, this deposit is mainly composed of well sorted sand particles with small portion of fines and high relative density due to mild diagenesis.

3.3 Micro-structural and Mineralogical Studies

Figure 4 shows a selected SEM image for particles layout. Particles are characterized by high angularity and irregular shapes, existence of fine materials in a coating and a pore filling forms. Particle contacts are best described as tangential, long and short straight with some concavo-convex (interpenetrative) and sutured contacts. The severe chemical dissolution experienced during diagenesis is shown through existence of pitting cavities.

Collected grains under optical microscope with different colors were numbered and labeled as clear, white, red, purple, and black for mineralogical composition identification using Energy dispersive X-ray spectrum analysis. 95% of clear and white particles are basically silicates with insignificant aluminates portion (quartz mineral) which is one of the featured



Figure 2 Outcrop of High Hill River locked sand. Note the steepness of the natural slope, approximately 60 degrees.

characteristics of orthoquartzite family (Pettijohn et al. 1972), see Figure 5a. Only 5% of white

Table 2 Tripartite classification, Depositional environment, and Lithology of McMurray sand Formation members (Carrigy 1967; Hein et al. 2000).

McMurray Formation members	Depositional Environment	Lithology of McMurray sand formation
Upper	Shore-face (coastal plain sediments)	-Fine-grained quartzose sands to clayey silts. - Doesn't belong to the orthoquartzite family due to existence of significant percentage of clay minerals and non-quartz clastics
Middle	Estuarine channel complex	-well sorted fine to medium grained quartzose sands.
Lower	Braided to meandering fluvial sand channels	- medium grained sand and pebble conglomerates

particles are kaolinite, as evidenced by the existence of nearly equal ratio between silicates and aluminates and existence of potassium, which is a typical clay mineral in sandstones of lower Cretaceous age in Alberta (Carrigy and Mellon 1964); Figure 5b. The red particle is a feldspar particle, which is presumably the source of the kaolinite, and reflecting the source of the parent rock as either igneous or metamorphic rock. The

purple one is identified as belonging to the amphibole family. One of the black ones is a black chert identified by the existence of multiple quartz crystals while the other one is a titanium particle typically found in sedimentary rocks (Pettijohn et al. 1972).

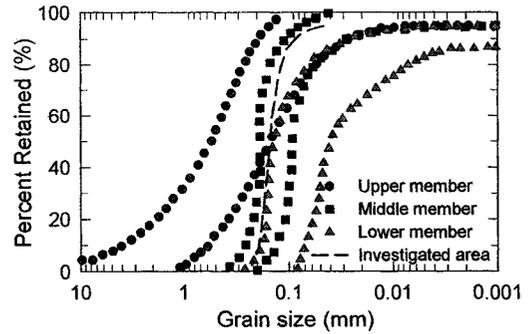


Figure 3 Grain size distributions of McMurray Formation and the investigated Formation.

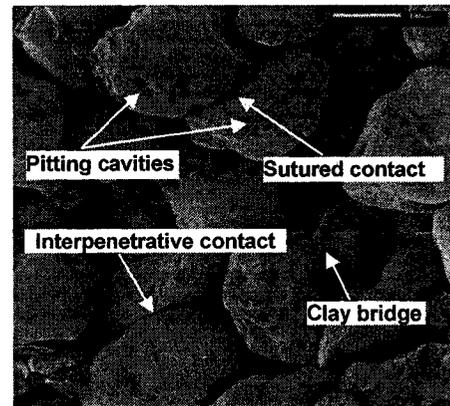


Figure 4 Selected SEM image illustrating contacts nature, pitting cavities, and clay bridges between detrital grains.

The results of X-ray analysis suggested that the mineralogy is dominated by the existence of quartz and that the black bands are basically fine rock fragments and chert particles, reflecting the depositional layering common to sedimentary environments. In order to confirm these observations thin sections were also prepared to provide additional insights into the depositional history of this deposit. Thin sections were prepared with dimensions of 75x55mm after impregnations with a blue-dyed epoxy resin under high vacuum and examined using a petrographic microscope equipped with a high resolution Nikon camera to record observation. In well sorted bands, grain particles range from angular to very angular.

Approximately 90% of the particles were quartz particles with 5% polycrystalline quartz originating from a metamorphic rock, 4% very fine grained rock fragments (mostly chert) and less than 1% of clay minerals, bridging quartz particles. The observed contact nature is generally short and long contacts with few tangential and interpenetrative contacts. In the less-sorted bands most of particles were angular in shape. Variable fine rock fragments are present in the band under examination, implying the variability of parent rocks. Less-sorted band comprises approximately 60% monocristalline quartz, 5% polycrystalline quartz, 10% siliceous siltstone, 20% chert, 5% feldspar and clay minerals. Clean fractured quartz grain are also evident in most of the sections.

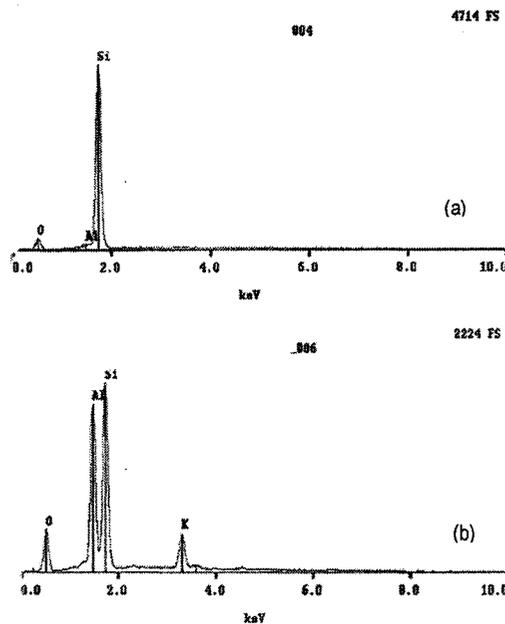


Figure 5 Example of X-ray results: (a) quartz white and clear particles; (b) white Kaolinite particle.

It appears that deposit was subjected to overburden loads that were adequate to cause this fracturing and hence influence the structure of the deposit. Existence of clean cracks within quartz grains implies that clay minerals were deposited following the compaction stage or its weak nature so that to prevent loss of porosity due to compaction by early cementation. In poorly sorted layers, there is evidence that compaction embedded quartz between the fine grained rock fragments reflecting the influence of burial depth after deposition. Quartz overgrowth was also observed in the thin sections (Figure 6a). This overgrowth was identified via distinguished dust lines as one of the diagnostic aids for identifying overgrowth (Pittman 1972). A fresh skeletal grain derived by dissolution is also apparent in Figure

6b, suggesting the process of dissolution occurred after the compaction stage; otherwise the compaction process would have destroyed this skeleton and that the dissolution is still in process

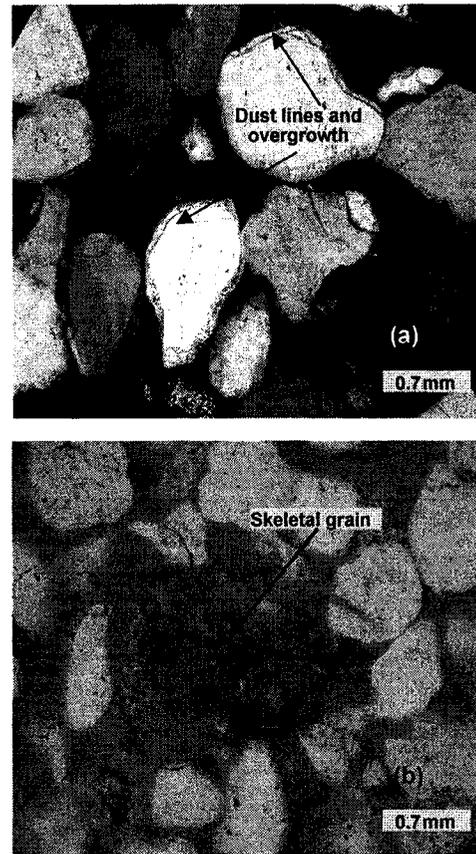


Figure 6 (a): Micrograph under CPL showing quartz overgrowth, quartz grains as white and light to dark grey and pores as black;(b) Micrograph under PPL demonstrating existence of a skeletal grain due to dissolution during diagenesis.

3.4 Uniaxial Strength Characteristics

Cylindrical specimens were prepared with dimensions of 63mm diameter x 63mm high (1:1 sample) with extreme care due to its friable nature and brittleness. Local strains were monitored through a commercial system provided by SBEL (Structural Behavior Engineering Laboratories Ltd.) and modified to incorporate an additional vertical LVDT as well as two horizontal ones. The local strain measuring system comprised three vertical miniature LVDTs for axial deformation as well as a miniature spring-loaded LVDT attached to a light aluminum chain at the mid-height of specimen for lateral deformation along with two horizontally mounted LVDTs. The system was

also augmented with an external LVDT for the sake of comparing local and external axial strain measurements. End effects were alleviated via applying a thick layer of silicone grease between two latex membranes placed at platens/specimen interface. Upon completion of setup, the specimen was sheared under a constant strain-controlled rate of 0.05%/hr using an ELE loading frame with a capacity of 17 MPa (for a 63mm diameter specimen). Measurement of axial loads was achieved using a Sensotech load cell with a low capacity of 7 MPa (for a 63mm diameter specimen). In the following sections, a portion of the experimental results is presented and discussed to assess the value of local strain measurements. Details of experimental setup, LVDTs calibration, and proof testing are not presented herein.

3.4.1 Uniaxial strength and axial stiffness

Figure 7 shows the deviator stress-axial strain relationship based on local and external axial strain measurements. Local axial strain is derived from the average of the three vertical LVDTs readings. In general the specimen attained peak strength of 2.0 MPa at axial strains of 0.32% and 0.57% based on local and external measurements respectively. According to ISRM (1978b), it is classified as a very low strength rock (weak rock) situated closer to the soil boundaries (characterized by a uniaxial compressive strength < 0.25 MPa) and constitutes the boundary between arenaceous soils and rocks. The source of high uniaxial strength presumably stems from the relatively extensive area of contacts as well as the increase in coordination number; number of contacts per particle. That is, particles interact as surfaces rather than point contacts (as in dense sands) and hence the rotational resistance is enhanced. As a consequence, in absence of bonding individual contacts can sustain considerable contact forces, moments, and stress obliquity due to interlocking (Creswell and Powrie 2004).

Both strain systems showed an extremely brittle post-peak response. External measurements had a tendency to overestimate axial strain at any stage during shear evolution and particularly in pre-peak regime. Only at very large axial strains do internal and external measurements converge together. In the pre-peak regime, both stress-strain curves were initially concave followed by a linear region. The size of the initial concave portion of the stress-strain is significantly overestimated by the external strain measurement as compared to local one. Within this zone, the nonlinear behavior suggests the closure of existing fractures within fractured grains identified in micro-structure studies. In the linear portion of the stress-strain curves, where the elastic zone is

traditionally identified, the axial stiffness based on local strain measurement is approximately 0.87 GPa and roughly 2 times the externally measured axial stiffness. Though both measuring systems indicate strain weakening in the post-peak regime, the local one exhibits a slightly less abrupt reduction in strength.

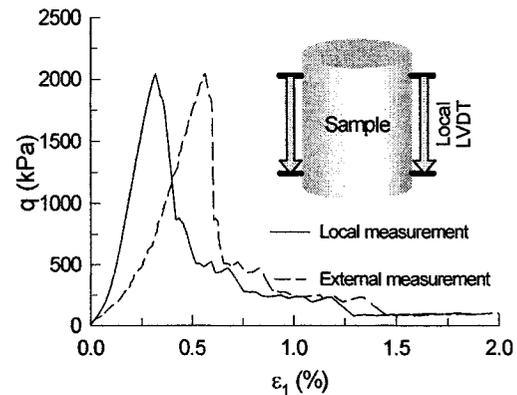


Figure 7 Local and external uniaxial stress curves.

3.4.2 Significance of local axial strain rate

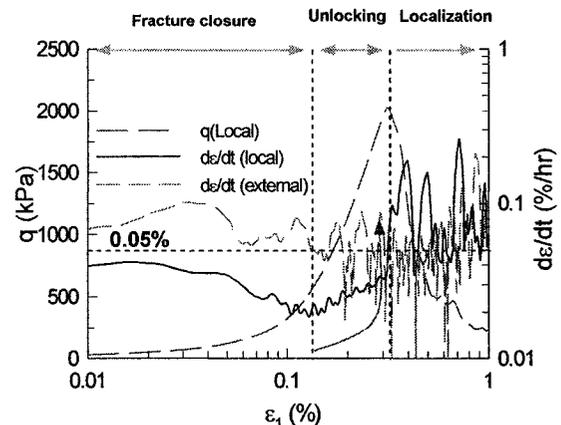


Figure 8 Comparison of external and local axial strain's rate.

Figure 8 shows the variation of instantaneous local and external strain rates in logarithmic scale with shearing evolution as well as the local deviator stress-axial strain curve. Despite assigning a controlled constant rate of strain (0.05%), the external and local strain rate deviate from that value. The external strain rate value is higher than the assigned value at the initial stages due to bedding effects and coincides with the

specified values upon closure of fractures (identified via initial nonlinearity in axial stiffness) throughout shearing process. On the other hand, the local strain rate value records smaller values, up to 40% of the assigned value, before most fractures within specimen are closed. Upon attaining this level, most of the original fractures are closed and particles become in full contact, a linear increase in axial strain rate occurs as a reflection of unlocking mechanism; destruction of interlocked fabric; followed by a sudden increase around peak strength (around 64% increase in axial strain rate). This signifies unlocking of most sample interlocked elements and initiation of localization in post-peak regime confirmed via divergence of individual axial strains from its mean value. In post-peak regime, strain rate oscillates above the specified value reflecting in average an increase in strain rate as well as non-uniformity of strain rate as an indicative of progressive rupture of interlocking structure and localization. As consequent, local strain rates are indicative of the failure process at all stages of shearing as compared to external measurements strain rates. This issue is more significant when assigning strain rates for the sake of comparing geomechanical response at different strain rates or where creep rates are of particular interest.

4. CONCLUSIONS

This paper provided an up to date synopsis of the common characteristics of locked sands to put together a framework to identify such deposits. Based on this framework, experimental work was presented to identify the fitting of High-Hill oil-free sand formation to locked sands. The experimental work comprised micro-structural observation, mineralogical composition analysis. Uniaxial compressive strength test was also carried out to reveal the strength of its unbonded coherent mass and to explore the significance of local strain measurements.

Based on the test results, High-Hill river sand formation can be described as an overconsolidated lower Cretaceous McMurray sand formation that possesses characteristics similar to oil sands middle members. The dense nature of the samples, mineralogical composition, and micro-structural observations are in keeping with the characteristics of locked sands. The intimate contact between grains forming a coherent mass stems from authigenic quartz, identified through overgrowth and dissolution, rather than cementation encountered in either sandstones or hard rocks. The existence of minor percentage of kaolinite in the form of coating or bridges is presumably responsible for the friable nature of this deposit. The fine particles presumably originated from either the diagenesis

of the overlying shale formation or feldspar inherited from parent rocks.

With regard to strength, the formation exhibits a uniaxial compressive strength of 2 MPa and hence may be classified as a very weak rock. The interlocking between particles is the main source of such high uniaxial strength and axial stiffness in the absence of virtual bonding between sand particles. The axial stiffness based on local strain measurement is approximately 0.87 GPa and almost 2 times the externally measured axial stiffness. In the post-peak regime both measuring systems indicate strain weakening, however, the local one exhibits a slightly less abrupt reduction in strength. For proper characterization of such geomaterials and its oil-sand counterpart, experimental results interpretations founded on commonplace external measurements are in error and particularly when coupling strength and deformation is targeted. In terms of failure mechanism, variation in local axial strain rate provides more insight into failure evolution stages as compared to externally measured axial strain rate.

5. ACKNOWLEDGMENT

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