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CO_2 solubility in brine in silica nanopores in relation to geological CO_2 sequestration in tight formations: Effect of salinity and pH

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ABSTRACT

Geological CO₂ sequestration during CO₂ enhanced oil recovery in tight formations is a technically and economically viable option to alleviate carbon emission. In tight formations, there existsare enormous number of nano-scale pores, which can be filled with connate and injected water. In addition, the salinity and pH of the formation water vary regionally. In this work, we used molecular dynamics simulations to study the effects of salinity and pH on CO₂ solubility in brine in silica nanopores under typical geological conditions (353 K and ~ 175 bar). The pH effect is characterized by the deprotonation degree of silanol on the silica surface. We find that water mainly distributes around the silanol groups and CO₂ mainly enriches in the areas where silanol groups are vacant. Na⁺ ions are generally depleted from the non-deprotonated silica surface, whereas they are strongly attracted to the pore surfaces in the deprotonated cases. The different water hydration structures around the nondeprotonated and deprotonated silanols arise from the accumulation of Na⁺ ions in the vicinity of =SiO⁻ groups. As salinity increases, the average densities of CO₂ and water decrease in all silica nanopores and CO₂ solubility in brine in silica nanopores decreases. On the other hand, as pH increases, water density increases but CO₂ density decreases, resulting in a decrease of CO₂ solubility in brine in silica nanopores. CO₂ solubility in brine with a low pH range (\sim 2–5) can be as high as 1.3–1.6 times of that in bulk, while it is comparable with that in bulk at a high

solubility trapping in tight formations.

1. Introduction

Carbon capture and storage (CCS) has been broadly recognized as a viable method to mitigate the carbon emissions due to the continuous consumption of fossil fuels [1]. Long-term storage of carbon dioxide (CO₂) is the last step for the CCS chain [2]. Among all the storage methods, geological CO₂ sequestration has been proven to be technically and economically viable [2–5]. According to the Global Status of CCS (2019) [6], there are 19 industrial level CCS facilities in operation globally. These projects can capture and permanently store \sim 37 million tonne (Mt) of CO₂ annually, among which \sim 30 Mt are stored in the form of CO₂ enhanced oil recovery (EOR) which can effectively offset the financial burdens associated with the geological CO₂ sequestration [1]. On the other hand, according to the Annual Energy Outlook (2019) [7] by US Energy Information Administration, tight oil (oil deposited in tight formations) is an important US energy supply, accounting for 61% of total US oil production in 2018, while its production continues to

increase through 2030. One of the most promising methods to recover tight oil is CO₂-EOR, which has been successfully implemented in many tight oil fields accompanied with CO₂ sequestration [8–10]. Therefore, CO₂-EOR accompanied with CO₂ sequestration in tight formations becomes environmentally and economically attractive to energy producers and policy makers.

pH range (\sim 7–9). Overall, low salinity and low pH conditions are favored for geological CO₂ sequestration by

One of the widely used CO₂-EOR methods for tight formations is water alternating gas (WAG) flooding, in which massive amount of water is injected into the formations [9,11]. In addition, the formations usually contain a large amount of connate water originally [12–14]. In tight formations, there exists a considerable number of nanoscale pores, which is comparable to fluid molecular size [15,16], while silica is one of the most common minerals [3,4], which is generally hydrophilic [17]. Therefore, silica nanopores in tight formations can be saturated with the formation water during CO₂-EOR and CO₂ sequestration processes. As the CO₂ dissolution in brine at high pressures underground is one of CO₂ trapping mechanisms in geological CO₂ sequestration, it is imperative to

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study CO₂ solubility in brine in silica nanopores.

Several studies have shown that gas solubility in solvent in nanopores is either enhanced or suppressed compared to that in bulk [18-38], which depends on the substrate properties, pore size, molecular configurations of gas and solvent, and the ratio of molecular size between gas and solvent, etc. [26,30,34,36-38]. The detailed review of the previous works can be found in our recent work [37]. Under nanoconfinement, gas and solvent molecules can form completely different solvation structures and distributions from those in bulk [26,30,36]. Botan et al. reported that CO₂ solubility in water in montmorillonite clay interlayers can be one order of magnitude larger than that in bulk [24]. Our recent study indicated that, in kaolinite nanopores, the kaolinite substrate wettability has a dominant effect on the distributions of CO₂ and water [37]. CO₂ molecules form a strong adsorption layer on the silica facet of kaolinite, enhancing its solubility in water in kaolinite with the silica facet substrates. On the other hand, CO₂ is generally depleted from the gibbsite facet of kaolinite, thus rendering a reduced solubility in water in kaolinite with the gibbsite facet substrates [37]. In addition to the nano-confinement effect, salt ions, which are omnipresent in the formation water, can also play an important role on CO₂ solubility in brine [39–41]. Generally, CO₂ solubility in brine decreases as salinity increases due to the so-called "salting out" effect [39-42]. However, under nano-confinement, the hydration structure of salt ions is altered by the presence of substrates and fluid-solid interactions [34,43]. Therefore, the combined effect of nano-confinement and salt ions on CO₂ and water distributions as well as CO₂ solubility in brine in silica nanopores is imperative to geological CO₂ sequestration in tight formations.

On the other hand, the pH value of an aquifer varies regionally, imposing varying influences on the substrate surface chemistry [17,44,45]. For example, silica as one of the major constituents in tight formations [3,4], has varying surface chemistries as the pH of the formation water varies [17,44-47]. All silanol groups (=SiOH) on the silica surface are subject to deprotonation-protonation equilibria [17,44,45]. The neutral silanol terminated silica surface can be found at a pH between 2 and 5 [17,44,45]. As pH increases, part of silanol groups are gradually deprotonated. The deprotonation degree can reach up to 20%, when the pH is between 7 and 9 [44,45]. As a result, the pH of the formation water can dictate silica surface chemistry, which varies from a neutrally-charged surface to a negatively-charged one [17,44,45]. Such surface chemistry alteration can further influence CO₂, water, and salt ion distributions. Prior works studied water structures and diffusion in silica nanopores as a function of surface charge [46,47]. These works show that water orientation close to the silica surface is altered, and its diffusion coefficient in silica nanopores is reduced [46,47]. Bonnaud et *al.* [48] investigated Ca^{2+} ion solvation in charged silica nanopores. They found that most Ca^{2+} ions are attracted to the silica surface, while their hydration structure close to the surface is different from that in bulk. Renou et al. [49] studied water structural and dynamic properties in three different silica nanopores (protonated surface, deprotonated surface with charge re-distribution, and deprotonated surface with Na⁺ charge compensation). They found that the water distribution and diffusion are greatly affected by surface chemistry. Haria and Lorenz [50,51] studied NaCl and CaCl₂ solution flow through the charged silica nanopores, in which the effect of pore size, cation type, and salinity were considered. They observed that pore size imposes varying influences on the current of NaCl and CaCl2 solutions as charge inversion was observed in the CaCl₂ system which does not occur in the NaCl system. However, to the best of our knowledge, the pH effect on CO₂, water, and salt ion distributions as well as CO₂ solubility in brine in silica nanopores under the geological CO₂ sequestration condition are still not clear.

In this work, we study the coordinated effect of salinity and pH on CO_2 and water distributions in silica nanopores by molecular dynamics (MD) simulations. MD simulation is a computer simulation method for analyzing the physical movements of atoms and molecules by explicitly considering the intermolecular interactions. The trajectories of atoms

and molecules are determined by solving the Newton's equations of motion, and further analyzed to obtain structural and thermodynamic properties. Therefore, MD simulation is a powerful tool to explore the microscopic phenomena and mechanisms, especially for the scale and conditions where experiments are hardly accessible (e.g., the high-pressure and high-temperature geological conditions as well as the nanoscale confinement) [52]. Na⁺ and Cl⁻ are used to represent salt ions and six salt concentrations up to ~12 wt% are used to study the salinity effect. Three deprotonation degrees (0%, 8.3%, and 16.7%) of silanol groups are devised to denote the effect of pH. Temperature and pressure are set as 353 K and ~ 175 bar, respectively, which are the typical geological conditions for CO₂ sequestration [3,4].

The remainder of this paper is organized as following. In Section 2, the simulation methodology including simulation system, molecular models and validation, and simulation details is presented. In Section 3, the salinity and pH effects on the fluid distributions are analyzed by presenting the density distributions normal to the pore surface, 2-D density contour plot parallel to the pore surface, and the radial density distributions around the silanol and siloxide groups on the pore surface. Then, we present the effect of salinity and pH on CO_2 solubility in brine in silica nanopores. In Section 4, the key findings and potential implications are summarized.

2. Simulation methodology

2.1. Simulation system

An example of simulation systems is depicted as Fig. 1(a). The center of the simulation system is brine confined in a silica slit nanopore with a length of ~12 nm in the x-direction. On both sides of the pore, two ~6 nm brine slabs and two ~6 nm CO₂ slabs are placed symmetrically. The brine and CO₂ slabs serve as the outside bulk brine reservoirs and bulk CO₂ phase, respectively. CO₂ and water molecules as well as salt ions can freely move within the system via molecular diffusion. Once the system reaches equilibrium, the fluid distributions in the nanopore and bulk can be obtained in the corresponding regions. The salt ions are Na⁺ and Cl⁻, which are the most common ions in brine [39–41]. The salinities in the outside bulk brine reservoirs range from 0 to ~12 wt%, which covers the typical salinity range of formation water underground [39–41]. The initial setting for each system can be found in **Table S1** in the **Supporting Information (SI)**.

The silica nanopores consist of two identical amorphous silica sheets with hydroxylated inner surface. Amorphous silica is chosen because it is one of the most abundant constituents in tight sandstone formations [3,4]. The amorphous silica is constructed based on Emami et al. [44]. The surface area of the silica sheet is $\sim 12 \times 8.3 \text{ nm}^2$ (the x-y plane), and the thickness of each silica sheet is around 1.5 nm. The roughness of the inner surface is illustrated in Fig. 1(b). The amplitude, namely, the vertical deviation from the mean line, is around 0.03 nm. The silica slab consists of 3×2 periodic cells in the x- and y-directions as shown in Fig. 1(c1)–(c3), where the blue and green solid circles are Si atoms in \equiv SiOH and in \equiv SiO⁻, respectively. \equiv SiO⁻ groups in Fig. 1(c2) and (c3) are randomly picked and converted from \equiv SiOH groups in Fig. 1(c1), which account for 8.3% and 16.7% of the total number of surface groups, respectively. These three configurations of silica surface correspond to the pH values of \sim 2–5, \sim 5–7, and \sim 7–9, respectively [17,44,45], which cover the typical pH ranges in tight formations [53,54]. While the dissolved CO₂ in brine is subject to the reaction with water to form carbonic acid, the concentration of H₂CO₃ is three orders of magnitude smaller than the dissolved CO₂ concentration [55]. Therefore, the formation of carbonic acid is not considered in this work. The density of \equiv SiOH groups in Fig. 1(c1) is 4.74/nm², which is in line with the experimental measurements (average at $4.6-4.9/nm^2$) [56–58]. The pore size is around 2.5 nm, representing a typical nanopore in tight sandstone formations [15,16].

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Fig. 1. (a) Schematic representation of simulation system. The yellow, red, white, purple, and cyan spheres represent Si, O, H, Na⁺, and Cl⁻, respectively; the blue and pink dots are water and CO₂ molecules, respectively. (b) The surface roughness characterization: zero in the legend represent the average position of the surface atoms (excluding H atoms) in the *z*-direction, and the positive and negative values represent the coordinate deviation of surface atoms from the average position. (c1)-(c3) The distribution of Si atoms in the groups of \equiv SiO⁻ on the surface of a periodic cell (the surface consists of 3 × 2 periodic cells in the *x*- and *y*-directions); the blue and green circles represent Si atoms in \equiv SiO⁻, respectively. \equiv SiO⁻ groups in (c1), (c2), and (c3) account for the deprotonation degree of 0.0%, 8.3%, and 16.7%, respectively, which corresponds to pH of ~2–5, ~5–7, and ~7–9, respectively. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

2.2. Molecular models and validation

Silica is described by the parameters from Emami *et al* [44]. Water and CO₂ molecules are modeled by SPC/E [59] and EPM2 [60] models, respectively. Salt ions (Na⁺ and Cl⁻) parameters are adopted from Smith and Dang (SD) [61]. The force field parameters can be found in Tables S2 and S3 in the SI. The choice of such force field combination is based on the following consideration: First, regarding the silica, the parameters developed by Emami et al. resolve numerous limitations of prior silica parameters and reduce the uncertainties in the calculated interfacial properties comparing to experimental data [44]. It explicitly considers the full range of variable surface properties (such as zeta potential and silanol surface density) and pH (silanol deprotonation degree) [44]. It has been proven to accurately predict water contact angles in the water-gas (CO2 or air)-silica systems [17,44]. Moreover, this silica model is compatible with the commonly used water models (such as SPC series and TIP series) [44]. Second, the water and CO₂ models are carefully chosen among various combinations by comparing simulation results with experimental measurements in terms of CO₂ bulk density [62], CO₂-water interfacial tension (IFT) [63], CO₂ solubility in bulk water [41], and CO₂ diffusion coefficient in bulk water [64] over a broad range of pressure. The results indicate that the optimized SPC/E and EPM2 [65] can quantitatively reproduce the above-mentioned properties and outperform other force field combinations (see the results in Fig. S1 and the detailed discussion in context in the SI). Therefore, the optimized SPC/E and EPM2 are selected to model water and CO2 molecules. Third, because SPC/E model [59] has been selected for water, the salt ion force field should be compatible with SPC/E water. SD model for NaCl [61] is a widely accepted one to describe brine coupled with SPC/E water [61,66-69]. Finally, we also compare the CO₂ solubility in brine [40] and CO₂-brine IFT [70] to experimental data over a wide range of salinity (see the results in Fig. S2 and the detailed discussion in context in the SI). The semi-quantitative agreement demonstrates that the water, CO₂, and salt ion models can mimic the salinity effect with a reasonable

accuracy.

2.3. Simulation details

All simulations are conducted by the GROMACS package (version 2019.5) [71,72]. The systems are first relaxed by the steepest descent algorithm until the maximum force is less than 1000 kJ·mol^{-1} \cdot nm⁻¹. Afterward, a 50-ns NVT simulation is conducted for equilibration in each case with a time step of 2 fs. The equilibrium is carefully checked by comparing CO₂, water, and salt ion density distributions in the silica nanopores, CO₂ solubility in the outside bulk brine reservoirs and in the silica nanopores, and CO₂ density in the bulk CO₂ phase every 5-ns in each case (see Figs. S3 and S4 in the SI as an example). We find that these properties stabilize in all cases within this time period, suggesting that the systems are equilibrated in 50 ns. The sampling stage is conducted following the equilibration stage with either 10-ns or 20-ns NVT simulations for each case (refer to Table S1 for details). The trajectory in the sampling stage is saved every 100 steps (200 fs) for analysis. As a result, the ensemble averaged properties are obtained over 50,000 or 100,000 configurations. Three-dimensional periodic boundary conditions (PBCs) are applied. The cut-off distance for Lenard-Jones potential is 1.2 nm and compensated by analytical tail corrections [73]. The longrange Columbic interaction is addressed by the particle-mesh Ewald (PME) algorithm [74]. The silicon and bulk oxygen atoms in the silica substrate are fixed throughout the simulation, whereas the functional groups (hydroxyls and deprotonated hydroxyls) on the silica surface are allowed to rotate around silicon atoms (see Table S3 in the SI). Water and CO2 molecules are treated as rigid bodies by the SETTLE algorithm [75] and by introducing two virtual atoms [76], respectively. The system pressure (174.5 \pm 1.5 bar, see Table S1 for details) is dictated by CO2 density in the bulk CO2 phase from the NIST Chemistry Webbook [62] as the CO₂ model used in this work can accurately reproduce its bulk density in the studied range of pressures (see Fig. S1(a) in the SI). The system temperature (set as 353 K) is controlled by the velocity-

rescale thermostat [77].

3. Results and discussion

In this section, we first present the 1-D density profiles normal to the silica surface (along the *z*-direction), and then the 2-D density contour plots in the adsorption layer parallel to the silica surface (the *x*-*y* plane) are illustrated. Then, the radial density distributions around the silanol and siloxide groups are presented to show the effect of surface chemistry on fluid distributions. Finally, the CO_2 solubility in brine in silica nanopores and in bulk under various salinities and pH values are calculated.

3.1. Density distributions in the z-direction

In Fig. 2, we present the number density distributions of each element in silica nanopores along the z-direction in the nondeprotonated cases with various salinities. The oxygen and hydrogen atoms in water molecules are denoted as Ow and Hw, respectively; the oxygen and carbon atoms in CO2 molecules are denoted as Oc and Cc, respectively. The data are extracted from the central region of the silica nanopores (from x = 14 to 22 nm) to avoid the pore end effect as in our previous work [37]. While the statistical errors in density profiles are not shown in Fig. 2 for clarity, they would be reflected in the solubility calculations (see Section 3.4). We also present an example to show the oscillations of each elemental density profile in Fig. S5. As seen from Fig. 2, water has a layering structure close to the silica surface, forming two adsorption layers for both Ow and Hw on each surface. In addition, CO₂ also forms a single adsorption peak on each surface, while salt ions are generally depleted from the surface, which is in line with the previous studies [43,46,47]. As salinity increases, the densities of water and CO₂ monotonically drop in both the adsorption layer and bulk region. In Figs. S6 and S7, we also depict the number density distributions of each element in the silica nanopores with deprotonation degrees of 8.3% and 16.7%, respectively. As pH increases (the silica surface deprotonation

degree increases), the water adsorption layer becomes increasingly prominent, whereas CO_2 adsorption layer gradually weakens. As deprotonation degree increases from 0.0%, Na⁺ ion distributions change from the depletion at the pore surface to the enrichment due to the electrostatic interactions, generating several noticeable spikes in density profiles, while Cl⁻ ions are always depleted from the silica surface. The increase in water density in the vicinity of the pore surfaces as the deprotonation degree increases is probably because as more Na⁺ ions are attracted to the pore surfaces, more water molecules are needed to hydrate these counter-ions. Overall, all water, CO_2 , and salt ion density distributions converge to their respective bulk values away from the silica surfaces (see Fig. S8 in the SI).

In Fig. 3, we present the number density distributions, reduced density distributions as well as orientation parameters in silica nanopores with deprotonation degree of 0.0% at salinity of 7.33 wt%. The reduced density distributions represent the number density distributions normalized by the density of the corresponding component in the outside brine reservoirs. The orientation parameters S_z of water and CO₂ are given as [37]:

$$S_z = 1.5 \times \left\langle \cos^2 \theta_z \right\rangle - 0.5 \tag{1}$$

where θ_z is the angle between the *z*-axis and molecular axis (the lines connecting two Hw atoms for water and two Oc atoms for CO₂, respectively, as shown in the inset of Fig. 3(b)), and $\langle \cdots \rangle$ implies the ensemble average. Positive, zero, and negative orientation parameters represent perpendicular, random, and parallel alignments of molecular axis to the silica surface, respectively. Only the data in the range of 1.5 nm < x < 4.0 nm are presented. Beyond this range, too few water and CO₂ molecules are found therein because of the rough surface. Fig. 3 shows that CO₂ molecules co-adsorb on the silica surface with water, while salt ions are depleted. Regarding the molecular orientation, from the pore surface first, then gradually become randomly oriented and followed by a weakly-parallel alignment at its density peak position. On the other hand, CO₂ molecules gradually change from a perpendicular



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Fig. 2. Number density distributions of each element in silica nanopores in the z-direction with a deprotonation degree of 0.0% at various salinities. Ow, Hw, Oc, and Cc represent oxygen and hydrogen atoms of water, oxygen and carbon atoms of CO₂, respectively.



Fig. 3. (a) Number density distributions; (b) reduced density distributions as well as the orientation parameters in silica nanopores at salinity of 7.33 wt% and deprotonation degree of 0.0%. Ow, Hw, Oc, and Cc represent oxygen and hydrogen atoms of water, oxygen and carbon atoms of CO₂, respectively.

alignment at the pore surface to a parallel alignment in the CO_2 adsorption layer. In Figs. S9 and S10, we present, respectively, the number density distributions, reduced density distributions as well as orientation parameters in silica nanopores with deprotonation degree of 8.3% at salinity of 7.37 wt% and deprotonation degree of 16.7% at salinity of 7.70 wt%. Unlike the non-deprotonated cases as shown in Fig. 3, water orientation in those deprotonated cases is weakly perpendicular in the vicinity of the pore surface. CO_2 orientations and distributions, however, are similar to those in the non-deprotonated cases. The differences in water orientations indicate that surface charge can alter water structures close to the pore surface in line with

prior results [46,47,49]. Meanwhile, as the deprotonation degree increases, due to the more prominent Na⁺ ion accumulation close to the pore surface, Cl⁻ ions can also enrich beyond Na⁺ layers, showing a characteristic of electrical double layer (EDL). We also note that salinity has an insignificant effect on water and CO₂ orientations (see Fig. S11). The snapshots in Fig. S12 present the typical configurations of the cases in Figs. 3, S9 and S10 by VMD [78]. As deprotonation degree increases, more Na⁺ ions but fewer CO₂ molecules are found close to the pore surfaces, which is consistent with Figs. 2, S6, and S7.



Fig. 4. 2-D density contour plot in the adsorption layer parallel to the pore surface (the *x*-*y* plane) at the salinity of 7.70 wt% and deprotonation degree of 16.7%. Blue and green dots are Si atoms in Ξ Si(OH) and in Ξ Si(O⁻) groups, respectively. Ow, Hw, Oc, and Cc represent oxygen and hydrogen atoms of water, oxygen and carbon atoms of CO₂, respectively. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

W. Li et al.

3.2. Density distributions in the x-y plane

While the results in Section 3.1 can provide the density distributions of each component along the *z*-direction, they are averaged in the *x*-*y* plane, which cannot reflect the fluid distributions in the *x*-*y* plane. Therefore, we present the 2-D density contour plot in the adsorption layer parallel to the pore surface (the *x*-*y* plane) in Fig. 4 at salinity of 7.70 wt% and deprotonation degree of 16.7%. The adsorption layer for each element is defined as the region from the position where its density is 10% of its bulk density (that in the outside brine reservoirs) to the first local minimum in the *z*-direction.

As shown in Fig. 4, water molecules mainly accumulate around the \equiv SiOH and \equiv SiO⁻ groups, whereas CO₂ can only enrich in the areas where these surface groups are scarce. Salt ions, especially Na⁺, are strongly attracted by the \equiv SiO⁻ groups owing to the strong electrostatic interaction. However, salt ions are repelled from the \equiv SiOH groups. The other cases with deprotonation degrees of 0.0% and 8.3% are depicted in Figs. S13 and S14, respectively. As deprotonation degree increases, water density increases, while the opposite is true for CO₂. In addition, Na⁺ ions become gradually enriched at the surface. These results are consistent with those in Section 3.1.

3.3. Radial density surrounding the surface groups

To further explore the spatial distributions of water, CO_2 and salt ions around the surface groups (\equiv SiOH and \equiv SiO⁻), we present the radial distribution densities of each element around the O atom in \equiv SiOH and \equiv SiO⁻ groups in Figs. 5 and 6, respectively. The deprotonation degree is 16.7% and the salinity ranges from 0 to 11.89 wt%. The data are extracted from the central region of the nanopores as well. The radial density is counted as the average number of target elements in the volume of hemi-toroidal shells around the O atom in the surface groups as shown in the schematic insets in Figs. 5(b) and 6(b). The results for the cases with deprotonation degree of 0.0% and 8.3% are shown in Figs. S15-S17.

As shown in Figs. 5 and 6, the Ow distributions have a single peak

Chemical Engineering Journal xxx (xxxx) xxx

and the Hw distributions have two peaks around the O atom in \equiv SiOH groups, while around the O atom in \equiv SiO⁻ groups, both Ow and Hw distributions have a significant density peak at the contact distances followed by a few small oscillations. Besides, CO₂ distribution has a peak around the \equiv SiOH groups, while depleted from the \equiv SiO⁻ groups. As a result, as deprotonation degree increases, the number of CO₂ molecules in silica nanopores decreases. In addition, as expected, Na⁺ ion distributions have a much more prominent peak around the O atom in \equiv SiO⁻ groups than those in \equiv SiOH groups due to the net negative charge of \equiv SiO⁻ groups. Cl⁻ ions enrich at the local minimum of Na⁺ ion distributions beyond their peak around the O atom in \equiv SiOH groups, while it is depleted from \equiv SiO⁻ groups. The increasing salinity generally lowers the radial densities of water and CO₂ around the O atom in both surface groups.

Fig. 7 presents the water and Na⁺ radial distributions around \equiv SiOH and \equiv SiO⁻ groups as well as the typical configurations of water molecules and Na⁺ ions within a hemi-sphere (radius of 0.6 nm) around the O atom in the surface groups. Fig. 7(a) depicts that the hydration structure of water and Na⁺ around \equiv SiOH groups, where the Ow–Hw bonds of water molecules are pointing toward the O atom in \equiv SiOH groups. Na⁺ ions are generally depleted from the O atom in \equiv SiOH groups as shown in Fig. 7(c). On the other hand, Fig. 7(b) illustrates the hydration structure of water and Na⁺ around \equiv SiO⁻ groups, in which the peak position of Na⁺ is between those of Hw and Ow. Na⁺ ions are attracted to the \equiv SiO⁻ groups are close (see Fig. 7(d) and 7(e)). The peak value in Hw distributions is approximately twice of that in Ow distributions, suggesting the Hw–Ow–Hw orientations of water molecules around the \equiv SiO⁻ groups shown in Fig. 7(d) and 7(e).

Fig. 8 presents the radial density distributions of water, CO_2 and salt ions around the O atom in the \equiv SiOH and in \equiv SiO⁻ groups, respectively, at various deprotonation degrees. As deprotonation degree increases, water and salt ion radial density profiles around both \equiv SiOH and \equiv SiO⁻ groups increase, while those of CO_2 decrease. The increase in water and Cl⁻ radial density distributions can be partially attributed to the increased number of Na⁺ ions close to \equiv SiO⁻ groups. The increases



Fig. 5. Radial density distributions of each element around the O atom in \equiv SiOH groups at various salinities and deprotonation degree of 16.7%. Ow, Hw, Oc, and Cc represent oxygen and hydrogen atoms of water, oxygen and carbon atoms of CO₂, respectively.

Chemical Engineering Journal xxx (xxxx) xxx



Fig. 6. Radial density distributions of each element around the O atom in Ξ SiO⁻ groups at various salinities and deprotonation degree of 16.7%. Ow, Hw, Oc, and Cc represent oxygen and hydrogen atoms of water, oxygen and carbon atoms of CO₂, respectively.



Ow-Hw bonds point to OsioH

Hw-Ow-Hw angles surround Osio-

Fig. 7. Radial distribution for Hw (hydrogen of water), Ow (oxygen of water), and Na⁺ around the O atom in (a) \equiv SiOH groups with deprotonation degree of 0.0% and salinity of 7.33 wt%; (b) \equiv SiO⁻ groups surface deprotonation degree = 16.7% and salinity = 7.70 wt%, respectively. Snapshots of water and Na⁺ around (c) \equiv SiOH group, (d) \equiv SiO⁻ group at a low local density of \equiv SiO⁻ groups, and (e) \equiv SiO⁻ group at a high local density of \equiv SiO⁻ groups, respectively.

in water and salt ion radial density distributions around \equiv SiOH groups are due to an increasing number of \equiv SiO⁻ groups which can greatly attract salt ions and water molecules in their vicinities.

3.4. CO₂ solubility in brine in silica nanopores

 CO_2 solubility in brine S in silica nanopores is determined by:

Chemical Engineering Journal xxx (xxxx) xxx



Fig. 8. Radial density distributions of each element around O atom in surface groups under different deprotonation degrees @ salinity = 7.51 ± 0.19 wt%. Ow, Hw, Oc, and Cc represent oxygen and hydrogen atoms of water, oxygen and carbon atoms of CO₂, respectively.

$$S = N_{CO_2} / V_{pore} \tag{2}$$

where N_{CO_2} and V_{pore} are the molar number of CO₂ molecules and the effective pore volume, respectively. The effective pore volume is obtained from helium adsorption by the grand canonical Monte Carlo (GCMC) simulations [37,79,80]. The CO₂ solubility in brine in silica nanopores is presented in Fig. 9.

 CO_2 solubility in bulk brine is averaged over all the cases in the outside brine reservoirs (see Fig. 1(a)). It is observed that presence of salt ions generally reduces the CO_2 solubility in bulk brine and in brine in silica nanopores. At low pH values (pH values in the ranges of \sim 2–5 and \sim 5–7), CO_2 solubility in brine in silica nanopores is generally higher



Fig. 9. CO_2 solubility in brine in silica nanopores under various conditions. For comparison, we also present CO_2 solubility in bulk brine.

than that in bulk due to the strong CO₂ adsorption on the silica surfaces (see **Figs. 2 and S5**). CO₂ solubility in brine in silica nanopores in the pH range of ~2–5 is 1.3–1.6 times of that in bulk. As pH further increases (pH values in the range of ~7–9), CO₂ solubility in brine in silica nanopores becomes comparable to that in bulk brine. Collectively, low salinity and low pH conditions are favorable for CO₂ storage by solubility trapping in silica nanopores in tight formations. We note that in our previous work [37], CO₂ solubility in water is reduced in kaolinite nanopores with gibbsite facet as inner-surface (strong hydrophilic), while enhanced in kaolinite nanopores with silica facet as inner-surface (less hydrophilic). The results in this work are consistent with our previous one [37] as silica surface becomes more hydrophilic as its surface deprotonation degree increases [17].

4. Conclusion

In this work, MD simulation is used to study the coordinated effect of pH and salinity on the fluid distributions and CO_2 solubility in brine in silica nanopores, which have been poorly understood but can be of practical significance for geological CO_2 sequestration in tight formations. The number density distributions normal to the silica surface, 2-D density contour plots parallel to the surface, and radial distributions around the surface groups are presented to obtain a comprehensive understanding of fluid structural properties under various conditions.

In non-deprotonated cases, water has a layered structure close to the silica surface, and CO_2 molecules co-adsorb on the silica surface with water, while ions are depleted from the silica surface. As deprotonation degree gradually increases (from 0% to 16.7%), water adsorption becomes increasingly significant and CO_2 adsorption gradually decreases, while Na⁺ ions are strongly attracted by the silica surface, forming several density spikes. The 2-D density contour plots reveal that water molecules enrich around the surface groups on the surface, while CO_2 molecules accumulate in the areas away from the surface groups. Na⁺ ions are strongly attracted to the deprotonated surface groups. The water orientation around \equiv SiOH and \equiv SiO⁻ groups are different due to the different water hydration structures. The radial distributions of CO_2

have a peak around the \equiv SiOH groups, while they are generally depleted from the \equiv SiO⁻ groups. Overall, the increase in salinity generally lowers both water and CO₂ density distributions in silica nanopores. Collectively, CO₂ solubility in brine in silica nanopores decreases as salinity and pH increases, and CO₂ solubility in brine in silica nanopores can be 30–60% higher than that in bulk for low pH cases in the salinity range of 0–12 wt%. Therefore, low salinity and low pH environments are favored for geological CO₂ sequestration by solubility trapping in silica nanopores in tight formations. Considering that in tight formations, there are a large number of nanoscale pores, CO₂ dissolution in brine in silica nanopores may play an important role in CO₂ sequestration during CO₂-EOR in tight formations. Our work provides some important insights into the CO₂ sequestration during CO₂-EOR in tight formations and the optimization of the process.

The dissolved CO₂ in brine is subject to the reaction with water to form carbonic acid, which is a weak acid. The pH value of CO₂ saturated water solution is \sim 3.2 under the conditions in this work according to experimental measurements [81]. This suggests that the pH of brine decreases as CO₂ dissolves into the brine. The acidic environment is beneficial for CO₂ solubility in brine in silica nanopores, but this process (CO₂ dissolution in brine in silica nanopores and silica surface chemistry alteration) might take a long time, especially in tight formations. However, the solubility trapping mechanism has the time scale of thousands of years [1] and the dissolved CO₂ in brine in silica nanopores can incur the subsequent mineral trapping [82]. On the other hand, the findings in this work are only applicable for an equilibrium system, while the dynamic properties (e.g., the process of CO₂ dissolving into brine by advection or diffusion) are not discussed, which are also important to the CCS procedures. In our future work, we would explicitly explore the dynamic process of geological CO₂ sequestration.

Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Appendix A. Supplementary data

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Chemical Engineering Journal xxx (xxxx) xxx

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Chemical Engineering Journal xxx (xxxx) xxx