

**University of Alberta**

Effect of Chlorine on the Melting of the Subcratonic Lithospheric Mantle

by

Linglin Chu

A thesis submitted to the Faculty of Graduate Studies and Research  
in partial fulfillment of the requirements for the degree of

Master of Science

Department of Earth & Atmospheric Sciences

©Linglin Chu  
Spring 2010  
Edmonton, Alberta

Permission is hereby granted to the University of Alberta Libraries to reproduce single copies of this thesis and to lend or sell such copies for private, scholarly or scientific research purposes only. Where the thesis is converted to, or otherwise made available in digital form, the University of Alberta will advise potential users of the thesis of these terms.

The author reserves all other publication and other rights in association with the copyright in the thesis and, except as herein before provided, neither the thesis nor any substantial portion thereof may be printed or otherwise reproduced in any material form whatsoever without the author's prior written permission.

## **Examining Committee**

Jeff Gu, Department of Physics

Tom Chacko, Department of Earth & Atmospheric Sciences

Robert W. Luth, Department of Earth & Atmospheric Sciences

## Abstract

The presence of chlorine in the subcratonic lithospheric mantle (SCLM) has been evaluated by compiling the compositional data of fluid inclusions in fibrous diamonds. Chlorine associates with potassium, dissolving in water and forming a KCl-bearing brine with the  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  molar ratio of 0.05-0.68.

To examine the effect of such a KCl-bearing brine on the melting behavior of the SCLM, we conducted experiments in the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  and  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  systems at 5 GPa and 1100-1700 °C. In the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system, the solidus temperature of forsterite+enstatite is ~1230 °C. In the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  systems with molar  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  ratios of 0.2, 0.4 and 0.6, the solidus temperatures are ~1430 °C, ~1530 °C and ~1580 °C, respectively. The increase in the temperature of the solidus demonstrates that KCl elevates the solidus of the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system. Therefore, KCl in the SCLM can prevent melting at the  $\text{H}_2\text{O}$ -saturated solidus, and a KCl-bearing fluid can be a robust agent for mantle metasomatism.

## **Acknowledgment**

Establishing the method of study and writing this thesis were not easy tasks, and I would like to express thanks to my supervisor, Dr. Robert W. Luth, a professor of Igneous Petrology and Experimental Geochemistry in the Department of Earth & Atmospheric Sciences at the University of Alberta, for giving me his continuous guidance, supervision and advice on this research. I am grateful for his wisdom and academic strength, and I will carry the knowledge gained into future studies.

I would like to acknowledge Diane Caird, the Superpress and X-ray Technician in the Department of Earth & Atmospheric Sciences at the University of Alberta, for teaching me sample assembly and how to operate the Superpress. But more than that, I'd like to express appreciation for her motherly care while I was in the lab.

I would like to acknowledge Dr. Sergi Matveev, the Microprobe Technician in the Department of Earth & Atmospheric Sciences at the University of Alberta, and Dr. Tom Chacko, a professor in the Department of Earth & Atmospheric Sciences at the University of Alberta, for helping me develop skills and solve problems when using the microprobe.

Finally, I would like to thank my buddies, Andreas Enggist and Xianmin Hu. "A friend in need is a friend indeed."

## Table of Contents

Chapter 1. Introduction .....	1
References .....	3
Chapter 2. Chlorine-bearing fluids in the subcratonic lithospheric mantle .....	6
Introduction .....	6
Data compilation .....	7
Limitations.....	13
Composition of chlorine-bearing fluids .....	17
Conclusion.....	23
References .....	23
Chapter 3. Effect of KCl on melting of the $Mg_2SiO_4$ - $MgSiO_3$ - $H_2O$ system at 5 GPa.....	25
Introduction .....	25
Experimental techniques .....	26
Results .....	31
Discussion .....	41
Conclusion.....	42
References .....	42
Chapter 4. General Discussion and Conclusions .....	45
References .....	46
Appendix. EPMA analysis of run products .....	47

## **List of Tables**

Table 2-1. Original compositional data of fluid inclusions in fibrous diamonds.	11
Table 2-1 (continued). Original compositional data of fluid inclusions in fibrous diamonds. ....	12
Table 2-2. Calculated composition of fluid inclusions in fibrous diamonds. ....	15
Table 2-2 (continued). Calculated composition of fluid inclusions in fibrous diamonds. ....	16
Table 3-1. Starting materials. ....	29
Table 3-2. Experimental conditions and results. ....	34

## List of Figures

Figure 2-1. Concentrations of elements in fluid inclusions in fibrous diamonds. .....	19
Figure 2-2. Distribution of the concentration of Cl in fluid inclusions in fibrous diamonds. ....	20
Figure 2-3. Projection of Cl, K and Na in fluid inclusions in fibrous diamonds.	21
Figure 2-4. Projection of the calculated $[\text{SiO}_4]^{4-}$ , $[\text{CO}_3]^{2-}$ , $[\text{Cl}]^-$ and $\text{H}_2\text{O}$ for fluid inclusions in diamonds. ....	22
Figure 3-1. Thermogravimetric analysis of starting material of $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{H}_2\text{O}$ system. ....	30
Figure 3-2. Backscattered electron images of run products. ....	35
Figure 3-3. Ca contents of forsterite (filled circle) and enstatite (open circle) in the $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{H}_2\text{O}$ system as a function of temperature. ....	38
Figure 3-4. Ca contents of forsterite in the $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{H}_2\text{O}$ system and $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{KCl}$ - $\text{H}_2\text{O}$ systems. ....	39
Figure 3-5. Phase relations of the $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{H}_2\text{O}$ system and $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{KCl}$ - $\text{H}_2\text{O}$ systems at 5 GPa. ....	40

## Abbreviations

°C	degree celsius
cm	centimeter
GPa	gigapascal
h	hour
km	kilometer
kV	kilovolt
min	minute
mm	millimeter
mol%	molar percent
nA	nanoampere
P	pressure
ppm	part per million
s	second
T	temperature
wt%	weight percent



## Chapter 1. Introduction

“Volatiles” in the petrological literature are constituents that are liquid or gaseous at normal Earth’s surface conditions, and they mainly include H<sub>2</sub>O, CO<sub>2</sub> and sulfur-bearing and halogen-bearing species (Luth, 2003). Present in the Earth’s mantle, these constituents can have significant effects on the stability and physical-chemical properties of partial melts and the mantle (Dasgupta and Dixon, 2009; and references therein). Of these volatiles, halogens are rarely studied from mantle samples, and their effects on mantle processes are poorly constrained by experiments. Based on studies of fluid inclusions in fibrous diamonds (Johnson et al., 2000; Burgess et al., 2002; Burgess et al., 2009), chlorine has much higher concentrations than other halogens in the mantle-derived fluids, which indicates that chlorine is the dominant halogen in the upper mantle and may play a more significant role than fluorine, bromine or iodine.

The existence of chlorine in the upper mantle has been documented in different tectonic settings over the past 30 years. In subduction zones, chlorine exists as solutes in subduction-related magmas (Perfit et al., 1980; Poorter et al., 1989; Kent et al., 2002) and as highly saline fluid inclusions in subduction zone metasediments (Philippot, 1993) and eclogite-facies rocks (Scambelluri and Philippot, 2001). In cratonic areas, the existence of chlorine in the upper mantle is evidenced by alkali chlorides in kimberlite groundmass (Kamenetsky et al., 2004), chlorine-rich fluids trapped in fibrous diamonds (Izraeli et al., 2001; Tomlinson et al., 2006; Klein-BenDavid et al., 2007) and chlorine-bearing apatite and mica in mantle xenoliths (Smith et al., 1981; Smith, 1981). In a Group I

kimberlite described as “exceptionally fresh” from the Udachnaya-East pipe in Siberia, alkali chlorides, alkali carbonates and sulfates (ratio 5:3:1) were found as >8 wt% of the groundmass (Kamenetsky *et al.*, 2004). A mantle origin of these chlorides and carbonates has been confirmed by Sr, Nd and Pb isotope studies (Kamenetsky *et al.*, 2004; Maas *et al.*, 2005). In fibrous diamonds from the Diavik mine in Canada, fluid inclusions can have chlorine contents up to 34.2 wt% (Klein-BenDavid *et al.*, 2007). In mantle xenoliths, chlorine content can be up to 1.0 wt% in apatite and 0.05 wt% in mica (Smith, 1981; Smith *et al.*, 1981).

If widespread in the upper mantle, chlorine could potentially play important roles in affecting partial melting and metasomatism of the mantle and also in affecting physical-chemical properties of resulting melts or fluids. Experiments at pressures of <0.9 GPa show that chlorine reduces the activity of water in hydrous systems and increases the melting temperature of albite relative to the pure water case (Shmulovich and Graham, 1996). If this reduction of water activity still happens under high-pressure conditions, we anticipate that chlorine will increase the vapor-saturated solidus of hydrous peridotites and thereby prevent melting and stabilize water-rich fluids, which are effective agents for mantle metasomatism. Moreover, Keppler (1996) proposed that chloride would favor complexing with ions, such as  $\text{Rb}^{2+}$ ,  $\text{Ba}^{2+}$ ,  $\text{Sr}^{2+}$  and  $\text{Pb}^{2+}$ , rather than with more highly charged cations, such as  $\text{Nb}^{5+}$  and  $\text{Ta}^{5+}$ . Chemical transport by the chlorine-rich fluid will generate the relative enrichment of large ionic lithophile elements, such as Pb and U, and the relative depletion in high field strength elements, such as Nb and Ta, which has been invoked to explain the trace element pattern for calc-alkaline

magmas in subduction zones where chlorine can be transported by subducted slab into the mantle and then released to the mantle wedge.

This thesis is composed of two sections. The first section ([Chapter 2](#)) compiles compositional data of fluid inclusions in fibrous diamonds in order to evaluate the composition of chlorine-bearing fluids in the subcratonic lithospheric mantle. The second section ([Chapter 3](#)) reports the results of high pressure and temperature experiments in the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  and  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  systems to reveal the effects of chlorine on the melting of the hydrous Earth's mantle.

## References

- Burgess, R., Layzelle, E., Turner, G. and Harris, J.W., 2002, Constraints on the age and halogen composition of mantle fluids in Siberian coated diamonds: *Earth and Planetary Science Letters*, v. 197, p. 193-203.
- Burgess, R., Cartigny, P., Harrison, D., Hobson, E. and Harris, J., 2009, Volatile composition of microinclusions in diamonds from the Panda kimberlite, Canada: Implications for chemical and isotopic heterogeneity in the mantle: *Geochimica et Cosmochimica Acta*, v. 73, p. 1779-1794.
- Dasgupta, R., and Dixon, J.E., 2009, Volatiles and volatile-bearing melts in the Earth's interior: *Chemical Geology*, v. 262, p.1-3.
- Izraeli, E.S., Harris, J.W. and Navon, O., 2001, Brine inclusions in diamonds: a new upper mantle fluid: *Earth and Planetary Science Letters*, v. 187, p. 323-332.
- Johnson, L.H., Burgess, R., Turner, G., Hariss, J.H. and Milledge, H.J., 2000, Noble gas and halogen geochemistry of mantle fluids in diamonds: comparison of African and Canadian stones: *Geochimica et Cosmochimica Acta*, v. 64, p. 717-732.
- Kamenetsky, M.B., Sobolev, A.V., Kamenetsky, V.S., Maas, R., Danyushevsky, L.V., Thomas, R., Pokhilenko, N.P. and Sobolev, N.V., 2004, Kimberlite melts rich in alkali chlorides and carbonates: A potent metasomatic agent in the mantle: *Geology*, v 32, p. 845-848.
- Kent, A.J.R., Peate, D.W., Newman, S., Stolper, E.M. and Pearce, J.A., 2002, Chlorine in submarine glasses from the Lau Basin: seawater contamination and

constraints on the composition of slab-derived fluids: *Earth and Planetary Science Letters*, v. 202, p. 361-377.

Keppler, H., 1996, Constraints from partitioning experiments on the composition of subduction-zone fluids: *Nature*, v. 380, p. 237-240.

Klein-BenDavid, O., Izraeli, E.S., Hauri, E. and Navon, O., 2007, Fluid inclusions in diamonds from the Diavik mine, Canada and the evolution of diamond-forming fluids: *Geochimica et Cosmochimica Acta*, v. 71, p. 723-744.

Luth, R.W., 2003, Mantle volatiles-distribution and consequences. *in* Carlson R.W., Holland, H.D. and Turekian, K.K., ed., *The mantle and core: Treatise on Geochemistry*, v. 2, p. 319-361.

Maas, R., Kamenetsky, M.B., Sobolev, A.V., Kamenetsky, V.S., and Sobolev, N.V., 2005, Sr, Nd and Pb isotope evidence for a mantle origin of alkali chlorides and carbonates in the Udachnaya kimberlite, Siberia: *Geology*, v. 33, p. 549-552.

Perfit, M.R., Gust, D.A., Bence, A.E., Arculus, R.J. and Taylor, S.R., 1980, Chemical characteristics of island-arc basalts: Implications for mantle sources: *Chemical Geology*, v. 30, p. 227-265.

Philippot, P., 1993, Fluid-melt-rock interaction in mafic eclogites and coesite-bearing metasediments: Constraints on volatile recycling during subduction: *Chemical Geology*, v. 108, p. 93-112.

Poorter, R.P.E., Varekamp, J.C., Vanbergen, M.J., Kreulen, R., Sriwana, T., Vroon, P.Z. and Wirakusumah, A.D., 1989, The Sirung volcanic boiling spring: An extreme chloride-rich, acid brine on Pantar (Lesser Sunda Islands, Indonesia): *Chemical Geology*, v. 76, p. 215-228.

Scambelluri, M. and Philippot, P., 2001, Deep fluids in subduction zones: *Lithos*, v. 55, p. 213-227.

Shmulovich, K.I. and Graham, C.M., 1996, Melting of albite and dehydration of brucite in H<sub>2</sub>O-NaCl fluids to 9 kbars and 700-900°C: implications for partial melting and water activities during high pressure metamorphism: *Contributions to Mineralogy and Petrology*, v. 124, p. 370-382.

Smith, J.V., 1981, Halogen and phosphorus storage in the Earth: *Nature*, v. 289, p. 762-765.

Smith, J.V., Delaney, J.S., Herving, R.L. and Dawson, J.B., 1981, Storage of F and Cl in the upper mantle: geochemical implications: *Lithos*, v. 14, p. 133-147.

Tomlinson, E.L., Jones, A.P. and Harris, J.W., 2006, Co-existing fluid and silicate inclusions in mantle diamond: *Earth and Planetary Science Letters*, v. 250, p. 581-595.

## Chapter 2. Chlorine-bearing fluids in the subcratonic lithospheric mantle

### Introduction

The presence of chlorine-bearing fluids in the subcratonic lithospheric mantle (SCLM) has been documented by fluid inclusions in fibrous diamonds (Izraeli et al., 2001; Tomlinson et al., 2006; Klein-BenDavid et al., 2007). Because of their mechanical strength and chemical inertness, diamonds are probably the best material for transporting pristine fluid samples from the mantle to the surface (Navon et al., 1988). According to studies of syngenetic mineral inclusions in fibrous diamonds from both Koffiefontein in South Africa (Izraeli et al., 2004) and the Panda mine in Canada (Tomlinson et al., 2006), fibrous diamonds formed mainly at the pressure of 4-6 GPa and the temperature of 900-1200 °C. Therefore, the fluids trapped in the diamonds can represent fluids at depths of 130-200 km in the SCLM.

Compositions of the trapped fluids in the fibrous diamonds from different cratons have been reported by Schrauder and Navon (1994), Izraeli et al. (2001), Shiryayev et al. (2005), Tomlinson et al. (2006), Klein-BenDavid et al. (2006), Klein-BenDavid et al. (2007) and Klein-BenDavid et al. (2009). In the reported compositions, some components were analyzed but others were calculated based on various assumptions. Therefore, these data incorporate both analytical uncertainties and uncertainties associated with the assumptions inherent to the calculations, which introduces limitations in applying these data to reveal the compositions of fluids in the SCLM. Moreover, both H<sub>2</sub>O and Cl have been found as common components in the trapped fluids and the Cl/(Cl+H<sub>2</sub>O) molar ratio

may control the compositions of the trapped fluids, however, this Cl/(Cl+H<sub>2</sub>O) molar ratio has not been calculated in the previous studies.

This chapter compiles the reported compositional data of fluid inclusions in fibrous diamonds and examines their limitations, with the aim to achieve a more accurate understanding of the compositions of chlorine-bearing fluids in the SCLM. Based on the more accurate understanding of compositions, the range of Cl/(Cl+H<sub>2</sub>O) molar ratio for the fluids is calculated in this chapter, which will guide the preparation of starting materials in the next chapter ([Chapter 3](#)).

### **Data compilation**

The compositions of sub-micrometer fluid inclusions have been reported in cubic fibrous diamonds, fibrous coats of octahedral diamonds and some undefined diamond fragments from Jwaneng in Botswana ([Schrauder and Navon, 1994](#)), Koffiefontein in South Africa ([Izraeli et al., 2001](#)), Brazil ([Shiryaev et al., 2005](#)), Diavik and Panda mines in Canada ([Tomlinson et al., 2006](#); [Klein-BenDavid et al., 2006](#); [Klein-BenDavid et al., 2007](#)) and Yakutia diamond field in Russia ([Klein-BenDavid et al., 2006](#); [Klein-BenDavid et al., 2009](#)). Elements such as Si, Ti, Al, Cr, Fe, Mg, Ca, Ba, Sr, Na, K, P, S and Cl, were analyzed by electron probe microanalysis (EPMA) for individual inclusions. Volumes activated by the electron beam are bigger than inclusion volumes, and therefore the totals of the analyses are significantly lower than 100%. [Schrauder and Navon \(1994\)](#), being aware of the uncertainty arising from variable inclusion depths and the assumption of a uniform matrix, estimated precision of 10-40% for the various oxides in individual analyses and less than 10% for the sample averages. They

also estimated that the accuracy of the results obtained by EPMA is in the range of 10-15% by comparing the EPMA data with proton-probe and neutron activation data of the same diamonds. [Izraeli et al. \(2004\)](#) estimated the EPMA precision and accuracy by examination of olivine microinclusions in clouds of fibrous diamonds. They estimated the precision to be better than ~15% by analyzing the same mineral inclusions at different dates and with different EDS systems, and they estimated the accuracy to be in the range of only a few percent for the most abundant elements by comparing the range of Mg/(Mg+Fe) molar ratio and Si/(Mg+Fe) molar ratio of olivine microinclusions with those of large olivine inclusions in diamonds.

Infrared absorption spectra of the fibrous diamonds were collected with Fourier transform infrared spectroscopy (FTIR). The absorption of the trapped materials in fluid inclusions was obtained after correcting the spectrum for baseline and diamond absorption. The most common bands after the corrections are those of carbonate and water. The  $[\text{CO}_3]^{2-}/([\text{CO}_3]^{2-}+\text{H}_2\text{O})$  molar ratio was estimated from the maximum height of net absorption bands of carbonate and water at  $\sim 1430$  and  $\sim 3420 \text{ cm}^{-1}$  and Beer's law. The entire wafer thickness of diamond was sampled by IR spectroscopy, but it can be cancelled in calculating the  $[\text{CO}_3]^{2-}/([\text{CO}_3]^{2-}+\text{H}_2\text{O})$  molar ratio, because the thickness is the same for both  $[\text{CO}_3]^{2-}$  and  $\text{H}_2\text{O}$  for each analysis if assuming that  $\text{H}_2\text{O}$  and  $[\text{CO}_3]^{2-}$  are in same inclusions. Because different absorption coefficients were used in the different studies, the  $[\text{CO}_3]^{2-}/([\text{CO}_3]^{2-}+\text{H}_2\text{O})$  molar ratio has been recalculated in this thesis using consistent absorption coefficients ( $\epsilon_{\text{calcite}} = 250 \text{ L/mol cm}^{-1}$  ( $\sim 1430 \text{ cm}^{-1}$ ))



(Navon et al., 1988),  $\epsilon_{\text{liquid water}} = 80 \text{ L/mol cm}^{-1}$  ( $3420 \text{ cm}^{-1}$ ) (Thompson, 1965), which are believed by Navon et al. (1988) to be representative of the absorption by the carbonate and hydrous species in the inclusions). As stated by Navon et al. (1988), absorption coefficients of other carbonates (for example,  $\text{Na}_2\text{CO}_3$ , gaylucite) differ from that of calcite by less than 30%, and the water absorption coefficient is accurate to  $\sim 30\%$ <sup>1</sup>.

Based on results of a transmission electron microscopy study on four diamonds from Canada and Siberia, together with the tight range of composition detected by EPMA and the volatiles detected by FTIR, Klein-BenDavid et al. (2006) suggested that the fluids trapped in inclusions were uniform, dense and supercritical and then evolved to secondary phases, including crystallized minerals and low-density fluids, during cooling. Most elements in fluid inclusions are hosted in the crystallized minerals such as silicates, carbonates, phosphate, sulfates and chlorides, which is supported by the agreement of the composition of the whole inclusion to that of its daughter crystals (Klein-BenDavid et al., 2006). Based on these observations, Klein-BenDavid et al. (2006) postulated that  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ ,  $\text{Ba}^{2+}$ ,  $\text{Na}^+$  and  $\text{K}^+$  were balanced by  $[\text{Cl}]^-$ ,  $[\text{PO}_4]^{3-}$  and  $[\text{CO}_3]^{2-}$  to form chloride, phosphate and carbonate and that  $\text{Si}^{4+}$ ,  $\text{Ti}^{4+}$  and  $\text{Al}^{3+}$  were balanced by  $\text{O}^{2-}$  to form silicates, and they calculated the  $[\text{CO}_3]^{2-}$  content by the formula:

$$[\text{CO}_3]^{2-} = \text{Ca}^{2+} + \text{Mg}^{2+} + \text{Fe}^{2+} + \text{Ba}^{2+} + 0.5\text{Na}^+ + 0.5\text{K}^+ - 0.5[\text{Cl}]^- - 1.5[\text{PO}_4]^{3-}.$$

Based on studies of minerals crystallized from trapped fluids in more fibrous

---

<sup>1</sup> The original paper states this accuracy to be “-30%”. The assignment of accuracy as a negative percentage makes no sense; I believe the authors meant “ $\sim 30\%$ ”.

diamonds, Klein-BenDavid et al. (2007) assumed that  $\text{Ti}^{4+}$ ,  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ ,  $\text{Ba}^{2+}$ ,  $\text{Na}^+$  and  $\text{K}^+$  were balanced by  $[\text{Cl}]^-$ ,  $[\text{PO}_4]^{3-}$ ,  $[\text{CO}_3]^{2-}$  and  $[\text{SiO}_4]^{4-}$  and that  $\text{Si}^{4+}$  and  $\text{Al}^{3+}$  were associated with  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$  and  $\text{K}^+$  in proportions to form mica, and they changed the formula to be  $[\text{CO}_3]^{2-} = 2\text{Ti}^{4+} + \text{Ca}^{2+} + \text{Mg}^{2+} + \text{Fe}^{2+} + \text{Ba}^{2+} + 0.5\text{Na}^+ + 0.5\text{K}^+ - 0.5[\text{Cl}]^- - 1.5[\text{PO}_4]^{3-} - [\text{SiO}_4]^{4-}$ . In this formula,  $[\text{SiO}_4]^{4-}$  is subtracted with the factor of 1, which implies for the silicates with the  $(2\text{Ti} + \text{Ca} + \text{Mg} + \text{Fe} + \text{Ba} + 0.5\text{Na} + 0.5\text{K})/\text{Si}$  molar ratio of 1, such as pyroxenes. Following the calculation of the  $[\text{CO}_3]^{2-}$  content, the  $\text{H}_2\text{O}$  content can be calculated by using the  $[\text{CO}_3]^{2-}$  content and the  $[\text{CO}_3]^{2-}/([\text{CO}_3]^{2-} + \text{H}_2\text{O})$  molar ratio.

In this thesis, only Si, Al, Fe, Mg, Ca, Na, K and Cl are considered (Table 2-1) because the other elements such as Ti, Cr, Ba, Sr, P and S are minor (<5mol% for each) and not always reported. These elements, together with the  $[\text{CO}_3]^{2-}/([\text{CO}_3]^{2-} + \text{H}_2\text{O})$  molar ratio, are calculated into four components with a total of 100 mol%, including silica tetrahedron ( $[\text{SiO}_4]^{4-}$ ), carbonate ion ( $[\text{CO}_3]^{2-}$ ), chloride ion ( $[\text{Cl}]^-$ ) and water molecule ( $\text{H}_2\text{O}$ ), where  $[\text{SiO}_4]^{4-} = \text{Si}^{4+}$ ,  $[\text{CO}_3]^{2-} = \text{Ca}^{2+} + \text{Mg}^{2+} + \text{Fe}^{2+} + 0.5\text{Na}^+ + 0.5\text{K}^+ - 0.5[\text{Cl}]^-$ ,  $[\text{Cl}]^- = [\text{Cl}]^-$ , and  $\text{H}_2\text{O} = [\text{CO}_3]^{2-}(1/\alpha - 1)$  where  $\alpha = [\text{CO}_3]^{2-}/([\text{CO}_3]^{2-} + \text{H}_2\text{O})$  from FTIR.

Table 2-1. Original compositional data of fluid inclusions in fibrous diamonds.

Sample	Location	Num	wt%										C/CH	References
			SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Cl	tot			
ON-DVK-269	Diavik, Canada	12	4.1	0.7	6.1	2.3	4.8	9.5	31.9	34.2	3.1	0.15	Klein-BenDavid et al., 2007	
ON-DVK-272	Diavik, Canada	18	3.8	0.9	7.5	3.9	7.5	8.3	29.1	3.8	3.5	0.21	Klein-BenDavid et al., 2007	
ON-DVK-275	Diavik, Canada	19	5.2	0.7	5.9	4.5	6.4	11.8	24.7	30.7	3.6	0.13	Klein-BenDavid et al., 2007	
ON-DVK-276	Diavik, Canada	11	2.8	0.5	7.1	3.83	2.5	11.1	27.1	32.4	4.4	0.27	Klein-BenDavid et al., 2007	
ON-DVK-280	Diavik, Canada	33	3.8	3	5.5	2.2	4	13	23.7	32.1	3.7	0.33	Klein-BenDavid et al., 2007	
ON-DVK-282	Diavik, Canada	13	2.9	0.7	6.6	3	4.8	8.8	29.9	32.8	4.1	0.3	Klein-BenDavid et al., 2007	
ON-DVK-283	Diavik, Canada	20	4.2	0.5	5.9	4	5.6	11.5	27.5	30	3.6	0.39	Klein-BenDavid et al., 2007	
ON-DVK-286	Diavik, Canada	23	3.4	0.8	6.8	3.8	5.8	12.8	24.7	31.8	5.4	0.22	Klein-BenDavid et al., 2007	
ON-DVK-287	Diavik, Canada	20	3.5	0.5	5.3	5.9	7.3	13.9	25.4	28.9	4.2	0.36	Klein-BenDavid et al., 2007	
ON-DVK-288	Diavik, Canada	20	5	0.6	5.2	5.2	6.7	10.2	26.8	30.3	3.4	0.19	Klein-BenDavid et al., 2007	
ON-DVK-281	Diavik, Canada	48	27.5	3.6	12.6	6.2	14.3	3.2	22.5	1.6	3.4	0.27	Klein-BenDavid et al., 2007	
PAN2	Panda, Canada	43	4.6	1.4	12.2	5.7	7.1	3.9	22.4	21.3	13.5	0.34	Tomlinson et al., 2006	
PAN3	Panda, Canada	17	3.1	0.8	6.9	3.3	4.9	4.6	27.9	29.2	6.1	0.35	Tomlinson et al., 2006	
PAN4	Panda, Canada	13	15.1	1.7	9.8	1.2	4.9	4.1	26	22.2	8	0.30	Tomlinson et al., 2006	
PAN5	Panda, Canada	22	5.8	1.9	26.8	7.8	12.7	5.9	12.7	13.1	7.9	0.39	Tomlinson et al., 2006	
PAN6	Panda, Canada	27	4	1.3	16.9	4.6	7.6	7.6	21.9	23.4	7.9	0.41	Tomlinson et al., 2006	
PAN7	Panda, Canada	28	3.5	0.8	18.9	4.9	9	9	19.9	21.4	23.2	0.42	Tomlinson et al., 2006	
PAN8	Panda, Canada	17	4.5		13.6	5.5	7.4	11.8	22.4	26.6	5.1	0.41	Tomlinson et al., 2006	
JWN90	Jwaneng, Botswana	50	25.1	2.5	16.1	10.1	15.6	2.6	18.3	1.3	0.7	Schrauder and Navon, 1994		
JWN87	Jwaneng, Botswana	92	25.6	2	17.5	10.8	16.3	2.2	15.4	1	0.5	Schrauder and Navon, 1994		
JWN106	Jwaneng, Botswana	16	28.4	2.2	14.9	10.6	12.6	3	18.4	1.7	0.4	Schrauder and Navon, 1994		
JWN89	Jwaneng, Botswana	27	29.4	2.9	16.1	8.2	13.4	2.3	18.1	0.9	0.6	Schrauder and Navon, 1994		
JWN112	Jwaneng, Botswana	70	36.5	4.2	13.5	7.8	9.1	2.5	16.9	0.7	0.5	Schrauder and Navon, 1994		
JWN110	Jwaneng, Botswana	42	41.3	4.9	11.1	5.1	6.8	2.4	18.7	1.1	0.4	Schrauder and Navon, 1994		
JWN99	Jwaneng, Botswana	34	44.1	4.6	10.5	5.8	5.6	2.2	17.8	1	0.1	Schrauder and Navon, 1994		
JWN91	Jwaneng, Botswana	34	45.1	5.4	10.7	5.7	5.1	1.6	16.4	0.8	0.3	Schrauder and Navon, 1994		
JWN92	Jwaneng, Botswana	42	47.8	5.5	9.9	4.8	4.7	3	15.3	0.8	0.4	Schrauder and Navon, 1994		
JWN88	Jwaneng, Botswana	12	50.1	6.2	9.8	4.6	3.2	0.3	15.9	1.4	0.2	Schrauder and Navon, 1994		
JWN115	Jwaneng, Botswana	76	50.3	5.3	12.2	4.4	3.3	2.8	12.1	1.2	0.2	Schrauder and Navon, 1994		
JWN93	Jwaneng, Botswana	36	52.3	5.6	8.8	5.1	5.5	2	11.5	0.7	0.2	Schrauder and Navon, 1994		
155	Koffiefontein, South Africa	2	5.1	0	20.6	1.7	7.7	5.2	26.8	27.8	3.5	0.08	Izraeli et al., 2001	

Note: C/CH represents  $[\text{CO}_3]^{2-}/([\text{CO}_3]^{2-} + \text{H}_2\text{O})$  molar ratio.

Table 2-1 (continued). Original compositional data of fluid inclusions in fibrous diamonds.

Sample	Location	Num	wt%											C/CH	References
			SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	Cl	tot				
L56	Koffiefontein, South Africa	15	5.1	0.5	14.5	1.1	11.1	4.6	28.8	38.5	4.2	0.09	Izraeli et al., 2001		
L59	Koffiefontein, South Africa	7	7.2	0.7	14.2	1.8	4.7	1.1	36.5	43.2	4.1	0.02	Izraeli et al., 2001		
L62	Koffiefontein, South Africa	13	5.9	0.3	20.3	1.8	9.6	4.4	26.7	31.2	4.3	0.23	Izraeli et al., 2001		
L65	Koffiefontein, South Africa	31	7.1	0.5	10.1	1.1	3.7	2.8	39	42.2	4.2	0.004	Izraeli et al., 2001		
L75	Koffiefontein, South Africa	13	3.7	0.4	20	2.2	7.7	2	30.6	35.7	6.4	0.18	Izraeli et al., 2001		
L76	Koffiefontein, South Africa	11	8.9	0.6	22	2	7.5	4.1	25.5	29.3	3.7	0.055	Izraeli et al., 2001		
UB 5-41	Yakutia, Russia	2	7.3	1.5	9.5	10.6	15.3	12.1	24.5	7.8	3.6	0.56	Klein-BenDavid et al., 2009		
UDC 239	Yakutia, Russia	52	9.8	1.1	7.3	25.5	20.5	9.5	17.3	2.5	5.4	0.64	Klein-BenDavid et al., 2009		
UDC 241	Yakutia, Russia	35	8.5	0.6	7.9	19.1	17.7	10.5	24	5.4	4.4	0.64	Klein-BenDavid et al., 2009		
UDC 243	Yakutia, Russia	7	8.5	1.5	11.2	17.7	15.4	8.8	22	4.7	2.2	0.62	Klein-BenDavid et al., 2009		
UDC 244	Yakutia, Russia	27	9.2	0.7	7.1	22.8	18.8	12.1	18.6	3.7	3.3	0.67	Klein-BenDavid et al., 2009		
UDC 246	Yakutia, Russia	4	7.6	1.4	6.4	17.5	12.5	14.2	23.4	7.2	2.9	0.56	Klein-BenDavid et al., 2009		
UDC 252	Yakutia, Russia	28	8	0.6	8.5	22	20.6	8.2	23.5	4	4.3	0.66	Klein-BenDavid et al., 2009		
ZAR 263	Yakutia, Russia	25	9.5	1.7	6.3	23.1	16.9	9.7	23	5	3.3	0.68	Klein-BenDavid et al., 2009		
ZAR 265	Yakutia, Russia	14	8.8	1.7	8	22.6	18.6	4	25.6	4.3	3	0.63	Klein-BenDavid et al., 2009		
UB 81	Yakutia, Russia	23	8.8	0.5	7.4	24.1	18.2	9.4	20.6	4.1	5.1	0.65	Klein-BenDavid et al., 2009		
FK 925	Yakutia, Russia	29	8.6	0.5	10	23.6	24.9	6.9	19.8	2.4		0.67	Klein-BenDavid et al., 2009		
FK 951	Yakutia, Russia	50	8.6	0.6	8.2	25.3	26.3	7.5	17.3	3.2		0.66	Klein-BenDavid et al., 2009		
FK 952	Yakutia, Russia	15	9.8	0.3	10.4	21.6	16.1	6.3	27.2	5.2		0.49	Klein-BenDavid et al., 2009		
AIK 297	Yakutia, Russia	3	11.9	3.6	14	10.7	21.2	8.1	18.7	1.9	4.3	0.32	Klein-BenDavid et al., 2009		
EG 133	Yakutia, Russia	15	9.8	1.3	17.5	12	21.3	8.3	22.5	2.7		0.66	Klein-BenDavid et al., 2009		
AIK 298	Yakutia, Russia	20	16.3	3.1	11.9	10.9	17.1	7	23.6	3	4.5	0.56	Klein-BenDavid et al., 2009		
EG 138	Yakutia, Russia	33	37.1	6.6	8.4	3.8	16.7	0.6	17.3	5.9		0.3	Klein-BenDavid et al., 2009		
FK 927	Yakutia, Russia	17	62.8	12.5	2.9	1.6	3.4	0.01	15.1	0.9		0.09	Klein-BenDavid et al., 2009		
UDC 242	Yakutia, Russia	5	57.4	13	3.8	1.4	5.8	0	12.8	0.7	5.6	0.28	Klein-BenDavid et al., 2009		
UDC 255	Yakutia, Russia	31	65.2	13.3	1.2	0.8	2.2	2.3	10.7	1	14.2	0.19	Klein-BenDavid et al., 2009		
Sp.BR-1	Brazil	84	38.6	7.4	6.4	3	11.7	3.1	16.1	4.9		0.28	Shiryaev et al., 2005		
Sp.BR-2	Brazil	41	25.6	2.8	13.2	7.1	16.2	3.9	15.2	4.8		0.56	Shiryaev et al., 2005		
Sp.BR-5inner	Brazil	45	45.4	5.2	9.7	4	6.7	2.8	14.9	3.7		0.43	Shiryaev et al., 2005		
Sp.BR-5outer	Brazil	23	25.4	2.6	13.5	7	16.9	4.5	15.7	4.8		0.68	Shiryaev et al., 2005		

Note: C/CH represents  $[\text{CO}_3]^{2-}/([\text{CO}_3]^{2-}+\text{H}_2\text{O})$  molar ratio.

## Limitations

The calculation of  $[\text{CO}_3]^{2-}$  in this thesis is based on the assumptions used by [Klein-BenDavid et al. \(2007\)](#) that  $\text{Ti}^{4+}$ ,  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ ,  $\text{Ba}^{2+}$ ,  $\text{Na}^+$  and  $\text{K}^+$  are balanced by  $[\text{Cl}]^-$ ,  $[\text{PO}_4]^{3-}$ ,  $[\text{CO}_3]^{2-}$  and  $[\text{SiO}_4]^{4-}$  and that  $\text{Si}^{4+}$  and  $\text{Al}^{3+}$  are associated with  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$  and  $\text{K}^+$  in proportions to form mica. This assumption considers that  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$  and  $\text{K}^+$  not only combine with  $[\text{CO}_3]^{2-}$  and  $[\text{Cl}]^-$  but also combine with  $[\text{SiO}_4]^{4-}$ , which is supported by the discovery of silicate minerals such as mica, pyroxenes and olivine in fluid inclusions. According to this assumption, the calculation of  $[\text{CO}_3]^{2-}$  should include the subtraction of  $[\text{SiO}_4]^{4-}$  with a factor. In the calculation by [Klein-BenDavid et al. \(2007\)](#), this factor is 1, which is only suitable for the fluid inclusions where the crystallized silicate minerals have the  $(2\text{Ti} + \text{Ca} + \text{Mg} + \text{Fe} + \text{Ba} + 0.5\text{Na} + 0.5\text{K})/\text{Si}$  molar ratio of 1. However, most crystallized silicate minerals in fluid inclusions do not have this molar ratio of 1. For example, quartz has this molar ratio of 0, and olivine has this molar ratio of 2. For fluid inclusions where the crystallized silicate mineral is only quartz, the calculation by [Klein-BenDavid et al. \(2007\)](#) will generate a negative value for  $[\text{CO}_3]^{2-}$ . Therefore, the calculation of  $[\text{CO}_3]^{2-}$  in this thesis does not include the subtraction of  $[\text{SiO}_4]^{4-}$ . Besides this, the calculation of  $[\text{CO}_3]^{2-}$  also does not consider the minor elements such as Ti, Cr, Ba, Sr, P and S.

Without subtracting  $[\text{SiO}_4]^{4-}$ ,  $[\text{CO}_3]^{2-}$  can be overestimated for Si-rich fluid inclusions, and  $\text{H}_2\text{O}$  can also be overestimated because  $\text{H}_2\text{O}$  is calculated from  $[\text{CO}_3]^{2-}$  using the formula:  $\text{H}_2\text{O} = [\text{CO}_3]^{2-}/(\alpha - 1)$ . Therefore, the  $[\text{Cl}]^-/([\text{Cl}]^- + \text{H}_2\text{O})$  molar ratio can be underestimated for Si-rich fluid inclusions because of the

overestimation of H<sub>2</sub>O, and the  $[\text{Cl}^-]/([\text{Cl}^-] + \text{H}_2\text{O})$  molar ratio for Si-rich fluid inclusions cannot represent their actual Cl/(Cl + H<sub>2</sub>O) molar ratio. However, the effect of not subtracting  $[\text{SiO}_4]^{4-}$  is negligible for Si-poor fluid inclusions. Furthermore, the neglect of Ti, Cr, Ba, Sr, P and S could cause uncertainty in the calculated  $[\text{CO}_3]^{2-}$ . For example, neglecting Ti, Cr, Ba and Sr underestimates  $[\text{CO}_3]^{2-}$ , and neglecting P and S overestimates  $[\text{CO}_3]^{2-}$ . Because these elements are minor in fluid inclusions and the underestimation and overestimation of  $[\text{CO}_3]^{2-}$  can be cancelled by each other to some extent, the uncertainty caused by neglecting these elements is negligible. Therefore, the calculation without subtracting  $[\text{SiO}_4]^{4-}$  and without considering the minor elements can still yield a relatively accurate Cl/(Cl + H<sub>2</sub>O) molar ratio for the Si-poor fluid inclusions.

Table 2-2. Calculated composition of fluid inclusions in fibrous diamonds.

Sample	Location	mol %										mol%			
		Si	Al	Fe	Mg	Ca	Na	K	Cl	[SiO <sub>4</sub> ] <sup>4-</sup>	[CO <sub>3</sub> ] <sup>2-</sup>	[Cl]	H <sub>2</sub> O		
ON-DVK-269	Diavik, Canada	3.0	0.6	3.8	2.5	3.8	13.6	30.0	42.7	2.61	9.07	36.90	51.41		
ON-DVK-272	Diavik, Canada	4.5	1.3	7.4	6.9	9.5	19.0	43.9	7.6	1.75	20.01	2.96	75.28		
ON-DVK-275	Diavik, Canada	4.0	0.6	3.8	5.1	5.2	17.5	24.1	39.7	2.49	9.43	24.94	63.13		
ON-DVK-276	Diavik, Canada	2.2	0.5	4.6	4.4	2.1	16.7	26.9	42.7	2.48	13.21	48.61	35.70		
ON-DVK-280	Diavik, Canada	2.9	2.7	3.6	2.5	3.3	19.5	23.4	42.1	3.93	13.13	56.29	26.65		
ON-DVK-282	Diavik, Canada	2.2	0.6	4.3	3.4	4.0	13.2	29.4	42.9	2.68	13.80	51.33	32.20		
ON-DVK-283	Diavik, Canada	3.2	0.5	3.8	4.6	4.6	17.2	27.0	39.1	3.93	18.89	47.63	29.55		
ON-DVK-286	Diavik, Canada	2.6	0.7	4.3	4.3	4.7	18.8	23.8	40.8	2.38	13.17	37.77	46.68		
ON-DVK-287	Diavik, Canada	2.6	0.4	3.3	6.6	5.9	20.2	24.3	36.7	2.79	20.93	39.07	37.21		
ON-DVK-288	Diavik, Canada	3.8	0.5	3.3	5.9	5.5	15.2	26.2	39.4	3.04	12.50	31.18	53.29		
ON-DVK-281	Diavik, Canada	26.3	4.1	10.1	8.8	14.7	5.9	27.5	2.6	12.51	23.29	1.23	62.97		
PAN2	Panda, Canada	4.4	1.6	9.7	8.1	7.3	7.2	27.3	34.4	3.88	22.22	30.46	43.44		
PAN3	Panda, Canada	2.7	0.8	5.1	4.3	4.6	7.8	31.2	43.4	3.40	14.74	54.22	27.64		
PAN4	Panda, Canada	13.6	1.8	7.4	1.6	4.7	7.2	29.9	33.9	13.83	15.55	34.46	36.16		
PAN5	Panda, Canada	5.5	2.1	21.2	11.0	12.9	10.8	15.4	21.0	3.65	31.74	13.99	50.62		
PAN6	Panda, Canada	3.4	1.3	12.1	5.9	7.0	12.6	23.9	33.9	3.40	26.02	33.68	36.90		
PAN7	Panda, Canada	3.0	0.8	13.6	6.3	8.3	15.0	21.8	31.2	2.80	28.84	29.02	39.34		
PAN8	Panda, Canada	3.5		8.8	6.4	6.2	17.8	22.2	35.1	3.64	24.79	36.42	35.15		
JWN90	Jwaneng, Botswana	24.2	2.8	13.0	14.5	16.1	4.9	22.5	2.1	22.69	52.72	1.99	22.60		
JWN87	Jwaneng, Botswana	25.2	2.3	14.4	15.8	17.2	4.2	19.3	1.7	17.54	40.65	1.16	40.65		
JWN106	Jwaneng, Botswana	27.1	2.5	11.9	15.1	12.9	5.5	22.4	2.7	16.83	32.58	1.71	48.87		
JWN89	Jwaneng, Botswana	28.8	3.4	13.2	12.0	14.1	4.4	22.7	1.5	24.63	44.45	1.28	29.64		
JWN112	Jwaneng, Botswana	35.9	4.9	11.1	11.4	9.6	4.8	21.2	1.2	28.46	35.31	0.93	35.31		
JWN110	Jwaneng, Botswana	40.6	5.7	9.1	7.5	7.2	4.6	23.5	1.8	30.17	27.39	1.36	41.08		
JWN99	Jwaneng, Botswana	43.4	5.3	8.6	8.5	5.9	4.2	22.3	1.7	10.85	8.87	0.42	79.86		
JWN91	Jwaneng, Botswana	45.2	6.4	9.0	8.5	5.5	3.1	21.0	1.4	28.08	21.32	0.84	49.75		
JWN92	Jwaneng, Botswana	47.1	6.4	8.2	7.1	5.0	5.7	19.2	1.3	36.68	24.91	1.04	37.37		
JWN88	Jwaneng, Botswana	50.5	7.4	8.3	6.9	3.5	0.6	20.5	2.4	26.21	14.51	1.24	58.04		
JWN115	Jwaneng, Botswana	50.4	6.3	10.2	6.6	3.5	5.4	15.5	2.0	25.04	14.79	1.01	59.16		
JWN93	Jwaneng, Botswana	52.6	6.6	7.4	7.6	5.9	3.9	14.7	1.2	26.00	14.68	0.59	58.73		
155	Koffiefontein, South Africa	4.1	0.0	13.8	2.0	6.6	8.1	27.5	37.8	1.33	6.91	12.25	79.51		

Table 2-2 (continued). Calculated composition of fluid inclusions in fibrous diamonds.

Sample	Location	mol %										mol%			
		Si	Al	Fe	Mg	Ca	Na	K	Cl	[SiO <sub>4</sub> ] <sup>4-</sup>	[CO <sub>3</sub> ] <sup>2-</sup>	[Cl]	H <sub>2</sub> O		
156	Kofffontein, South Africa	3.6	0.4	8.5	1.2	8.4	6.3	25.8	45.9	2.07	6.43	26.46	65.04		
159	Kofffontein, South Africa	4.8	0.6	7.9	1.8	3.4	1.4	31.1	49.0	1.61	1.64	16.37	80.38		
162	Kofffontein, South Africa	4.5	0.3	12.9	2.0	7.8	6.5	25.9	40.2	3.54	14.89	31.73	49.84		
165	Kofffontein, South Africa	4.8	0.4	5.7	1.1	2.7	3.7	33.5	48.2	0.46	0.38	4.62	94.55		
175	Kofffontein, South Africa	2.7	0.3	12.3	2.4	6.1	2.9	28.7	44.5	2.15	11.29	35.11	51.45		
176	Kofffontein, South Africa	6.9	0.5	14.2	2.3	6.2	6.2	25.2	38.4	1.75	4.87	9.74	83.65		
UB 5-41	Yakutia, Russia	6.2	1.5	6.8	13.5	14.0	20.0	26.7	11.3	5.65	47.11	10.23	37.01		
UDC 239	Yakutia, Russia	8.0	1.1	5.0	31.2	18.0	15.1	18.1	3.5	6.73	57.83	2.91	32.53		
UDC 241	Yakutia, Russia	6.9	0.6	5.4	23.1	15.4	16.5	24.8	7.4	6.31	55.62	6.79	31.28		
UDC 243	Yakutia, Russia	7.4	1.5	8.1	22.8	14.3	14.8	24.3	6.9	6.50	54.19	6.09	33.21		
UDC 244	Yakutia, Russia	7.4	0.7	4.8	27.5	16.3	19.0	19.2	5.1	6.78	59.36	4.62	29.24		
UDC 246	Yakutia, Russia	6.1	1.3	4.3	21.1	10.8	22.3	24.1	9.9	5.42	48.09	8.70	37.79		
UDC 252	Yakutia, Russia	6.5	0.6	5.8	26.6	17.9	12.9	24.3	5.5	5.78	58.95	4.90	30.37		
ZAR 263	Yakutia, Russia	7.5	1.6	4.2	27.3	14.4	14.9	23.3	6.7	7.19	58.75	6.41	27.65		
ZAR 265	Yakutia, Russia	7.4	1.7	5.6	28.4	16.8	6.5	27.5	6.1	6.37	55.66	5.28	32.69		
UB_81	Yakutia, Russia	7.2	0.5	5.1	29.3	15.9	14.9	21.5	5.7	6.31	57.66	4.98	31.05		
FK 925	Yakutia, Russia	7.0	0.5	6.8	28.8	21.8	11.0	20.7	3.3	6.00	61.08	2.84	30.08		
FK 951	Yakutia, Russia	6.9	0.6	5.5	30.4	22.7	11.7	17.8	4.4	5.82	59.74	3.67	30.78		
FK 952	Yakutia, Russia	7.9	0.3	7.0	26.0	13.9	9.8	28.0	7.1	5.56	43.82	5.00	45.61		
AIK 297	Yakutia, Russia	10.9	3.9	10.7	14.6	20.8	14.4	21.8	2.9	5.19	29.89	1.40	63.52		
EG 133	Yakutia, Russia	8.4	1.3	12.6	15.4	19.7	13.9	24.7	3.9	7.61	58.63	3.56	30.20		
AIK 298	Yakutia, Russia	14.4	3.2	8.8	14.3	16.2	12.0	26.6	4.5	12.04	47.15	3.76	37.05		
EG 138	Yakutia, Russia	34.1	7.2	6.5	5.2	16.5	1.1	20.3	9.2	21.68	21.74	5.84	50.73		
FK 927	Yakutia, Russia	58.8	13.8	2.3	2.2	3.4	0.0	18.0	1.4	24.44	6.75	0.59	68.22		
UDC 242	Yakutia, Russia	56.4	15.1	3.1	2.1	6.1	0.0	16.1	1.2	45.33	15.05	0.94	38.69		
UDC 255	Yakutia, Russia	62.0	14.9	1.0	1.1	2.2	4.2	13.0	1.6	48.64	9.52	1.26	40.58		
Sp.BR-1	Brazil	36.9	8.3	5.1	4.3	12.0	5.7	19.6	7.9	24.48	19.97	5.27	50.29		
Sp.BR-2	Brazil	24.9	3.2	10.7	10.3	16.9	7.3	18.8	7.9	21.30	40.26	6.77	31.67		
Sp.BR-5inner	Brazil	43.9	5.9	7.8	5.8	6.9	5.2	18.4	6.1	37.03	24.75	5.12	33.10		
Sp.BR-5outer	Brazil	24.1	2.9	10.7	9.9	17.2	8.3	19.0	7.7	23.61	46.60	7.56	22.23		



## Composition of chlorine-bearing fluids

Concentrations of the elements Si, Al, Fe, Mg, Ca, Na, K and Cl are calculated into molar percentages and shown in [Table 2-2](#) and [Figure 2-1](#). From the concentration of Cl, fluids in fibrous diamonds divide into high-Cl fluids and low-Cl fluids ([Figure 2-1a and 2-1b, respectively](#)), which are shown as the two peaks in the histogram of Cl mol% in [Figure 2-2](#). The high-Cl fluids are rich in Cl and K but poor in Si, Mg and Ca, and they are from Diavik and Panda mines in Canada and Koffiefontein in South Africa. In contrast, the low-Cl fluids are poor in Cl but rich in Si, Mg, Ca and K, and they are from Brazil, the Diavik mine in Canada, Jwaneng in Botswana and Yakutia in Russia. The Cl/(Cl+K) molar ratio is around 0.6 for the high-Cl fluids but <0.35 for the low-Cl fluids, and the K/(K+Na) molar ratio is >0.5 for all fluids ([Figure 2-3](#)). Overall, the concentration of elements indicates that fluids trapped in fibrous diamonds can be divided into the high-Cl fluids and the low-Cl fluids, and Cl is mainly associated with K.

The four components  $[\text{SiO}_4]^{4-}$ ,  $[\text{CO}_3]^{2-}$ ,  $[\text{Cl}]^-$  and  $\text{H}_2\text{O}$  are calculated and shown in [Table 2-2](#) and projected in [Figure 2-4](#). This projection of the four components shows that the high-Cl fluids, except those in fibrous diamonds from the Panda mine, fall along a line and the low-Cl fluids distribute close to the  $[\text{SiO}_4]^{4-}$ - $[\text{CO}_3]^{2-}$ - $\text{H}_2\text{O}$  plane. The fluid inclusions in the fibrous diamonds from the Panda mine fall between the line and the plane, which may be caused by mixing of different types of fluids. As listed in [Table 2-2](#), the  $[\text{Cl}^-]/([\text{Cl}^-]+\text{H}_2\text{O})$  molar ratio of the high-Cl fluids is 0.05-0.68, and the  $[\text{Cl}^-]/([\text{Cl}^-]+\text{H}_2\text{O})$  molar ratio of the low-Cl fluids is 0.01-0.25. Because the chlorine-rich fluids are poor in Si, the

calculated  $[\text{CO}_3]^{2-}$  and  $\text{H}_2\text{O}$  are not affected by the calculation without subtracting  $[\text{SiO}_4]^{4-}$  and therefore, their calculated  $[\text{Cl}^-]/([\text{Cl}^-+\text{H}_2\text{O}]$  molar ratio would represent the actual  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  molar ratio of fluids. However, because the low-Cl fluids are rich in Si, the calculated  $[\text{Cl}^-]/([\text{Cl}^-+\text{H}_2\text{O}]$  molar ratio underestimates the  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  molar ratio of the fluids. Therefore, the  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  molar ratio of chlorine-bearing fluids in fibrous diamonds is more accurately represented by the  $[\text{Cl}^-]/([\text{Cl}^-+\text{H}_2\text{O}]$  molar ratio of 0.05-0.68 for the high-Cl fluids.

In previous works, [Klein-BenDavid et al. \(2007\)](#) concluded that there is a continuous compositional evolution between saline-HDFs and carbonatitic-HDFs by plotting the composition of fluid inclusions on the (Si+Al)-(K+Na)-(Ca+Mg+Fe) triangular diagram, and, similarly, [Safonov et al. \(2007\)](#) also concluded a continuous compositional evolution between the Si-bearing chloride-carbonate brine and the chloride-bearing carbonatitic liquid by plotting compositions of fluid inclusions on the (Si+Al+Ti+P)-Cl-(Ca+Mg+Fe+Ba+Sr) triangular diagram. Because the high-Cl fluids in this thesis are equivalents to the saline-HDFs by [Klein-BenDavid et al. \(2007\)](#) and the Si-bearing chloride-carbonate brine by [Safonov et al. \(2007\)](#) and the low-Cl fluids in this thesis include the carbonatitic-HDFs by [Klein-BenDavid et al. \(2007\)](#) and the chloride-bearing carbonatitic liquid by [Safonov et al. \(2007\)](#), the division of high-Cl fluids and low-Cl fluids in this thesis is opposed to the continuous compositional evolution by [Klein-BenDavid et al. \(2007\)](#) and [Safonov et al. \(2007\)](#).

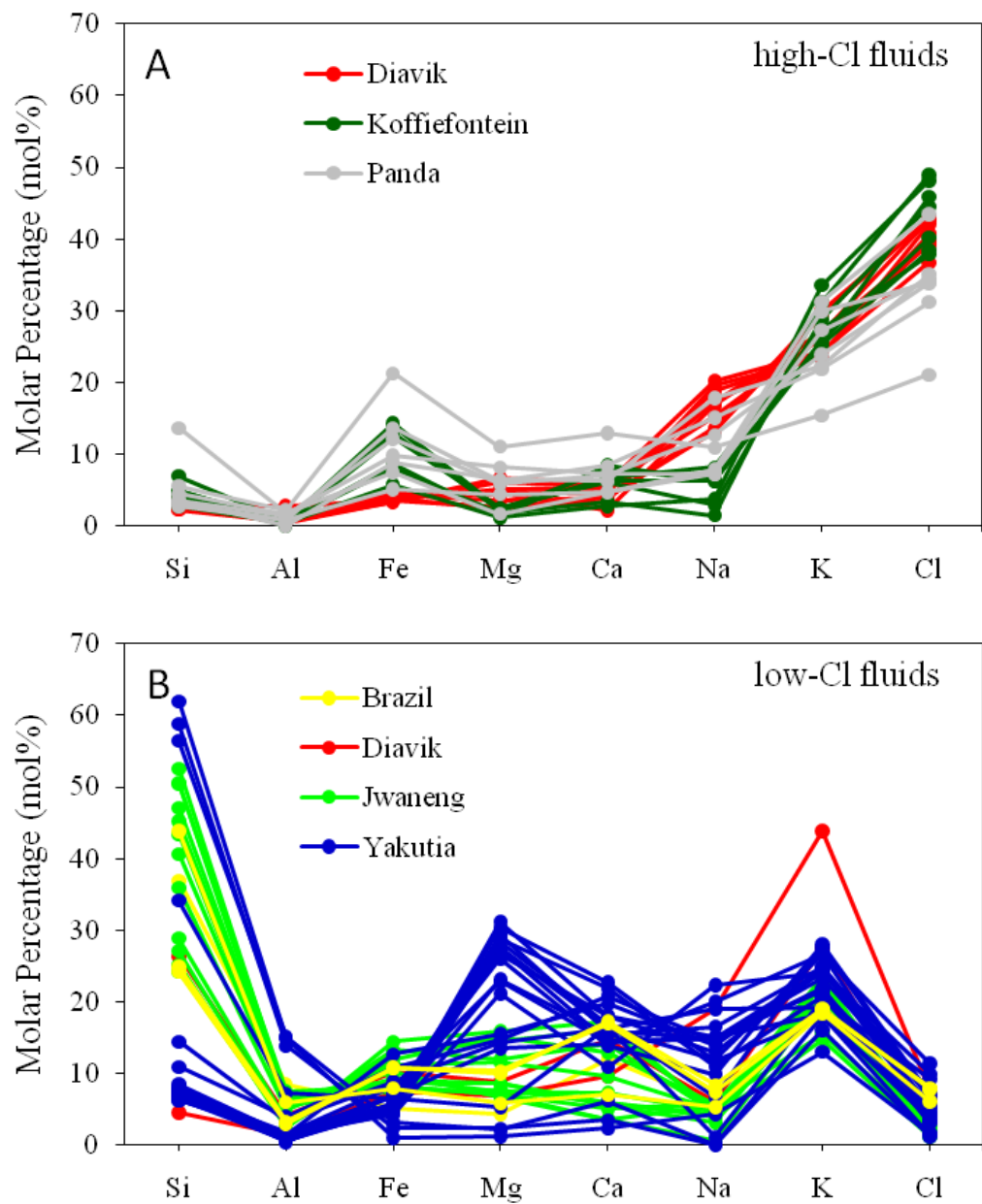


Figure 2-1. Concentrations of elements in fluid inclusions in fibrous diamonds.

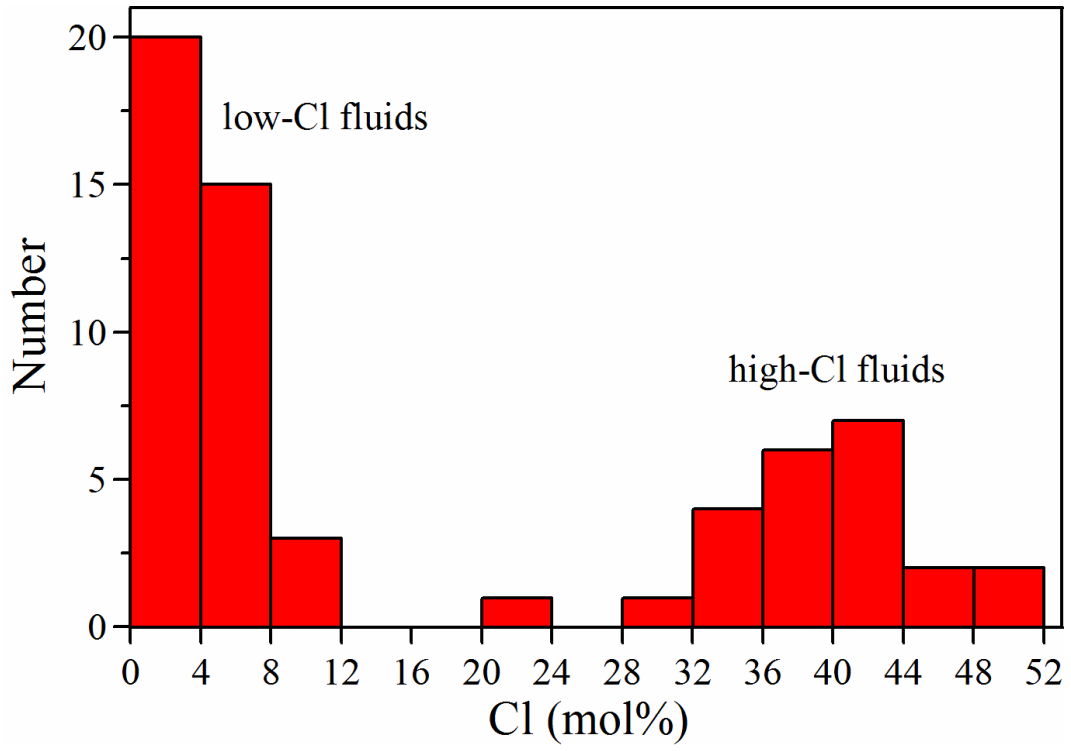


Figure 2-2. Distribution of the concentration of Cl in fluid inclusions in fibrous diamonds.

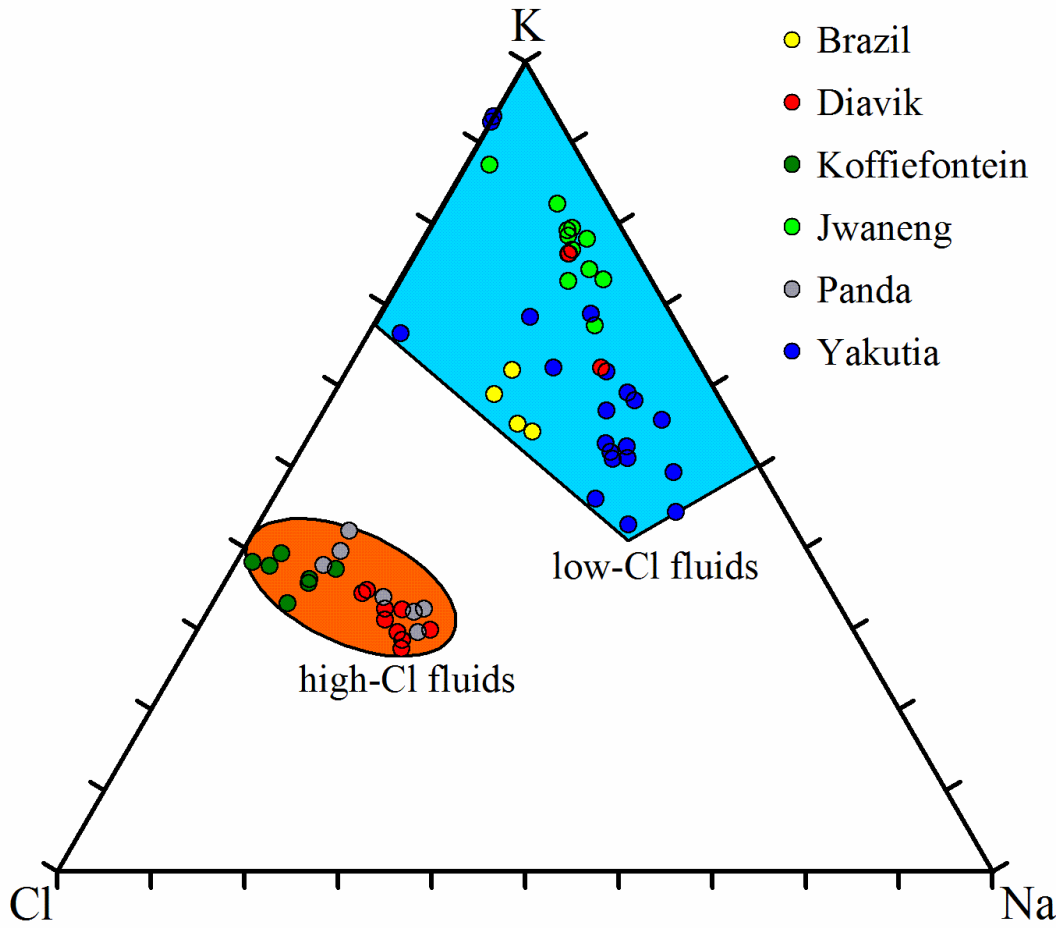


Figure 2-3. Projection of Cl, K and Na in fluid inclusions in fibrous diamonds.

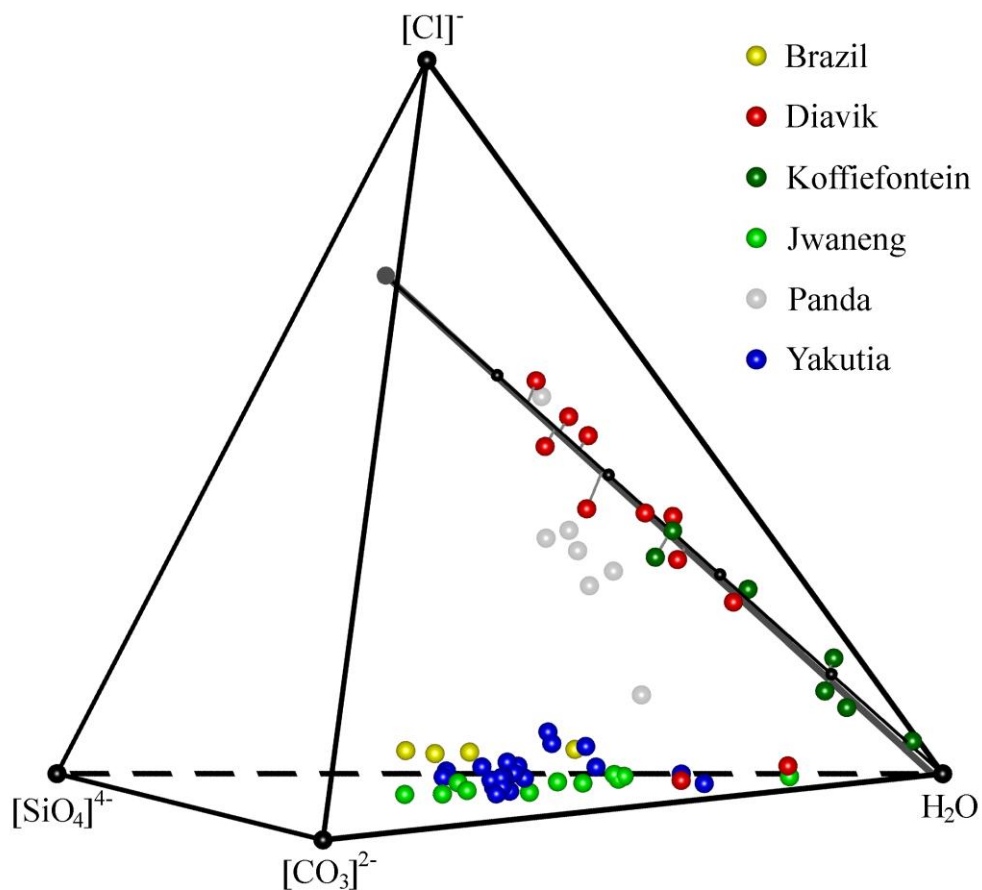


Figure 2-4. Projection of the calculated  $[\text{SiO}_4]^{4-}$ ,  $[\text{CO}_3]^{2-}$ ,  $[\text{Cl}]^-$  and  $\text{H}_2\text{O}$  for fluid inclusions in diamonds.

Data are in [Table 2-2](#). The formula to calculate  $[\text{CO}_3]^{2-}$  is in the text. The gray solid line is the regression of chlorine-rich fluid inclusions in diamonds from Diavik and Koffiefontein mines, which can be expressed as the equations:  $[\text{SiO}_4]^{4-} = -0.0378\text{H}_2\text{O} + 4.5295$  ( $r^2 = 0.73$ ),  $[\text{CO}_3]^{2-} = -0.2387\text{H}_2\text{O} + 23.789$  ( $r^2 = 0.78$ ) and  $[\text{Cl}]^- = -0.7236\text{H}_2\text{O} + 71.682$  ( $r^2 = 0.97$ ), and which intersects the  $[\text{SiO}_4]^{4-}$ - $[\text{CO}_3]^{2-}$ - $[\text{Cl}]^-$  plane at the point ( $[\text{SiO}_4]^{4-}$ ,  $[\text{CO}_3]^{2-}$ ,  $[\text{Cl}]^-$ ) = (4.5, 23.8, 71.7). The black solid line ties the point of the intersection and the  $\text{H}_2\text{O}$  vertex, and it almost superposes the gray regressed line. Small black spheres on the black solid line show the water contents of 20, 40, 60, and 80 mol%.

This significant difference is due to considering H<sub>2</sub>O as an end member in plotting, in this thesis. Klein-BenDavid et al. (2007) and Safonov et al. (2007) did not consider H<sub>2</sub>O as an end member of fluid compositions in plotting, therefore, their results of plotting are equal to that of projecting the compositions of fluid inclusions from the “H<sub>2</sub>O” vertex to the [SiO<sub>4</sub>]<sup>4-</sup>-[CO<sub>3</sub>]<sup>2-</sup>-[Cl]<sup>-</sup> triangular plane in Figure 2-4. This projection can cause an intersection or connection of the group of high-Cl fluids and the group of low-Cl fluids in the [SiO<sub>4</sub>]<sup>4-</sup>-[CO<sub>3</sub>]<sup>2-</sup>-[Cl]<sup>-</sup> triangular plane. Therefore, the division of high-Cl fluids and low-Cl fluids, revealed by plotting with considering H<sub>2</sub>O in compositions of fluid inclusions in this thesis, is more reliable than the continuous compositional evolution of fluids by Klein-BenDavid et al. (2007) and Safonov et al. (2007). This division is possibly caused by fluid immiscibility between the high-Cl fluids and the low-Cl fluids because both of these two kinds of fluid inclusions exist in fibrous diamonds from the Diavik mine in Canada.

## **Conclusion**

Fluids trapped in fibrous diamonds represent fluids in the SCLM, and they divide into high-Cl fluids and low-Cl fluids. Chlorine in the fluids is mainly associated with K and dissolves in H<sub>2</sub>O with the Cl/(Cl+H<sub>2</sub>O) molar ratio of 0.05-0.68.

## **References**

- Izraeli, E.S., Harris, J.W. and Navon, O., 2004, Fluid and mineral inclusions in cloudy diamonds from Koffiefontein, South Africa: *Geochimica et Cosmochimica Acta*, v. 68, p. 2561-2575.
- Izraeli, E.S., Harris, J.W. and Navon, O., 2001, Brine inclusions in diamonds: a new upper mantle fluid: *Earth and Planetary Science Letters*, v. 187, p. 323-332.

Klein-BenDavid, O., Izraeli, E.S., Hauri, E.H. and Navon, O., 2007, Fluid inclusions in diamonds from the Diavik Mine, Canada and the evolution of diamond-forming fluids: *Geochimica et Cosmochimica Acta*, v. 71, p. 723-744.

Klein-BenDavid, O., Logvinova, A.M., Schrauder, M., Spetius, Z.V., Weiss, Y., Hauri, E., Kaminsky, F.V., Sobolev, N.V. and Navon, O., 2009, High-Mg carbonatitic microinclusions in some Yakutian Diamonds- a new type of diamond-forming fluid: *Lithos*, In Press.

Klein-BenDavid, O., Richard, W. and Navon, O., 2006, TEM imaging and analysis of microinclusions in diamonds; a close look at diamond-growing fluids: *American Mineralogist*, v. 91, p. 353-365.

Navon, O., Hutcheon, I.D., Rossman, G.R. and Wasserburg, G.J., 1988, Mantle-derived fluids in diamond micro-inclusions: *Nature*, v. 335, p. 784-789.

Schrauder, M. and Navon, O., 1994, Hydrous and carbonatitic mantle fluids in fibrous diamonds from Jwaneng, Botswana: *Geochimica et Cosmochimica Acta*, v. 58, p. 761.

Shiryaev, A.A., Izraeli, E.S., Hauri, E.H., Zakharchenko, O.D. and Navon, O., 2005, Chemical, optical and isotopic investigation of fibrous diamonds from Brazil: *Russian Geology and Geophysics*, v. 46, p. 1185-1201.

Thompson, W.K., 1965, Infra-red spectroscopic studies of aqueous systems. Part 1. Molar extinction coefficients of water, deuterium oxide, deuterium hydrogen oxide, aqueous sodium chloride and carbon disulphide: *Transactions of the Faraday Society*, v. 61, p. 2635-2640.

Tomlinson, E.L., Jones, A.P. and Harris, J.W., 2006, Co-existing fluid and silicate inclusions in mantle diamond: *Earth and Planetary Science Letters*, v. 250, p. 581-595.



## Chapter 3. Effect of KCl on melting of the $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$ system at 5 GPa

### Introduction

Chlorine has been found as a major constituent in fluids trapped in fibrous diamonds (Navon et al., 1988; Izraeli et al., 2001; Tomlinson et al., 2006; Klein-BenDavid et al., 2006; Klein-BenDavid et al., 2007). It is mainly associated with potassium, forming a KCl-bearing brine with the  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  molar ratio of 0.05-0.68 (see Chapter 2). Syngenetic mineral inclusions in the fibrous diamonds from Koffiefontein in South Africa and from the Panda mine in Canada revealed that these fibrous diamonds formed at pressures of 4-6 GPa and temperatures of 900-1200 °C (Izraeli et al., 2004; Tomlinson et al., 2006). The high end of this temperature range is above the anticipated water-saturated peridotite solidus of Grove et al. (2006). Because these fibrous diamonds originate from the subcratonic lithospheric mantle (SCLM), the KCl-bearing brine in fluid inclusions of these diamonds indicates the existence of KCl in the SCLM at the depth relevant to the pressure 4-6 GPa. KCl in the brine may reduce the activity of  $\text{H}_2\text{O}$  sufficiently to stabilize a fluid, rather than a melt, under these conditions.

Chlorides have been shown to affect melting behavior at lower pressure. The melting of albite and dehydration of brucite in  $\text{H}_2\text{O}\text{-NaCl}$  fluids to 0.9 GPa and 700-900 °C shows that NaCl increases the solidus temperature of albite+fluid relative to the pure water case by reducing the activity of water in hydrous systems (Shmulovich and Graham, 1996). The  $\text{H}_2\text{O}$  activity measurement in concentrated NaCl, KCl and KCl-NaCl solutions to 1.5 GPa and 550-900 °C

shows that KCl reduces the activity of water more than NaCl does at the same condition (Aranovich and Newton, 1996; Aranovich and Newton, 1997; Aranovich and Newton, 1998), which means that KCl can increase melting temperature more than the same amount of NaCl. However, whether this effect exists at high pressures, corresponding to the depth of the SCLM, is unknown. This study conducts experiments to examine the effect of KCl on melting of the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system, a simple model for  $\text{H}_2\text{O}$ -bearing harzburgitic Earth's mantle, at 5 GPa (~160 km).

### **Experimental techniques**

Starting materials (Table 3-1) were prepared in four different compositions by mixing  $\text{Mg}(\text{OH})_2$  ( $\geq 95\%$  purity),  $\text{SiO}_2$  (99.999% purity) and KCl (99.999% purity) that were dried at  $120^\circ\text{C}$  for  $>12$  h. Material 1 is a mixture of  $\text{Mg}(\text{OH})_2$  and  $\text{SiO}_2$  in 3:2 molar proportion, generating the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system with  $\text{Mg}_2\text{SiO}_4/\text{MgSiO}_3$  molar ratio of 1 and water content of 18.3 wt%. Materials 2, 3 and 4, generated by mixing Material 1 (Mg:Si=1.5) with KCl in different proportions, have  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  molar ratios of 0.2, 0.4 and 0.6, respectively, which are consistent with those of chlorine-rich fluid inclusions in fibrous diamonds from the Diavik mine in Canada (Klein-BenDavid et al., 2007), ranging from 0.05 to 0.68 (see in Chapter 2). All starting materials were ground in an agate mortar under ethanol for 1 h and then dried at  $120^\circ\text{C}$  and stored in a desiccator. After this drying, the water content of Material 1 was checked by thermogravimetric analysis (Figure 3-1).

Approximately 3-5 mg of the appropriate starting material was loaded into an open-ended Pt capsule for each experiment. After drying at 120°C for >12 h, the capsule was welded shut quickly, within 3 min, and compressed gently into a cylindrical shape in a steel die. The compressed capsules had a diameter of 1.5 mm and a length ranging from 1.0 to 2.3 mm.

All experiments were run in the 2000 ton uniaxial split-sphere multi-anvil apparatus in the C.M. Scarfe Laboratory for Experimental Petrology at the University of Alberta. The 18mm sample assembly with high-T stepped graphite furnace, described by [Walter et al. \(1995\)](#), was used in all experiments. A semi-sintered MgO-5%Cr<sub>2</sub>O<sub>3</sub> octahedron was used as a pressure cell, and the sample assembly was inserted in a cylindrical hole in the octahedron. A W<sub>95</sub>Re<sub>5</sub>-W<sub>74</sub>Re<sub>26</sub> thermocouple was inserted axially into the cylindrical sample assembly, and it was in contact with the Pt capsule. Each sample assembly was dried at 120°C for >10 h, then fired at 1000°C in (2% H<sub>2</sub>) N<sub>2</sub> gas flow for 1 h. The sample capsule was removed from the assembly for this latter step. After firing, the sample capsule was re-inserted in the assembly, and then the assembly was kept in 120°C until use. The dried sample assembly was placed in the octahedral cavity formed by triangular truncations of 11 mm of eight 32.5 mm edge-length WC-cubes. These cubes formed a big cube placed in the cubic cavity formed by six anvils of the uniaxial press. Calibrations of pressure and thermal gradient have been described in detail by [Walter et al. \(1995\)](#). The pressure calibration is accurate to within ±0.5 GPa, and the thermal gradient of the high-T stepped graphite furnace is +25°C mm<sup>-1</sup>, away from the furnace midline ([Walter et al.](#),

1995). Experiments were brought to 5 GPa, and then heated at  $\sim 70^\circ\text{C s}^{-1}$  to 1100-1700°C, and maintained at  $\pm 5^\circ\text{C}$ . No pressure correction to the emf of the thermocouple was applied. Durations for runs at  $\leq 1600^\circ\text{C}$  were 8-12 h, and those for runs at  $> 1600^\circ\text{C}$  were 2 h. Experiments were quenched by turning off the power to the furnace, resulting in a temperature drop to  $< 300^\circ\text{C}$  in 2-5 s, and then decompressed over 3 h. After the runs, the capsules were weighed and then punctured with a razor blade. The punctured capsules were dried at  $120^\circ\text{C}$  for  $> 12$  h and then reweighed. If there was no weight loss after the drying, the capsule was presumed to have leaked, and the experiment was repeated. The punctured capsules in successful experiments were boiled in distilled water for 20 min twice to remove water-soluble material, such as chloride, because the presence of this material caused plucking of run products during polishing. The boiled capsules were dried and mounted in an epoxy plug using Petropoxy 154. After mounting, sample mounts were ground, polished with progressively finer suspensions of alumina, ultrasonically washed and then dried at  $120^\circ\text{C}$  for  $> 12$  h before being coated with carbon.

Table 3-1. Starting materials.

System	Mg <sub>2</sub> SiO <sub>4</sub> -MgSiO <sub>3</sub> -H <sub>2</sub> O		Mg <sub>2</sub> SiO <sub>4</sub> -MgSiO <sub>3</sub> -KCl- H <sub>2</sub> O	
Starting Material	1	2	3	4
<i>(wt%)</i>				
SiO <sub>2</sub>	40.72	34.23	27.05	19.06
Mg(OH) <sub>2</sub>	59.28	49.84	39.38	27.74
KCl	0	15.93	33.56	53.20
<i>Calculated molar ratio</i>				
Mg/Si	1.5	1.5	1.5	1.5
Cl/(Cl+H <sub>2</sub> O)	0	0.2	0.4	0.6
<i>Calculated water content (wt%)</i>				
H <sub>2</sub> O	18.31	15.40	12.17	8.57
<i>Calculated KCl content in generated aqueous solutions (wt%)</i>				
KCl	0	50.85	73.40	86.13

Material 1 represents the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O system. Material 2, 3 and 4 represent the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl- H<sub>2</sub>O systems with the Cl/(Cl+H<sub>2</sub>O) molar ratio of 0.2, 0.4 and 0.6, respectively.

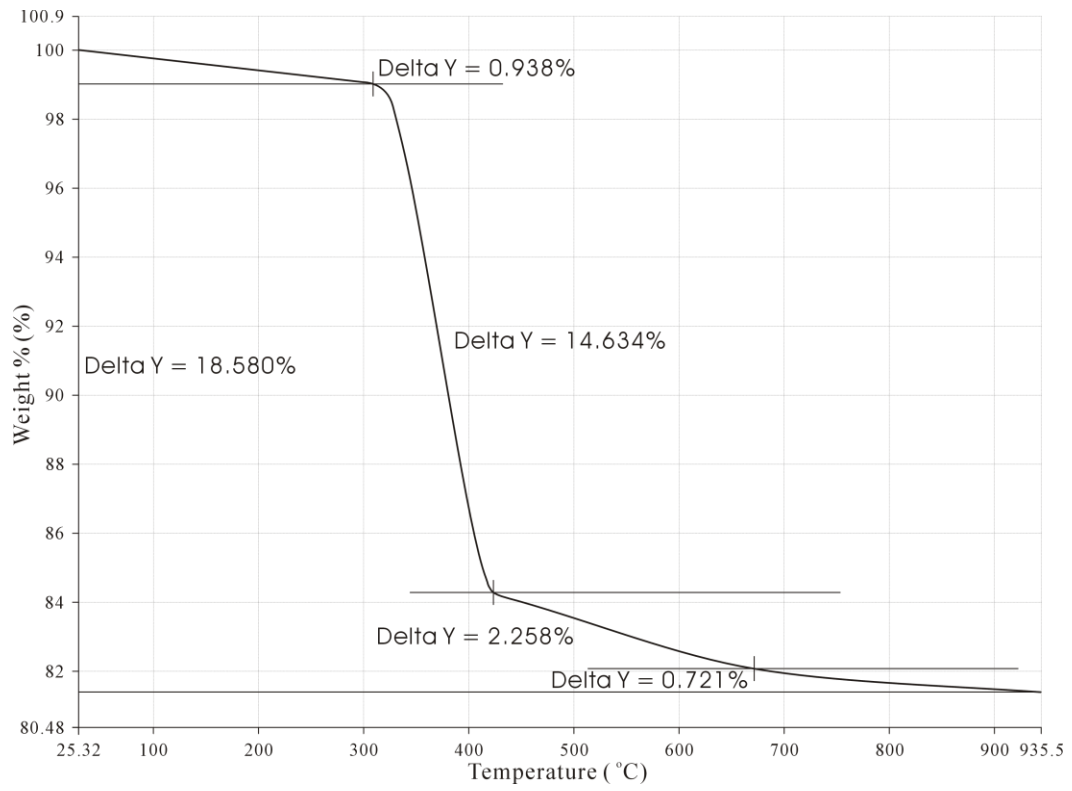


Figure 3-1. Thermogravimetric analysis of starting material of  $Mg_2SiO_4$ - $MgSiO_3$ - $H_2O$  system.

The total weight loss (Delta Y) of 18.58 wt% is close to the calculated water content of 18.31 wt% for the Material 1 in [Table 3-1](#).

Examination and analysis of the run products were carried out with the JEOL 8900R electron microprobe at the University of Alberta. The textural relationships between coexisting phases were examined using backscattered electron (BSE) imaging at least twice by repolishing the mounts. To examine the compositional changes at the solidus, the Ca contents of forsterite and enstatite were analyzed using 20 kV accelerating voltage, and the signal was detected by wavelength dispersive spectrometers with the PETH crystal. Because the Ca contents of forsterite and enstatite were much lower than that of diopside standard, the sample current for the diopside standard was 15 nA, and that for forsterite and enstatite samples was 150 nA. Counting time was 120 s on peaks and 60 s on background at each side. Si and Mg were also analyzed for the ZAF correction. In analyzing Ca, the detection limit and the standard deviation reported by microprobe were 5 ppm and ~10%, respectively. [Potts \(1986\)](#) pointed out that the reported detection limit ( $2 \sigma_{\text{background}}$ ) is an over-optimistic estimate of the lower limit of detection and the limit of determination for quantitative measurement is  $6 \sigma_{\text{background}}$ . Therefore, the limit of determination ( $6 \sigma_{\text{background}}$ ) used in this thesis was calculated from the reported counts and counting time using the formula of [Scott and Love \(1983\)](#). The calculated limit of determination ( $6 \sigma_{\text{background}}$ ) in analyzing Ca was 15 ppm. The EPMA analysis of run products is listed in the Appendix.

## **Results**

Experimental conditions and results are summarized in [Table 3-2](#). Because the fluids and melts are unquenchable, the texture of quenched phases cannot reveal the relationships, such as immiscibility, between the water/brine and

silicate melts under experimental conditions. “Vapor” and “Liquid” are used to describe subsolidus fluids and suprasolidus melts, respectively. The solidus is determined by examining the amount of quenched phases and the change of Ca content in the primary phases. Because of the thermal gradient in the capsule, subsolidus vapor and suprasolidus liquid can appear in one capsule. According to the study of [Walter et al., 1995](#), the thermal gradient for the high-T stepped graphite furnace is  $+25^{\circ}\text{C mm}^{-1}$  away from the furnace midline. Based on this thermal gradient, the temperature at the center of the capsule ( $T'$ ) was corrected from the run temperature which was measured by a thermocouple at the top, and the temperature range ( $R$ ) in the capsule was calculated from the inner height of the capsule.

Textures of the run products are shown in [Figure 3-2](#). For all systems, phases interpreted to be primary phases are equant to sub-equant and typically have large grain size, whereas quenched phases are subhedral to euhedral and typically display elongate prismatic habit. The primary phases in the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system were typically smaller than those in  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  systems, and the quenched phases in the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system are generally larger than those in  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  systems. In the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system, primary forsterite and enstatite coexist in all runs; the quenched phases have compositions similar to enstatite and forsterite, and their amount increases abruptly at the run temperature of  $1250^{\circ}\text{C}$ . In  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  systems with  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  ratio of 0.2, 0.4 and 0.6, the primary phases changes



from coexisting forsterite and enstatite at lower temperatures to only forsterite at higher temperatures; the quenched phases contain Si, Mg, K and Cl, and the amount of quenched phases increases abruptly at the run temperatures of 1450°C, 1600 °C and 1650°C, respectively. Quenched liquid (qL) and quenched liquid (qL) are very similar in texture, but they can be differentiated from the amount of quenched phases (Figure 3-2). The quenched phases in small amounts are determined as qV and those in large amounts are determined as qL. The quenched phases in moderate amounts are difficult to distinguish as either quenched vapor or quenched liquid, so they are noted as quenched vapor and liquid (qV/qL). The coexistence of both vapor and liquid in one capsule is explained to be due to the existence of thermal gradient in the capsule.

Table 3-2. Experimental conditions and results.

Run #	molar Cl/(Cl+H <sub>2</sub> O)	T(°C)	Duration (h)	Results	T correction (°C)		Ca in Ol (ppm)			Ca in Opx (ppm)		
					T'	R	N	Av	SD	N	Av	SD
3722	0	1100	12	En+Fo+V	1080	14	27	30	11	26	191	11
3710	0	1200	8	En+Fo+V	1178	12	20	29	7	15	220	29
3711	0	1250	12	En+Fo+V/L	1229	12	10	49	8	7	191	11
3709	0	1300	8	En+Fo+L	1285	11	16	29	7	14	128	6
3755	0	1350	12	En+Fo+L	1324	16	7	26	8	10	143	19
3765	0	1400	12	En+Fo+L	1379	9	15	74	43	18	160	26
3762	0	1450	12	En+Fo+L	1428	11	12	71	37	19	256	97
3771	0	1450	12	En+Fo+L	1425	11	48	26	6			
3770	0	1500	12	En+Fo+L	1478	13	14	31	5	37	193	140
3715	0.2	1300	12	En+Fo+V	1277	18	11	28	6	16	234	9
3753	0.2	1300	12	En+Fo+V	1276	11	19	30	9	9	223	16
3721	0.2	1400	12	En+Fo+V/L	1377	18	48	26	6	33	195	11
3750	0.2	1450	12	Fo+L	1428	12	19	47	12			
3727	0.2	1500	12	Fo+L	1480	14	56	20	6			
3737	0.2	1600	12	Fo+L	1580	10	11	25	4			
3761	0.2	1650	2	L	1629	10						
3748	0.4	1400	12	En+Fo+V	1376	14	11	31	6	14	266	27
3728	0.4	1500	12	En+Fo+V	1476	17	28	37	4	11	219	7
3759	0.4	1550	12	En+Fo+V/L	1529	11	19	58	18	14	217	18
3744	0.4	1600	12	Fo+L	1584	8	26	31	11			
3763	0.4	1650	2	Fo+L	1627	15	14	83	16			
3747	0.4	1700	2	L	1676	13						
3754	0.6	1500	12	En+Fo+V	1475	12	6	30	5	9	261	20
3739	0.6	1600	10	En+Fo+V/L	1581	10	17	40	7	9	250	60
3772	0.6	1650	2	Fo+L	1624	12	57	26	5			
3740	0.6	1700	2	Fo+L	1682	10	15	43	6			
3758	0.6	1700	2	Fo+L	1677	11	16	40	7			

Abbreviations: *T* temperature, *N* number of analysis, *R* temperature range in capsule, *Av* average, *SD* standard deviation, *En*

enstatite, *Fo* forsterite, *V* vapor, *L* liquid.

T correction is done by  $T' = T - \frac{1}{2}gh$  and  $R = \frac{1}{2}gl$ , where *g* is the thermal gradient in graphite furnace, *h* is the outer height of capsule, and *l* is the inner height of capsule.

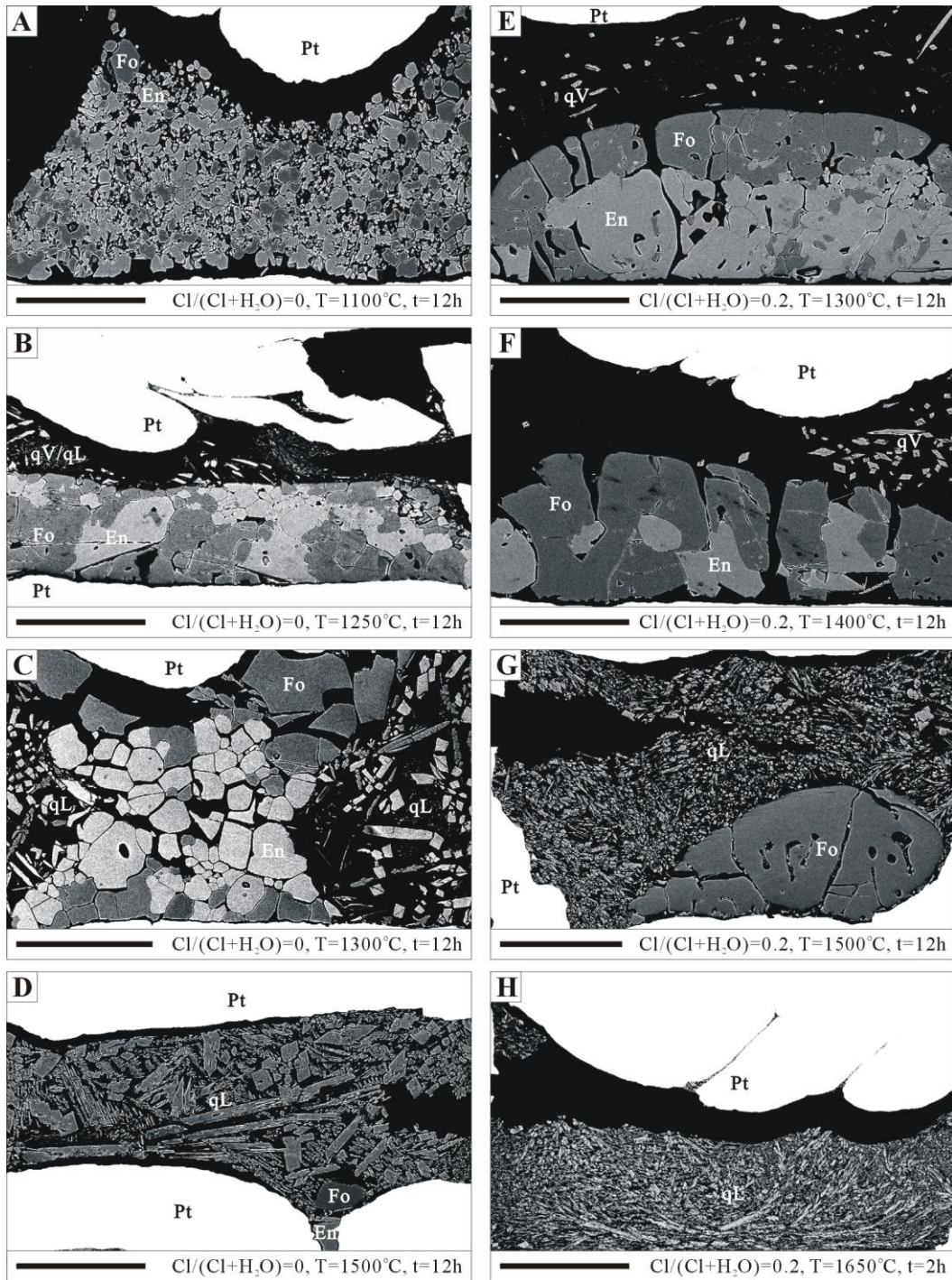


Figure 3-2. Backscattered electron images of run products.

A-D Run products of  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system, primary phases are forsterite and enstatite. E-F Run products of the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  system with  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  ratio of 0.2, subsolidus primary phases are forsterite and enstatite, suprasolidus primary phases are forsterite, and the enstatite/forsterite modal ratio decreases as the temperature increases.

Abbreviations: *Pt*, platinum capsule, *Fo*, forsterite, *En*, enstatite, *qV*, quenched vapor, *qL*, quenched liquid. Black scale bars at left bottom represent 400  $\mu\text{m}$ .

For all systems, primary forsterite has Ca content of 20-83 ppm, and primary enstatite has Ca content of 128-226 ppm. The calcium in these experiments was probably derived from the reagent-grade  $\text{Mg}(\text{OH})_2$  starting material. In the  $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{H}_2\text{O}$  system, forsterite has an obvious jump in Ca content around the temperature of  $1229^\circ\text{C}$ , and enstatite has a big drop in Ca content at the same temperature (Figure 3-3). In  $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{KCl}$ - $\text{H}_2\text{O}$  systems with  $\text{Cl}/(\text{Cl}+\text{H}_2\text{O})$  ratio of 0.2, 0.4 and 0.6, forsterite has jumps in Ca content at temperatures of  $1428^\circ\text{C}$ ,  $1529^\circ\text{C}$  and  $1581^\circ\text{C}$ , respectively (Figure 3-4).

Abrupt increases in the amount of quenched phases and abrupt changes in Ca content of primary phases are interpreted to indicate the solidus, because transition from vapor to liquid will cause abrupt changes in the solubility of silicate in fluid and partitioning of trace elements between solid and vapor is different from that between solid and liquid. In the experiments on melting of peridotite- $\text{CO}_2$  system by Dasgupta and Hirschman (2006), the sodium content of the clinopyroxene dropped abruptly at the solidus, which suggests that the drop in content of minor element in pyroxene can indicate the position of solidus. In the  $\text{Mg}_2\text{SiO}_4$ - $\text{MgSiO}_3$ - $\text{H}_2\text{O}$  system (Figure 3-3), the drop of Ca content in enstatite indicates the temperature of the solidus, and the jump of Ca content in olivine also happened at the same temperature. That forsterite and enstatite have different behaviors in changing Ca content at the temperature of solidus may be due to their different crystal structures, and both the drop of Ca content in enstatite and the jump of Ca content in olivine can be used in determine the solidus. Both the textural relationships and the Ca peaks give the same result for the

Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O system and the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O systems with Cl/(Cl+H<sub>2</sub>O) molar ratio of 0.2 and 0.4. Solidus of the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O system with Cl/(Cl+H<sub>2</sub>O) molar ratio of 0.6 cannot be determined from textural relationships because the quenched material collapsed when large amount of water-soluble material was removed, and it is only determined by the Ca peak of forsterite.

The temperature of the solidus changing with Cl/(Cl+H<sub>2</sub>O) molar ratio is shown in [Figure 3-5](#). Above the solidus, both forsterite and enstatite coexist with melt in the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O system, which is consistent with a eutectic melting relationship of forsterite and enstatite at 5 GPa. In contrast, only forsterite coexists with melt in Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O systems, which may reveal the incongruent melting of enstatite.

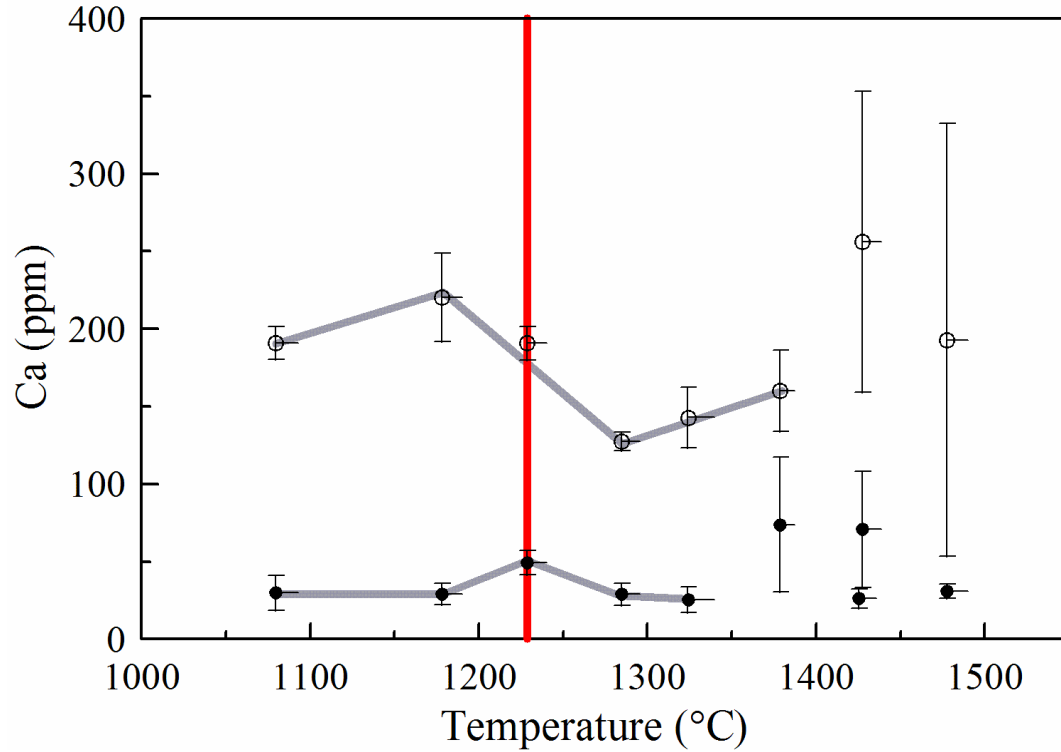


Figure 3-3. Ca contents of forsterite (filled circle) and enstatite (open circle) in the  $Mg_2SiO_4$ - $MgSiO_3$ - $H_2O$  system as a function of temperature.

Error bars show temperature uncertainties in horizontal and 2 standard deviation of Ca content in vertical. The vertical line shows the inferred position of solidus. The big error bars for the Ca contents of forsterite and enstatite from high temperature experiments are because of the hardness in distinguishing primary phases from quenched phases.

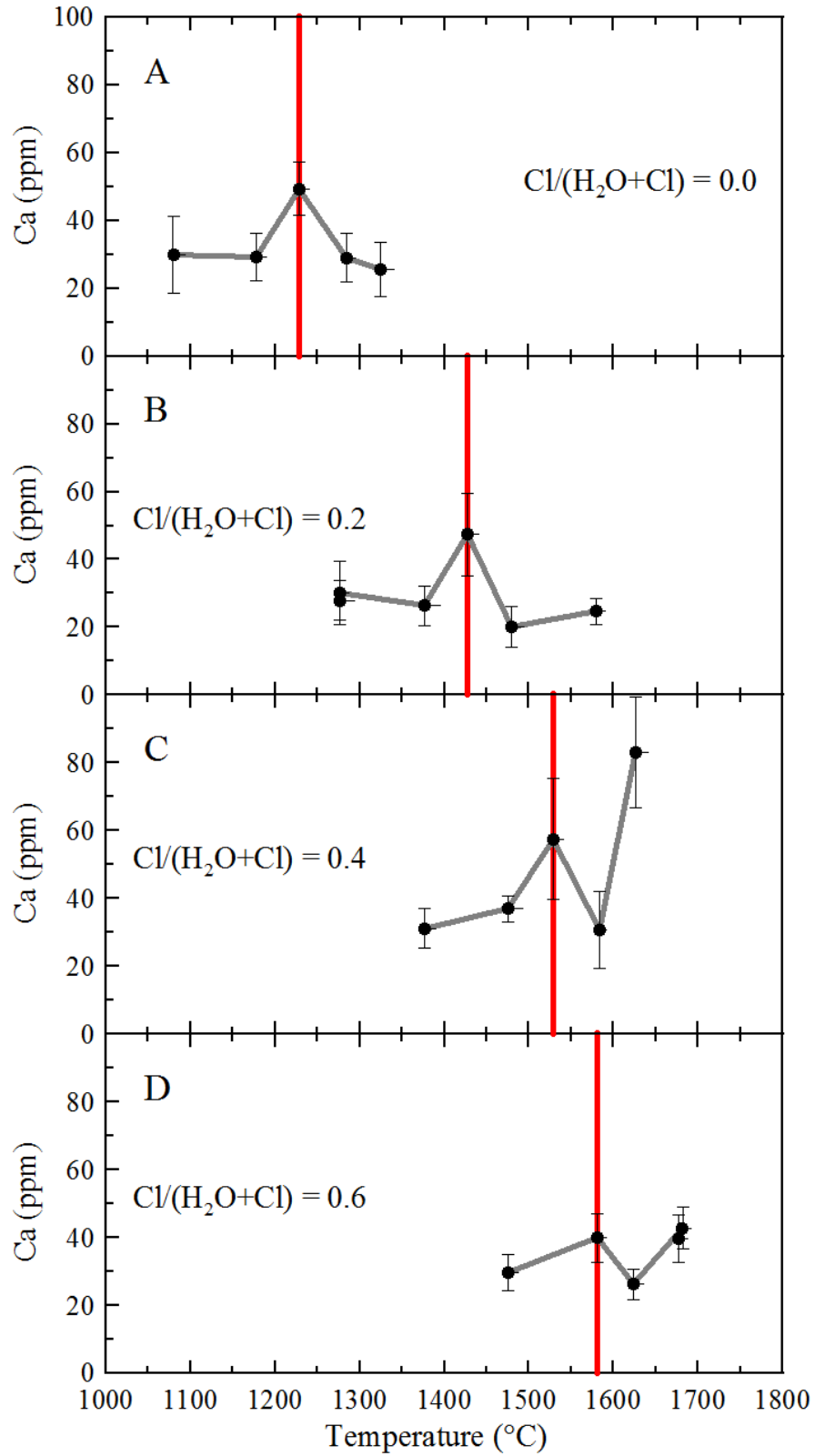


Figure 3-4. Ca contents of forsterite in the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O system and Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O systems.

Error bars show 2 standard deviation of Ca content. Vertical lines show positions of solids.

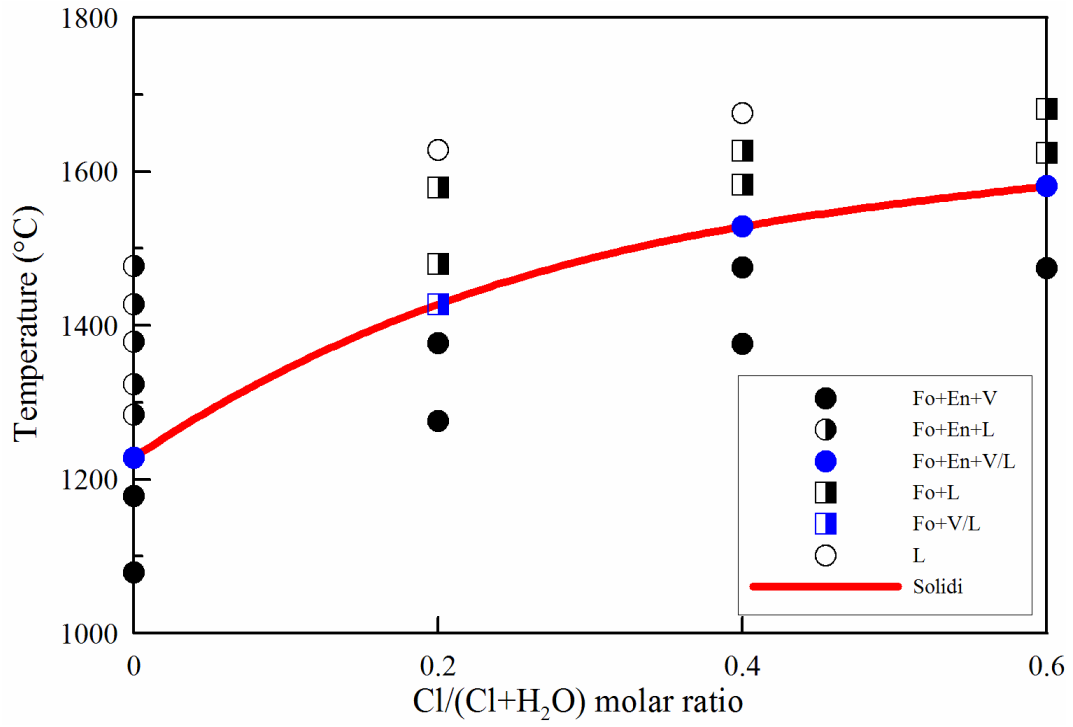


Figure 3-5. Phase relations of the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O system and Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O systems at 5 GPa.

The curve shows the relation between solidi and Cl concentrations.



## Discussion

Chloride can significantly reduce H<sub>2</sub>O activity in aqueous solutions, which was evidenced by measuring H<sub>2</sub>O activity in NaCl, KCl and NaCl-KCl solutions by the brucite-periclase equilibrium (Aranovich and Newton, 1996; Shmulovich and Graham, 1996; Aranovich and Newton, 1997). The experiments by Aranovich and Newton (1997) revealed that the H<sub>2</sub>O activity is nearly coincident with the square of its mole fraction in KCl solution at the pressure range from 4 to 15 kbar. Melting experiments of albite in NaCl solution by Shmulovich and Graham (1996) showed that the temperature of the vapor-saturated solidus of albite in a solution with NaCl mole fraction of 0.2 is ~100 °C higher than that in pure water at 9 kbar. Experiments on Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O systems at 5 GPa in this study (Figure 3-5) show that temperatures of solidi of the systems with the Cl/(Cl+H<sub>2</sub>O) molar ratio of 0.2, 0.4 and 0.6 are, respectively, ~200 °C, ~300 °C and ~350 °C higher than that with pure water, which indicates that chloride can also significantly reduce H<sub>2</sub>O activity in aqueous solution at pressure relevant to depth of ~160km in the Earth's mantle. Therefore, chlorine in the SCLM, evidenced by fluid inclusions in fibrous diamonds, can inhibit melting of the mantle by reducing the H<sub>2</sub>O activity in fluids, and the KCl-bearing fluids can be a robust agent of mantle metasomatism.

The absence of primary enstatite under suprasolidus conditions in Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O systems may be due to the incongruent melting of enstatite. The incongruent melting of enstatite, discovered by Bowen and Andersen (1914) at 1 atm, has a significant bearing on the generation of

silica-saturated magmas from silica-undersaturated mantle rocks by partial melting or from silicate-undersaturated magmas by fractional crystallization (Kushiro et al., 1968), and thus potentially plays an important role in the Mg/Si differentiation of the planet. The incongruent melting of enstatite disappears at pressures higher than ~0.5 GPa under anhydrous conditions (Boyd et al., 1964) and at some pressure between 3 and 5.5 GPa under hydrous conditions (Kushiro et al., 1968; Inoue, 1994). The eutectic melting inferred in this study in the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system at 5 GPa reveals that the incongruent melting of enstatite disappears at pressures less than 5 GPa. If the absence of primary enstatite under suprasolidus conditions in  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  systems is due to the incongruent melting of enstatite, it will mean that the chlorine-rich brine can make the incongruent melting of enstatite happen at the pressure of 5 GPa or more.

## **Conclusion**

KCl significantly elevates the solidus temperature of the  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-H}_2\text{O}$  system at 5 GPa by reducing the activity of  $\text{H}_2\text{O}$  in the fluid. This indicates that chlorine can prevent melting in the SCLM and the KCl-bearing fluids can be a robust agent of mantle metasomatism. In  $\text{Mg}_2\text{SiO}_4\text{-MgSiO}_3\text{-KCl-H}_2\text{O}$  systems, the absence of primary enstatite under suprasolidus conditions is probably due to the incongruent melting of enstatite, which needs to be defined by further experiments on the melting of the  $\text{MgSiO}_3\text{-KCl-H}_2\text{O}$  system.

## **References**

- Aranovich, L.Y. and Newton, R.C., 1996, H<sub>2</sub>O activity in concentrated NaCl solutions at high pressures and temperatures measured by the brucite-periclase equilibrium: *Contributions to Mineralogy and Petrology*, v. 125, p. 200-212.
- Aranovich, L.Y. and Newton, R.C., 1997, H<sub>2</sub>O activity in concentrated KCl and KCl-NaCl solutions at high temperatures and pressures measured by the brucite-periclase equilibrium: *Contributions to Mineralogy and Petrology*, v. 127, p. 261-271.
- Aranovich, L.Y. and Newton, R.C., 1998, Reversed determination of the reaction: phlogopite + quartz = enstatite + potassium feldspar + H<sub>2</sub>O in the ranges 750-875 °C and 2-12 kbar at low H<sub>2</sub>O activity with concentrated KCl solutions: *American Mineralogist*, v. 83, p. 193-204.
- Bowen, N.L. and Andersen, O., 1914, The binary system MgO-SiO<sub>2</sub>: The *American Journal of Science*, v. 37, p.487-500.
- Boyd, F.R., England, J.L. and Davis, B.T.C., 1964, Effects of pressure on the melting and polymorphism of enstatite, MgSiO<sub>3</sub>: *Journal of Geophysical Research*, v. 69, p. 2101-2109.
- Dasgupta, R. and Hirschman, M.M., 2006, Melting in the Earth's deep upper mantle caused by carbon dioxide: *Nature*, v. 440, p. 659-662.
- Grove, T.L., Chatterjee, N., Parman, S.W. and Médard, E., 2006, The influence of H<sub>2</sub>O on mantle wedge melting: *Earth and Planetary Science Letters*, v. 249, p. 74-89.
- Inoue, T., 1994, Effect of water on melting phase relations and melt composition in the system Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O up to 15 GPa: *Physics of the Earth and Planetary Interiors*, v. 85, p. 237-263.
- Izraeli, E.S., Harris, J.W. and Navon, O., 2001, Brine inclusions in diamonds: a new upper mantle fluid: *Earth and Planetary Science Letters*, v. 187, p. 323-332.
- Izraeli, E.S., Harris, J.W. and Navon, O., 2004, Fluid and mineral inclusions in cloudy diamonds from Koffiefontein, South Africa: *Geochimica et Cosmochimica Acta*, v. 68, p. 2561-2575.
- Klein-BenDavid, O., Izraeli, E.S., Hauri, E.H. and Navon, O., 2007, Fluid inclusions in diamonds from the Diavik Mine, Canada and the evolution of diamond-forming fluids: *Geochimica et Cosmochimica Acta*, v. 71, p. 723-744.
- Klein-BenDavid, O., Richard, W. and Navon, O., 2006, TEM imaging and analysis of microinclusions in diamonds; a close look at diamond-growing fluids: *American Mineralogist*, v. 91, p. 353-365.

- Kushiro, I., Yoder, H.S. and Nishikawa, M., 1968, Effect of water on the melting of enstatite: *Bulletin of the Geological Society of America*, v. 79, p. 1685-1692.
- Navon, O., Hutcheon, I.D., Rossman, G.R. and Wasserburg, G.J., 1988, Mantle-derived fluids in diamond micro-inclusions: *Nature*, v. 335, p. 784-789.
- Potts, P.J., 1986, *A handbook of silicate rock analysis*, Chapman and Hall, New York, NY, p. 17.
- Scott, V.D, Love, G., 1983, *Quantitative Electron-Probe Microanalysis*, Ellis Horwood Limited, p. 137-142.
- Shmulovich, K.I. and Graham, C.M., 1996, Melting of albite and dehydration of brucite in H<sub>2</sub>O-NaCl fluids to 9 kbars and 700-900°C: implications for partial melting and water activities during high pressure metamorphism: *Contributions to Mineralogy and Petrology*, v. 124, p. 370-382.
- Tomlinson, E.L., Jones, A.P. and Harris, J.W., 2006, Co-existing fluid and silicate inclusions in mantle diamond: *Earth and Planetary Science Letters*, v. 250, p. 581-595.
- Walter, M.J., Thibault, Y., Wei, K. and Luth, R.W., 1995, Characterizing experimental pressure and temperature conditions in multi-anvil apparatus: *Canadian Journal of Physics*, v. 73, p. 273-286.

## Chapter 4. General Discussion and Conclusions

The presence of chlorine in the subcratonic lithospheric mantle (SCLM) has been evaluated by compiling the compositional data of fluid inclusions in fibrous diamonds. It mainly associates with potassium, dissolving in water and forming a KCl-bearing brine with the Cl/(Cl+H<sub>2</sub>O) molar ratio of 0.05-0.68 mol%. To examine the effect of such a KCl-bearing brine on the melting behavior of the Earth's mantle, we conducted experiments in the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O and Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O systems at 5 GPa and 1100-1700 °C. In the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O system, the temperature of the fluid-saturated solidus is ~1230 °C, and both forsterite and enstatite coexist with the liquid under suprasolidus conditions. In the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O systems with molar Cl/(Cl+H<sub>2</sub>O) ratios of 0.2, 0.4 and 0.6, the temperatures of the fluid-saturated solidus are ~1430 °C, ~1530 °C and ~1580 °C, respectively, and only forsterite coexists with liquid under suprasolidus conditions. The increase in the temperature of the solidus demonstrates the significant effect of KCl on elevating the solidus of the Mg<sub>2</sub>SiO<sub>4</sub>-MgSiO<sub>3</sub>-H<sub>2</sub>O system by reducing the activity of H<sub>2</sub>O in the fluid. The change in the melting residues indicates that the incongruent melting of enstatite (enstatite = forsterite + silica-rich melt) could happen at pressures over 5 GPa in KCl-bearing systems, which needs to be verified by experiments on the MgSiO<sub>3</sub>-KCl-H<sub>2</sub>O system in future work.

If KCl is present in the Earth's mantle, it will elevate the melting temperature of the H<sub>2</sub>O-bearing mantle. In subduction zones, brine inclusions, with ~50 wt% NaCl and KCl, have been found in omphacite in eclogite ([Scambelluri and](#)

Philippot, 2001), which indicates the existence of chloride-rich brines in subducting slabs. Release of these brines, from subducting slabs to mantle wedges, will inhibit melting of the mantle wedges. In cratonic areas, fluid inclusions, with ~47 wt% KCl and NaCl, have been found in fibrous diamonds (Klein-BenDavid *et al.*, 2007), which indicates the existence of chloride-rich fluids in the diamond stability field of the cratonic mantle. These fluids will inhibit melting at the H<sub>2</sub>O-saturated solidus in the mantle, and the KCl-bearing brine can be a robust agent for mantle metasomatism at temperatures greater than that of the H<sub>2</sub>O-saturated solidus. Accordingly, the brine-bearing fibrous diamonds are inferred to form from fluid rather than melt in the SCLM.

## References

- Klein-BenDavid, O., Izraeli, E.S., Hauri, E.H. and Navon, O., 2007, Fluid inclusions in diamonds from the Diavik Mine, Canada and the evolution of diamond-forming fluids: *Geochimica et Cosmochimica Acta*, v. 71, p. 723-744.
- Scambelluri, M. and Philippot, P., 2001, Deep fluids in subduction zones: *Lithos*, v. 55, p. 213-227.

### Appendix. EPMA analysis of run products

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3709	LC3-1	Fo	56.027	40.022	0.0038	96.053
3709	LC3-2	Fo	54.814	41.515	0.0033	96.332
3709	LC3-3	En	39.339	59.713	0.0190	99.071
3709	LC3-4	En	38.564	59.424	0.0164	98.005
3709	LC3-5	En	37.989	58.777	0.0179	96.784
3709	LC3-6	En	39.330	57.759	0.0180	97.107
3709	LC3-7	En	38.469	57.237	0.0175	95.724
3709	LC3-8	En	38.005	58.020	0.0165	96.042
3709	LC3-9	Fo	52.589	39.617	0.0026	92.209
3709	LC3-10	Fo	49.574	39.690	0.0032	89.267
3709	LC3-11	Fo	49.974	41.070	0.0030	91.047
3709	LC3-12	En	37.197	58.304	0.0188	95.520
3709	LC3-13	En	37.139	59.185	0.0183	96.343
3709	LC3-14	En	39.216	60.001	0.0163	99.234
3709	LC3-15	En	39.937	61.031	0.0168	100.985
3709	LC3-16	En	39.969	59.925	0.0162	99.911
3709	LC3-17	En	40.183	60.432	0.0185	100.634
3709	LC3-18	En	40.173	60.448	0.0179	100.639
3709	LC3-19	En	39.443	59.216	0.0171	98.676
3709	LC3-20	Fo	55.249	42.112	0.0031	97.364
3709	LC3-21	Fo	47.822	37.529	0.0031	85.354
3709	LC3-22	Fo	54.906	43.307	0.0042	98.217
3709	LC3-23	Fo	54.451	41.216	0.0052	95.672
3709	LC3-24	Fo	53.831	42.644	0.0039	96.479
3709	LC3-25	Fo	53.932	41.960	0.0043	95.896
3709	LC3-26	Fo	52.279	40.467	0.0050	92.751
3709	LC3-27	Fo	54.593	41.903	0.0053	96.501
3709	LC3-28	Fo	52.719	43.015	0.0059	95.740
3709	LC3-29	Fo	54.627	41.846	0.0047	96.478
3709	LC3-30	Fo	53.980	41.531	0.0050	95.516
3709	LC3-31	q	39.091	57.950	0.0431	97.085
3709	LC3-32	q	40.121	60.723	0.0453	100.890
3709	LC3-33	q	39.519	58.801	0.0439	98.365
3709	LC3-34	q	39.512	58.800	0.0493	98.362
3710	LC4-1	Fo	53.724	40.576	0.0027	94.303
3710	LC4-2	Fo	50.015	40.339	0.0045	90.359
3710	LC4-3	Fo	53.463	41.314	0.0060	94.783
3710	LC4-4	En	38.120	60.024	0.0272	98.172
3710	LC4-5	Fo	50.316	42.232	0.0048	92.553
3710	LC4-6	Fo	54.015	42.658	0.0031	96.676
3710	LC4-7	En	36.849	60.911	0.0263	97.787
3710	LC4-8	En	36.969	60.723	0.0278	97.721
3710	LC4-9	Fo	53.967	43.200	0.0040	97.171
3710	LC4-10	Fo	51.588	43.219	0.0041	94.811

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3710	LC4-11	En	36.163	60.589	0.0303	96.783
3710	LC4-12	Fo	54.304	42.670	0.0034	96.978
3710	LC4-13	En	37.255	61.269	0.0311	98.556
3710	LC4-14	En	36.851	61.162	0.0390	98.053
3710	LC4-15	Fo	53.283	43.330	0.0039	96.617
3710	LC4-16	Fo	53.908	42.507	0.0039	96.419
3710	LC4-17	En	38.122	59.661	0.0317	97.816
3710	LC4-18	Fo	60.803	43.467	0.0042	104.274
3710	LC4-19	En	40.225	60.182	0.0287	100.436
3710	LC4-20	En	39.100	58.884	0.0281	98.013
3710	LC4-21	En	39.650	59.839	0.0283	99.518
3710	LC4-22	Fo	56.599	41.916	0.0044	98.519
3710	LC4-23	En	39.624	59.285	0.0280	98.938
3710	LC4-24	Fo	56.504	42.257	0.0033	98.764
3710	LC4-25	Fo	56.291	41.728	0.0037	98.023
3710	LC4-26	Fo	56.040	41.317	0.0035	97.361
3710	LC4-27	Fo	55.966	41.777	0.0021	97.745
3710	LC4-28	Fo	55.405	41.279	0.0044	96.688
3710	LC4-29	En	41.037	55.965	0.0377	97.041
3710	LC4-30	En	40.239	57.684	0.0263	97.950
3710	LC4-31	Fo	58.579	40.356	0.0038	98.939
3710	LC4-32	Fo	58.844	40.113	0.0063	98.963
3710	LC4-33	En	40.239	58.608	0.0275	98.875
3710	LC4-34	Fo	57.917	41.642	0.0047	99.564
3710	LC4-35	En	41.600	57.815	0.0326	99.448
3710	LC4-36	q	42.048	58.456	0.0892	100.595
3710	LC4-37	q	40.597	57.402	0.1065	98.108
3710	LC4-38	q	40.499	59.781	0.0834	100.365
3711	LC5-1	Fo	60.803	40.439	0.0053	101.247
3711	LC5-2	En	42.377	60.767	0.0273	103.172
3711	LC5-3	En	40.993	61.267	0.0272	102.288
3711	LC5-4	Fo	58.254	41.232	0.0061	99.492
3711	LC5-5	Fo	57.690	42.293	0.0067	99.990
3711	LC5-6	En	39.067	57.574	0.0792	96.722
3711	LC5-7	En	39.839	60.377	0.0269	100.244
3711	LC5-8	En	39.449	59.583	0.0731	99.107
3711	LC5-9	Fo	54.003	40.050	0.0069	94.060
3711	LC5-10	Fo	56.331	41.856	0.0090	98.196
3711	LC5-11	Fo	59.100	44.480	0.0065	103.587
3711	LC5-12	En	41.317	62.502	0.0257	103.845
3711	LC5-13	En	56.406	42.404	0.0070	98.817
3711	LC5-14	En	41.053	61.398	0.0263	102.478
3711	LC5-15	En	39.992	59.799	0.0256	99.817
3711	LC5-16	En	39.735	58.141	0.0342	97.911
3711	LC5-17	En	40.180	60.995	0.0236	101.199
3711	LC5-18	Fo	53.836	39.800	0.0079	93.644



Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3711	LC5-19	Fo	58.654	44.506	0.0060	103.166
3711	LC5-20	Fo	57.277	44.540	0.0069	101.824
3711	LC5-21	Fo	50.662	40.957	0.0101	91.629
3711	LC5-22	En	39.036	59.428	0.0298	98.495
3711	LC5-23	Fo	52.753	38.998	0.0120	91.763
3711	LC5-24	En	36.946	58.592	0.0324	95.571
3711	LC5-25	Fo	52.786	40.491	0.0343	93.312
3711	LC5-26	En	34.417	52.401	0.0324	86.851
3711	LC5-27	q	37.931	61.210	0.2651	99.413
3711	LC5-28	q	41.120	57.988	0.3266	99.442
3711	LC5-29	q	40.980	56.517	0.5652	98.075
3711	LC5-30	q	40.940	55.553	0.5518	97.058
3715	LC7-1	En	39.536	55.522	0.0311	95.090
3715	LC7-2	Fo	59.846	40.587	0.0034	100.436
3715	LC7-3	Fo	61.756	42.280	0.0033	104.039
3715	LC7-4	Fo	60.022	42.267	0.0037	102.293
3715	LC7-5	Fo	60.257	42.675	0.0027	102.935
3715	LC7-6	Fo	58.546	42.735	0.0043	101.285
3715	LC7-7	En	40.341	58.614	0.0310	98.987
3715	LC7-8	En	40.286	59.470	0.0334	99.790
3715	LC7-9	En	40.338	59.860	0.0310	100.230
3715	LC7-10	En	39.805	59.294	0.0334	99.133
3715	LC7-11	En	39.852	58.885	0.0324	98.770
3715	LC7-12	En	39.853	57.344	0.0313	97.229
3715	LC7-13	Fo	60.201	42.707	0.0028	102.911
3715	LC7-14	En	41.312	60.003	0.0318	101.348
3715	LC7-15	En	41.008	59.735	0.0318	100.776
3715	LC7-16	En	39.810	57.613	0.0319	97.456
3715	LC7-17	En	40.385	58.329	0.0299	98.745
3715	LC7-18	En	39.657	57.942	0.0309	97.631
3715	LC7-19	En	38.676	56.932	0.0308	95.640
3715	LC7-20	En	39.449	60.229	0.0301	99.709
3715	LC7-21	En	38.149	57.395	0.0535	95.599
3715	LC7-22	En	39.145	59.791	0.0331	98.970
3715	LC7-23	Fo	55.826	42.213	0.0050	98.044
3715	LC7-24	Fo	52.476	43.403	0.0053	95.884
3715	LC7-25	Fo	53.554	44.052	0.0042	97.610
3715	LC7-26	Fo	58.172	44.417	0.0044	102.593
3715	LC7-27	Fo	60.290	45.462	0.0079	105.760
3715	LC7-28	Fo	60.268	45.465	0.0046	105.738
3715	LC7-29	En	40.184	60.351	0.0350	100.571
3715	LC7-30	q	38.796	62.798	0.1160	101.713
3715	LC7-31	q	39.831	58.967	0.1375	98.939
3715	LC7-32	q	40.993	60.788	0.1182	101.902
3715	LC7-33	q	41.080	61.168	0.1198	102.371
3721	LC10-1	En	40.157	57.522	0.0269	97.707

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3721	LC10-2	En	39.931	57.678	0.0271	97.637
3721	LC10-3	En	39.741	57.974	0.0273	97.743
3721	LC10-4	En	39.585	58.072	0.0239	97.681
3721	LC10-5	En	39.429	57.927	0.0251	97.382
3721	LC10-6	En	39.421	57.647	0.0260	97.095
3721	LC10-7	En	39.514	57.467	0.0268	97.008
3721	LC10-8	En	39.628	57.197	0.0297	96.855
3721	LC10-9	En	39.238	58.613	0.0281	97.880
3721	LC10-10	En	39.288	58.631	0.0274	97.947
3721	LC10-11	En	39.327	58.820	0.0281	98.176
3721	LC10-12	En	39.168	58.738	0.0283	97.935
3721	LC10-13	En	39.346	59.063	0.0254	98.435
3721	LC10-14	En	39.388	58.684	0.0273	98.100
3721	LC10-15	En	39.356	58.613	0.0284	97.998
3721	LC10-16	En	39.337	58.520	0.0275	97.885
3721	LC10-17	En	39.249	58.495	0.0273	97.772
3721	LC10-18	En	38.945	58.316	0.0356	97.297
3721	LC10-19	En	39.206	58.555	0.0262	97.788
3721	LC10-20	En	39.328	58.726	0.0262	98.081
3721	LC10-21	En	39.396	58.761	0.0261	98.184
3721	LC10-22	En	39.395	58.741	0.0269	98.164
3721	LC10-23	En	39.315	59.485	0.0362	98.837
3721	LC10-24	En	39.109	59.105	0.0356	98.250
3721	LC10-25	En	39.105	58.942	0.0324	98.080
3721	LC10-26	En	39.649	59.868	0.0276	99.545
3721	LC10-27	En	39.596	59.668	0.0265	99.291
3721	LC10-28	En	39.527	59.638	0.0264	99.192
3721	LC10-29	En	39.379	59.545	0.0273	98.952
3721	LC10-30	En	39.217	59.403	0.0274	98.648
3721	LC10-31	En	38.663	58.291	0.0278	96.982
3721	LC10-32	En	38.661	58.870	0.0325	97.564
3721	LC10-33	En	38.472	58.701	0.0263	97.200
3721	LC10-34	En	38.950	59.282	0.0259	98.259
3721	LC10-35	En	39.192	59.568	0.0241	98.785
3721	LC10-36	En	39.353	59.755	0.0233	99.132
3721	LC10-37	En	39.463	59.938	0.0236	99.425
3721	LC10-38	En	39.577	60.023	0.0242	99.625
3721	LC10-39	Fo	57.546	41.739	0.0033	99.288
3721	LC10-40	Fo	57.134	41.753	0.0038	98.891
3721	LC10-41	Fo	56.738	41.576	0.0028	98.317
3721	LC10-42	Fo	56.139	41.412	0.0027	97.554
3721	LC10-43	Fo	56.049	41.275	0.0029	97.327
3721	LC10-44	Fo	55.576	41.019	0.0032	96.598
3721	LC10-45	Fo	55.091	40.788	0.0033	95.882
3721	LC10-46	Fo	55.132	41.157	0.0020	96.291
3721	LC10-47	Fo	55.093	41.285	0.0031	96.381

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3721	LC10-48	Fo	55.410	41.596	0.0034	97.010
3721	LC10-49	Fo	56.029	41.820	0.0040	97.853
3721	LC10-50	Fo	56.533	42.099	0.0035	98.636
3721	LC10-51	Fo	56.699	42.113	0.0036	98.816
3721	LC10-52	Fo	57.322	42.401	0.0057	99.729
3721	LC10-53	Fo	57.463	42.699	0.0047	100.167
3721	LC10-54	Fo	57.676	42.861	0.0035	100.541
3721	LC10-55	Fo	57.199	42.762	0.0047	99.966
3721	LC10-56	Fo	56.731	42.572	0.0029	99.306
3721	LC10-57	Fo	56.811	42.915	0.0029	99.729
3721	LC10-58	Fo	56.748	42.732	0.0035	99.484
3721	LC10-59	Fo	56.548	42.645	0.0029	99.196
3721	LC10-60	Fo	56.446	42.471	0.0036	98.921
3721	LC10-61	Fo	56.357	42.297	0.0061	98.660
3721	LC10-62	Fo	56.174	42.228	0.0029	98.405
3721	LC10-63	Fo	57.367	42.716	0.0037	100.087
3721	LC10-64	Fo	57.935	43.350	0.0041	101.289
3721	LC10-65	Fo	57.156	43.066	0.0084	100.231
3721	LC10-66	Fo	57.095	42.647	0.0040	99.746
3721	LC10-67	Fo	56.994	42.718	0.0052	99.717
3721	LC10-68	Fo	56.816	42.713	0.0037	99.533
3721	LC10-69	Fo	56.299	42.569	0.0037	98.872
3721	LC10-70	Fo	56.122	42.602	0.0036	98.728
3721	LC10-71	Fo	55.751	42.457	0.0037	98.212
3721	LC10-72	Fo	56.553	42.553	0.0033	99.109
3721	LC10-73	Fo	56.876	42.544	0.0027	99.423
3721	LC10-74	Fo	57.246	42.639	0.0031	99.888
3721	LC10-75	Fo	57.496	42.752	0.0032	100.251
3721	LC10-76	Fo	57.088	43.135	0.0028	100.226
3721	LC10-77	Fo	56.862	42.868	0.0040	99.734
3721	LC10-78	Fo	56.313	42.904	0.0028	99.220
3721	LC10-79	Fo	55.928	42.341	0.0167	98.286
3721	LC10-80	Fo	55.787	42.168	0.0043	97.959
3721	LC10-81	Fo	55.465	42.134	0.0042	97.603
3721	LC10-82	Fo	57.244	42.982	0.0028	100.229
3721	LC10-83	Fo	56.782	43.376	0.0038	100.162
3721	LC10-84	Fo	56.318	43.561	0.0038	99.883
3721	LC10-85	Fo	56.139	42.033	0.0027	98.175
3721	LC10-86	Fo	55.599	41.861	0.0029	97.463
3721	LC10-87	Fo	55.628	41.810	0.0039	97.442
3721	LC10-88	Fo	55.133	42.094	0.0036	97.231
3722	LC11-1	Fo	58.039	42.151	0.0034	100.193
3722	LC11-2	Fo	57.506	42.019	0.0022	99.527
3722	LC11-3	En	40.092	58.928	0.0235	99.044
3722	LC11-4	Fo	56.752	42.085	0.0037	98.841
3722	LC11-5	Fo	56.176	42.106	0.0054	98.287

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3722	LC11-6	En	40.186	60.477	0.0251	100.689
3722	LC11-7	En	40.001	60.706	0.0250	100.733
3722	LC11-8	Fo	57.018	42.984	0.0050	100.007
3722	LC11-9	En	40.028	60.449	0.0268	100.504
3722	LC11-10	En	39.793	60.394	0.0255	100.213
3722	LC11-11	Fo	57.286	42.835	0.0050	100.126
3722	LC11-12	Fo	57.294	42.906	0.0044	100.204
3722	LC11-13	Fo	57.169	43.156	0.0049	100.330
3722	LC11-14	En	39.742	60.384	0.0263	100.153
3722	LC11-15	En	39.651	59.698	0.0262	99.376
3722	LC11-16	En	39.420	60.155	0.0264	99.602
3722	LC11-17	En	39.527	59.794	0.0279	99.350
3722	LC11-18	En	40.128	59.957	0.0270	100.113
3722	LC11-19	En	40.173	59.624	0.0372	99.835
3722	LC11-20	En	39.894	59.639	0.0239	99.558
3722	LC11-21	En	38.893	59.721	0.0449	98.660
3722	LC11-22	En	39.940	59.656	0.0254	99.622
3722	LC11-23	En	39.780	59.671	0.0285	99.480
3722	LC11-24	Fo	57.922	42.161	0.0054	100.089
3722	LC11-25	En	40.006	59.255	0.0237	99.285
3722	LC11-26	En	40.383	59.280	0.0234	99.687
3722	LC11-27	Fo	56.961	41.782	0.0036	98.747
3722	LC11-28	Fo	56.959	42.158	0.0028	99.120
3722	LC11-29	Fo	56.465	42.424	0.0016	98.891
3722	LC11-30	Fo	56.899	42.479	0.0030	99.381
3722	LC11-31	Fo	57.715	42.021	0.0033	99.739
3722	LC11-32	En	39.337	60.170	0.0286	99.536
3722	LC11-33	Fo	56.833	42.059	0.0027	98.895
3722	LC11-34	Fo	56.583	42.241	0.0033	98.827
3722	LC11-35	En	40.112	59.669	0.0271	99.809
3722	LC11-36	En	40.032	59.293	0.0262	99.352
3722	LC11-37	En	39.976	59.339	0.0266	99.342
3722	LC11-38	En	40.521	60.086	0.0258	100.633
3722	LC11-39	En	39.409	59.243	0.0283	98.681
3722	LC11-40	Fo	56.308	42.326	0.0043	98.638
3722	LC11-41	Fo	56.439	42.442	0.0070	98.888
3722	LC11-42	Fo	56.426	42.271	0.0029	98.700
3722	LC11-43	En	40.356	59.331	0.0265	99.714
3722	LC11-44	En	39.368	58.728	0.0266	98.123
3722	LC11-45	Fo	56.281	42.228	0.0029	98.512
3722	LC11-46	Fo	56.515	42.309	0.0025	98.827
3722	LC11-47	Fo	56.520	42.164	0.0028	98.687
3722	LC11-48	Fo	55.958	41.938	0.0032	97.899
3722	LC11-49	En	39.372	59.407	0.0251	98.805
3722	LC11-50	Fo	55.741	41.231	0.0059	96.978
3722	LC11-51	Fo	55.791	41.704	0.0064	97.502

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3722	LC11-52	En	39.403	59.516	0.0269	98.947
3722	LC11-53	Fo	57.234	41.305	0.0068	98.546
3722	LC11-54	Fo	56.692	41.228	0.0065	97.927
3722	LC11-55	En	40.374	58.488	0.0256	98.888
3727	LC12-1	Fo	56.406	38.017	0.0048	94.428
3727	LC12-2	Fo	54.903	38.382	0.0030	93.288
3727	LC12-3	Fo	56.155	39.387	0.0041	95.546
3727	LC12-4	Fo	55.753	39.397	0.0026	95.153
3727	LC12-5	Fo	55.003	39.310	0.0028	94.316
3727	LC12-6	Fo	55.068	39.776	0.0034	94.847
3727	LC12-7	Fo	54.506	39.107	0.0020	93.615
3727	LC12-8	Fo	54.720	39.578	0.0031	94.301
3727	LC12-9	Fo	54.676	39.852	0.0031	94.531
3727	LC12-10	Fo	53.879	39.894	0.0034	93.776
3727	LC12-11	Fo	53.370	40.241	0.0032	93.614
3727	LC12-12	Fo	52.305	39.356	0.0032	91.664
3727	LC12-13	Fo	51.496	37.980	0.0031	89.479
3727	LC12-14	Fo	52.330	38.401	0.0024	90.733
3727	LC12-15	Fo	59.577	40.494	0.0027	100.074
3727	LC12-16	Fo	59.214	40.635	0.0029	99.852
3727	LC12-17	Fo	58.626	40.721	0.0034	99.350
3727	LC12-18	Fo	58.064	40.856	0.0036	98.924
3727	LC12-19	Fo	57.672	41.022	0.0014	98.695
3727	LC12-20	Fo	57.881	41.380	0.0014	99.262
3727	LC12-21	Fo	57.260	41.484	0.0024	98.746
3727	LC12-22	Fo	57.319	41.641	0.0021	98.962
3727	LC12-23	Fo	57.086	41.979	0.0017	99.067
3727	LC12-24	Fo	57.125	42.149	0.0024	99.276
3727	LC12-25	Fo	56.888	42.151	0.0036	99.043
3727	LC12-26	Fo	56.626	42.240	0.0018	98.868
3727	LC12-27	Fo	56.492	42.421	0.0017	98.915
3727	LC12-28	Fo	56.458	42.848	0.0011	99.307
3727	LC12-29	Fo	54.880	41.578	0.0035	96.462
3727	LC12-30	Fo	54.615	41.134	0.0024	95.751
3727	LC12-31	Fo	54.313	40.851	0.0025	95.167
3727	LC12-32	Fo	54.087	40.568	0.0033	94.658
3727	LC12-33	Fo	53.866	40.358	0.0022	94.226
3727	LC12-34	Fo	53.813	40.163	0.0032	93.979
3727	LC12-35	Fo	52.776	39.978	0.0029	92.757
3727	LC12-36	Fo	53.136	40.370	0.0035	93.510
3727	LC12-37	Fo	54.251	40.860	0.0026	95.114
3727	LC12-38	Fo	54.285	41.062	0.0020	95.349
3727	LC12-39	Fo	54.068	41.425	0.0023	95.495
3727	LC12-40	Fo	53.774	41.613	0.0031	95.390
3727	LC12-41	Fo	53.412	41.861	0.0022	95.275
3727	LC12-42	Fo	54.423	42.150	0.0027	96.576

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3727	LC12-43	Fo	54.712	42.565	0.0022	97.279
3727	LC12-44	Fo	52.321	43.219	0.0027	95.543
3727	LC12-45	Fo	51.810	43.580	0.0025	95.393
3727	LC12-46	Fo	52.445	44.083	0.0030	96.531
3727	LC12-47	Fo	49.880	44.260	0.0033	94.143
3727	LC12-48	Fo	48.873	43.275	0.0050	92.153
3727	LC12-49	Fo	48.814	42.841	0.0031	91.658
3727	LC12-50	Fo	48.986	42.049	0.0023	91.037
3727	LC12-51	Fo	47.681	41.122	0.0033	88.806
3727	LC12-52	Fo	44.727	39.094	0.0028	83.824
3727	LC12-53	Fo	44.607	38.985	0.0027	83.595
3727	LC12-54	Fo	44.862	38.087	0.0036	82.953
3727	LC12-55	Fo	50.542	41.357	0.0033	91.902
3727	LC12-56	Fo	52.329	41.121	0.0026	93.453
3727	LC12-57	q	38.335	60.043	0.0419	98.421
3727	LC12-58	q	36.385	61.785	0.0433	98.214
3727	LC12-59	q	37.214	62.076	0.0496	99.341
3728	LC13-1	Fo	59.947	41.131	0.0110	101.089
3728	LC13-2	Fo	60.225	41.248	0.0102	101.483
3728	LC13-3	Fo	58.992	41.441	0.0096	100.443
3728	LC13-4	Fo	57.048	41.268	0.0097	98.326
3728	LC13-5	Fo	57.461	41.672	0.0098	99.143
3728	LC13-6	Fo	56.183	40.217	0.0057	96.406
3728	LC13-7	Fo	56.760	40.323	0.0056	97.089
3728	LC13-8	Fo	57.131	40.459	0.0054	97.595
3728	LC13-9	Fo	57.017	40.622	0.0055	97.645
3728	LC13-10	Fo	57.028	40.746	0.0052	97.779
3728	LC13-11	Fo	57.476	40.768	0.0054	98.250
3728	LC13-12	Fo	58.192	40.794	0.0049	98.991
3728	LC13-13	Fo	57.197	41.166	0.0053	98.368
3728	LC13-14	Fo	56.653	40.588	0.0047	97.246
3728	LC13-15	Fo	56.683	40.915	0.0053	97.603
3728	LC13-16	Fo	56.776	41.079	0.0050	97.860
3728	LC13-17	Fo	56.773	41.116	0.0046	97.894
3728	LC13-18	Fo	57.153	41.295	0.0044	98.453
3728	LC13-19	Fo	56.519	41.700	0.0048	98.224
3728	LC13-20	Fo	56.654	41.893	0.0048	98.552
3728	LC13-21	Fo	56.895	41.979	0.0051	98.879
3728	LC13-22	Fo	57.531	41.992	0.0047	99.528
3728	LC13-23	Fo	57.591	42.235	0.0049	99.831
3728	LC13-24	Fo	57.252	42.394	0.0056	99.652
3728	LC13-25	Fo	57.300	42.600	0.0044	99.904
3728	LC13-26	Fo	57.842	42.181	0.0044	100.027
3728	LC13-27	Fo	57.922	42.125	0.0046	100.052
3728	LC13-28	Fo	57.540	42.492	0.0058	100.038
3728	LC13-29	Fo	57.232	41.297	0.0052	98.534

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3728	LC13-30	Fo	57.173	41.929	0.0052	99.107
3728	LC13-31	Fo	57.767	41.377	0.0061	99.150
3728	LC13-32	Fo	57.892	41.684	0.0052	99.581
3728	LC13-33	Fo	58.272	41.355	0.0045	99.632
3728	LC13-34	En	41.027	56.190	0.0297	97.247
3728	LC13-35	En	41.047	56.044	0.0300	97.122
3728	LC13-36	En	40.836	55.528	0.0308	96.396
3728	LC13-37	En	40.838	55.656	0.0297	96.524
3728	LC13-38	En	40.776	55.407	0.0306	96.214
3728	LC13-39	En	40.874	55.527	0.0300	96.432
3728	LC13-40	En	41.054	57.350	0.0281	98.433
3728	LC13-41	En	41.768	57.333	0.0307	99.132
3728	LC13-42	En	42.015	56.989	0.0315	99.036
3728	LC13-43	En	41.230	57.193	0.0295	98.453
3728	LC13-44	En	41.227	57.008	0.0296	98.265
3737	LC17-1	Fo	56.529	42.277	0.0039	98.810
3737	LC17-2	Fo	55.962	42.233	0.0043	98.199
3737	LC17-3	Fo	56.331	42.178	0.0030	98.512
3737	LC17-4	Fo	55.841	42.128	0.0041	97.973
3737	LC17-5	Fo	55.718	42.085	0.0034	97.806
3737	LC17-6	Fo	55.674	42.159	0.0029	97.836
3737	LC17-7	Fo	55.155	41.648	0.0034	96.806
3737	LC17-8	Fo	55.495	41.580	0.0031	97.078
3737	LC17-9	Fo	55.631	41.616	0.0030	97.250
3737	LC17-10	Fo	55.660	41.470	0.0040	97.134
3737	LC17-11	Fo	55.752	41.579	0.0035	97.335
3739	LC18-2	Fo	57.202	41.876	0.0064	99.085
3739	LC18-3	Fo	57.382	42.069	0.0069	99.458
3739	LC18-4	Fo	57.225	42.527	0.0053	99.757
3739	LC18-6	Fo	56.002	41.620	0.0074	97.630
3739	LC18-7	Fo	56.046	41.976	0.0051	98.027
3739	LC18-8	Fo	56.527	41.386	0.0046	97.918
3739	LC18-9	Fo	56.648	41.739	0.0042	98.391
3739	LC18-10	Fo	54.942	41.733	0.0047	96.680
3739	LC18-12	Fo	55.324	42.019	0.0057	97.349
3739	LC18-13	Fo	55.214	42.127	0.0048	97.346
3739	LC18-14	Fo	54.731	42.086	0.0050	96.822
3739	LC18-15	Fo	55.754	42.922	0.0047	98.681
3739	LC18-16	Fo	56.593	42.648	0.0053	99.246
3739	LC18-17	Fo	57.103	42.344	0.0055	99.453
3739	LC18-18	Fo	55.846	42.098	0.0047	97.949
3739	LC18-19	Fo	57.112	42.206	0.0059	99.324
3739	LC18-22	Fo	56.058	42.942	0.0054	99.006
3740	LC19-5	Fo	55.856	42.653	0.0072	98.516
3740	LC19-6	Fo	56.193	42.567	0.0060	98.766
3740	LC19-7	Fo	56.601	42.666	0.0054	99.273

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3740	LC19-8	Fo	56.919	42.535	0.0066	99.461
3740	LC19-9	Fo	56.956	42.495	0.0060	99.457
3740	LC19-10	Fo	56.370	42.423	0.0055	98.799
3740	LC19-11	Fo	56.322	42.578	0.0059	98.906
3740	LC19-12	Fo	55.733	42.545	0.0061	98.284
3740	LC19-13	Fo	57.139	42.745	0.0044	99.889
3740	LC19-14	Fo	57.145	42.476	0.0078	99.629
3740	LC19-19	Fo	56.143	42.922	0.0059	99.071
3740	LC19-20	Fo	56.224	42.702	0.0049	98.931
3740	LC19-21	Fo	56.907	42.876	0.0045	99.788
3740	LC19-22	Fo	57.118	42.622	0.0046	99.745
3740	LC19-23	Fo	57.367	42.380	0.0055	99.753
3744	LC21-1	Fo	58.771	41.917	0.0042	100.692
3744	LC21-2	Fo	58.670	42.200	0.0023	100.872
3744	LC21-3	Fo	58.402	42.299	0.0031	100.704
3744	LC21-4	Fo	58.128	42.460	0.0031	100.591
3744	LC21-5	Fo	58.147	42.620	0.0037	100.771
3744	LC21-9	Fo	56.095	41.358	0.0045	97.458
3744	LC21-10	Fo	55.881	41.478	0.0043	97.363
3744	LC21-11	Fo	55.881	41.515	0.0035	97.400
3744	LC21-13	Fo	59.095	42.283	0.0059	101.384
3744	LC21-14	Fo	58.699	42.494	0.0074	101.201
3744	LC21-15	Fo	58.841	42.943	0.0062	101.790
3744	LC21-16	Fo	56.313	41.340	0.0060	97.659
3744	LC21-17	Fo	56.867	41.666	0.0078	98.541
3744	LC21-22	Fo	55.440	41.217	0.0042	96.661
3744	LC21-23	Fo	55.052	41.470	0.0039	96.526
3744	LC21-24	Fo	57.564	42.649	0.0024	100.215
3744	LC21-25	Fo	57.711	43.136	0.0030	100.850
3744	LC21-26	Fo	56.611	42.709	0.0034	99.323
3744	LC21-27	Fo	55.810	42.498	0.0030	98.311
3744	LC21-28	Fo	55.445	42.564	0.0035	98.013
3744	LC21-29	Fo	55.537	42.826	0.0030	98.366
3744	LC21-30	Fo	57.303	41.897	0.0036	99.204
3744	LC21-31	Fo	57.007	42.539	0.0038	99.550
3744	LC21-32	Fo	56.689	42.345	0.0042	99.038
3744	LC21-33	Fo	56.613	42.634	0.0054	99.253
3744	LC21-34	Fo	56.115	42.606	0.0050	98.726
3748	LC24-4	Fo	58.128	41.442	0.0050	99.575
3748	LC24-5	En	40.209	60.308	0.0355	100.553
3748	LC24-6	En	39.324	59.865	0.0378	99.228
3748	LC24-7	En	39.489	59.427	0.0378	98.955
3748	LC24-8	En	39.612	59.236	0.0387	98.888
3748	LC24-9	En	39.971	59.076	0.0384	99.086
3748	LC24-10	En	40.460	59.347	0.0332	99.841
3748	LC24-11	En	40.241	59.718	0.0326	99.992



Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3748	LC24-12	En	39.989	60.148	0.0328	100.171
3748	LC24-13	En	39.779	60.339	0.0320	100.151
3748	LC24-14	En	38.855	59.326	0.0424	98.224
3748	LC24-15	En	39.847	59.408	0.0436	99.300
3748	LC24-16	En	39.385	59.966	0.0335	99.385
3748	LC24-17	En	39.483	59.541	0.0364	99.061
3748	LC24-18	En	39.787	58.636	0.0353	98.459
3748	LC24-19	Fo	56.803	41.560	0.0040	98.367
3748	LC24-22	Fo	55.988	42.487	0.0091	98.484
3748	LC24-23	Fo	56.239	42.121	0.0108	98.371
3748	LC24-24	Fo	56.334	41.793	0.0044	98.131
3748	LC24-25	Fo	57.989	41.360	0.0042	99.353
3748	LC24-26	Fo	57.742	41.857	0.0038	99.603
3748	LC24-27	Fo	56.999	42.403	0.0058	99.408
3748	LC24-28	Fo	55.791	42.006	0.0031	97.800
3748	LC24-29	Fo	56.389	41.297	0.0046	97.691
3748	LC24-32	Fo	57.950	42.060	0.0039	100.014
3748	LC24-33	Fo	57.092	42.522	0.0049	99.619
3748	LC24-34	Fo	56.920	42.544	0.0042	99.468
3750	LC26-9	Fo	56.674	43.638	0.0044	100.317
3750	LC26-10	Fo	55.839	43.563	0.0049	99.407
3750	LC26-11	Fo	55.391	42.906	0.0054	98.303
3750	LC26-12	Fo	55.811	42.460	0.0050	98.276
3750	LC26-13	Fo	55.961	42.645	0.0043	98.610
3750	LC26-14	Fo	56.823	42.111	0.0048	98.939
3750	LC26-15	Fo	57.335	42.079	0.0063	99.420
3750	LC26-16	Fo	56.536	43.752	0.0081	100.296
3750	LC26-17	Fo	55.792	43.172	0.0066	98.971
3750	LC26-18	Fo	55.571	42.592	0.0069	98.170
3750	LC26-19	Fo	56.025	42.020	0.0068	98.052
3750	LC26-20	Fo	56.898	42.521	0.0056	99.425
3750	LC26-21	Fo	57.545	42.559	0.0086	100.113
3750	LC26-22	Fo	58.146	42.694	0.0069	100.847
3750	LC26-23	Fo	57.006	41.649	0.0109	98.666
3750	LC26-24	Fo	56.165	41.724	0.0085	97.898
3750	LC26-25	Fo	57.093	42.595	0.0064	99.695
3750	LC26-26	Fo	57.163	42.109	0.0057	99.278
3750	LC26-27	Fo	57.381	42.523	0.0060	99.910
3750	LC26-28	q	39.695	59.044	0.0396	98.780
3750	LC26-29	q	39.526	59.262	0.0412	98.830
3750	LC26-30	q	55.737	42.034	0.0527	97.825
3750	LC26-31	q	40.269	59.437	0.0265	99.733
3750	LC26-32	q	40.142	59.366	0.0343	99.543
3750	LC26-33	q	39.672	59.316	0.0416	99.031
3750	LC26-34	q	40.774	59.137	0.0389	99.951
3753	LC29-1	Fo	57.525	42.504	0.0038	100.033

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3753	LC29-2	Fo	56.338	42.772	0.0037	99.114
3753	LC29-3	Fo	56.302	43.231	0.0030	99.536
3753	LC29-4	Fo	55.968	42.315	0.0029	98.286
3753	LC29-5	Fo	55.937	42.410	0.0034	98.350
3753	LC29-6	Fo	56.569	42.502	0.0033	99.074
3753	LC29-7	Fo	57.064	41.955	0.0045	99.024
3753	LC29-8	Fo	56.668	41.779	0.0041	98.451
3753	LC29-9	Fo	56.194	42.383	0.0040	98.581
3753	LC29-10	Fo	55.622	42.211	0.0034	97.836
3753	LC29-11	Fo	56.370	41.439	0.0030	97.812
3753	LC29-12	Fo	55.648	42.454	0.0025	98.105
3753	LC29-13	Fo	56.879	42.786	0.0035	99.669
3753	LC29-14	Fo	56.803	42.985	0.0040	99.792
3753	LC29-15	Fo	56.707	42.591	0.0045	99.303
3753	LC29-16	Fo	56.995	43.059	0.0053	100.059
3753	LC29-17	Fo	57.317	42.218	0.0056	99.541
3753	LC29-18	Fo	56.438	42.760	0.0066	99.205
3753	LC29-19	Fo	57.214	42.512	0.0067	99.733
3753	LC29-20	En	39.690	59.489	0.0307	99.210
3753	LC29-21	En	40.059	59.986	0.0330	100.079
3753	LC29-22	En	40.279	59.897	0.0309	100.208
3753	LC29-23	En	40.287	60.200	0.0288	100.517
3753	LC29-24	En	39.880	60.001	0.0259	99.908
3753	LC29-25	En	39.707	60.315	0.0287	100.051
3753	LC29-26	En	39.494	59.816	0.0309	99.342
3753	LC29-27	En	39.096	60.417	0.0309	99.545
3753	LC29-28	En	39.501	60.203	0.0332	99.738
3754	LC30-1	En	39.739	59.948	0.0374	99.725
3754	LC30-2	En	39.341	60.085	0.0358	99.463
3754	LC30-3	En	39.007	59.503	0.0340	98.545
3754	LC30-4	En	39.338	59.754	0.0331	99.126
3754	LC30-5	En	39.096	59.964	0.0341	99.095
3754	LC30-6	En	39.559	60.350	0.0327	99.943
3754	LC30-7	En	39.415	60.567	0.0422	100.025
3754	LC30-8	En	39.916	60.291	0.0356	100.243
3754	LC30-9	En	39.628	60.693	0.0359	100.358
3754	LC30-10	Fo	57.349	42.560	0.0048	99.914
3754	LC30-11	Fo	57.254	43.174	0.0039	100.432
3754	LC30-12	Fo	56.477	43.403	0.0036	99.884
3754	LC30-13	Fo	56.382	43.087	0.0026	99.472
3754	LC30-14	Fo	56.075	43.479	0.0045	99.559
3754	LC30-15	Fo	56.444	42.835	0.0042	99.283
3758	LC31-1	Fo	56.813	41.706	0.0039	98.523
3758	LC31-2	Fo	56.518	41.925	0.0043	98.447
3758	LC31-3	Fo	56.711	42.116	0.0040	98.831
3758	LC31-4	Fo	57.430	41.845	0.0056	99.281

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3758	LC31-5	Fo	57.531	41.918	0.0062	99.455
3758	LC31-6	Fo	57.280	41.944	0.0058	99.230
3758	LC31-7	Fo	56.962	42.056	0.0052	99.023
3758	LC31-8	Fo	57.051	42.097	0.0058	99.154
3758	LC31-9	Fo	57.397	41.862	0.0044	99.264
3758	LC31-10	Fo	58.479	42.379	0.0058	100.864
3758	LC31-11	Fo	57.008	41.953	0.0062	98.967
3758	LC31-12	Fo	58.045	41.886	0.0050	99.936
3758	LC31-13	Fo	58.311	41.673	0.0056	99.990
3758	LC31-14	Fo	56.863	42.318	0.0066	99.188
3758	LC31-15	Fo	57.654	42.329	0.0070	99.990
3758	LC31-16	Fo	58.173	42.774	0.0061	100.953
3755	LC32-1	En	39.711	61.251	0.0189	100.981
3755	LC32-2	En	40.007	60.250	0.0184	100.276
3755	LC32-3	En	40.423	60.088	0.0184	100.530
3755	LC32-4	En	40.058	60.234	0.0171	100.310
3755	LC32-5	En	38.924	59.361	0.0215	98.307
3755	LC32-6	En	39.095	60.687	0.0174	99.800
3755	LC32-7	En	38.715	59.832	0.0181	98.566
3755	LC32-8	En	39.326	60.318	0.0260	99.671
3755	LC32-9	En	39.924	59.823	0.0387	99.787
3755	LC32-10	Fo	56.302	42.364	0.0026	98.669
3755	LC32-11	Fo	56.158	42.656	0.0027	98.817
3755	LC32-12	Fo	56.206	42.728	0.0025	98.937
3755	LC32-13	Fo	56.393	42.745	0.0039	99.142
3755	LC32-14	Fo	55.858	42.687	0.0048	98.550
3755	LC32-15	Fo	56.578	43.104	0.0054	99.687
3755	LC32-16	Fo	55.708	43.332	0.0024	99.042
3755	LC32-17	En	38.896	60.076	0.0181	98.991
3755	LC32-18	En	38.854	59.609	0.0444	98.508
3755	LC32-19	En	39.682	59.257	0.0194	98.959
3759	LC35-1	Fo	58.413	41.838	0.0070	100.258
3759	LC35-2	Fo	58.657	41.744	0.0076	100.409
3759	LC35-3	Fo	58.895	41.826	0.0069	100.728
3759	LC35-4	Fo	58.146	42.493	0.0074	100.646
3759	LC35-5	Fo	57.825	42.356	0.0078	100.189
3759	LC35-6	Fo	57.314	41.971	0.0069	99.292
3759	LC35-7	Fo	57.615	42.636	0.0063	100.257
3759	LC35-8	Fo	57.507	42.349	0.0056	99.862
3759	LC35-9	Fo	56.838	41.961	0.0074	98.807
3759	LC35-10	Fo	57.267	42.082	0.0118	99.361
3759	LC35-11	En	40.523	60.136	0.0299	100.690
3759	LC35-12	En	40.485	59.711	0.0289	100.226
3759	LC35-13	En	40.623	59.836	0.0311	100.491
3759	LC35-14	En	40.374	59.935	0.0274	100.337
3759	LC35-15	En	40.323	60.110	0.0254	100.459

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3759	LC35-16	Fo	56.466	42.821	0.0063	99.293
3759	LC35-17	En	40.261	60.201	0.0269	100.490
3759	LC35-18	En	40.274	59.966	0.0273	100.268
3759	LC35-19	En	40.483	59.875	0.0321	100.391
3759	LC35-20	En	40.615	59.411	0.0317	100.058
3759	LC35-21	En	40.312	59.080	0.0317	99.425
3759	LC35-22	En	40.106	58.275	0.0302	98.412
3759	LC35-23	En	39.583	57.682	0.0325	97.298
3759	LC35-24	Fo	56.285	41.194	0.0090	97.488
3759	LC35-25	En	39.531	58.842	0.0324	98.406
3759	LC35-26	En	39.308	59.199	0.0260	98.534
3759	LC35-27	Fo	56.072	41.219	0.0050	97.296
3759	LC35-28	Fo	55.301	41.812	0.0063	97.119
3759	LC35-29	Fo	56.067	41.806	0.0051	97.878
3759	LC35-30	Fo	56.041	42.177	0.0089	98.227
3759	LC35-31	Fo	55.890	42.227	0.0114	98.129
3759	LC35-32	Fo	56.409	41.949	0.0113	98.370
3759	LC35-33	Fo	56.207	42.142	0.0130	98.362
3759	LC35-34	Fo	54.316	42.137	0.0178	96.471
3759	LC35-35	En	39.428	59.358	0.0417	98.829
3763	LC36-1	Fo	55.054	41.784	0.0137	96.852
3763	LC36-2	Fo	56.490	42.171	0.0118	98.673
3763	LC36-3	Fo	56.705	42.141	0.0127	98.859
3763	LC36-4	Fo	56.857	42.423	0.0139	99.294
3763	LC36-5	Fo	57.228	42.397	0.0125	99.638
3763	LC36-6	Fo	57.904	42.263	0.0101	100.177
3763	LC36-7	Fo	56.945	42.801	0.0118	99.758
3763	LC36-8	Fo	56.984	42.639	0.0110	99.634
3763	LC36-9	Fo	57.456	42.547	0.0093	100.012
3763	LC36-10	Fo	57.357	42.268	0.0095	99.635
3763	LC36-11	Fo	57.616	42.703	0.0103	100.330
3763	LC36-12	Fo	56.068	41.953	0.0155	98.037
3763	LC36-13	Fo	58.867	42.797	0.0072	101.671
3763	LC36-14	Fo	57.866	42.361	0.0206	100.248
3763	LC36-15	Fo	56.989	41.808	0.0112	98.808
3763	LC36-16	Fo	56.900	41.994	0.0030	98.897
3763	LC36-17	Fo	57.975	42.336	0.0020	100.313
3762	LC37-1	En	40.178	58.729	0.0286	98.936
3762	LC37-2	En	40.624	58.968	0.0217	99.614
3762	LC37-3	En	40.945	58.988	0.0227	99.956
3762	LC37-4	En	41.244	57.855	0.0232	99.123
3762	LC37-5	En	40.788	58.466	0.0192	99.274
3762	LC37-6	En	39.963	58.695	0.0230	98.682
3762	LC37-7	En	39.688	58.444	0.0259	98.159
3762	LC37-8	En	39.231	58.505	0.0263	97.763
3762	LC37-9	En	39.514	58.814	0.0237	98.352

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3762	LC37-10	Fo	55.234	41.906	0.0221	97.163
3762	LC37-11	Fo	56.642	42.118	0.0157	98.776
3762	LC37-12	Fo	57.477	42.087	0.0111	99.575
3762	LC37-13	Fo	58.246	42.127	0.0123	100.385
3762	LC37-14	Fo	57.332	42.445	0.0112	99.788
3762	LC37-15	Fo	58.575	42.581	0.0066	101.163
3762	LC37-16	Fo	57.174	42.642	0.0065	99.823
3762	LC37-17	Fo	57.368	42.971	0.0083	100.347
3762	LC37-18	Fo	58.260	42.719	0.0073	100.986
3762	LC37-19	Fo	57.859	42.934	0.0073	100.800
3762	LC37-20	Fo	58.209	42.598	0.0061	100.813
3762	LC37-21	Fo	57.284	42.798	0.0043	100.086
3762	LC37-22	En	40.738	60.132	0.0515	100.923
3762	LC37-23	En	40.644	60.188	0.0514	100.885
3762	LC37-24	En	40.771	60.216	0.0522	101.040
3762	LC37-25	En	41.155	59.709	0.0501	100.915
3762	LC37-26	En	41.209	59.725	0.0485	100.984
3762	LC37-27	En	41.426	59.581	0.0448	101.053
3762	LC37-28	En	41.368	59.649	0.0458	101.064
3762	LC37-29	En	41.439	59.276	0.0426	100.759
3762	LC37-30	En	41.515	59.261	0.0461	100.823
3762	LC37-31	En	41.393	59.109	0.0167	100.519
3762	LC37-32	En	41.620	59.164	0.0147	100.799
3762	LC37-33	q	41.463	59.466	0.0251	100.955
3762	LC37-34	q	39.990	58.741	0.0342	98.766
3762	LC37-35	q	40.832	59.792	0.0491	100.674
3762	LC37-36	q	39.823	58.933	0.0331	98.790
3762	LC37-37	q	39.242	58.741	0.0356	98.019
3762	LC37-38	q	39.461	58.649	0.0245	98.135
3762	LC37-39	q	40.381	58.226	0.0267	98.634
3762	LC37-40	q	39.481	58.609	0.0281	98.119
3762	LC37-41	q	39.933	57.704	0.0300	97.668
3762	LC37-42	q	56.779	40.664	0.0092	97.452
3762	LC37-43	q	40.731	57.785	0.0278	98.544
3765	LC38-1	En	40.899	59.571	0.0199	100.490
3765	LC38-2	En	40.700	58.696	0.0200	99.417
3765	LC38-3	En	40.926	58.335	0.0207	99.282
3765	LC38-4	En	40.546	58.083	0.0220	98.652
3765	LC38-5	En	39.698	58.249	0.0226	97.970
3765	LC38-6	En	40.327	58.790	0.0204	99.138
3765	LC38-7	En	39.815	58.719	0.0217	98.556
3765	LC38-8	Fo	57.741	42.094	0.0110	99.846
3765	LC38-9	En	39.910	58.560	0.0226	98.493
3765	LC38-10	En	40.137	59.352	0.0221	99.512
3765	LC38-11	En	39.976	59.677	0.0261	99.680
3765	LC38-12	En	39.039	58.214	0.0208	97.274

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3765	LC38-13	En	39.292	58.853	0.0229	98.168
3765	LC38-14	En	39.547	59.250	0.0222	98.820
3765	LC38-15	En	39.383	58.851	0.0170	98.251
3765	LC38-16	En	39.407	57.877	0.0278	97.312
3765	LC38-17	En	39.957	59.817	0.0313	99.806
3765	LC38-18	En	39.750	59.975	0.0178	99.743
3765	LC38-19	En	39.152	59.289	0.0177	98.459
3765	LC38-20	Fo	55.254	41.367	0.0036	96.625
3765	LC38-21	Fo	55.831	41.156	0.0057	96.993
3765	LC38-22	Fo	55.365	42.039	0.0231	97.428
3765	LC38-23	Fo	55.323	41.789	0.0187	97.131
3765	LC38-24	Fo	55.467	41.339	0.0081	96.814
3765	LC38-25	Fo	57.095	41.378	0.0054	98.479
3765	LC38-26	Fo	56.532	41.629	0.0050	98.166
3765	LC38-27	Fo	56.426	42.035	0.0076	98.469
3765	LC38-28	Fo	55.374	41.456	0.0029	96.833
3765	LC38-29	Fo	55.718	41.196	0.0048	96.919
3765	LC38-30	Fo	56.122	41.178	0.0150	97.315
3765	LC38-31	Fo	56.740	41.722	0.0137	98.476
3765	LC38-32	Fo	55.367	41.198	0.0139	96.579
3765	LC38-33	Fo	57.157	41.753	0.0129	98.923
3765	LC38-34	q	37.210	61.305	0.0317	98.548
3765	LC38-35	q	38.896	60.545	0.0460	99.488
3765	LC38-36	q	39.052	59.810	0.0295	98.892
3765	LC38-37	q	39.405	59.500	0.0334	98.939
3770	LC39-1	En	38.234	59.395	0.0146	97.644
3770	LC39-2	En	37.813	58.953	0.0150	96.781
3770	LC39-3	En	39.155	58.820	0.0142	97.989
3770	LC39-4	En	39.596	58.913	0.0133	98.523
3770	LC39-5	En	39.603	59.529	0.0131	99.145
3770	LC39-6	En	40.410	59.608	0.0128	100.031
3770	LC39-7	En	39.172	60.069	0.0131	99.254
3770	LC39-8	En	39.360	60.252	0.0138	99.626
3770	LC39-9	En	38.575	59.984	0.0135	98.573
3770	LC39-10	En	38.251	59.247	0.0141	97.512
3770	LC39-11	En	38.398	59.246	0.0136	97.658
3770	LC39-12	En	38.444	59.531	0.0145	97.990
3770	LC39-13	En	38.478	59.538	0.0140	98.030
3770	LC39-14	En	38.528	59.594	0.0157	98.138
3770	LC39-15	En	38.744	59.688	0.0162	98.449
3770	LC39-16	En	39.404	59.736	0.0162	99.157
3770	LC39-17	En	39.833	59.440	0.0164	99.290
3770	LC39-18	Fo	56.596	41.596	0.0042	98.196
3770	LC39-19	Fo	55.413	42.006	0.0051	97.424
3770	LC39-20	Fo	54.836	42.089	0.0036	96.929
3770	LC39-21	Fo	55.058	42.561	0.0030	97.622

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3770	LC39-22	Fo	55.516	42.796	0.0054	98.317
3770	LC39-23	Fo	55.763	42.668	0.0043	98.435
3770	LC39-24	Fo	56.250	42.825	0.0047	99.080
3770	LC39-25	Fo	56.771	42.725	0.0043	99.500
3770	LC39-26	Fo	57.467	42.825	0.0038	100.296
3770	LC39-27	Fo	57.642	42.432	0.0038	100.078
3770	LC39-28	Fo	57.295	43.069	0.0048	100.369
3770	LC39-29	Fo	56.673	43.197	0.0049	99.875
3770	LC39-30	Fo	56.346	43.203	0.0049	99.554
3770	LC39-31	Fo	55.809	43.172	0.0037	98.985
3770	LC39-32	q	40.488	59.741	0.0261	100.256
3770	LC39-33	q	40.023	55.048	0.0927	95.166
3770	LC39-34	q	39.655	59.032	0.0273	98.715
3770	LC39-35	q	40.091	60.536	0.0218	100.649
3770	LC39-36	q	39.985	61.839	0.0367	101.862
3770	LC39-37	q	39.530	60.028	0.0232	99.582
3770	LC39-38	q	39.378	58.750	0.0719	98.202
3770	LC39-39	q	39.581	58.781	0.0209	98.383
3770	LC39-40	q	40.283	61.809	0.0389	102.132
3770	LC39-41	q	39.422	61.361	0.0273	100.811
3770	LC39-42	q	38.457	60.851	0.0202	99.329
3770	LC39-43	q	39.566	60.576	0.0649	100.208
3770	LC39-44	q	39.178	61.108	0.0659	100.354
3770	LC39-45	q	40.106	61.405	0.0190	101.531
3770	LC39-46	q	39.741	61.163	0.0476	100.953
3770	LC39-47	q	39.088	61.286	0.0211	100.396
3770	LC39-48	q	39.725	60.745	0.0206	100.491
3770	LC39-49	q	39.838	60.495	0.0218	100.355
3770	LC39-50	q	40.805	59.468	0.0287	100.302
3770	LC39-51	q	39.976	59.330	0.0341	99.341
3771	LC40-1	Fo	52.987	41.263	0.0031	94.253
3771	LC40-2	Fo	53.841	42.015	0.0023	95.858
3771	LC40-3	Fo	53.707	42.416	0.0026	96.126
3771	LC40-4	Fo	54.036	42.665	0.0027	96.704
3771	LC40-5	Fo	54.469	43.213	0.0030	97.685
3771	LC40-6	Fo	55.827	42.440	0.0022	98.269
3771	LC40-7	q	38.539	60.513	0.0834	99.137
3771	LC40-8	Fo	56.168	42.730	0.0026	98.901
3771	LC40-9	Fo	54.885	42.124	0.0035	97.013
3771	LC40-10	Fo	55.355	42.288	0.0027	97.646
3771	LC40-11	Fo	55.638	42.440	0.0030	98.081
3771	LC40-12	Fo	55.836	42.461	0.0033	98.300
3771	LC40-13	Fo	56.218	42.594	0.0031	98.815
3771	LC40-14	Fo	56.621	42.461	0.0036	99.086
3771	LC40-15	Fo	56.981	42.452	0.0036	99.437
3771	LC40-16	Fo	57.428	42.446	0.0036	99.878

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3771	LC40-17	Fo	57.588	42.709	0.0035	100.301
3771	LC40-18	Fo	57.148	42.973	0.0038	100.125
3771	LC40-19	Fo	56.181	43.209	0.0029	99.393
3771	LC40-20	Fo	55.733	43.165	0.0042	98.902
3771	LC40-21	Fo	55.164	43.147	0.0033	98.314
3771	LC40-22	Fo	54.539	43.595	0.0028	98.137
3771	LC40-23	Fo	55.132	43.882	0.0032	99.017
3771	LC40-24	Fo	55.601	43.992	0.0040	99.597
3771	LC40-25	q	38.592	61.471	0.0601	100.125
3771	LC40-26	q	38.681	61.525	0.0572	100.265
3771	LC40-27	q	38.561	61.262	0.0757	99.901
3771	LC40-28	Fo	55.584	42.323	0.0049	97.912
3771	LC40-29	Fo	55.667	42.840	0.0039	98.511
3771	LC40-30	Fo	55.811	42.919	0.0039	98.734
3771	LC40-31	Fo	55.775	42.999	0.0036	98.778
3771	LC40-32	Fo	55.156	43.121	0.0054	98.283
3771	LC40-33	Fo	56.853	41.893	0.0045	98.751
3771	LC40-34	Fo	56.102	42.439	0.0036	98.545
3771	LC40-35	Fo	56.848	42.647	0.0038	99.499
3771	LC40-36	Fo	56.933	42.207	0.0032	99.143
3771	LC40-37	Fo	57.130	42.279	0.0035	99.413
3771	LC40-38	Fo	56.502	42.586	0.0032	99.091
3771	LC40-39	Fo	56.631	42.686	0.0030	99.320
3771	LC40-40	Fo	56.694	42.836	0.0022	99.532
3771	LC40-41	Fo	56.747	43.005	0.0035	99.756
3771	LC40-42	Fo	56.806	42.888	0.0038	99.698
3771	LC40-43	Fo	56.384	43.041	0.0033	99.428
3771	LC40-44	Fo	56.464	43.076	0.0039	99.544
3771	LC40-45	Fo	56.711	42.985	0.0038	99.700
3771	LC40-46	Fo	56.635	43.159	0.0043	99.798
3771	LC40-47	Fo	56.593	43.155	0.0055	99.754
3771	LC40-48	Fo	56.905	43.148	0.0050	100.058
3771	LC40-49	Fo	56.786	43.230	0.0042	100.020
3771	LC40-50	Fo	56.647	43.335	0.0034	99.985
3771	LC40-51	Fo	56.167	43.528	0.0041	99.699
3771	LC40-52	Fo	57.625	43.416	0.0047	101.046
3771	LC40-53	q	40.209	59.541	0.0892	99.841
3771	LC40-54	q	39.857	59.696	0.0939	99.649
3771	LC40-55	q	40.891	59.689	0.0782	100.660
3771	LC40-56	q	43.776	53.220	0.0612	97.059
3771	LC40-57	q	39.790	59.328	0.0607	99.180
3771	LC40-58	q	40.399	60.736	0.0217	101.157
3771	LC40-59	q	41.346	60.158	0.0309	101.536
3771	LC40-60	q	40.178	60.472	0.0222	100.673
3771	LC40-61	q	41.289	60.328	0.0222	101.640
3771	LC40-62	q	39.495	61.010	0.0727	100.579



Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3772	LC41-1	Fo	58.131	44.784	0.0038	102.919
3772	LC41-2	Fo	57.574	44.600	0.0043	102.178
3772	LC41-3	Fo	57.137	44.741	0.0036	101.882
3772	LC41-4	Fo	56.476	44.717	0.0039	101.197
3772	LC41-5	Fo	55.871	44.495	0.0034	100.369
3772	LC41-6	Fo	55.176	44.384	0.0029	99.563
3772	LC41-7	Fo	54.170	44.358	0.0034	98.531
3772	LC41-8	Fo	53.888	44.273	0.0036	98.165
3772	LC41-9	Fo	59.359	43.615	0.0039	102.978
3772	LC41-10	Fo	59.661	43.364	0.0040	103.029
3772	LC41-11	Fo	59.135	43.201	0.0042	102.340
3772	LC41-12	Fo	58.813	43.077	0.0043	101.894
3772	LC41-13	Fo	60.326	42.929	0.0041	103.259
3772	LC41-14	Fo	60.214	42.842	0.0050	103.061
3772	LC41-15	Fo	60.042	42.715	0.0035	102.761
3772	LC41-16	Fo	60.089	42.637	0.0043	102.730
3772	LC41-17	Fo	60.247	42.467	0.0046	102.719
3772	LC41-18	Fo	60.501	42.102	0.0044	102.607
3772	LC41-19	Fo	58.941	42.111	0.0037	101.056
3772	LC41-20	Fo	59.237	41.468	0.0033	100.708
3772	LC41-21	Fo	58.585	41.799	0.0036	100.388
3772	LC41-22	Fo	57.451	42.156	0.0040	99.611
3772	LC41-23	Fo	56.066	41.883	0.0028	97.952
3772	LC41-24	Fo	56.656	42.466	0.0036	99.126
3772	LC41-25	Fo	56.831	42.531	0.0033	99.365
3772	LC41-26	Fo	56.935	42.521	0.0044	99.460
3772	LC41-27	Fo	57.030	42.469	0.0039	99.503
3772	LC41-28	Fo	56.696	42.449	0.0042	99.149
3772	LC41-29	Fo	56.830	42.557	0.0034	99.390
3772	LC41-30	Fo	57.333	42.781	0.0030	100.117
3772	LC41-31	Fo	56.489	42.844	0.0028	99.336
3772	LC41-32	Fo	56.987	42.992	0.0037	99.983
3772	LC41-33	Fo	56.749	43.160	0.0043	99.913
3772	LC41-34	Fo	52.059	42.210	0.0044	94.273
3772	LC41-35	Fo	51.200	41.812	0.0039	93.016
3772	LC41-36	Fo	56.234	42.610	0.0032	98.847
3772	LC41-37	Fo	56.082	42.208	0.0043	98.294
3772	LC41-38	Fo	56.977	44.071	0.0030	101.051
3772	LC41-39	Fo	56.268	43.735	0.0032	100.006
3772	LC41-40	Fo	56.588	43.691	0.0037	100.283
3772	LC41-41	Fo	56.854	43.532	0.0019	100.388
3772	LC41-42	Fo	57.102	43.167	0.0039	100.273
3772	LC41-43	Fo	57.318	42.999	0.0029	100.320
3772	LC41-44	Fo	56.968	42.827	0.0032	99.798
3772	LC41-45	Fo	57.004	42.327	0.0036	99.335
3772	LC41-46	Fo	56.760	42.996	0.0033	99.759

Run #	Analysis #	Phase	MgO (wt%)	SiO <sub>2</sub> (wt%)	CaO (wt%)	Total (wt%)
3772	LC41-47	Fo	56.965	42.763	0.0031	99.731
3772	LC41-48	Fo	55.983	42.489	0.0031	98.475
3772	LC41-49	Fo	55.395	42.287	0.0035	97.686
3772	LC41-50	Fo	57.093	42.129	0.0041	99.226
3772	LC41-51	Fo	56.725	42.055	0.0045	98.785
3772	LC41-52	Fo	57.277	41.878	0.0030	99.158
3772	LC41-53	Fo	57.497	42.731	0.0033	100.231
3772	LC41-54	Fo	58.002	42.180	0.0027	100.185
3772	LC41-55	Fo	57.757	41.568	0.0041	99.329
3772	LC41-56	Fo	57.337	40.896	0.0040	98.237
3772	LC41-57	Fo	57.257	40.891	0.0046	98.153

Abbreviations: *En* enstatite in primary phase, *Fo* forsterite in primary phase, *q* quenched phases.

Note: CaO (wt%) is calculated from counts reported by microprobe.